

22 **Abstract**

23 Cities in South and Southeast Asia are developing rapidly without routine, up-to-date
24 knowledge of air pollutant precursor emissions. This data deficit can potentially be addressed
25 for nitrogen oxides (NO_x) by deriving city NO_x emissions from satellite observations of
26 nitrogen dioxide (NO_2) sampled under windy conditions. NO_2 plumes of isolated cities are
27 aligned along a consistent wind-rotated direction and a best-fit Gaussian is applied to estimate
28 emissions. This approach currently relies on non-standardized selection of the area to sample
29 around the city centre and Gaussian fits often fail or yield non-physical parameters. Here, we
30 automate this approach by defining many (54) sampling areas that we test with Tropospheric
31 Monitoring Instrument (TROPOMI) NO_2 observations for 2019 over 19 cities in South and
32 Southeast Asia. Our approach is efficient, adaptable to many cities, standardizes and eliminates
33 sensitivity of the Gaussian fit to sampling area choice, and increases success of deriving annual
34 emissions from 40-60% with one sampling area to 100% (all 19 cities) with 54. The annual
35 emissions we estimate range from $16 \pm 5 \text{ mol s}^{-1}$ for Yangon (Myanmar) and Bangalore (India)
36 to $125 \pm 41 \text{ mol s}^{-1}$ for Dhaka (Bangladesh). With the enhanced success of our approach, we
37 find evidence from comparison of our top-down emissions to past studies and to inventory
38 estimates that the wind rotation and EMG fit approach may be biased, as it does not adequately
39 account for spatial and seasonal variability in NO_x photochemistry. Further methodological
40 development is needed to enhance its accuracy and to exploit it to derive sub-annual emissions.

41

42 **Plain Language Summary**

43 Cities are a large source of nitrogen oxides (NO_x) that go on to form many types of air pollutants
44 of harm to human health. City NO_x emissions estimated with observations from space-based
45 instruments are vital in regions that lack access to up-to-date, locally developed inventories.
46 Success of obtaining satellite-derived emissions hinges on user selection of a sampling area
47 around each city centre. Here we present an automated, efficient method that uses many (54)
48 sampling areas. When tested on 19 cities in South and Southeast Asia, annual NO_x emissions
49 are obtained for all 19 cities compared to about half the selected cities when using a single
50 sampling area. With this updated approach, we estimate total NO_x emissions in 2019 that range
51 from 23 kilotonnes for Yangon and Bangalore to almost 10-times more (181 kilotonnes) for
52 Dhaka. The greater success of our updated approach also helps us identify that the accuracy of
53 emissions derivation from satellite observations should be further improved by accounting for
54 the influence of spatial and seasonal variability in NO_x photochemistry.

55

56 **1 Introduction**

57 Nitrogen oxides ($\text{NO}_x \equiv \text{NO}_2 + \text{NO}$) react to form particulate nitrate and tropospheric
58 ozone and deposit to sensitive habitats (Luo et al., 2019; Sillman, 1999), thus degrading air
59 quality, altering climate, and adversely affecting human health and the environment (Grulke &
60 Heath, 2020; Lelieveld et al., 2015; Yue et al., 2017; Marais et al., 2023). Controls targeting
61 anthropogenic sources of NO_x have been extensively implemented in cities in Europe, the US
62 and China (Curier et al., 2014; de Foy et al., 2016; Silvern et al., 2019). In cities in other parts
63 of the world, particularly South and Southeast Asia, NO_x is increasing rapidly due to fast
64 economic development and limited or absent air quality policies (Vohra et al., 2021; 2022).
65 Vohra et al. (2022) used 14 years of satellite observations of NO_2 from the Ozone Monitoring
66 Instrument (OMI) to infer increases of $\sim 1\text{-}14 \text{ \% a}^{-1}$ in surface NO_2 pollution in almost all rapidly
67 developing large cities in South and Southeast Asia. Only in Jakarta did NO_2 decline due to
68 emission controls applied to vehicles (Vohra et al., 2022). Population projections suggest that,

69 by 2100, one-fifth of the world's most populous cities will be in Southeast Asia (Hoornweg &
70 Pope, 2017), necessitating reliable and up-to-date NO_x emissions estimates for assessing the
71 impact of this growth on urban air quality and for informing air quality policies.

72 Bottom-up inventories provide estimates of anthropogenic NO_x emissions, but publicly
73 available versions for South and Southeast Asia do not adequately represent contemporary
74 local conditions, as these are derived using outdated activity data, are resource-intensive to
75 produce so lag the present day, are at spatial resolutions that are coarser than many cities in the
76 region, and data needed to compile the inventories do not exist for many countries (Kurokawa
77 & Ohara, 2020). The two most used bottom-up inventories for these regions are the Regional
78 Emission inventory in Asia (REAS) (Kurokawa & Ohara, 2020) and the inventory known as
79 MIX, a mosaic of REAS and other regional inventories (Li et al., 2017). REAS and MIX are
80 at ~25 km resolution, MIX only covers 2 years of data, and the most recent years are 2015 for
81 REAS and 2010 for MIX. Still, REAS and inventories used to create MIX are routinely
82 incorporated in global inventories such as the Community Emissions Data System (CEDS_{GBD-}
83 _{MAPS}) (McDuffie et al., 2020), and Hemispheric Transport of Air Pollution (HTAP) (Crippa et
84 al., 2023).

85 Independent and contemporary estimates of city NO_x emissions can be derived with
86 satellite observations of tropospheric NO₂ vertical column densities (VCDs) without the need
87 for resource-intensive computer models. A method first proposed by Beirle et al. (2011)
88 involves selecting isolated cities and treating these as large point sources of NO_x. In this
89 approach, individual satellite pixels within a target domain centred on a city centre were split
90 into eight major wind directions to resolve the city plume in each direction. A mathematical
91 function was then fit to the plume to account for its Gaussian shape and exponential decay of
92 NO₂. This fit, referred to as an Exponential Modified Gaussian (EMG), yields parameters that
93 are then used to estimate NO_x emissions. It also yields an effective lifetime of NO_x for the city
94 plume that is dominated by dispersion for the windy conditions sampled. As dispersion
95 dominates, the derived lifetime is much shorter than the chemical lifetime of NO_x that includes
96 conversion to nitric acid (HNO₃) or organic nitrates (de Foy et al., 2014; Laughner & Cohen,
97 2019) and, to a lesser extent, dry deposition of NO₂ (Zhang et al., 2012). Beirle et al. (2011)
98 used OMI observations of NO₂ to derive NO_x emissions for eight global megacities. The Beirle
99 et al. (2011) approach required many (four) years of OMI data to achieve distinct plumes in
100 each wind direction.

101 Valin et al. (2013) expanded on the approach developed by Beirle et al. (2011) by
102 demonstrating that all satellite data can instead be aligned along a single upwind-downwind
103 direction relative to the city centre. This approach reduced the number of observations needed
104 to distribute the data by wind direction and so extended application to a greater number of
105 geographically isolated cities over shorter sampling periods. Wind rotation of OMI
106 observations and the EMG fit have since been used to calculate city NO_x emissions
107 predominantly in the US (de Foy et al., 2014; Goldberg et al., 2019a; Lu et al., 2015) and for
108 select cities worldwide (Goldberg et al., 2021). Following the 2017 launch of the higher spatial
109 resolution TROPOspheric Monitoring Instrument (TROPOMI), the wind rotation, EMG fit,
110 and related approaches have been extended to smaller isolated cities and shorter sampling
111 periods than was possible with OMI. Applications include cities in western Europe (Lorente et
112 al., 2019; Pope et al., 2022), China (Wu et al., 2021), the US (Goldberg et al., 2019b), and
113 worldwide (Lange et al., 2022), as well as investigating changes in NO_x emissions due to
114 COVID-19 lockdown measures in the New York Metropolitan Area (Tzortziou et al., 2022)
115 and for select cities in India, Argentina, and Spain (Lange et al., 2022). So far, the wind rotation
116 and EMG fit has only been applied to 5-13 cities in South and Southeast Asia as part of global
117 studies (Goldberg et al., 2021; Lange et al., 2022).

118 Even though there has been substantial development and use of the EMG fit, it still
119 requires that a user define a sampling area around the city that effectively captures the wind
120 rotated plume. The area selected varies with city size and plume length (Lu et al., 2015;
121 Goldberg et al., 2019a; Lange et al., 2022). This approach often yields no or poor EMG fits
122 and non-physical best-fit parameters (Laughner & Cohen, 2019), decreasing the likelihood of
123 deriving top-down emissions. Selecting appropriate city-specific areas for the wide-ranging
124 city sizes in South and Southeast Asia is also time consuming and not standardized.

125 Here we develop a near-automated and efficient EMG fitting routine for deriving
126 annual city NO_x emissions, demonstrate the utility of this automation by applying it to
127 TROPOMI NO₂ observations over isolated cities in South and Southeast Asia with wide-
128 ranging city sizes, compare our top-down emissions to past studies and a global bottom-up
129 inventory, and exploit the greater success of our updated sampling to identify opportunities to
130 further develop the EMG fit approach.

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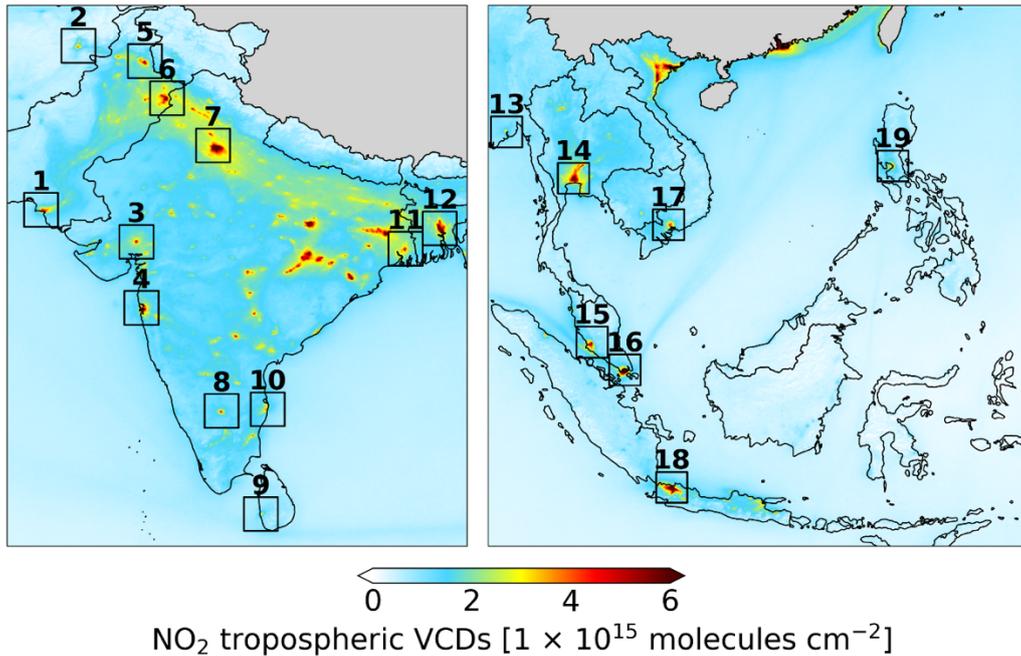
132 **2 Materials and Methods**

133 **2.1 TROPOMI NO₂ and City Selection**

134 We use Level 2 TROPOMI NO₂ tropospheric column VCDs for 2019 from the
135 Sentinel-5P Products Algorithm Laboratory (S5P-PAL) portal ([https://data-portal.s5p-
136 pal.com/](https://data-portal.s5p-pal.com/); last acquired 30 January 2022). These data have been retrieved with a consistent
137 algorithm (version 02.03.01) and corrected for a low bias in NO₂ over polluted scenes (Eskes
138 et al., 2021). TROPOMI achieves daily global coverage with a swath width of 2600 km, an
139 equator crossing time of 13:30 local solar time (LST), and a nadir pixel resolution that increased
140 on 5 August 2019 from 7 km × 3.5 km to 5.5 km × 3.5 km. We use cloud-free, high-quality
141 data identified with a quality flag ≥ 0.75 (van Geffen et al., 2021).

142 To identify isolated cities appropriate for top-down estimate of NO_x emissions, we first
143 oversample TROPOMI NO₂ to obtain high-resolution gridded annual means (0.05° × 0.05°;
144 ~6 km latitude × ~5 km longitude) by weighting areas of overlap between the satellite pixels
145 and cells on a fixed latitude-longitude grid using tessellation (Sun et al., 2018). We use the
146 resultant gridded TROPOMI NO₂ shown in Figure 1 to manually select 19 cities that are
147 isolated hotspots. The 19 selected cities are Karachi, Islamabad, and Lahore in Pakistan; Kabul
148 in Afghanistan; Ahmedabad, Mumbai, Delhi, Bangalore, Chennai, and Kolkata in India;
149 Colombo in Sri Lanka; Dhaka in Bangladesh; Yangon in Myanmar; Bangkok in Thailand;
150 Kuala Lumpur in Malaysia; the sovereign city Singapore; Ho Chi Minh City in Vietnam;
151 Jakarta in Indonesia; and Manila in the Philippines. Other hotspots in Figure 1 are either not
152 cities, such as the coal-fired power plants concentrated in eastern India, or are not isolated, such
153 as Hanoi, Haiphong and Nam Dinh in northern Vietnam.

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155

156 **Figure 1.** Annual mean TROPOMI tropospheric NO₂ VCDs over South and Southeast Asia in
 157 2019. Maps show South (left) and Southeast (right) Asia TROPOMI NO₂ oversampled to 0.05°
 158 × 0.05°. The 19 selected cities, numbered from east to west, are Karachi (1), Islamabad (5),
 159 and Lahore (6) in Pakistan; Kabul (2) in Afghanistan; Ahmedabad (3), Mumbai (4), Delhi (7),
 160 Bangalore (8), Chennai (10), and Kolkata (11) in India; Colombo (9) in Sri Lanka; Dhaka (12)
 161 in Bangladesh; Yangon (13) in Myanmar; Bangkok (14) in Thailand; Kuala Lumpur (15) in
 162 Malaysia; the sovereign city Singapore (16); Ho Chi Minh City (17) in Vietnam; Jakarta (18)
 163 in Indonesia; and Manila in the Philippines (19).

164

165 2.2 Wind Rotation and EMG Fit

166 Figure 2 illustrates the major steps involved in the wind rotation and EMG fit to derive
 167 annual NO_x emissions for Singapore. The wind fields we use to calculate wind direction and
 168 speed to retain TROPOMI NO₂ observations under windy conditions are the fifth generation
 169 European ReAnalysis (ERA5) 3D hourly *u* and *v* wind components
 170 (<https://cds.climate.copernicus.eu/cdsapp#!/home>; last acquired 18 March 2022) provided at
 171 0.25° × 0.25° resolution. At each TROPOMI NO₂ pixel, we compute collocated mean ERA5
 172 wind speeds and directions 30 min around 13:30 LST, the TROPOMI overpass time, in the
 173 lowest 5 layers (≥ 900 hPa) to capture dispersion of mixed-layer near-surface NO₂ plumes.
 174 Within a 4° × 4° domain around each city centre, we isolate TROPOMI pixels with coincident
 175 wind speeds > 2 m s⁻¹, the threshold typically used for windy conditions (Beirle et al., 2011;
 176 Pope et al., 2022). We rotate each TROPOMI NO₂ pixel by the angle of its wind direction,
 177 preserving the distance of the pixel from the city centre. This aligns all pixels along the same
 178 “upwind-downwind” direction that in our work is from north to south (Figure 2(a)). After wind
 179 rotating all pixels in a year (as in Figure 2), we grid pixels onto a uniform 0.05° × 0.05° grid
 180 using simple point-in-box averaging (Figure 2(a)) and fill empty grid cells (grey squares in
 181 Figure 2(a)) using nearest-neighbour interpolation to reduce low biases in the steps that follow.

182 Next, the 2D map in Figure 2(b) is converted to 1D line densities by summing all grid
 183 cells in the across-wind (east-to-west) direction in 0.05° upwind-downwind (north-to-south)
 184 increments. In the standard approach, a single area smaller than the 4° × 4° domain is used,

185 defined by the distance upwind, downwind, and across-wind of the city centre. Instead of using
 186 a single area, we define multiple areas that encompass the range of sizes typically used in past
 187 studies (Goldberg et al., 2021; Lange et al., 2022; Laughner & Cohen, 2019). These, defined
 188 as distances from the city centre, are 0.5°, 0.75°, and 1° upwind, 0.5°, 0.75°, 1.0°, 1.25°, 1.5°,
 189 1.75°, 2.0° downwind, and 0.5°, 0.75°, and 1.0° across-wind, with the requirement that the
 190 distance downwind of the city centre is \geq the distance upwind to capture the extent of the city
 191 plume. This yields 54 areas and associated line densities. The sizes of the smallest and largest
 192 areas sampled and the across-wind 0.05° increments summed to obtain line densities in the
 193 smallest area sampled are shown in Figure 2(b).

194 The EMG model we use to fit to the observed 1D line densities is the Laughner &
 195 Cohen (2019) formulation:

$$196 \quad F(x|a, x_0, \mu_x, \sigma_x, B) = \frac{a}{2x_0} \exp\left(\frac{\mu_x}{x_0} + \frac{\sigma_x^2}{2x_0^2} - \frac{x}{x_0}\right) \operatorname{erfc}\left(-\frac{1}{\sqrt{2}}\left[\frac{x-\mu_x}{\sigma_x} - \frac{\sigma_x}{x_0}\right]\right) + B \quad (1),$$

197 where x is the distance of each line density upwind and downwind of the city centre (Figure
 198 2(c)) and a , x_0 , μ_x , σ_x and B are best-fit parameters. Of these, a is total NO₂ in the plume (in
 199 moles), x_0 is the e -folding distance or length scale of NO₂ decay (in km), μ_x is the location of
 200 the apparent source relative to the city centre (in km) or the peak of the Gaussian fit that in
 201 Figure 2(c) is located \sim 20 km downwind or south of the city centre, σ_x is the Gaussian
 202 smoothing length scale (in km) that is \sim 2.355 \times the Full Width at Half Maximum (FWHM),
 203 and B is background NO₂ (in moles m⁻¹).

204 We use initial guesses for the best-fit parameters in Equation (1) that are similar to those
 205 from Laughner & Cohen (2019), but our fitting procedure differs. Laughner & Cohen (2019)
 206 used a non-linear interior point minimization algorithm (the *fmincon* function in MATLAB) to
 207 optimize model parameters with 10 iterations per line density. Instead, we perform the fit with
 208 the *scipy.optimize.curve_fit* module from SciPy Python package version 1.7.3 and iterate on
 209 the fit until the difference in fitting parameters between the current and previous iteration is
 210 negligible ($< 0.001\%$) for at most 10 iterations. Fit convergence is usually achieved after 3
 211 iterations. Only good-quality fits are retained, identified with goodness-of-fits (R^2) > 0.8 , as in
 212 Laughner & Cohen (2019). We further screen for physically implausible best-fit parameters
 213 using criteria similar to Laughner & Cohen (2019): a is positive, x_0 is at least 1.6 km
 214 (approximately $1/e$ of the grid resolution), μ_x is within the sampling area, the emission width
 215 is less than the e -folding distance ($\sigma_x < x_0$), background NO₂ is positive and less than the
 216 maximum line density value, and the e -folding distance occurs between the plume centre and
 217 the edge of the sampling area. We introduce an additional requirement to ensure that x_0 is within
 218 the sampling area ($x_0 < \text{length of sampling area downwind of the city centre}$).

219 The Singapore example in Figure 2 is an ideal city, as all 54 EMG fits are successful.
 220 Figure 2(c) shows that the observed line densities are most sensitive to the across-wind length,
 221 as this determines the amount of NO₂ summed to yield each line density. We will demonstrate
 222 in Section 3 that for many of the cities in Figure 1 a large number of EMG fits fail to meet the
 223 conditions for success, necessitating as many as 54 fits.

224 The successful EMG fits are used to calculate effective NO_x lifetimes (τ_{NO_x} ; reported
 225 in h) and midday NO_x emissions (E_{NO_x} ; in moles s⁻¹):

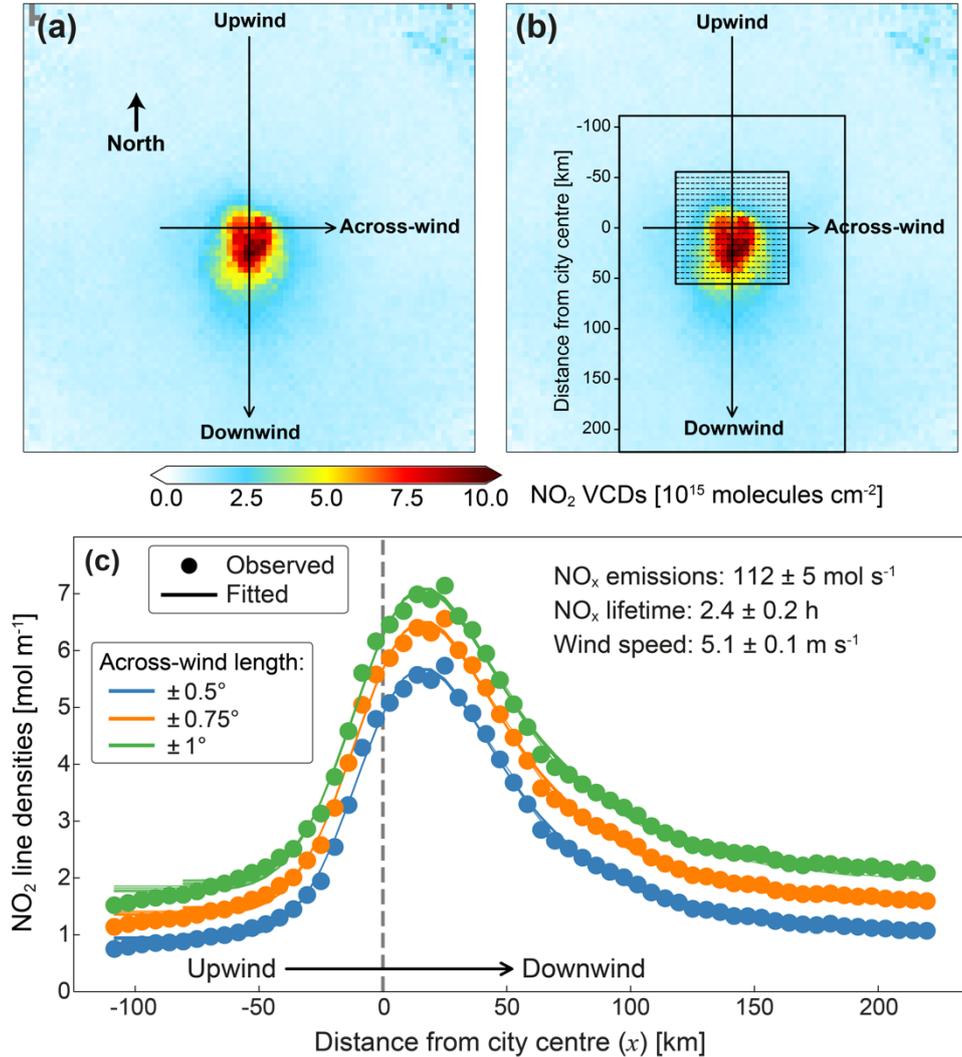
$$226 \quad \tau_{NO_x} = \frac{x_0}{\omega} \quad (2)$$

$$227 \quad E_{NO_x} = \gamma \times \frac{a}{\tau_{NO_x}} \quad (3),$$

228 where ω is the sampling area mean wind speed (in m s^{-1}) and γ is the unitless molar ratio of
 229 $[\text{NO}_x]/[\text{NO}_2]$ to convert moles NO_2 to moles NO_x . The up to 54 individual estimates of τ_{NO_x}
 230 and E_{NO_x} are averaged to obtain values for each city.

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232



233

234 **Figure 2.** Illustration of major steps in the wind rotation and EMG fit to derive annual NO_x
 235 emissions for Singapore. The main steps in each panel are wind rotate and grid windy scene
 236 TROPOMI NO_2 pixels to $0.05^\circ \times 0.05^\circ$ (a), fill data gaps (b), and fit the EMG function (Eq.
 237 (1)) (solid lines) to observed line densities (filled circles) (c). In (b), black rectangles show the
 238 extent of the largest and smallest sampling areas and dashed lines in the smallest area show the
 239 0.05° increments used to calculate the line densities in (c). All 54 successful EMG fits, 18 lines
 240 for each of the three across-wind lengths, are shown in (c). Values in (c) give the mean and
 241 standard deviation of the city NO_x emissions (Eq. (3)), effective NO_x lifetime (Eq. (2)), and
 242 sampling area ERA5 wind speed. The goodness-of-fit (R^2) is ≥ 0.99 for all fits in (c).

243

244 We use the same $[\text{NO}_x]/[\text{NO}_2] = 1.32$ value as Beirle et al. (2011) and subsequent
 245 studies to represent rapid cycling between NO and NO_2 . Liu et al. (2022) determined with
 246 synthetic experiments that city NO_x emissions are relatively unaffected by variability in
 247 $[\text{NO}_x]/[\text{NO}_2]$, but that study was for US cities. Surface measurements aid in determining

248 suitability of $[\text{NO}_x]/[\text{NO}_2] = 1.32$, but these are limited to cities in India and have data quality
249 issues (Vohra et al., 2021). Instead, we use the GEOS-Chem model to assess suitability of the
250 1.32 value. We simulate the model in 2019 and sample the lowest model layer around the
251 TROPOMI overpass time. We use output from a coarse and finer resolution version of GEOS-
252 Chem to also test sensitivity of this ratio to model resolution, especially given many of these
253 cities are coastal (Figure 1). We use the classical configuration of the model that operates on a
254 single computational node, called GEOS-Chem Classic (GCCClassic), and the high-
255 performance model configuration (GCHP) that is a parallelized across multiple computational
256 nodes to enable finer resolution global simulations (Eastham et al., 2018). GCCClassic is version
257 13.3.4 (<https://doi.org/10.5281/zenodo.5764874>) run on a fixed $2^\circ \times 2.5^\circ$ global grid and
258 GCHP is version 13.4.1 (<https://doi.org/10.5281/zenodo.6564711>) run on a C360 global grid
259 ($\sim 25 \text{ km} \times \sim 31 \text{ km}$). GCCClassic and GCHP use the same vertical grid and chemical mechanism.
260 For GCCClassic, grid squares that overlap with each city are sampled, whereas for GCHP, we
261 use city sampling extents determined from a combination of administrative and geographic
262 boundary shapefiles and Google Maps (Figure S1). Midday sampling is at 12:00 to 15:00 LST
263 from GCCClassic and 13:00 to 14:00 LST from GCHP. At midday, NO_x is in photochemical
264 steady state, so the relative abundance of NO and NO_2 is insensitive to the extent of the
265 sampling window around midday (Potts et al., 2021).

266 We calculate uncertainties in the NO_x emissions by adding individual errors in
267 quadrature. These include best-fit parameters x_0 and a , sampling area mean wind speed ω , the
268 TROPOMI NO_2 observations, and $[\text{NO}_x]/[\text{NO}_2]$. We use the relative standard deviation from
269 all successful EMG fits to calculate city-specific errors in x_0 and a . For ω , we consider errors
270 due to the choice of spatial and temporal sampling and the threshold used for windy conditions.
271 We use the Beirle et al. (2011) estimated 10% error in temporal sampling choice and 5% error
272 due to vertical sampling choice. We conduct our own tests of the sensitivity to threshold and
273 spatial sampling choice. For $[\text{NO}_x]/[\text{NO}_2]$ we assess whether the 10% error attributed to this
274 variable by Beirle et al. (2011) is appropriate by quantifying the percent deviation of GCCClassic
275 and GCHP $[\text{NO}_x]/[\text{NO}_2]$ from 1.32. Beirle et al. (2011) applied a 30% error to OMI that is also
276 appropriate for TROPOMI. Even though uncertainties in TROPOMI slant columns (NO_2 along
277 the viewing path) are much less than those from OMI (van Geffen et al., 2020), the air mass
278 factor used to convert slant columns to VCDs remains the largest contributor to errors in NO_2
279 VCDs and is similar for OMI and TROPOMI (van Geffen et al., 2021).

280 **2.3 Bottom-up Anthropogenic Emissions**

281 We compare our top-down estimates to anthropogenic NO_x emissions from the widely
282 used bottom-up HTAP inventory version 3 (HTAP_v3) (Crippa et al., 2023). HTAP_v3 has
283 high enough spatial resolution ($0.1^\circ \times 0.1^\circ$) to resolve cities selected in Figure 1. The most
284 recent year is 2018, achieved by extending emissions from the regional REAS inventory ending
285 in 2015 to the year 2018 with trends from the Emissions Database for Global Atmospheric
286 Research (EDGAR) inventory. The same sampling boundaries as GCHP are used (Section 2.2;
287 Figure S1). The HTAP_v3 NO_x emissions include contributions from aviation, transport (road,
288 rail, pipeline, inland waters), shipping, energy, industry, and residential sectors.

289 Cities targeted can be influenced by non-anthropogenic NO_x sources, such as open
290 burning of biomass (Marvin et al., 2021) and natural sources such as soils (Weng et al., 2020)
291 and lightning (Miyazaki et al., 2014). We assess suitability of comparing our top-down
292 emissions to anthropogenic bottom-up emissions only by determining the percent contribution
293 of anthropogenic emissions to total NO_x emissions. To do this, we simulate total NO_x emissions
294 with the Harmonized Emissions Component (HEMCO) standalone model version 3.0.0
295 (<https://zenodo.org/records/4984639>; last accessed 20 March 2022) (Lin et al., 2021) and

296 sample the same spatial extent as GCHP and HTAP_v3 (Figure S1). HEMCO is run at a spatial
297 resolution of $0.25^\circ \times 0.3125^\circ$ (~ 28 km latitude $\times \sim 33$ km longitude). HEMCO calculates open
298 biomass burning emissions using the Global Fire Emissions Database with small fires
299 (GFED4s) inventory (Randerson et al., 2017) and reads in and processes lightning and soil NO_x
300 from offline emissions at the same resolution as HEMCO (Murray et al., 2012; Weng et al.,
301 2020).

302 Bottom-up emissions from HTAP_v3 are 24-h means, whereas top-down estimates
303 derived using TROPOMI are representative of midday emissions. Goldberg et al. (2021)
304 multiplied satellite-derived midday NO_x emissions by 0.77 to convert midday top-down NO_x
305 emissions to 24-h means for comparison to bottom-up inventories. This value was inferred
306 from bottom-up emissions estimates for the Netherlands, so may not be suitable for the selected
307 cities in South and Southeast Asia. The hourly scaling factors used by HEMCO for the chosen
308 cities range from 0.70 to 1.16. These are for the year 2000 and are extrapolations of values for
309 conditions in Europe, so may not be suitable for the year and cities targeted in this study. Given
310 this, we do not scale top-down emissions and instead discuss whether differences in averaging
311 times contribute to discrepancies between top-down and bottom-up emissions estimates.

312

313 **3 Results and Discussion**

314 **3.1 Wind Rotation and EMG Fit Metrics**

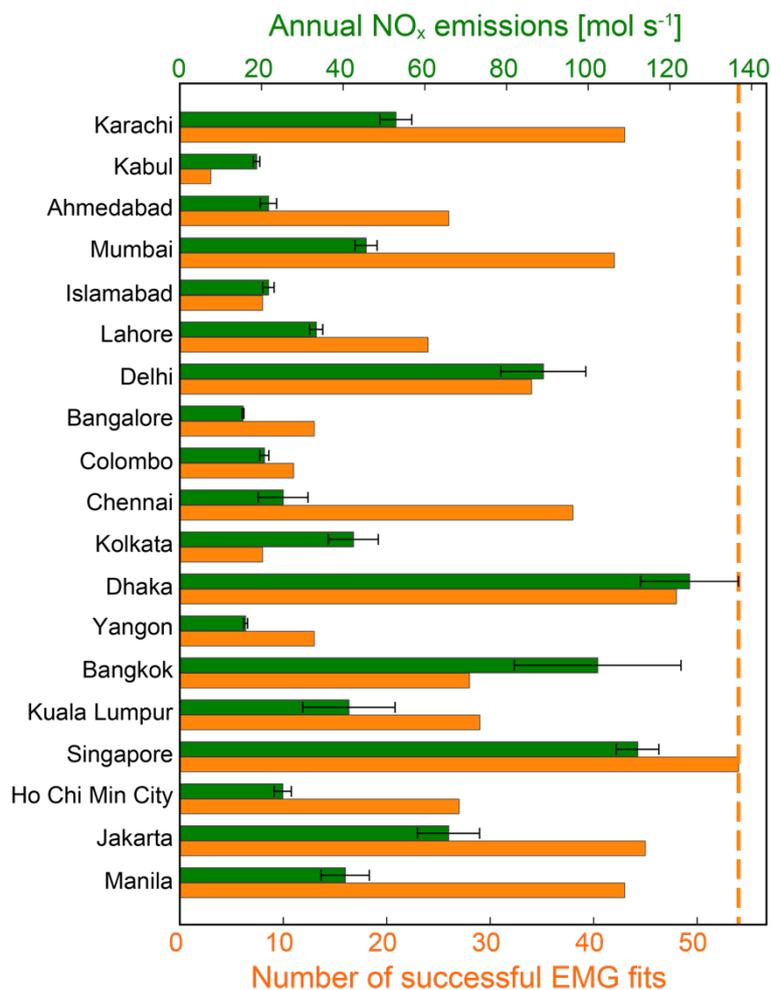
315 Isolating windy condition ($> 2 \text{ m s}^{-1}$) satellite pixels removes 8-34% of all 2019 quality-
316 and cloud-screened TROPOMI NO_2 pixels for most cities in Figure 1. Cities with greater data
317 loss are Lahore (43% data loss), Kabul (58%) and Islamabad (63%). No spatial data gap filling
318 (Section 2.2, Figure 2) is needed within the areas sampled, due to the high sampling frequency
319 of TROPOMI. If only a single domain size is selected, annual EMG fits meet all criteria for
320 success for 7 to 12 of the 19 cities in Figure 1, depending on the sampling area chosen. Using
321 our extended method, we successfully derive annual NO_x emissions for all 19 cities, due to the
322 enhanced probability of obtaining at least one successful EMG fit.

323 Figure shows the number of successful EMG fits (orange bars) range from 3 (Kabul) to
324 all 54 (Singapore). Singapore, Dhaka, Jakarta, Karachi, Manila, and Mumbai are least impacted
325 by the choice of sampling area. The 6 cities in Figure 3 with < 20 fits are most likely to fail if
326 only a single sampling area is used. For all retained EMG fits, differences between observed
327 and fitted NO_2 line densities, the fit residuals, are negligible. The most common causes for a
328 failed EMG fit rank as: background NO_2 (B in Equation (1)) $>$ maximum NO_2 line density (36%
329 of all fits conducted), $R^2 \leq 0.8$ (24%), emission width $>$ e -folding distance (19%), total plume
330 NO_2 (a in Equation (1)) $<$ 0 (13%), and e -folding distance $>$ the downwind length of the
331 sampling area (12%). Multiple causes can co-occur in a single fit, so cumulative percentages
332 exceed 100%.

333 We also test sensitivity of top-down NO_x emissions to the choice of wind speed
334 threshold and horizontal sampling extent to attribute an error to these. For this, we apply a
335 stricter wind speed threshold of 3 m s^{-1} and test the difference in NO_x emissions if instead of
336 filtering for windy conditions using pixel-mean wind fields, we calculate a sampling-area mean
337 wind speed to filter for windy conditions as in Goldberg et al. (2019a). We apply these
338 conditions to a mid-sized sampling area of 0.75° upwind, 1.5° downwind, and $\pm 0.75^\circ$ across-
339 wind. Variability in NO_x emissions for cities with successful EMG fits for all 4 wind sampling
340 conditions is at most 10% (Figure S2). Given these results, we attribute a 10% error to the
341 choice of horizontal sampling and to the wind speed threshold.

342 GClassic (coarse resolution) annual mean $[\text{NO}_x]/[\text{NO}_2]$ for the target cities ranges
 343 from 1.25 (Dhaka) to 1.41 (Kabul). The range in ratios from GCHP (finer resolution) is wider
 344 at 1.24 (Ahmedabad) to 1.64 (Kolkata). The difference in ratios between the coarse and fine
 345 resolution models is typically $\pm 10\%$, except for a few cities with ratios from the fine resolution
 346 model that exceed the coarse resolution model by 14% for Singapore, 16% for Lahore, 23%
 347 for Dhaka, and 23% for Kolkata. This is because the fine resolution model better resolves the
 348 city plume that includes a greater proportion of NO_x as NO from fresh emission sources. As
 349 the difference between the model city ratios and the 1.32 value is $\pm 10\%$ for most cities, we use
 350 the same 10% error for $[\text{NO}_x]/[\text{NO}_2]$ as Beirle et al. (2011).

351



352

353 **Figure 3.** Successful EMG fits and top-down NO_x emissions for the cities targeted in this study.
 354 Bars are emissions (green) and the corresponding number of successful fits (orange). Black
 355 error lines are NO_x emission standard deviations for all successful fits. The orange dashed line
 356 at 54 indicates the maximum possible EMG fits. Emissions multiplied by ~ 1.45 yields
 357 emissions in $\text{Gg NO}_2 \text{ a}^{-1}$.

358

359 3.2 Top-Down NO_x Emissions

360 Green bars in Figure 3 show the mean annual top-down NO_x emissions for all cities
 361 (values are in Table S1). These range from $\sim 16 \text{ mol s}^{-1}$ for Bangalore and Yangon to $\sim 125 \text{ mol}$
 362 s^{-1} for Dhaka. The range in the total mass of NO_x emitted for these cities, assuming the midday

363 emission rate is reasonably representative of the 24-h emission rate, is 23-181 Gg NO_x as NO₂.
364 Emissions for most cities are < 50 mol s⁻¹ (<73 Gg NO_x as NO₂ a⁻¹). Cities with emissions
365 between 50-100 mol s⁻¹ (73-145 Gg NO_x as NO₂ a⁻¹) include Karachi, Delhi, and Jakarta and >
366 100 mol s⁻¹ (> 145 Gg NO_x as NO₂ a⁻¹) include Bangkok, Singapore, and Dhaka. Emission
367 rates for Bangkok, Dhaka and Singapore are comparable to the range of top-down emissions
368 estimated for large, polluted cities in China using the EMG approach (Wu et al., 2021). The
369 effective lifetimes for the cities in Figure 1 (shown in Figure S3) range from 1.2 h for Colombo
370 to 6.3 h for Kuala Lumpur. Variability in effective lifetimes depends most strongly on the
371 downwind extent of the plume. The Pearson's correlation coefficient, R, between city mean
372 effective lifetimes and x₀ values is 0.90.

373 For the target cities, the relative standard deviations of annual NO_x emissions (black
374 error lines in Figure 3) range from just 1% for Bangalore to 27% for Kuala Lumpur. This is far
375 less than the equivalent Gaussian fit uncertainty of 10-50% estimated by Beirle et al. (2011)
376 for a single sampling area. The relatively large variability in Kuala Lumpur NO_x emissions is
377 because the smaller EMG sampling areas do not fully encompass the elongated wind rotated
378 city NO₂ plume, causing a low bias in NO_x emissions for the smaller areas sampled. The effect
379 of this is dampened by the almost 30 successful fits used to obtain mean NO_x emissions for this
380 city. The relative standard deviations of the NO_x lifetimes (Figure S3) range from 3% for
381 Bangalore to 37% for Chennai. The relative standard deviations of other parameters are ~6%
382 for wind speeds (Figure S4), 4% (Bangalore) to 38% (Chennai) for x₀, and 4% (Kabul and
383 Bangalore) to 37% (Bangkok) for *a*.

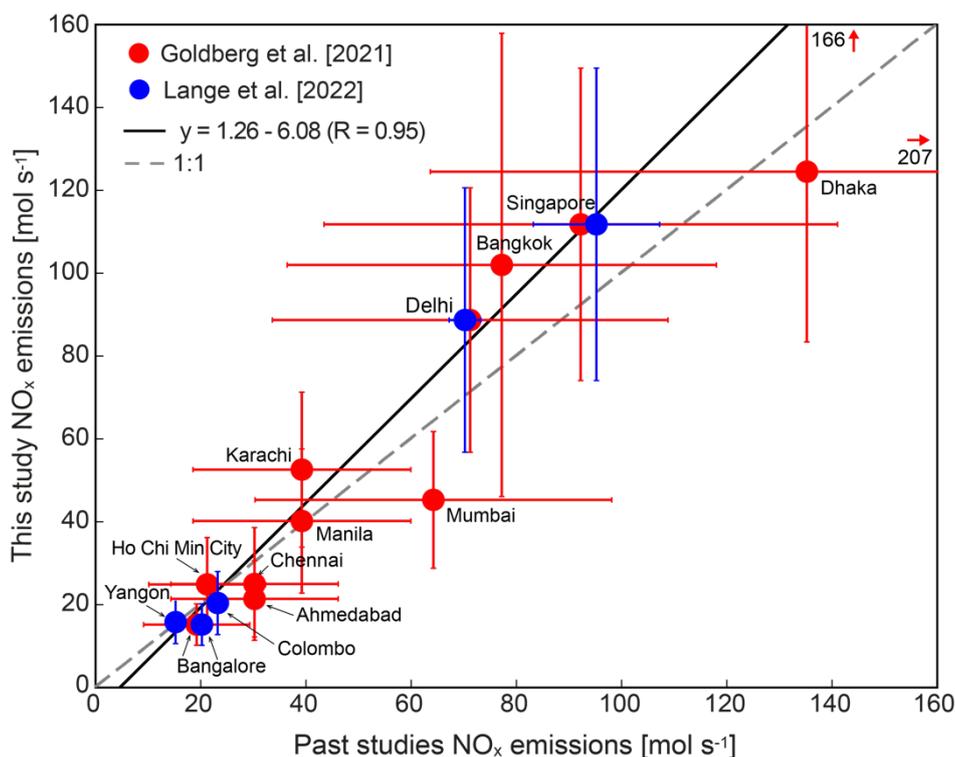
384 The overall uncertainty in annual NO_x emissions we obtain by adding all error
385 contributions in quadrature ranges from 32% for Bangalore and Yangon to 55% for Bangkok.
386 Values for all cities are in Table S1. The TROPOMI NO₂ VCDs make the largest contribution
387 to the overall uncertainty. The higher-end of our uncertainty estimates is similar to the typical
388 ~50% uncertainty reported in past studies (Beirle et al., 2011; Verstraeten et al., 2018; Goldberg
389 et al., 2021). We use our overall uncertainties in the comparison of our top-down emissions to
390 values from the literature and from HTAP in the sections that follows.

391

392 **3.3 Comparison to Top-Down Estimates from Past Studies**

393 To assess our approach, we compare in Figure 4 our annual NO_x emissions to values
394 from past studies that used similar sampling time periods and a single sampling area. These
395 include multiyear (2017-2019) mean emissions from Goldberg et al. (2021) obtained using the
396 OMI sensor and emissions from Lange et al. (2022) obtained with select days of TROPOMI
397 data from 2018 to 2020. Goldberg et al. (2021) estimated emissions for 10 of the 19 cities in
398 our study. These we read from their Figure S10 for Karachi, Figure S11 for 4 cities in India,
399 and Figure S13 for 5 cities in Southeast Asia and divide by the 0.77 midday to 24-h scaling
400 factor used in that study. Emissions are reported by Lange et al. (2022) for 5 of the 19 cities in
401 our study. Based on the regression statistics in Figure 4, our emissions are typically ~26% more
402 than estimates from these past studies. Exceptions are Mumbai, Ahmedabad, and Chennai that
403 in our study are 16-29% less than Goldberg et al. (2021). Lange et al. (2022) used an earlier
404 version of the TROPOMI data product that has a known low bias in NO₂ VCDs over very
405 polluted scenes (van Geffen et al., 2022). Differences in TROPOMI data products are the likely
406 cause for our higher Delhi (by 27%) and Singapore (by 18%) emissions. Relatively small error
407 estimates from Lange et al. (2022) are because they only propagate error contributions from
408 the wind speed data and the EMG fit.

409



410

411 **Figure 4.** Comparison of our and past top-down NO_x emissions. Symbols compare our
 412 emissions to those from Goldberg et al. (2021) (red) and Lange et al. (2022) (blue). Error bars
 413 are overall uncertainties for our study (Section 2.2, Table S1), the same 53% uncertainty
 414 applied to all cities by Goldberg et al. (2021) and the city-specific uncertainties for Lange et al.
 415 (2022). Lines are the Theil regression fit (solid black) and 1:1 relationship (dashed grey). Inset
 416 text gives the regression statistics and Pearson’s correlation coefficient (R). Arrows and inset
 417 text for Dhaka give the error values that extend beyond the plotting range.

418

419 Discrepancies between Goldberg et al. (2021) and our emissions are not as
 420 straightforward to diagnose, as Goldberg et al. (2021) use NO₂ VCDs from a different sensor
 421 (OMI) and apply a systematic 37% increase to NO_x emissions to correct for a low bias in OMI
 422 attributed to the coarse resolution a priori used in the NO₂ VCDs retrieval. Sampling area
 423 choice may also be a factor. For example, the smallest of our 54 areas yields NO_x emissions of
 424 102 mol s⁻¹ for Singapore that is 10 mol s⁻¹ less than the mean of all EMG fits. Goldberg et al.
 425 (2021) used year-round OMI data for all cities except Delhi and Karachi. As these cities are
 426 north of 25°N, only May-September observations were used by Goldberg et al. (2021). We find
 427 that Delhi and Karachi mean May-September TROPOMI NO₂ VCDs in 2019 averaged within
 428 the 4° × 4° domain selected for each city (Figure 2(a)-(b)) are 11-12% less than those in
 429 October-April, due to the shorter photochemical lifetime of NO_x in the warmer months. Open
 430 biomass burning emissions also influence seasonality in the TROPOMI NO₂ VCDs, but the
 431 EMG fit accounts for this by distinguishing background NO₂ (*B* in Equation (1)) from NO₂ in
 432 the city plume (*a* in Equation (1)).

433 We find that if we apply the EMG fit to individual months for Delhi and Karachi, all
 434 54 EMG fits fail for Delhi in July-August and yield spurious results in September due to large
 435 data loss resulting from persistent clouds during the monsoon season. All 12 months are
 436 retained for Karachi, Singapore and Manila. November-April mean values of *a* are 21% more
 437 than in May-October for Karachi, 9% more for Singapore, and 39% more for Manila. This
 438 suggests that using NO₂ VCDs for a portion of the year may yield systematic biases in

439 emissions that may not reflect seasonality in the underlying activities affecting the emissions.
440 Larger wintertime than summertime emissions have also been reported in the global study of
441 Lange et al. (2022). They quantified summer-to-winter emission ratios of ~ 0.5 for Colombo
442 and Delhi. The top-down emissions calculation (Equation (3)) does not fully account for
443 seasonality in photochemistry. The derived effective NO_x lifetimes used to calculate NO_x
444 emissions (Equation (2)) are mostly influenced by dispersion. As a result, the effective lifetimes
445 are much shorter than the expected chemical lifetimes of NO_x (de Foy et al., 2014). In the
446 synthetic experiment scenarios tested by de Foy et al. (2014), the EMG fit applied to wind
447 rotated data yielded an effective lifetime of 4 h for a 12-h chemical lifetime scenario. According
448 to Shah et al. (2020), the chemical lifetime of NO_x for central-eastern China centred at $\sim 35^\circ\text{N}$,
449 the northerly portion of our domain, ranges from ~ 6 h in summer to ~ 24 h in winter. None of
450 the monthly effective lifetimes for our target cities reproduces this seasonality and the longest
451 lifetime is 13.3 ± 3.7 h for Yangon in November. The implication is that the size of absolute
452 emissions derived with sub-annual satellite data may be biased, but should have negligible
453 effect if used to quantify relative trends, as in Goldberg et al. (2021) and Laughner & Cohen
454 (2019), for example.

455

456 3.4 Comparison to Bottom-up Emissions

457 Figure 5 compares annual top-down and bottom-up NO_x emissions. According to our
458 HEMCO simulations, anthropogenic sources account for most ($>87\%$) annual NO_x emissions.
459 The relative differences between our top-down estimates and the bottom-up inventory are
460 within 50% for Mumbai (1%), Bangkok (2%), Chennai (9%), Ahmedabad (11%), Kolkata
461 (21%), Singapore (21%), Bangalore (32%), Manila (35%), and Kuala Lumpur (46%). A 50-
462 100% difference occurs for Ho Chi Minh City (53%), Jakarta (54%), Delhi (64%), and
463 Colombo (91%). Even greater relative differences occur for Karachi (2.1 times), Islamabad
464 (2.1 times), Lahore (2.4 times), Yangon (3.3 times), Dhaka (6.9 times), and Kabul (11-fold).
465 The largest absolute discrepancies are for Dhaka and Jakarta. Bottom-up emissions are 107
466 mol s^{-1} less than the top-down values for Dhaka and 78 mol s^{-1} more for Jakarta. On a mass
467 basis, this is equivalent to a 155 Gg NO_x as NO_2 underestimate for Dhaka and a 113 Gg NO_x
468 as NO_2 overestimate for Jakarta.

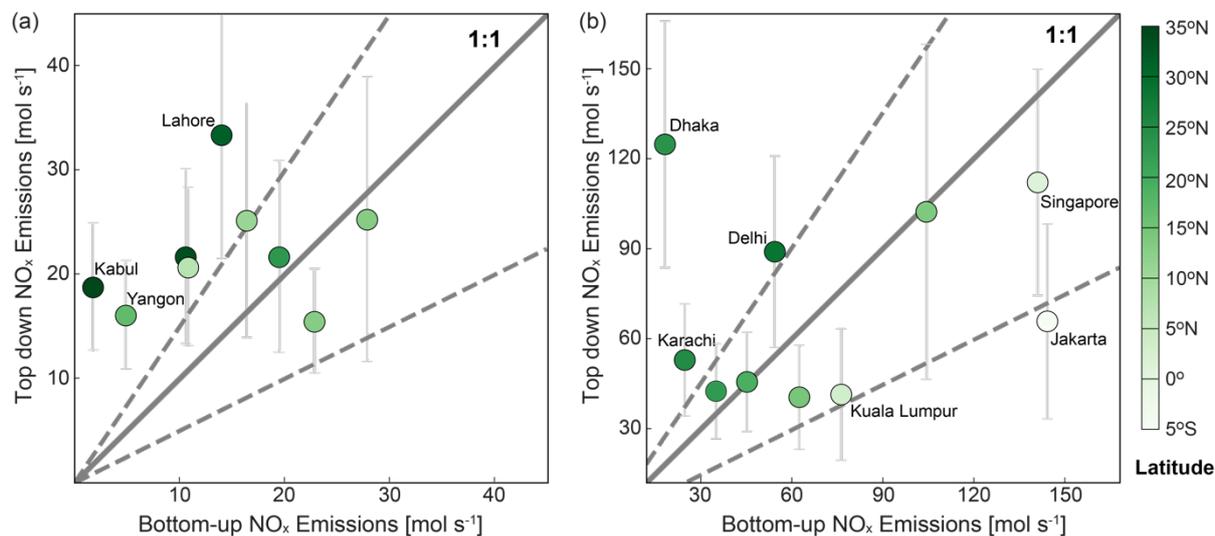
469 The different years used (2018 for HTAP, 2019 for TROPOMI) should at most account
470 for a 14% difference in emissions, based on the size of annual trends inferred by Vohra et al.
471 (2022) using long-term observations of OMI NO_2 VCDs over large and fast-growing cities in
472 South and Southeast Asia. Vohra et al. (2022) identified that emission inventories do not
473 capture the steep decline in NO_x emissions in Jakarta attributed to national policies targeting
474 vehicles. In addition to misrepresenting annual changes in underlying activities, the emission
475 factors are mostly informed by studies in China and Japan (Kurokawa & Ohara, 2020). The
476 bottom-up and top-down emissions differences for many cities also exceed the $\pm 30\%$
477 difference that results from the choice of bottom-up emissions grid sampling and the $\pm 30\%$
478 difference from the timing of the top-down (midday) and bottom-up (24-h) estimates inferred
479 by Goldberg et al. (2021).

480 Apparent in Figure 5 is a latitudinal pattern in the discrepancies. Top-down emissions
481 are greater than bottom-up emissions for cities to the north and vice versa for cities to the south,
482 so that in general top-down emissions exceed bottom-up emissions in South Asia and vice
483 versa in Southeast Asia. NO_x chemical loss varies with latitude, due to variability in the amount
484 of sunlight available to form hydroxyl and peroxy radicals required to form HNO_3 and organic
485 nitrates, the main daytime chemical loss pathway for NO_x . This latitudinal pattern is likely

486 because the EMG fit also does not fully account for spatial variability in NO_x photochemistry,
 487 imparting a bias in the top-down emissions. The size of this bias will depend on the relative
 488 contribution of NO_x chemical loss to total loss in the wind rotated plume.

489

490



491

492 **Figure 5.** Comparison of annual top-down and bottom-up NO_x emissions for target cities. Data
 493 are coloured by city centre latitude and split into top-down NO_x emissions < 40 mol s⁻¹ (a) and
 494 ≥ 40 mol s⁻¹ (b). Error bars are the overall uncertainty in top-down emissions estimates. Grey
 495 lines indicate 1:1 agreement (solid) and ±50% difference (dashed). The bottom-up emissions
 496 sampling extent of each city is in Figure S1. Data used to generate the figure are in Table S1.

497

498 4 Conclusions

499 City nitrogen oxides (NO_x) emissions can be derived with a now well-established
 500 approach using satellite observations of nitrogen dioxide (NO₂), wind rotation and a Gaussian
 501 fit to the city plume. Issues with this approach are that the choice of sampling area around the
 502 city centre is not standardized and so is prone to subjective area selection and the Gaussian fit
 503 often fails or yields non-physical best-fit parameters. Here we address these issues by applying
 504 54 sampling areas to isolated cities. We test our method with TROPospheric Monitoring
 505 Instrument (TROPOMI) NO₂ observations for 2019 over 19 large, isolated cities in South and
 506 Southeast Asia that lack contemporary, publicly available bottom-up emissions estimates.

507 Annual NO_x emissions, obtained for all 19 cities, are < 73 Gg NO_x as NO₂ a⁻¹ for most
 508 cities, between 73-145 Gg NO_x as NO₂ a⁻¹ for Karachi, Delhi, and Jakarta and > 145 Gg NO_x
 509 as NO₂ a⁻¹ for Bangkok, Dhaka, and Singapore. The overall uncertainty in the annual emissions
 510 is 30-60%. Our emissions estimates are in general ~27% more than past studies that use a single
 511 sampling area, due to differences in satellite data products and months targeted. The latter we
 512 suggest may lead to biases, as the top-down emissions estimate does not properly account for
 513 seasonality in photochemical loss of NO_x. Relative differences between our top-down estimates
 514 and a widely used bottom-up inventory are < 50% for 9 of the 19 cities, within 50-100% for
 515 Ho Chi Minh City, Jakarta, Delhi, and Colombo, and much greater for Karachi (2.1 times),
 516 Islamabad (2.1 times), Lahore (2.4 times), Yangon (3.3 times), Dhaka (6.9 times), and Kabul
 517 (11-fold). There is a latitudinal dependence of the size of these discrepancies that we suggest

518 is because the top-down approach also does not properly account for spatial variability in the
519 chemical lifetime of NO_x.

520 The increased success of deriving NO_x emissions with our updated approach enables
521 us to identify that further development is needed to account for time and space variability in
522 the chemical lifetime of NO_x to fully exploit the top-down approach to interrogate seasonality
523 in emissions, to validate bottom-up emissions, to exploit hourly observations from
524 geostationary instruments, and to inform air quality regulation.

525 **Data and Software Availability**

526 The TROPOMI tropospheric columns for 2019 are publicly available from the S5P-PAL Data
527 Portal (<https://data-portal.s5p-pal.com/>). GEOS-Chem source codes are preserved on Zenodo
528 by The International GEOS-Chem User Community (2021) for GCClassic version 13.3.4 and
529 by The International GEOS-Chem User Community (2022) for GCHP version 13.4.1.

530 **Author Contributions**

531 GL developed the methodology, GL and EAM processed, analysed and interpreted the data.
532 GL and EAM prepared the manuscript. KV assisted in data collection and analysis. RPH and
533 DZ conducted the GEOS-Chem simulations (RPH: GCClassic; DZ: GCHP). RVM contributed
534 to the methodology. SG contributed to interpretation of the results. All co-authors provided
535 editorial input.

536 **Conflicts of interest**

537 The authors declare there are no conflicts of interest.

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