Simulations of Energetic Neutral Atom sputtering from Ganymede in preparation for the JUICE mission

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Key Points:

- A new method for calculating sputtered fluxes at Ganymede is introduced
- The energy spectra of sputtered H_2O, O_2, and H_2 ENAs are calculated for the first time
- The Jovian Neutrals Analyzer on JUICE can remotely map ion precipitation at Ganymede

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Abstract

Jovian magnetospheric plasma irradiates the surface of Ganymede and is postulated to be the primary agent that changes the surface brightness of Ganymede, leading to asymmetries between polar and equatorial regions as well as between the trailing and leading hemispheres. As impinging ions sputter surface constituents as neutrals, ion precipitation patterns can be remotely imaged using the Energetic Neutral Atoms (ENA) measurement technique. Here we calculate the expected sputtered ENA flux from the surface of Ganymede to help interpret future observations by ENA instruments, particularly the Jovian Neutral Analyzer (JNA) onboard the JUpiter ICy moon Explorer (JUICE) spacecraft. We use sputtering models developed based on laboratory experiments to calculate sputtered fluxes of H_2O, O_2, and H_2. The input ion population used in this study is the result of test particle simulations using electric and magnetic fields from a hybrid simulation of Ganymede's environment. This population includes a thermal component (H^+ and O^+ from 10 eV to 10 keV) and an energetic component (H^+, O^{++} and S^{+++} from 10 keV to 10 MeV). We find a global ENA sputtering rate from Ganymede of 1.42×10^{27} s^{-1}, with contributions from H_2, O_2, and H_2O of 34%, 17%, and 49% respectively. We also calculate the energy distribution of sputtered ENAs, give an estimate of a typical JNA count rate at Ganymede, and investigate latitudinal variations of sputtered fluxes along a simulated orbit track of the JUICE spacecraft. Our results demonstrate the capability of the JNA sensor to remotely map ion precipitation at Ganymede.

Plain Language Summary

Particles trapped by Jupiter's magnetic field interact with Jupiter's moons. Ganymede, the largest of those moons, lacks a dense atmosphere to protect its surface from these energetic Jovian particles, but Ganymede's magnetic field is strong enough to influence their trajectory: charged particles are deflected away from equatorial regions to polar regions, resulting in uneven particle precipitation patterns at the surface of Ganymede. When ions hit the surface of Ganymede, they eject particles from the surface, in a process referred to as sputtering. Those particles are mostly neutral and therefore unaffected by Ganymede's magnetic fields, so we can image where ions hit the surface of Ganymede by measuring ejected neutral particles. The Jovian Neutrals Analyzer (JNA) will fly onboard the JUpiter ICy moon Explorer (JUICE) spacecraft and will measure sputtered neutrals in the vicinity of Ganymede. To help interpret the data to be collected by JNA, we used models derived from laboratory experiments to simulate what JNA will observe at Ganymede. Our results show that JNA will be able to show us where ions hit the surface of Ganymede, which is important as uneven ion precipitation is thought to explain why Ganymede's poles are brighter than its equatorial regions.

1 Introduction

Imaging plasma precipitation patterns at the surface of Ganymede is a key measurement for understanding the effect of Jovian plasma precipitation on the brightness and composition of the surface. Ganymede stands out as Jupiter's largest moon and also the only moon in the Solar System to feature an intrinsic magnetic field, causing the formation of a small magnetosphere inside Jupiters much larger magnetosphere. Ganymede's magnetic field locally impedes or enhances Jovian plasma access to its surface, resulting in variable precipitation patterns (Khurana et al., 2007; Fatemi et al., 2016; Poppe et al., 2018; Plainaki et al., 2020). Precipitating ions can be backscattered by the surface or cause surface constituents and radiolytic products to sputter. Backscattered and sputtered particles leave the surface mainly as neutral and with energies ranging from eV to MeV (Johnson, 1990).

Such neutral particles are usually referred to as Energetic Neutral Atoms (ENAs) (Gruntman, 1997). The trajectories of ENAs are not influenced by electric or magnetic fields and therefore preserve information about their original location and direction. Thus ENA measurements can and have been used to remotely map ion precipitation on airless bodies, where in-situ precipitating ion observation is not easily possible. For example, backscattered solar wind protons and sputtered oxygen atoms were observed at the Moon by the Interstellar Boundary EXplorer (IBEX) (McComas et al., 2009; Allegrini et al., 2013) and by CENA on Chandrayaan-1 (Wieser et al., 2009; Vorburger et al., 2014; Futaana et al., 2013). Ganymede will also be visited by an ENA instrument: the Jovian Neutrals Analyzer (JNA) will fly on-board the JUpiter ICy moon Explorer (JUICE) spacecraft.

To help interpret the data collected by JNA, estimates of ENA fluxes are needed, but require modelling of the sputtering process. The sputtering process has been widely studied because of its relevance for icy bodies such as Ganymede, Europa, and Enceladus. Their lack of a dense atmosphere leaves their surface exposed to ion precipitation, leading to the sputtering of surface constituents and radiolytic products. Along with other processes such as sublimation and photo-stimulated desorption, sputtering contributes to the creation of a neutral exosphere on several bodies (Cooper, 2001; Johnson et al., 2004; Marconi, 2007; Cassidy et al., 2010; Wurz et al., 2010).

While a comprehensive analytical description of the sputtering process is not currently available, sputtering has been extensively studied through laboratory experiments (Baragiola et al., 2003; Fam et al., 2008; Teolis et al., 2017; Galli et al., 2017, 2018). Several methods have been developed to calculate the sputtering yield of ions on icy surfaces as a function of projectile energy and species, incidence angle, and surface temperature (Johnson et al., 2004; Fam et al., 2008; Teolis et al., 2017). Such methods, or combinations of them, have been used extensively to simulate surface-plasma interactions at Europa (Vorburger, 2018; Plainaki et al., 2010, 2012) and Ganymede (Marconi, 2007; Turc et al., 2014; Plainaki et al., 2015; Shematovich, 2016; Leblanc et al., 2017; Poppe et al., 2018; Plainaki et al., 2020).

Here, we estimate sputtered ENA fluxes at the surface of Ganymede by applying models formulated by Fam et al. (2008), Johnson et al. (2004) and Teolis et al. (2017) to a population of incident Jovian plasma obtained through hybrid simulations by Poppe et al. (2018). This allows us to calculate the expected sputtered ENA fluxes of H_2O, H_2, and O_2 and to further apply the Thompson-Sigmund (Sigmund, 1969) law to calculate their energy distribution. By convolving JNA's estimated geometric factor with the energy distribution, we give an expected JNA count rate in the vicinity of Ganymede. Finally we investigate latitudinal variations of the sputtered ENA fluxes by simulating a simplified orbit of the JUICE spacecraft around Ganymede.

2 Materials and Methods

2.1 Input population

The incident ion population is taken from a combination of three-dimensional hybrid simulations of Ganymede's magnetosphere and subsequent backwards-Liouville particle tracing through the hybrid electromagnetic fields (Fatemi et al., 2016; Poppe et al., 2018). They simulated the ion velocity distribution for the Galileo G8 flyby, when Ganymede was in the Jovian plasma sheet. The incident population used as an input for calculating the sputtered ENA flux in this study is comprised of three-dimensional velocity distribution functions for thermal O^+ and H^+ from 10 eV to 10 keV and energetic H^+, O^{++} and S^{+++} from 10 keV to 10 MeV. The spatial resolution of the hybrid model is $1° \times 1°$ in latitude and longitude at the surface of Ganymede.

Figure 1 shows the resulting plasma precipitation pattern, i.e. a map of the ion flux integrated over all incident species, energies and angles. On the trailing hemisphere (180° - 360°W) of the equatorial regions, Ganymede's surface is shielded from Jovian plasma by Ganymede's intrinsic magnetic field. In contrast, intense precipitation is observed on the leading hemisphere of the equatorial regions (0° - 180°W), where plasma is accelerated back towards Ganymede by reconnection in the magnetotail (Fatemi et al., 2016; Poppe et al., 2018). The most intense flux is observed in the high-latitude cusp regions on the leading hemisphere, where open-closed magnetic field lines boundaries are located ($\pm 50° - 60°$ in latitude) (Poppe et al., 2018).
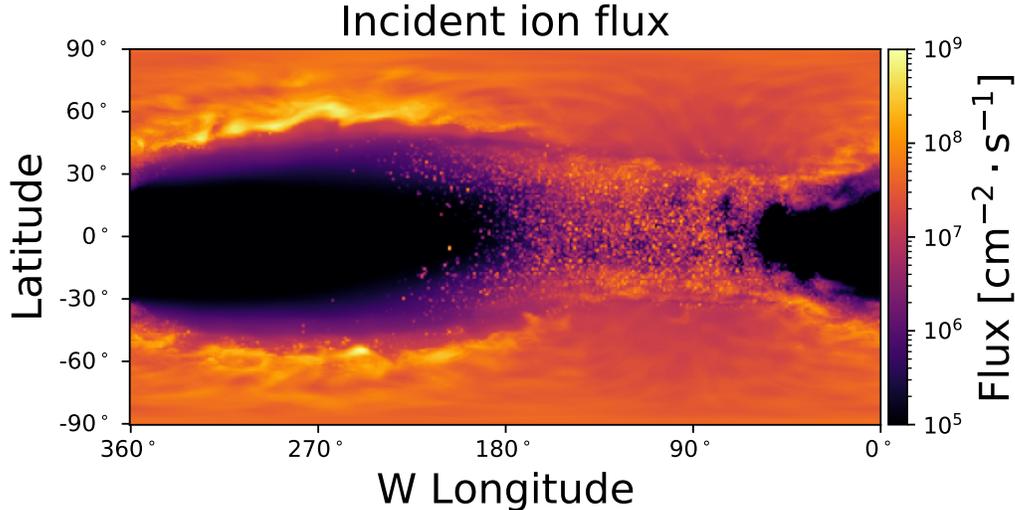

Figure 1. Incident ion flux at the surface of Ganymede, taken from Poppe et al. (2018) and integrated over all species, angles and energies. The leading hemisphere extends from 0°W to 180°W while the trailing hemisphere extends from 180°W to 360°W. For our study here, we choose a single period along Ganymede's orbit such that the sub-solar point is located at 270°W, i.e. the co-rotating plasma flow is aligned with the sunlight direction.

Figure 2 shows the energy distribution of the incident ion flux at Ganymede's surface resulting from Poppe's backwards-Liouville tracing model. The flux was integrated over all incident angles and averaged over the surface of Ganymede. Two components can be identified: i) the thermal component comprised of plasma from Io's torus diffusing outwards (Siscoe & Summers, 1981); ii) the energetic component originating from accelerated Io torus plasma and solar wind plasma diffusing inwards (Siscoe et al., 1981).
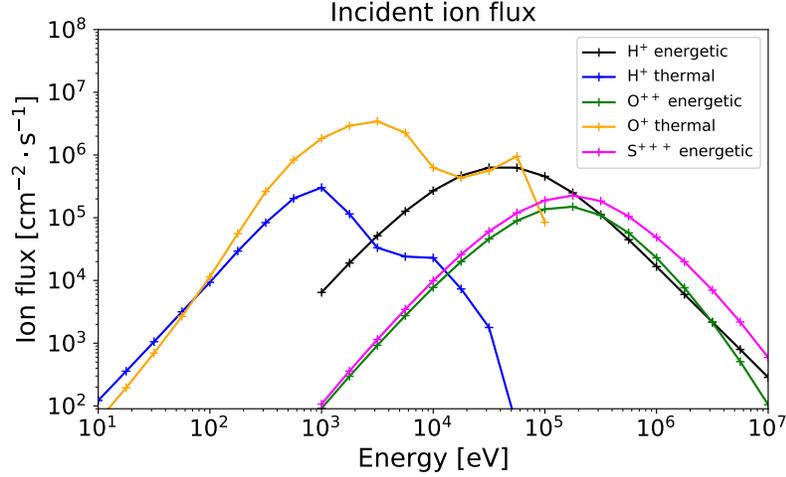
Figure 2. Energy spectra of the precipitating ions used as our input ion populations, integrated over all incident angles and averaged over the surface of Ganymede.

2.2 Sputtering yield

Ion sputtering on water ice has been extensively studied via laboratory experiments and the sputtered products are known to be comprised of H_2O, H_2, and O_2 (Johnson et al., 2004; Galli et al., 2017). While H_2O is a surface constituent directly sputtered by the impact of ions on water ice, H_2 and O_2 are radiolytic products generated in the material by the irradiation of water ice and subsequently sputtered by projectiles (Johnson et al., 2003; Paranicas et al., 2009; Teolis et al., 2017). Here we use three different functions to calculate the sputtering yield, depending on the energy of the incident and the sputtered species. Throughout the paper, the sputtering yield *of* a sputtered species *by* an incident species refers to the number of particles of the sputtered species released from the surface by one incident ion.

At incident ion energies higher than 100 keV, we use the model described in Johnson et al. (2004) to calculate $Y_{H_2O,high}$, the yield of H_2O as a function of the energy and species of the incident ion:

$$\frac{1}{Y_{H_2O,high}(v,Z)} = \frac{1}{11.2 \, Z^{2.8}(v/Z^{1/3})^{-2.24}} + \frac{1}{4.2 \, Z^{2.8}(v/Z^{1/3})^{2.16}}, \tag{1}$$

where Z is the atomic number of the incident particle and v the velocity of the incident particle in atomic units (1 au $= 2.18 \cdot 10^8$ cm $\cdot$ s^{-1}).

At incident ion energies lower than 100 keV, we use the model described in Fam et al. (2008) to calculate $Y_{H_2O,low}$, the yield of H_2O as a function of the energy, species, and incident angle of the incident particle:

$$Y_{H_2O,low}(E,m,Z,\beta) = \frac{1}{U_0}\left(\frac{3}{4\pi^2 C_0 \alpha S_n} + \eta S_e^2\right)\cos^{-f}(\beta), \tag{2}$$

where E, m, and Z are respectively the energy, mass, and atomic number of the projectile. β is the incidence angle, defined from the surface normal. U_0, C_0, α, S_n, η, S_e and f are constants; details can be found in Fam et al. (2008).

In eq. 2 we do not include the temperature-dependent component of Famá's model, as it is attributed to H_2 and O_2 produced by radiolysis, a temperature-dependent process. Instead, we calculate the yield of H_2 ($Y_{H_2}(E,T,\beta)$) and that of O_2 ($Y_{O_2}(E,T,\beta)$)

153 using the model derived by Teolis et al. (2017):

$$Y_{O_2}(E,T,\beta) = \frac{Y_{H_2}(E,T,\beta)}{2} = \epsilon g_{O_2}^0 x_o \left[1 - \exp\left(-\frac{r_o \cos(\beta)}{x_o}\right)\right] \left[1 + q_o \exp\left(-\frac{Q}{k_b T}\right)\right]/r_o \cos(\beta),$$

(3)

155 where T is the temperature of the surface of Ganymede, β the incidence angle of the pro-
156 jectile measured form the surface normal, and k_b is the Boltzmann constant. Details about
157 ϵ, $g_{O_2}^0$, x_0, r_0, q_0, and Q can be found in Teolis et al. (2017). The temperature model
158 used in this study was derived by Marconi (2007) based on data from the Galileo space-
159 craft (Orton et al., 1996). The dayside temperature is given by $T(\lambda) = 70 \ \cos(\lambda)^{0.75} +$
160 80 K (where λ is the sub-solar latitude) and the nightside temperature is a constant 80
161 K.
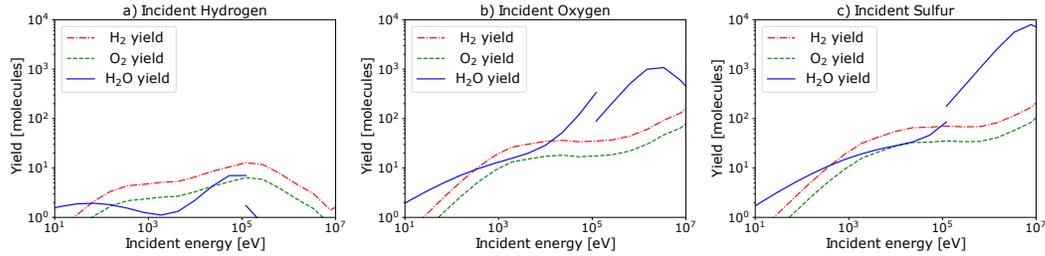
Figure 3. Sputtering yield of H_2O, H_2 and O_2 by incident O, H, and S ions. The disconti-
nuity at 100 keV for the H_2O yields is explained by the transition of the model from Famá's to
Johnson's. The yields of H_2 and O_2 are calculated using only Teolis' model.

162 Figure 3 shows the sputtering yield of H_2O, O_2 and H_2 by O, H and S ions imping-
163 ing on water ice. Generally, the H_2O yield by O and S is higher than 1 and increases with
164 energy for most of the energy range shown here. Because of its low atomic mass, the yield
165 by H is much lower. A surface temperature of 124 K was used to generate these figures,
166 which corresponds to an average daytime disk temperature (Grundy et al., 1999). As
167 mentioned above, the actual surface temperature used in our model varies between 80
168 K on the nightside and 150 K at the sub-solar point.

2.3 Sputtered energy distribution

170 We assume a Thompson-Sigmund law to calculate the probability distribution $S(K)$
171 of the energy of the sputtered particles (Sigmund, 1969).

$$S(K) \propto \frac{K}{(K + E_b)^3} \cdot \left(1 - \sqrt{\frac{K}{E_b + 4\,E_i(M_1 M_2)/(M_1 + M_2)^2}}\right),$$

(4)

173 where K is the energy of the sputtered neutral particle, E_b the binding energy of the sur-
174 face (0.054 eV as also used in Plainaki et al. (2015)), E_i the energy of the projectile, and
175 M_1 and M_2 are the masses of the projectile and sputtered neutral particle.

176 Figure 4a shows the energy distribution of H_2O molecules sputtered by H, O and
177 S with an incident energy of 100 keV. The main consequence of the difference in atomic
178 mass is the cutoff energy of sputtered particles: sputtered H has a cutoff energy of about
179 20 keV, much lower than the cutoff energy of heavier oxygen and sulfur at about 90 keV.
180 Figure 4b shows the energy distribution of H_2O, H_2 and O_2 molecules sputtered by 100
181 keV H. Higher masses result in lower cutoff energies, with respective cutoff energies for
182 O_2, H_2O, and H_2 of about 10 keV, 20 keV, and 80 keV.
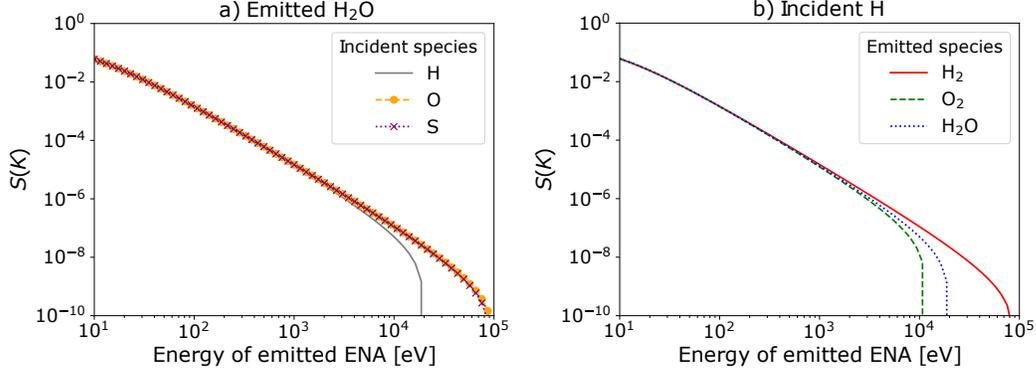
Figure 4. a) Probability distribution of the energy of H_2O sputtered by different incident species. b) Probability distribution of the energy of H_2O, H_2 and O_2 molecules sputtered by H, for incident energies of 100 keV.

2.4 Flux calculation

The differential flux of the sputtered neutrals, $j(K)$, is calculated from the combination of the above-mentioned parameters, where K is the energy of the sputtered particle. In our model, the incident plasma taken from hybrid simulations is $f(v, \theta, \phi)$, in units of $(m/s)^{-3}sr^{-1}cm^{-2}s^{-1}$, for each bin at the surface (Poppe et al., 2018), where v is the incident velocity, θ the incident elevation, and ϕ the incident azimuth.

The differential flux $j(K)$ (in units $cm^{-2}s^{-1}sr^{-1}eV^{-1}$) is calculated using the following expression:

$$j(K) = \frac{1}{2\pi} \int_v \int_\theta \int_\phi f(v, \theta, \phi) \ Y(E_i) \ S(K; E_i) \ v^2 \cos(\alpha) \sin(\theta) \ dv d\theta d\phi, \quad (5)$$

where E_i the energy of the incident species, $Y(E_i)$ the sputtering yield function, $S(K; E_i)$ the Thompson-Sigmund probability distribution function, and α is the angle between the velocity vector and the local normal vector pointing inward to the center of Ganymede at the corresponding latitude and longitude. We assume that sputtered neutrals are ejected isotropically and therefore divide the flux by 2π to get the flux per solid angle.

3 Results

3.1 ENA sputtering maps

Figure 5 shows the calculated flux maps of the sputtered ENAs (H_2, O_2, and H_2O) integrated over all incident species, incident angles, and energies. The colorbar ranges from 10^6 to 10^{10} $cm^{-2} \cdot s^{-1}$ for all three maps. H_2 and O_2 fluxes are about 5 times higher on the dayside than on the nightside due to the higher surface temperature on the dayside, which results in a higher yield of H_2 and O_2 (eq. 3). As the yield of H_2O is independent of temperature, no significant difference between the dayside and the nightside is observed other than that resulting from the input ion precipitation patterns.

Generally, similar patterns to the ion precipitation map (Figure 1) are observed for ENA sputtering. Indeed, the sputtered ENA flux is higher in the polar regions than in the equatorial regions, and the difference in ENA fluxes between the pole and the equator is more distinct in the trailing hemisphere. The similarity between ion precipitation patterns and sputtering rate patterns illustrates the relevance of the ENA imaging method to remotely map ion precipitation at Ganymede, as previously shown for terrestrial bod-
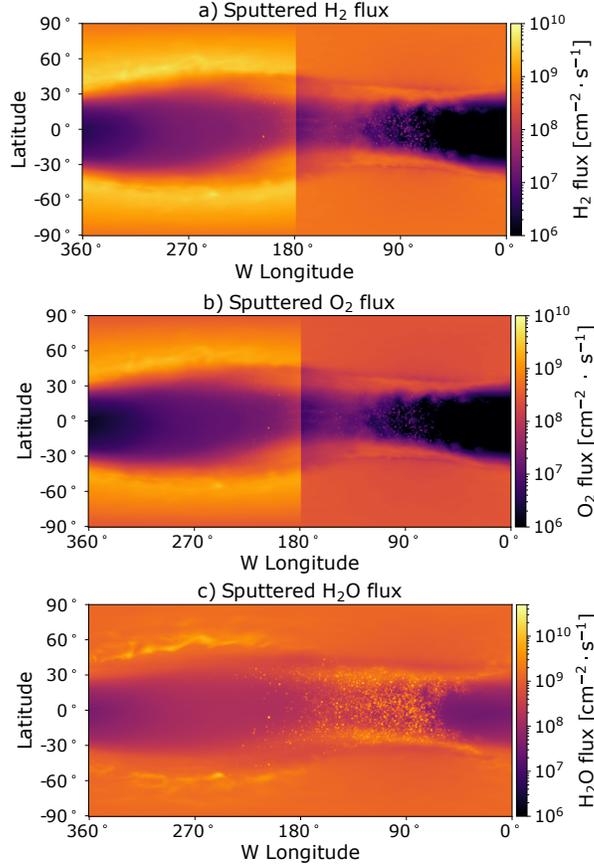
Figure 5. Maps of the sputtered fluxes of H_2, O_2 and H_2O, integrated over all incident species, energies, and angles.

ies (Futaana et al., 2006; Vorburger et al., 2014; Wieser et al., 2009; Allegrini et al., 2013; Futaana et al., 2013).

3.2 Sputtered energy distribution

Figure 6 shows the energy spectra of sputtered H_2O, H_2, and O_2. Because of the Thompson-Sigmund law (eq. 4) for the energy of sputtered ENAs, fluxes fall as energy increases. Although the incident ion population is dominated by the thermal O^+ component at low energies, the contribution of energetic S^{+++} ions to the sputtered H_2O, H_2, and O_2 ENA fluxes dominates over that of all other species across the entire energy range. This is likely a combined effect of the heavier mass of S^{+++} ions, resulting in high sputtering yields, and the fact that the energy distribution of S^{+++} is skewed towards higher energies.

The lowest contribution to sputtered ENAs comes from the thermal H^+ ion component because of their low incident flux (Figure 2) and low sputtering yield, lower than 10 across most of the energy range. For the same reason, energetic H^+ ions contribute the least out of the three energetic species to sputtering O_2 and H_2O. Energetic H^+ ions sputter less H_2O molecules than thermal O^+ ions up to 20 keV. At 20 keV, H_2O ENAs sputtered by O^+ cut off, due to O^+ ions being heavier than H^+ and also contributing more to the incident ion flux.
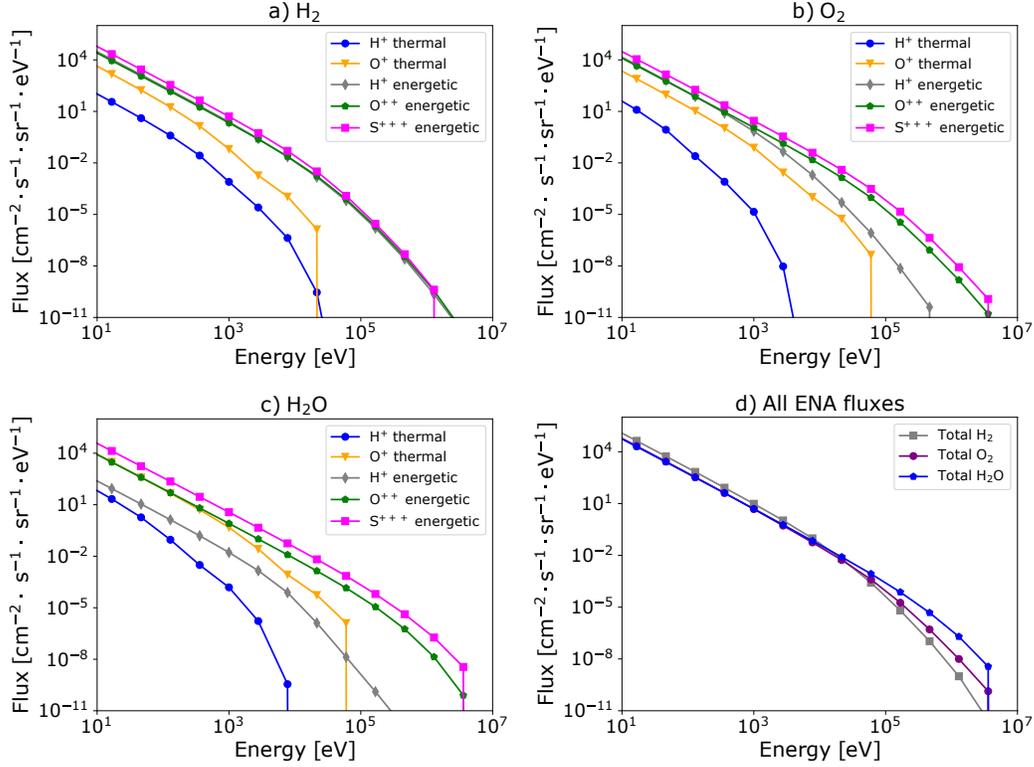
Figure 6. Globally averaged energy distributions of sputtered a) H_2, b) O_2, and c) H_2O, integrated over incident angles and energies. Different colors indicate the incident species which sputtered these ENAs. d) Energy distribution of all sputtered species, summed over incident species.

Total sputtered fluxes of H_2O, H_2, and O_2 are compared with one another in Figure 6d) and are shown to be within the same order of magnitude up to 10 keV. From 10 keV to 4 MeV (where all ENA fluxes cut off), H_2O fluxes dominate over H_2, and O_2.

3.3 Total sputtering rate

By integrating the sputtered ENA flux over energy and the entire surface we obtain global sputtering rates of 4.8, 2.4 and 7.0 $\times 10^{26}$ molecules per second for H_2, O_2, and H_2O respectively. This results in a total sputtering rate of 1.42×10^{27} molecules per second. Table 1 shows a comparison of our estimated sputtering rate with previous works. Given from left to right are the model reference, the input plasma model type (MHD or hybrid), the sputtering model(s), the energy range of the sputtered molecules, the species of the sputtered molecules, and the total sputtering rate.

All works give similar results within two orders of magnitude. The difference between our result and those of previous works can be qualitatively explained as follows. Plainaki et al. (2015) derived their ion population using electric and magnetic fields obtained with MHD simulations by Jia et al. (2008). Their energy range covered only 1-100 keV, which is narrower than in this study, leading to a total sputtering rate one order of magnitude lower than ours. They used Famá's model to calculate the yield of H_2O and O_2, but did not account for the sputtering of H_2.

Poppe et al. (2018), when they published the results of the hybrid simulation of Jovian plasma which were used in this study, took the opportunity to estimate the H_2O

Table 1. Total sputtering rate from the surface of Ganymede estimated by previously published works as well as this one.

Reference	Input	Sputtering models	Energy range	Sputtered species	Total sputtering rate [s^{-1}]
Plainaki et al., 2015	MHD[1]	Famá[3]	1-100 keV	H_2O, O_2	6.94×10^{25}
Poppe et al., 2018	Hybrid[2]	Johnson[4]	10 eV - 10 MeV	H_2O	7.5×10^{26}
Carnielli et al., 2020	MHD[1]	Famá[3], Johnson[4]	1 eV - 30 MeV	H_2O	2.25×10^{27}
This work	Hybrid[2]	Famá[3], Johnson[4], Teolis[5]	10 eV - 10 MeV	H_2O, O_2 H_2	1.42×10^{27}

[1] Jia et al. (2008)
[2] Fatemi et al. (2016)
[3] Fam et al. (2008)
[4] Johnson et al. (2004)
[5] Teolis et al. (2017)

ENA sputtering rate using Johnson's model. However, at incident energies lower than 100 keV, Johnson's model underestimates the yield, which is better reproduced by Famá's (Cassidy et al., 2013). Moreover, Poppe et al. (2018) considered only the sputtering of H_2O, whereas we considered O_2 and H_2 in addition to H_2O.

Carnielli et al. (2020) used the model in Jia et al. (2008) also used to derive their input ion population, but considered energies ranging from 1 eV to 30 MeV, a wider energy range than used here. Moreover, they considered the contribution of Ganymede's ionospheric ions, which they showed can contribute to up to 10% of the ENA sputtering rate. Their ionospheric ion population was comprised of O_2^+, O^+, H_2O, H_2^+, H^+, and OH^+ with energies ranging from 10 eV to 10 keV. As our input population did not include ionospheric ions and covered a narrower energy range, our total sputtering rate is expected to be lower than theirs.

Our results suggest that H_2 and O_2 account for half of the total neutral sputtering rate from the surface of Ganymede, showing that their contribution should be considered in addition to that of H_2O.

3.4 JNA count rate estimation

The JUICE spacecraft, planned to launch in 2022 and expected to reach Jupiter in the 2030s, carries the Particle Environment Package (PEP). PEP is comprised of six sensors tailored to study how Jovian plasma interacts with Ganymede's magnetosphere, tenuous atmosphere, and icy surface. In particular, the Jovian Neutrals Analyzer (JNA) will measure ENAs in the Jovian environment in the energy range between 10 eV to 3.3 keV, with a field-of-view of 15°×150° divided into 11 pixels (Shimoyama et al., 2018).

Here we estimate the count rate that JNA is expected to observe at Ganymede by multiplying the flux calculated in section 3.2 by JNA's estimated geometric factor, GF = 10^{-5} cm$^2 \cdot$ sr $\cdot$ eV/eV. Figure 7 shows simulated JNA count rates as a function of energy. The geometric factor we used is constant across the energy range, so the count rate distribution follows the Thompson-Sigmund law applied to the sputtered ENAs, result-
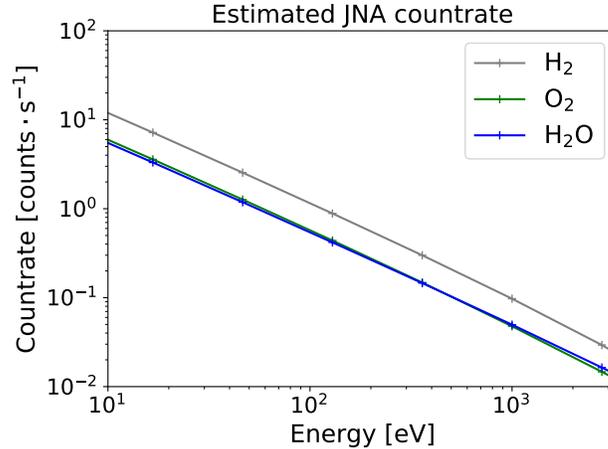
Figure 7. Simulated JNA count rate as a function of energy for sputtered H_2O, O_2 and H_2 in the energy range JNA can measure.

ing in count rates as high as 10^1 counts·s^{-1} at 20 eV and as low as 10^{-2} counts·s^{-1} at 3.3 keV. JNA is optimized to measure small fluxes, i.e. low count rates, even in the harsh radiation environment expected at Jupiter. However, such a spectra indicates that longer integration times are needed at high energies than at low energies, an important consideration for operations planning and future data analysis. As we did not account for any dissociative processes of the sputtered ENAs, we assume here that H_2, O_2, and H_2O are observed and detected by JNA as molecules. In reality, any molecule entering JNA would most likely be dissociated upon encountering JNA's conversion surface, as JNA uses a charge conversion surface to ionize ENAs in order to analyze their energy and guide them to JNA's detectors (Kazama et al., 2007). Since dissociated products would each leave the conversion surface with less energy than the original molecule, the assumption that JNA observes molecules likely gives an underestimation of the count rate JNA would measure at low energies.

3.5 JNA simulated observation

To illustrate JNA's ability to measure the variability of Jovian plasma precipitation at Ganymede, we calculate the differential ion flux at different latitudes on Ganymede. In Figure 8, the JUICE spacecraft is assumed to orbit Ganymede at an altitude of 1000km along the 90°W and 270°W meridians. At four locations along the orbital track, the flux was averaged over areas corresponding to the size of the footprint of JNA's center pixel. Those areas are referred to as zones. For each zone, the fluxes of sputtered H_2, O_2, and H_2O are shown, as well as the JNA one count level.

Zone 1 is situated at Ganymede's north pole and covers areas on both the dayside and the nightside. Zone 2 is centered around the sub-solar point. Zone 3 is located along the sub-solar longitude at latitude 60°N, near the open-closed field-line boundary where the ion flux peaks. Zone 4 is centered around the anti-solar point on the nightside.

Figure 8c shows that the highest flux is observed in Zone 3, reflecting the peak in ion flux at this location near the open-closed field-line boundary where both energetic and thermal plasma have easy access to the surface. The ion population there is dominated by thermal O^+ (see Poppe et al. (2018) for incident ion flux distributions at different regions on Ganymede). Zone 1, over the north pole, is exposed to an ion population similar to that in Zone 3, although precipitation at the poles is less intense. Con-
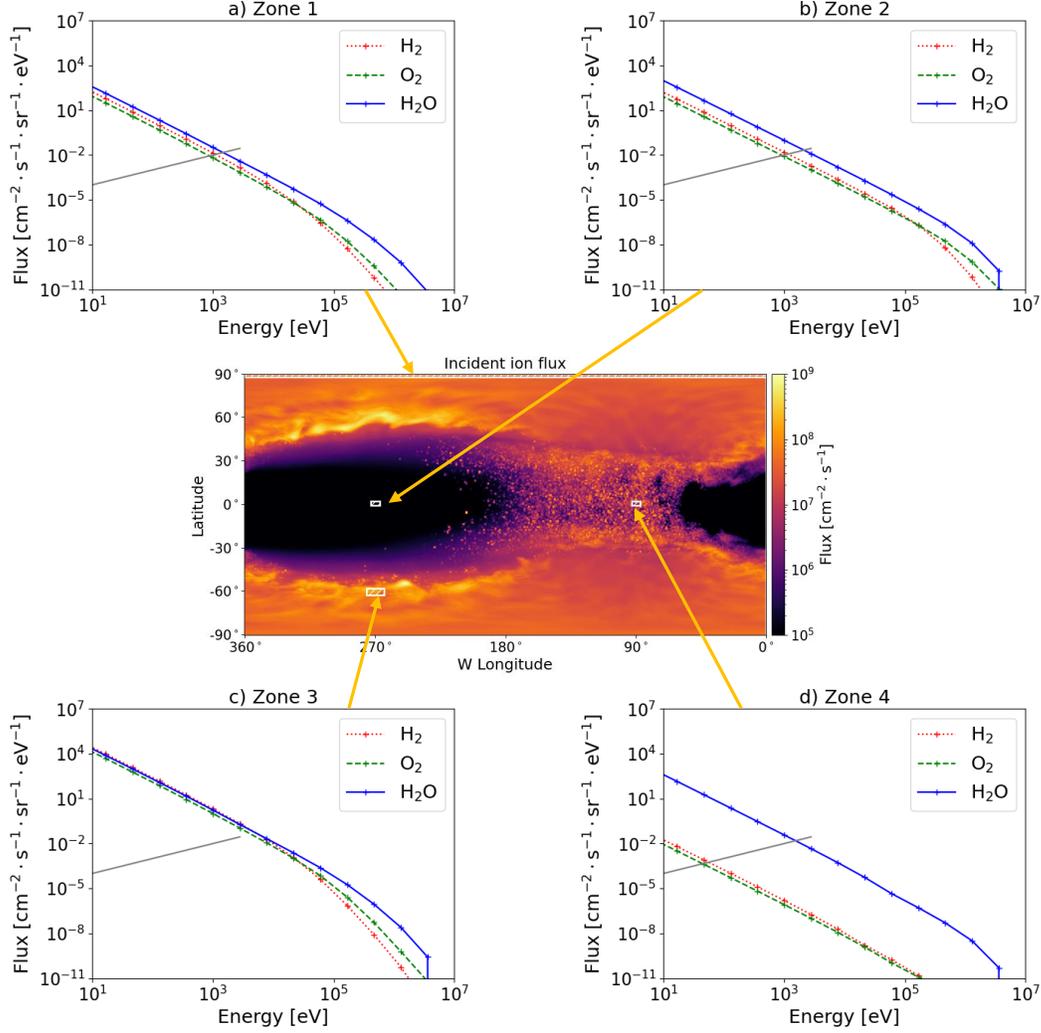
Figure 8. Flux of sputtered H_2, O_2 and H_2O at four locations on the surface of Ganymede, integrated over incident angles and energies. White rectangles show the area over which the flux was averaged and correspond to the footprint of JNA at 1000 km above the surface of Ganymede. The solid gray line indicates the one count level of the JNA instrument for the energy range that JNA can measure (10 eV to 3.3 keV).

sequently, the ENA flux is lower in Zone 1. Zone 2 is centered around the sub-solar point, where the incident ion flux is three orders of magnitude lower than in Zone 3, as Ganymede's magnetic field prevents low energy Jovian plasma from accessing the surface. The ion flux in Zone 2 is therefore dominated by energetic species, which are more efficient at sputtering H_2O than H_2 and O_2 (see Figure 3). This explains why the sputtered H_2O flux is about one order higher than that of H_2 and O_2. Figure 8d shows an even larger gap between H_2O fluxes and H_2 and O_2 fluxes but there the cause is different. In Zone 4, the ion flux includes contributions of both thermal and energetic species. However, Zone 4 is located on the nightside where the surface temperature is 80 K, which leads to much lower sputtering yields for H_2 and O_2 than for H_2O.

As sputtered fluxes of H_2O are not temperature-dependent, their variation is a direct result of the differences between the incident sputtering populations. We plot them
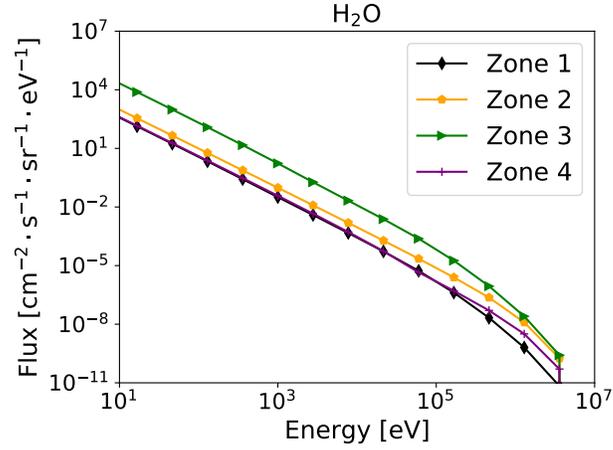
Figure 9. Sputtered H_2O fluxes for all zones, integrated over incident species, energies, and angles.

³²⁰ together in Figure 9 for easier comparison. The intense ion flux in Zone 3 yields the high-
³²¹ est sputtered H_2O flux at all energies. Despite their different incident ion populations,
³²² Zone 1 and Zone 4 shows similar sputtered H_2O fluxes, up to 100 keV, above which en-
³²³ ergetic ions in Zone 4 result in a higher sputtered H_2O flux. While the incident ion flux
³²⁴ is lowest in Zone 2, the sputtered H_2O flux is higher than in Zone 1 and 4, for the rea-
³²⁵ sons mentioned above.
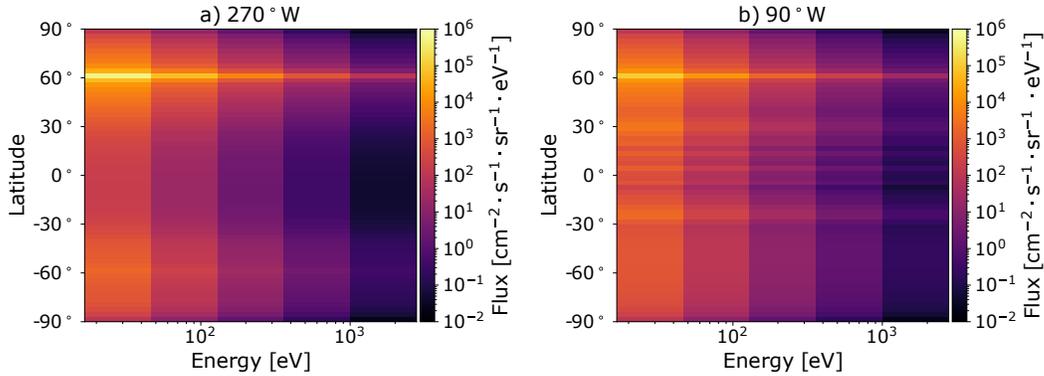
Figure 10. Flux of sputtered ENAs integrated over sputtered species and incident species, an-
gles, and energies, shown for each latitude zone corresponding to a JNA footprint along a) 270°W
and b) 90°W.

³²⁶   In Figure 10, JNA is also assumed to orbit Ganymede at 1000km along the 90°W
³²⁷ and 270°W meridians but here we divide the orbit track into 120 zones covering 3° each
³²⁸ in latitude (corresponding to the width of JNA's center pixel in elevation). The flux for
³²⁹ each latitude zone is shown, for the energy range that JNA can measure (between 10 eV
³³⁰ and 3.3 keV) and integrated over sputtered species and incident species, angles, and en-
³³¹ ergies.

³³²   Generally, Figure 10 shows that the flux of sputtered ENAs varies by about four
³³³ orders of magnitude along the simulated trajectory of JNA for all energy bins in the JNA

measuring range. Again we observe that the variability of the sputtered neutral flux reflects that of the incident ion flux. Along the 270°W meridian (on the dayside/trailing hemisphere), the flux gradually decreases from latitudes ±60° to the equator where the flux is minimal and four orders of magnitude lower than at +60°. We do note that while the results from Poppe's backwards-Liouville tracing model show a significantly higher flux at +60° than at −60° for the G8 flyby, a significant difference between the northern and southern hemispheres is not expected in reality.

On the 90°W meridian (on the nightside/leading hemisphere), the gradual decrease is interrupted at latitudes ±30° by narrow bands of intense sputtering, reflecting the narrow bands of intense ion precipitation in the incident flux (Figure 1). At these latitudes, the relative variation between adjacent zones is a factor of two to three. These large latitudinal variations show that ion precipitation patterns at the surface of Ganymede can be retrieved by remotely measuring ENAs.

4 Discussion and Conclusions

We presented a new method to simulate the sputtering process at Ganymede, in order to estimate sputtered ENA fluxes to be observed by the Jovian Neutrals Analyzer, an ENA sensor to be deployed at Ganymede by ESA's upcoming JUICE mission. Our method combines three sputtering yield models to calculate the yield of H_2, O_2, and H_2O separately. Our global sputtering rates show that H_2 and O_2 account for half of the total global sputtering rate from Ganymede. Our total global sputtering rate is in agreement with previous works, but by separating each species we were able to calculate their energy spectra, which is necessary in order to simulate JNA measurements. Indeed, JNA's mass resolution only allows it to distinguish between H and heavier species, but information about the mass and origin of heavier species may be retrieved by looking at their energy spectra. In this study we used a Thompson-Sigmund law to calculate the energy spectra of sputtered particles. In future work, the backscattering process should also be considered to more accurately simulate the energy spectra of ENAs to be observed at Ganymede. Backscattering is another process caused by precipitating ions, in which the impinging ion is neutralized (usually) and reflected by the surface. Measurements both in laboratories and in space suggest that backscattered particles would have energies in the range that JNA can measure, but distributed according to a Maxwell-Boltzmann-like law rather than the Thompson-Sigmund law applicable to sputtering (Futaana et al., 2012; Wieser et al., 2016). Backscattering yields are not well modeled, although studies by Wieser et al. (2016) and Futaana et al. (2012) suggest that a yield of about 0.1-0.2 can be applied for low ($\sim$keV) energies. The majority of ENAs in the 10 eV - 1 keV range are produced by the sputtering process (by high energy particles), so the backscattered contribution to the total ENA spectra is expected to be small. Nevertheless, the different shape of their spectra may allow us to distinguish backscattered ENAs from sputtered ENAs.

We also provided an estimate of expected JNA count rates and simulated the sputtered ENA flux at different locations along the track of a simplified orbit of the JUICE spacecraft. Our results show large latitudinal variations in sputtered ENA flux, demonstrating that JNA will be able to identify ion precipitation patterns by measuring ENAs. Future work will use realistic orbits of the JUICE spacecraft as well as JNA's calibrated instrument response, unavailable at the time of this study.

In conclusion, our results provide insight into the appearance of the data when JNA measures ENAs at Ganymede, as well as how the instrument should be operated optimally under limited power and data budget. The produced sputtering rate maps, energy spectra, and count rates in this study illustrate the capability of the ENA measuring technique to remotely map ion precipitation at Ganymede and provide clues for further potential ENA mapping in other icy bodies. Future work can easily use our model to pro-

duce more accurately simulated JNA spectra for different phases of the JUICE mission.
Such simulations are crucial for optimizing operations planning and make the most of
the limited integration time and data budget.

Acronyms

ENA Energetic Neutral Atoms
ESA European Space Agency
JNA Jovian Neutrals Analyzer
JUICE JUpiter ICy Moon Explorer
PEP Particle Environment Package

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Figure 1.

Incident ion flux
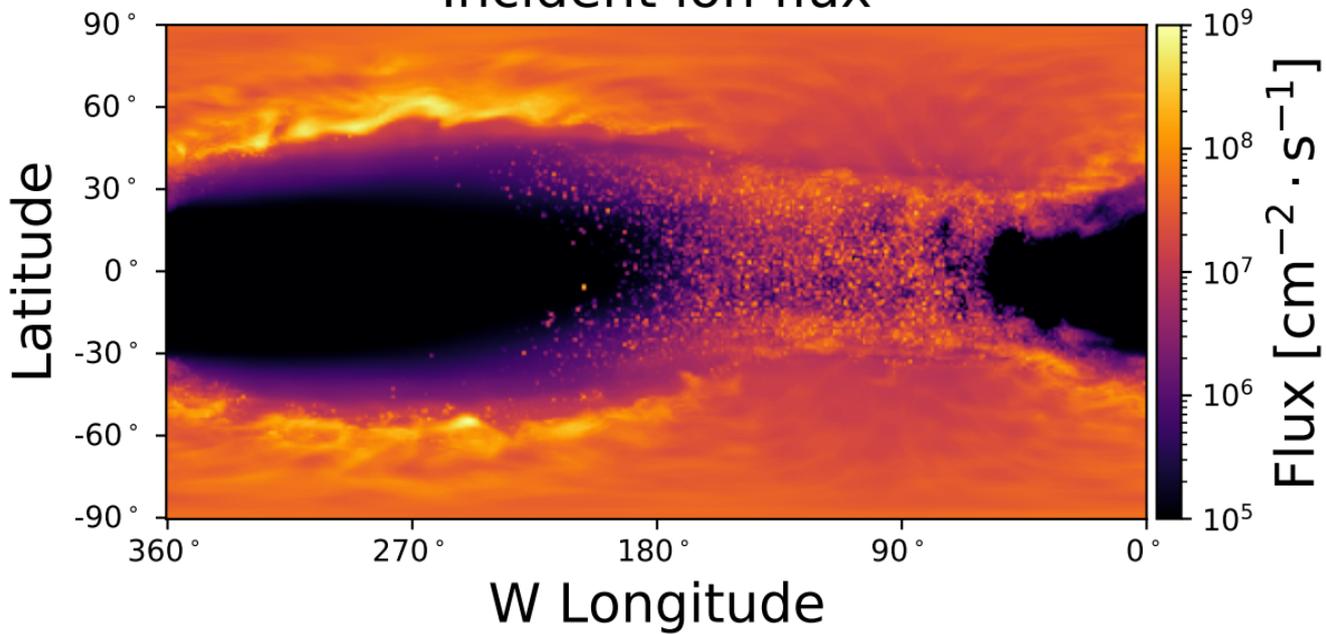
Figure 2.

Incident ion flux

Figure 3.

a) Incident Hydrogen   b) Incident Oxygen   c) Incident Sulfur

Figure 4.

a) Emitted H₂O — Incident species: H, O, S. b) Incident H — Emitted species: H₂, O₂, H₂O. Plots of $S(K)$ versus Energy of emitted ENA [eV].

Figure 5.

a) Sputtered H₂ flux

b) Sputtered O₂ flux

c) Sputtered H₂O flux
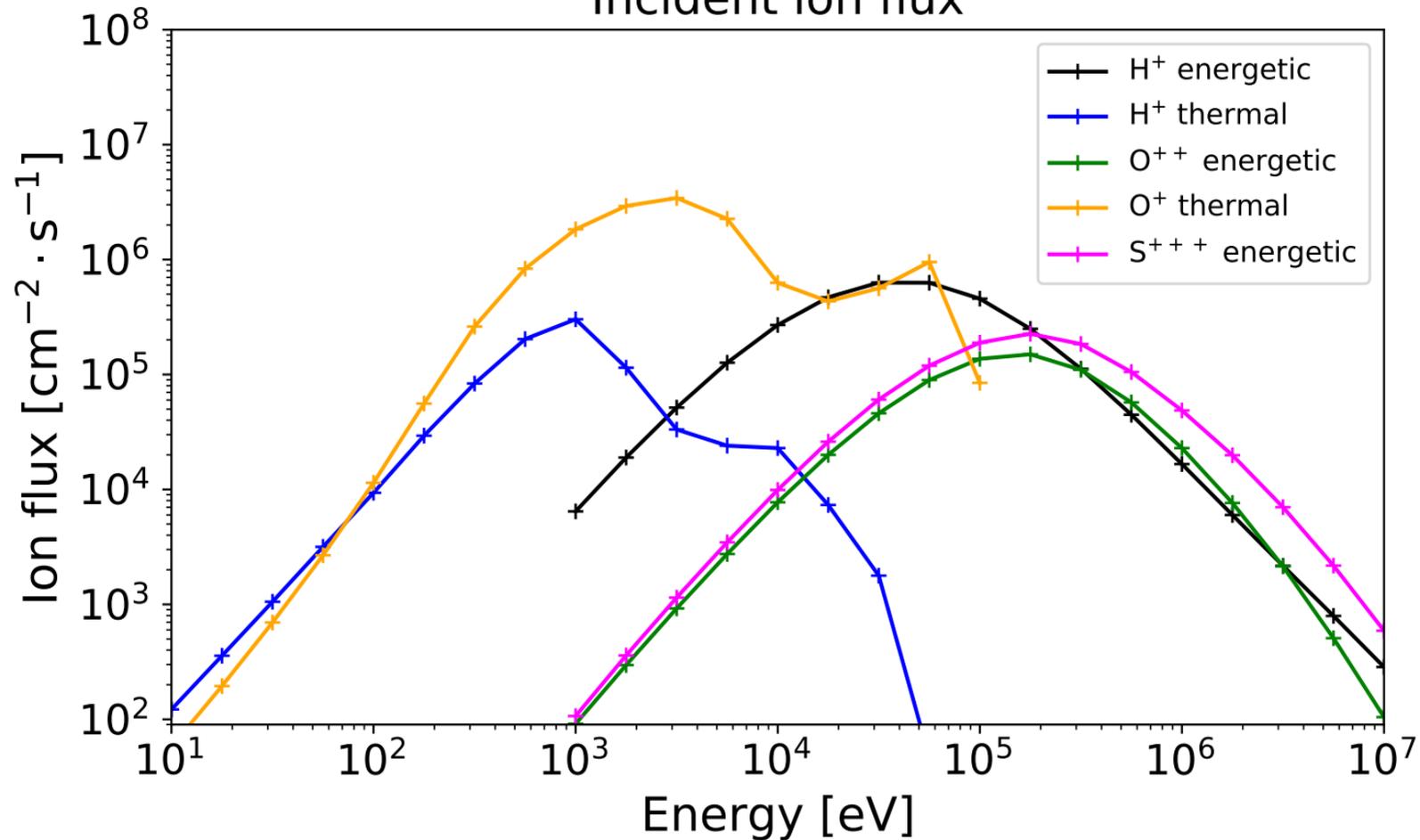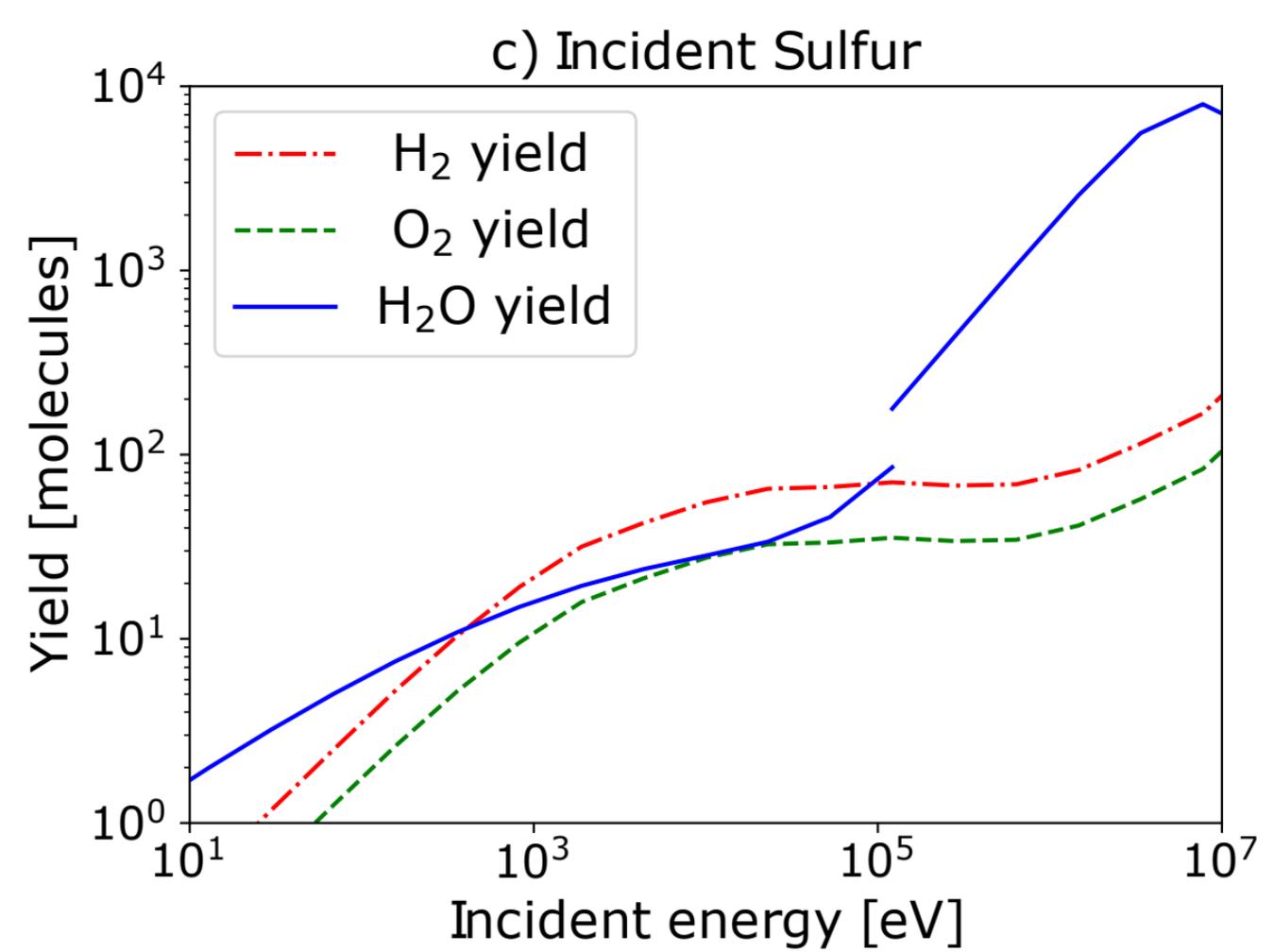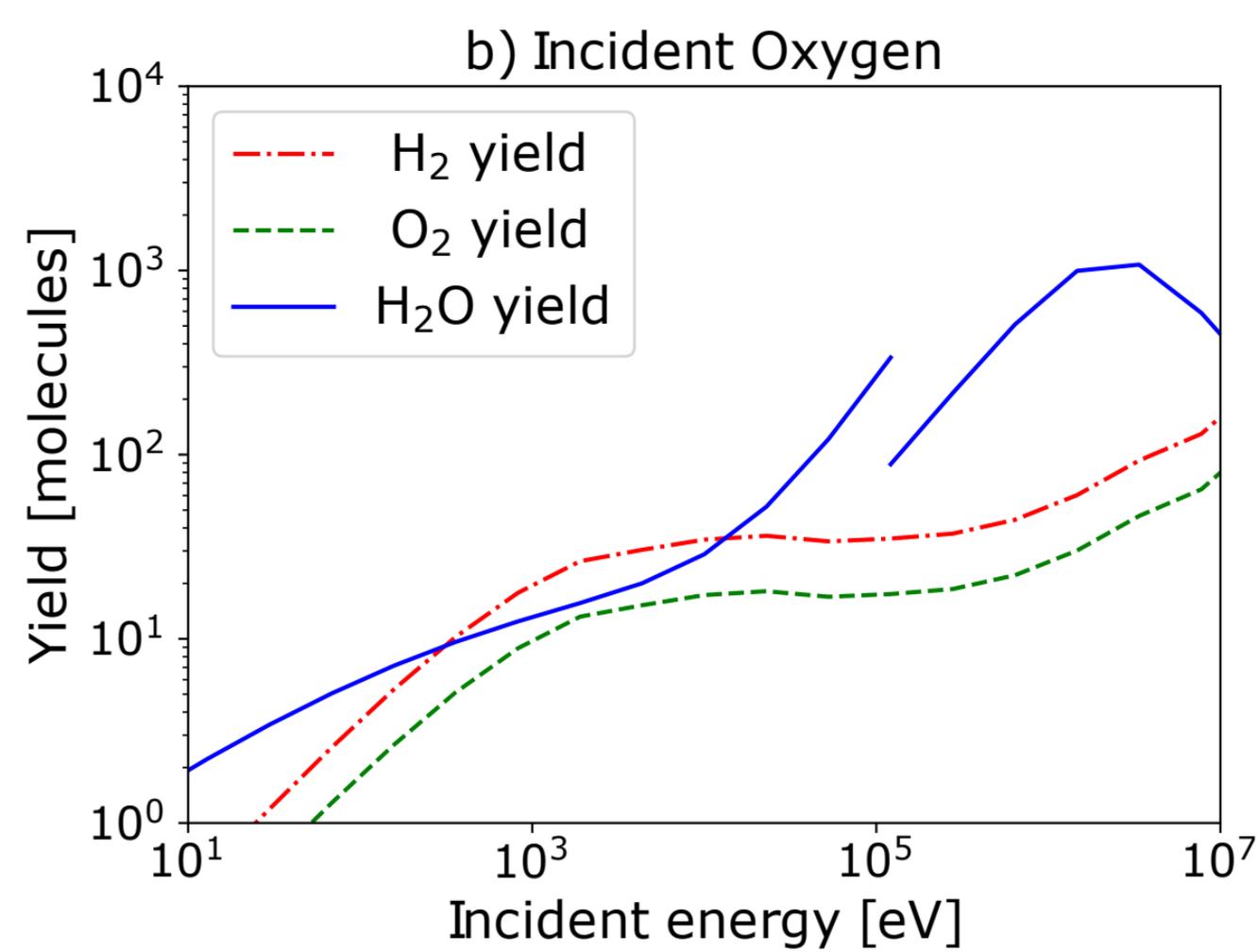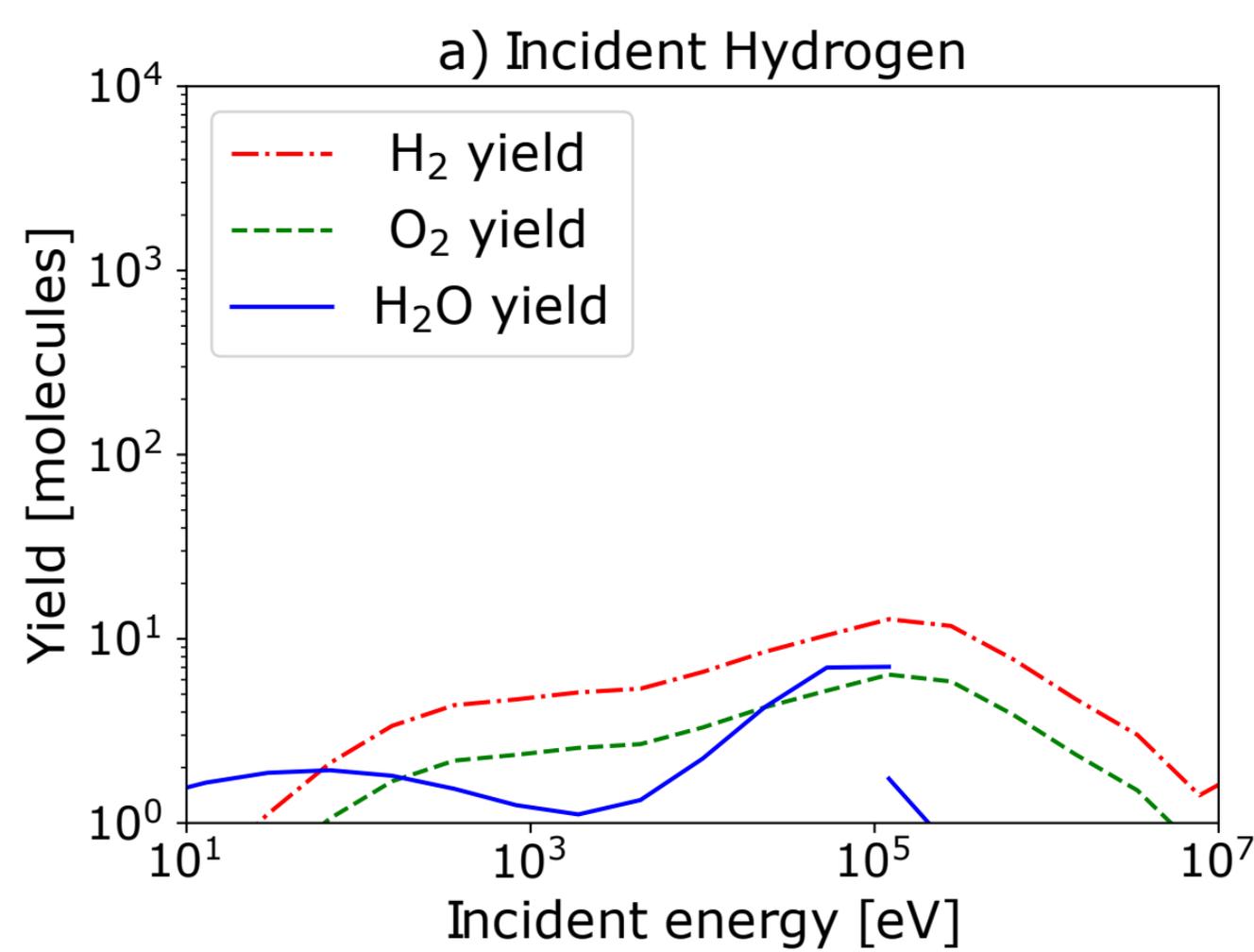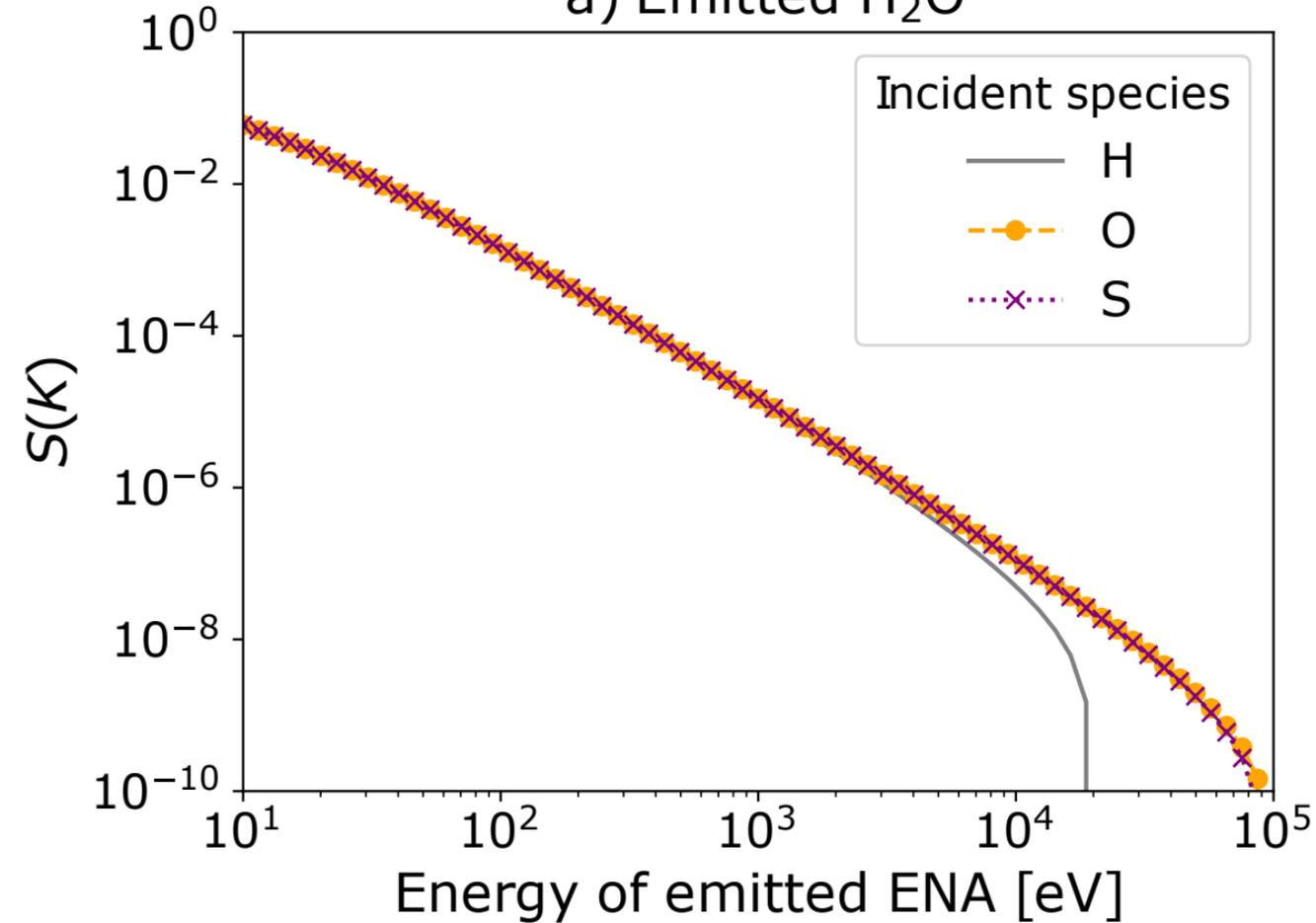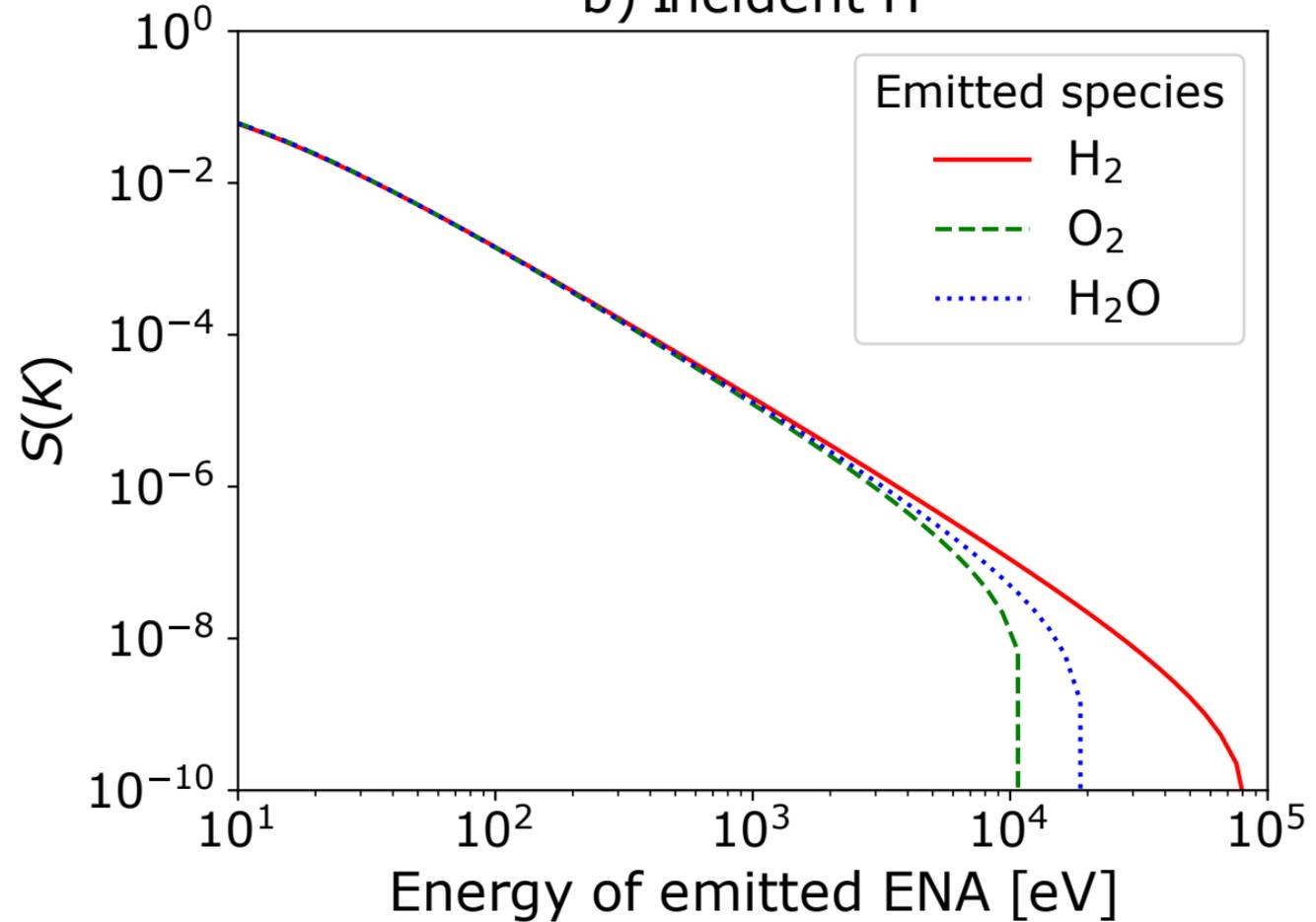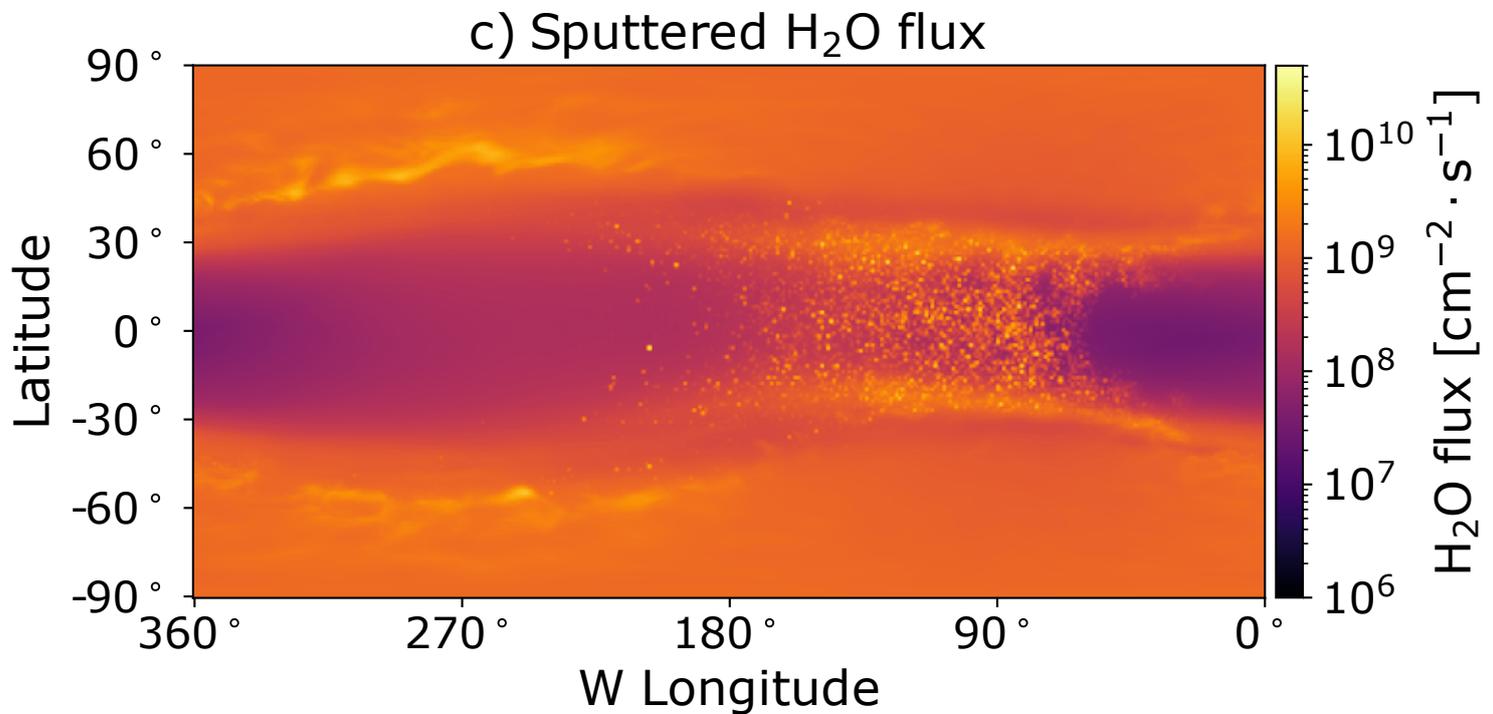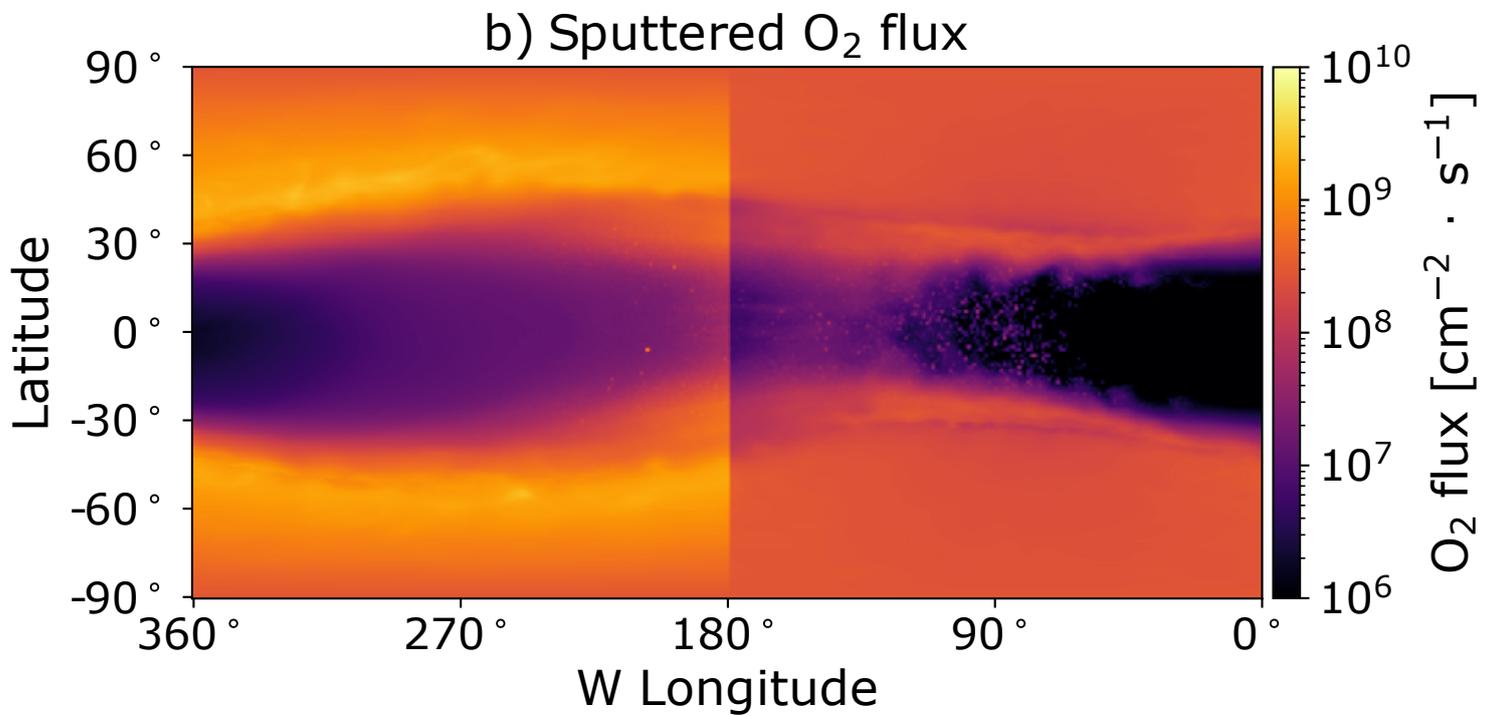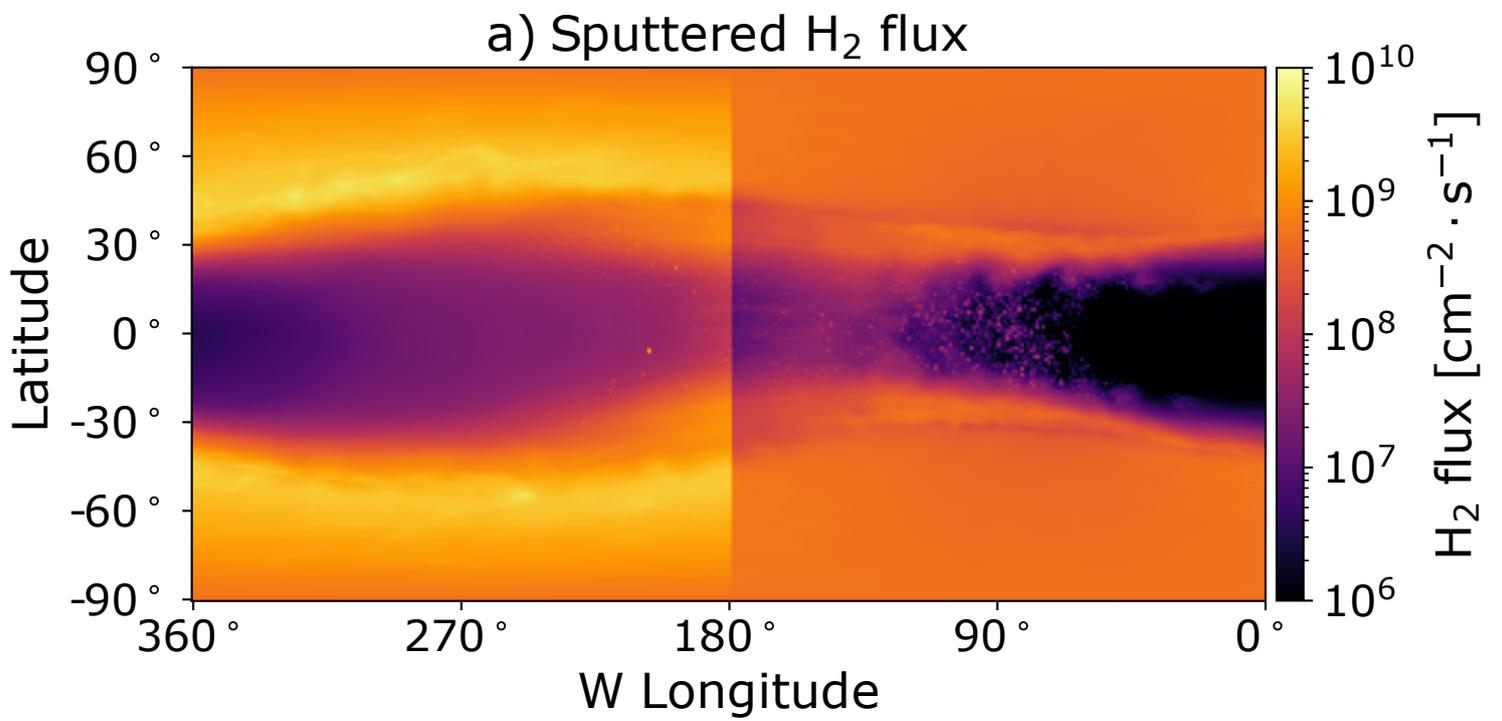
Figure 6.

a) H₂ — b) O₂ — c) H₂O — d) All ENA fluxes

Legend (panels a, b, c):
- H⁺ thermal
- O⁺ thermal
- H⁺ energetic
- O⁺⁺ energetic
- S⁺⁺⁺ energetic

Legend (panel d):
- Total H₂
- Total O₂
- Total H₂O

Axes: Flux [cm⁻²·s⁻¹·sr⁻¹·eV⁻¹] vs Energy [eV]

Figure 7.

Figure 8.

a) Zone 1   b) Zone 2

Incident ion flux

c) Zone 3   d) Zone 4

Figure 9.

Figure 10.

a) 270°W   b) 90°W