

Multi-solver spectral-element and adjoint methods

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Key Points:

- Simultaneous construction of Fréchet and Hessian kernels on the fly based upon spectral-element and adjoint methods.
- Only about a 2-fold computational cost required for the simultaneous computation when compared to the computation of Fréchet kernels.
- Truncated-Newton full-waveform inversion can be performed efficiently based upon the multi-solver spectral-element and adjoint methods.

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Abstract

The spectral-element method (SEM) for simulating wave propagation is widely used with adjoint methods for full-waveform inversion. Typically, SEM is used to compute forward and adjoint wavefields, which is then applied to evaluate the Fréchet derivatives for updating the seismic structural model. The Hessian is rarely computed as the high computational and storage costs, although it can improve the accuracy of the model update and model convergence. Instead the approximate Hessian is determined, which is obtained with less computational effort. We present a method for simultaneously constructing Fréchet and Hessian kernels on the fly, which we call Multi-solver spectral-element and adjoint methods (Multi-SEM). Rather than storing all the wavefields, Multi-SEM is computed on the fly and requires only about a 2-fold computational cost when compared to the computation of Fréchet kernels. Numerical examples demonstrate the functionality of the method and the computer codes are provided with this contribution.

Plain Language Summary

Recent advances in high-performance computing and quantum computing mean that full-waveform inversions (FWIs) are now routinely performed to achieve high-resolution imaging of the interior structure of the Earth. Typically, these are done using first-order derivatives, known as Fréchet kernels. Second-order derivatives, known as Hessian kernels, can be used to speed up convergence and to determine higher resolution of small-scale features. However, the Hessian is not commonly computed due to computational challenges such as high storage needs and long run times related to reading and writing. We present the Multi-solver spectral-element and adjoint methods (Multi-SEM), which generalizes

36 the conventional spectral-element and adjoint methods from the computation of Fréchet
37 kernels into the simultaneous computation of Fréchet and Hessian kernels. The kernels
38 are computed on the fly, which means that only a double computational cost is required
39 in comparison to the computation of Fréchet kernels only without the need to store sev-
40 eral 4-D wavefields, saving several TB of memory. We present the Hessian Kernels for
41 two different models to demonstrate their potential for achieving higher accuracy. Multi-
42 SEM improves the capability of FWI to image Earth structure, particularly in regions
43 characterized by small scale heterogeneities such as subductions zones.

44 **1 Introduction**

45 During the past twenty years the spectral-element method (SEM) (e.g., Patera, 1984;
46 Maday & Patera, 1989) has been widely used in the seismology community for simulat-
47 ing the propagation of surface and body waves in the Earth (e.g., Komatitsch & Tromp,
48 1999, 2002a, 2002b; Komatitsch et al., 2002c; Chaljub & Valette, 2004; Tromp et al., 2005;
49 Liu & Tromp, 2006; Chen et al., 2007; Tape et al., 2007; Chaljub et al., 2007; Liu & Tromp,
50 2008; Tromp et al., 2008; Fichtner et al., 2009; Tape et al., 2009; Peter et al., 2011; Liu
51 & Gu, 2012; Afanasiev et al., 2019), see Tromp (2020) for a review. Compared to other
52 solvers, the SEM is popular in seismology due to its great ability in handling complex
53 geometries and simulating surface waves with low numerical dispersion. Since 2005, the
54 adjoint method (e.g., Tarantola, 1984; Talagrand & Courtier, 1987) was successfully con-
55 nected with the SEM by Tromp et al. (2005), and has been used to compute the sensi-
56 tivity kernels with the forward and adjoint fields. For the elastic case, an implementa-
57 tion of little storage cost requires two simulations per event: a forward simulation of the

58 earthquake to the receivers, and another simulation carrying both the forward wavefield
59 and the adjoint wavefield simultaneously. In the latter simulation, the forward field is
60 reconstructed backward in time and the adjoint simulation is triggered by time-reversed
61 adjoint sources simultaneously at receivers. The computation of Fréchet kernels is achieved
62 via correlation of the reconstructed forward fields with the adjoint fields (e.g., Tromp et
63 al., 2008; Liu & Gu, 2012).

64 Computation and use of event-based Fréchet kernels from SEM and adjoint methods have
65 been performed in many studies. However, due to the high computational cost, the use
66 of Hessian kernels for one source and multiple receivers is not common even though the
67 theory was presented (e.g., Fichtner & Trampert, 2011). In practice, authors may use
68 the limited-memory Broyden-Fletcher-Goldfarb-Shanno (L-BFGS) algorithm (e.g., Noc-
69 cedal, 1989; Liu & Nocedal, 1989; Zou et al., 1993; Nocedal & Wright, 1999), which com-
70 putes the product of the inverse approximate Hessian and the gradient to estimate model
71 update using gradients and models from previous iterations. This solution is popular due
72 to its numerical efficiency. One competitive algorithm called truncated-Newton optimiza-
73 tion (e.g., Nash, 1985; Grippo et al., 1989; Nash & Nocedal, 1991; Nash, 2000) has been
74 well-documented in exploration seismology for full-waveform inversion (see e.g., Métivier
75 et al., 2014, 2017; Pan et al., 2017; Yang et al., 2018; Matharu & Sacchi, 2019), and it
76 has been demonstrated that it produces better results than the L-BFGS algorithm in
77 multi-parameter full-waveform inversion due to its mitigation in inter-parameter trade-
78 off, such as inversions for v_p , v_s , density, attenuation, and anisotropy or some of them.
79 Significant differences between the approximate Hessian and the full Hessian were ob-

80 served (Fichtner & Trampert, 2011). The truncated-Newton method is rarely used in
81 earthquake seismology due to the computational issue to construct the Hessian kernels.
82 However, efficient solutions constructing the Hessian kernels may make the truncated-
83 Newton method more appealing for full-waveform inversion (e.g., Tromp, 2020) or ad-
84 joint tomography (e.g., Tape et al., 2007, 2009).

85 The Hessian kernels can be computed by the method of Fichtner and Trampert (2011)
86 using pre-existing implementations of the adjoint tomography. One such approach in-
87 volves storing the forward and adjoint wavefields at all or sub-sampled time steps for later
88 determination of the Fréchet and Hessian kernels. This practically leads to big challenges
89 for the Hessian construction because of huge disk storage requirements in saving forward
90 and adjoint fields as well as their perturbations. Practical simulations may involve tens
91 to hundreds of millions of grid points and tens of thousands of time steps for each wave-
92 field. For computing the Hessian kernels, at least four sets of such wavefields are required
93 (Fichtner & Trampert, 2011). The disk storage may become a daunting issue even af-
94 ter sub-sampling schemes are introduced.

95 Another type of method to compute the Hessian is the scattering integral (SI) method
96 (e.g., Chen, Zhao, & Jordan, 2007; Chen, Jordan, & Li, 2007; Chen, 2011; Lee et al., 2014),
97 which is closed related to the adjoint methods (Tromp et al., 2005, 2008). The relative
98 computational efficiency of the two types of methods for the kernel calculation and in-
99 version depends on the overall problem geometry, in particular the ratio of the number
100 of sources to receivers (see Chen, Jordan, & Li, 2007; Lee et al., 2014). The SI method
101 may be more computationally efficient when the number of sources is comparable or larger

102 than the number of receivers. But when the number of receivers is large or the compu-
103 tation domain is expansive or shorter periods seismic waves are inverted, the computa-
104 tion and storage demand for the SI may become a daunting issue, in particular when the
105 updated structure is far away from the reference model where the Hessian for individ-
106 ual measurement needs to be recomputed in each iteration of the inversion. The disk stor-
107 age can be another challenging issue. For example in the Southern California crustal in-
108 version presented by Lee et al. (2014), the peak disk storage during the SI inversion was
109 about 39 Tb in addition to the huge input/output (I/O) overhead.

110 We present a numerically efficient method to compute Hessian kernels for one event, which
111 we call Multi-solver spectral-element and adjoint method (Multi-SEM). It is different from
112 the aforementioned wavefield storage techniques. Further developed from the adjoint meth-
113 ods in Tromp et al. (2005); Liu and Tromp (2006, 2008) where sensitivity kernels are cal-
114 culated from the simultaneous computation of adjoint wavefield and back-reconstructed
115 forward field, the Multi-SEM resolves the storage issue by constructing the Fréchet and
116 Hessian kernels on the fly for each or incremental time step through five SEM solvers.
117 Since only one time-step of both wavefields and the integrated kernels are kept in mem-
118 ory, the Multi-SEM is cheap in memory and easy to realize on present-day hardware with
119 only limited storage required as that of adjoint methods (Tromp et al., 2005; Liu & Tromp,
120 2006; Tromp et al., 2008), e.g., storing for the last frame of the forward fields. The com-
121 putation of the Hessian kernels by Multi-SEM requires only about two times the CPU
122 time compared to the computation of the Fréchet kernels alone. The Multi-SEM method
123 can be implemented on pre-existing spectral-element solvers such as the SPECFEM2D

124 (<https://github.com/geodynamics/specfem2d>), where one just slightly rearranges the
125 coding structure by coupling two solvers simultaneously for the forward simulation and
126 coupling five solvers simultaneously for the simultaneous backward and adjoint simula-
127 tion. Although five solvers are coupled and used, memory requirement could be designed
128 to be as small as possible since only one time-step of both wavefields and the integrated
129 kernels are kept in the temporary memory. The computational cost is slightly reduced
130 over individual five solver runs as all solvers share the same mesher database files except
131 those describing model material properties for the model and its update as discussed in
132 Section 3.

133 In this paper, we first review the theory on Fréchet and Hessian kernels and then present
134 the Multi-SEM method. Results for Fréchet and Hessian kernels are presented and dis-
135 cussed for 2-D synthetic models. The related codes are published in the public domain
136 for dissemination.

137 **2 Theory**

138 **2.1 Fréchet kernels**

139 Fréchet kernels, gradient or first-order derivatives of the seismic data functional, χ , can
140 be used to update the structural model from a chosen initial model via local optimiza-
141 tion rather than a costly global search. When the initial model is chosen sufficiently close
142 to the global minimum and when the source term is relatively accurate, the final model
143 from the local optimizations may also approach the true model. By perturbing the mea-
144 surements as $\delta\chi$ with respect to an isotropic model \mathbf{m} , we have (also see Tromp et al.,

145 2005)

$$146 \quad \delta\chi = \int_V \overline{K}_m \frac{\delta\mathbf{m}}{\mathbf{m}} d^3\mathbf{x} = \int_V K_m \delta\mathbf{m} d^3\mathbf{x}, \quad (1)$$

147 where $\overline{K}_m = K_m \mathbf{m}$. The \overline{K}_m or K_m denotes the *Fréchet* kernels and V denotes the
 148 model volume. Here we omit the spatial and temporal dependencies of the kernels for
 149 simplicity unless stated otherwise. In principle, the generic K_m can be expressed into
 150 different components depending on the choice of model parameterization (See Section
 151 1 of the Supporting Information). For simplicity, we only show the case for model pa-
 152 rameterization given by $\mathbf{m} = (\rho, \alpha, \beta)$, where ρ denotes the density and α and β denote
 153 the compressional and shear wave speeds. The kernel applied to the model perturbation
 154 in eq.(1) can be further expressed as

$$155 \quad K_m \delta\mathbf{m} = \begin{pmatrix} K'_\rho & K_\alpha & K_\beta \end{pmatrix} \begin{pmatrix} \delta\rho \\ \delta\alpha \\ \delta\beta \end{pmatrix}, \quad (2)$$

156 where $\delta\mathbf{m} = (\delta\rho, \delta\alpha, \delta\beta)^T$. As the computation of Fréchet kernels relies on the forward
 157 and the adjoint fields, we rewrite the Fréchet kernels as a function of the forward and
 158 adjoint fields

$$159 \quad \begin{pmatrix} K'_\rho \\ K_\alpha \\ K_\beta \end{pmatrix} = \begin{pmatrix} K'_\rho(\mathbf{s}^\dagger, \ddot{\mathbf{s}}) \\ K_\alpha(\mathbf{s}^\dagger, \mathbf{s}) \\ K_\beta(\mathbf{s}^\dagger, \mathbf{s}) \end{pmatrix}, \quad (3)$$

160 where \mathbf{s} and \mathbf{s}^\dagger are the forward and adjoint displacement fields, and $\ddot{\mathbf{s}}$ is the second-order
 161 time derivative of \mathbf{s} , i.e., the forward acceleration field. In practice, the field storage method

162 and/or the forward-field back-reconstruction method may be used to compute the Fréchet
 163 kernels (see Section 1 of the Supporting Information).

164 2.2 Hessian kernels

165 2.2.1 Components of Hessian kernels

166 Similar to the first-order form of the Fréchet kernels as shown in eq. (1), the second-order
 167 form or the Hessian operator can be written as (see Fichtner & Trampert, 2011)

$$168 \quad H(\delta\mathbf{m}_1, \delta\mathbf{m}_2) = \int_V K_m^1 \delta\mathbf{m}_2 d^3\mathbf{x} = \int_V (H_a + H_b + H_c) \delta\mathbf{m}_2 d^3\mathbf{x}, \quad (4)$$

169 where $K_m^1 = H_a + H_b + H_c$ denotes the Hessian kernels. Based upon the work of Fichtner
 170 and Trampert (2011), we rewrite each part of the product as

$$171 \quad H_a(\rho, \alpha, \beta) = \begin{pmatrix} K'_\rho(\mathbf{s}^\dagger, \delta\ddot{\mathbf{s}}) \\ K_\alpha(\mathbf{s}^\dagger, \delta\mathbf{s}) \\ K_\beta(\mathbf{s}^\dagger, \delta\mathbf{s}) \end{pmatrix}, H_b(\rho, \alpha, \beta) = \begin{pmatrix} K'_\rho(\delta\mathbf{s}^\dagger, \ddot{\mathbf{s}}) \\ K_\alpha(\delta\mathbf{s}^\dagger, \mathbf{s}) \\ K_\beta(\delta\mathbf{s}^\dagger, \mathbf{s}) \end{pmatrix}, \quad (5)$$

$$172 \quad H_c(\rho, \alpha, \beta) = \begin{pmatrix} \rho^{-1} K_\alpha(\mathbf{s}^\dagger, \mathbf{s}) \delta\alpha + \rho^{-1} K_\beta(\mathbf{s}^\dagger, \mathbf{s}) \delta\beta \\ \rho^{-1} K_\alpha(\mathbf{s}^\dagger, \mathbf{s}) \delta\rho + \alpha^{-1} K_\alpha(\mathbf{s}^\dagger, \mathbf{s}) \delta\alpha \\ \rho^{-1} K_\beta(\mathbf{s}^\dagger, \mathbf{s}) \delta\rho + \beta^{-1} K_\beta(\mathbf{s}^\dagger, \mathbf{s}) \delta\beta \end{pmatrix}. \quad (6)$$

where $\delta\mathbf{s}$ and $\delta\mathbf{s}^\dagger$ denote the perturbed forward and adjoint field due to model pertur-
 bation $\delta\mathbf{m}_1 = \delta\mathbf{m} = (\delta\rho, \delta\alpha, \delta\beta)^\top$. For simplicity, we use $\delta\mathbf{m}$ as the model perturbation
 from this point on. Eq. (5)-(6) show a link between the Hessian kernels (e.g., Fichtner
 & Trampert, 2011) and the Fréchet kernels (e.g., Tromp et al., 2005). It implies that the
 implementation framework for computing the Fréchet kernel can be used to compute the

Hessian kernels by replacing the regular field with its associated perturbed field. \mathbf{H}_a can be computed with the implementation of eq. (3) by replacing the forward fields with the perturbed forward fields. \mathbf{H}_b practically includes two contributions, i.e.,

$$\mathbf{H}_b = \mathbf{H}_b^{(m)} + \mathbf{H}_b^{(s)}, \quad (7)$$

where

$$\mathbf{H}_b^{(m)}(\rho, \alpha, \beta) = \begin{pmatrix} K'_\rho(\delta\mathbf{s}_m^\dagger, \mathfrak{S}) \\ K_\alpha(\delta\mathbf{s}_m^\dagger, \mathbf{s}) \\ K_\beta(\delta\mathbf{s}_m^\dagger, \mathbf{s}) \end{pmatrix}, \mathbf{H}_b^{(s)}(\rho, \alpha, \beta) = \begin{pmatrix} K'_\rho(\delta\mathbf{s}_s^\dagger, \mathfrak{S}) \\ K_\alpha(\delta\mathbf{s}_s^\dagger, \mathbf{s}) \\ K_\beta(\delta\mathbf{s}_s^\dagger, \mathbf{s}) \end{pmatrix}. \quad (8)$$

174 The former is due to the perturbation of the model, and the latter is due to the pertur-
 175 bation of the adjoint source which is defined as *approximate Hessian kernels* in Fichtner
 176 and Trampert (2011). Both the $\mathbf{H}_b^{(m)}$ and $\mathbf{H}_b^{(s)}$ can be computed with the implementa-
 177 tion of eq. (3) by replacing the adjoint fields with the associated perturbed adjoint fields.
 178 The construction for \mathbf{H}_c is straightforward based upon the Fréchet kernel K_m and the
 179 perturbation of the model $\delta\mathbf{m}$.

180 2.3 Perturbed fields and perturbed model

181 As eq. (5)-(8) show that the Hessian kernels can be computed with the same implemen-
 182 tation framework as that for the Fréchet kernels by adjoint methods in eq. (3), any spectral-
 183 element package for wavefield generation can be redesigned and adapted to compute the
 184 Hessian kernels just with additional efforts to compute the perturbed forward fields $\delta\mathbf{s}$
 185 and the perturbed adjoint field $\delta\mathbf{s}^\dagger$ due to a model perturbation $\delta\mathbf{m}$ and the perturbed
 186 adjoint source.

187 **2.3.1 Perturbed fields for H_a component**

188 The H_a component of the Hessian kernels accounts for the perturbation of the forward
 189 field, $\delta\mathbf{s}$. If we denote the wavefield generated due to the perturbed model $\mathbf{m}_r + v\delta\mathbf{m}$
 190 as $\mathbf{s}(\mathbf{m}_r + v\delta\mathbf{m}; \mathbf{x}, t)$, we may obtain the perturbed forward field due to $v\delta\mathbf{m}$ as (see also
 191 Fichtner & Trampert, 2011)

$$192 \quad \delta\mathbf{s} = \lim_{v \rightarrow 0} \frac{1}{v} [\mathbf{s}(\mathbf{m}_r + v\delta\mathbf{m}; \mathbf{x}, t) - \mathbf{s}(\mathbf{m}_r; \mathbf{x}, t)], \quad (9)$$

193 where \mathbf{m}_r denotes the reference model, $r = 0, 1, 2, \dots, N$ represents the iteration num-
 194 ber, and \mathbf{m}_0 means the initial model. The same consideration applies to the perturbed
 195 acceleration field $\delta\ddot{\mathbf{s}}$ for density kernel computation. In practical application such as full-
 196 waveform inversion, the model perturbation can be estimated by using truncated New-
 197 ton optimization (see e.g., Métivier et al., 2014, 2017; Pan et al., 2017; Yang et al., 2018;
 198 Matharu & Sacchi, 2019). In the first iteration, the steepest descent method may be used
 199 to compute the model update. For more details of the $v\delta\mathbf{m}$ determination, please refer
 200 to Fichtner and Trampert (2011). The computation of H_a is straightforward if we use
 201 the field storage method. However, storage and I/O demands may be quite significant
 202 when the model size or the number of sources is large.

203 **2.3.2 Perturbed fields for H_b component**

204 The H_b component consists of two contributions. One is from the approximate Hessian
 205 kernels $H_b^{(s)}$ due to the perturbation of the adjoint source, and the other is from the $H_b^{(m)}$
 206 due to the perturbation of the model. To compute $H_b^{(s)}$, the approximate perturbed ad-

207 joint field may be calculated as

$$208 \quad \delta \mathbf{s}_s^\dagger = \mathbf{s}_s^\dagger(\mathbf{m}_r; \mathbf{x}, T-t) - \mathbf{s}^\dagger(\mathbf{m}_r; \mathbf{x}, T-t). \quad (10)$$

209 where the $\mathbf{s}_s^\dagger(\mathbf{m}_r; \mathbf{x}, T-t)$ field is generated by the adjoint source $\mathbf{f}^\dagger(\mathbf{m}_r + v\delta\mathbf{m}; \mathbf{x}, T-t)$, and $\mathbf{s}^\dagger(\mathbf{m}_r; \mathbf{x}, T-t)$ is generated by the adjoint source $\mathbf{f}^\dagger(\mathbf{m}_r; \mathbf{x}, T-t)$. The only
 210 difference between the two adjoint fields is the adjoint sources used since the former ac-
 211 counts for the perturbation of the adjoint source as a result of $v\delta\mathbf{m}$.

212
 213 The perturbed adjoint field for the $H_b^{(m)}$ calculation may be given by

$$214 \quad \delta \mathbf{s}_m^\dagger = \lim_{v \rightarrow 0} \frac{1}{v} [\mathbf{s}_m^\dagger(\mathbf{m}_r + v\delta\mathbf{m}; \mathbf{x}, T-t) - \mathbf{s}^\dagger(\mathbf{m}_r; \mathbf{x}, T-t)], \quad (11)$$

215 where the two adjoint fields $\mathbf{s}_m^\dagger(\mathbf{m}_r + v\delta\mathbf{m}; \mathbf{x}, T-t)$ and $\mathbf{s}^\dagger(\mathbf{m}_r; \mathbf{x}, T-t)$ are generated
 216 through the perturbed and unperturbed model from the same adjoint source $\mathbf{f}^\dagger(\mathbf{m}_r; \mathbf{x}, T-t)$. The adjoint sources may be different based on the choices of seismic data functional
 217 χ as discussed in Tromp et al. (2005). Thereafter, the total perturbed adjoint field is
 218

$$219 \quad \delta \mathbf{s}^\dagger = \delta \mathbf{s}_s^\dagger + \delta \mathbf{s}_m^\dagger. \quad (12)$$

220 ***2.3.3 Perturbed model for H_c component***

221 From eq. (6), it is clear that the computation of H_c relies on the Fréchet kernels and model
 222 perturbation. It has also been shown that H_c is non-zero when the model is parametrized
 223 as ρ , α , and β but zero when the model is given in another two sets of parameterization
 224 (Fichtner & Trampert, 2011). See also Section 2 of the Supporting Information.

225 **3 Implementation**

226 The computation of Hessian kernels relies on the regular and perturbed forward and ad-
 227 joint fields. Its implementation is relatively straightforward based on the wavefield stor-
 228 age method (WSM) (see Section 3 of the Supporting Information), where for each time
 229 step or incremental time step, the associated stored fields are read into temporary mem-
 230 ory for the kernel calculation, and this process is repeated until the end of simulation.

231 In this section, we show how the Hessian kernels is computed on the fly by the Multi-
 232 SEM. For the following examples we only consider cases with purely elastic models.

233 **3.1 Forward simulation**

234 Figure 1 shows the comparison between the single-solver SEM and the Multi-SEM for
 235 forward simulations. The Multi-SEM carries wavefield simulations for two models simul-
 236 taneously, e.g., \mathbf{m}_1 and \mathbf{m}_2 , instead of one model used by the single-solver SEM, where
 237 $\mathbf{m}_2 = \mathbf{m}_1 + v\delta\mathbf{m}$. In this case, the wavefields, including displacement \mathbf{s} , velocity \mathbf{v} , ac-
 238 celeration $\ddot{\mathbf{s}}$, and the boundary contribution \mathbf{b} (we use \mathbf{b} for generality since it is typ-
 239 ically the velocity fields or the velocity and force fields when the SEM domain is cou-
 240 pled with an external model) are computed for the two models at each time step. The
 241 displacement seismograms $\mathbf{s}(\mathbf{x}_r, t)$ are computed by a spatial interpolation of fields at
 242 the receiver \mathbf{x}_r at each time step. The grid-point locations and mesh topology database
 243 files are shared by the two models used simultaneously in the forward simulation with
 244 Multi-SEM, and only arrays/files related to model material properties such as ρ , α , and
 245 β need to be defined separately for the two models. The CPU and memory requirements

246 for Multi-SEM are about twice the cost in the single-solver SEM simulation. The for-
247 ward simulations either for the single-solver SEM or the Multi-SEM are designed to pro-
248 vide the absorbing boundary fields, the last state of the forward field, and the seismo-
249 grams at receivers, for the subsequent simulations.

250 **3.2 Simultaneous backward and adjoint simulations**

251 Simultaneous backward and adjoint simulations are widely used in many SPECFEM pack-
252 ages (<https://geodynamics.org/cig/software/>) to construct the Fréchet kernels on
253 the fly. A workflow for computing the Fréchet kernels by conventional single-solver SEM
254 method is shown in Figure S1 of the Supporting Information. For purely elastic mod-
255 els, the backward simulation is a time-reversed reconstruction of the forward field us-
256 ing the last state of the forward field as a starting point. The absorbing boundary con-
257 tributions saved in the forward simulation are re-injected into the backward simulation
258 as the forward field is reconstructed backward in time. The simulations for backward re-
259 construction and adjoint wavefield are performed simultaneously so that the correspond-
260 ing time slices of forward and adjoint wavefield can be accessed both in memory in or-
261 der to calculate Fréchet kernels. The same course is used in the Multi-SEM with five SEM
262 solvers instead of two (see Figure 2 and Figure S2). In this case, the regular, perturbed
263 forward fields and the regular, perturbed adjoint fields for the two models are simulta-
264 neously reconstructed and computed for a time step, so that the Fréchet and Hessian ker-
265 nels can be calculated on the fly as wavefield products are computed and integrated over
266 time steps (see Figure 2 and Figure S2). Although the five SEM solver engines are cou-
267 pled and use the same mesh database excluding \mathbf{m}_1 and \mathbf{m}_2 loaded externally. The mem-

268 ory cost is small since only one time step of the various fields and the integrated kernels
 269 are kept in memory compared to the wavefield storage methods. Each Fréchet kernel needs
 270 3 (1 in forward and 2 in adjoint) simulations, while the Multi-SEM carries 7 (2 in for-
 271 ward and 5 in adjoint) simulations for the simultaneous computation of Fréchet and Hes-
 272 sian kernels. During the adjoint simulation, the memory is not 5/2 times that of a reg-
 273 ular kernel simulation due to the shared memory for the same mesh database (exclud-
 274 ing the two models' material properties). The CPU hours will be less than 2.5 (5/2) times
 275 due to the shared mesher for all SEM solver. Most of the CPU time is spent comput-
 276 ing the strain and stress calculations.

277 4 Numerical Examples

278 4.1 Models

279 To test the numerical implementation of Multi-SEM, three models are considered in this
 280 study. First, a homogeneous 2D model (*Model 1*) of the size of 800 *km* in the horizon-
 281 tal direction and 360 *km* in the vertical direction and with density $\rho=2900 \text{ kg/m}^3$, com-
 282 pressional wave speed $\alpha=8000 \text{ m/s}$, and shear wave speed $\beta=4800 \text{ m/s}$, is used as a start-
 283 ing background model to generate initial wavefields and waveforms. We use the inter-
 284 nal mesher of the SPECFEM2D package to mesh the model with 400 elements in the
 285 horizontal direction and 360 elements in the depth direction. With 5×5 Gauss-Lobatto-
 286 Legendre (GLL) points used for each element in 2D, this leads to $\sim 500 \text{ m}/250 \text{ m}$ hor-
 287 izontal/vertical grid-point spacing for the model. The second and the third model are
 288 perturbed versions of the homogeneous model. The second model (*Model 2*) has an ad-

289 ditional +10% perturbation in α and β over a 10 km \times 10 km squared area centered at
290 the horizontal location of 335 km and depth of 135 km (see Figure 3c for the perturba-
291 tion location indicated by H_c). The third model (*Model 3*) comprises three anomalies
292 of the size of 8 km \times 10 km, centered at the same depth of 115 km and horizontally at
293 120 km, 180 km, and 240 km, respectively, with +10% perturbations in α and β (see Fig-
294 ure 3f for the three perturbation locations indicated by H_c). No density perturbation is
295 considered for the second and third model. These models are chosen to illustrate the dif-
296 ferences in the calculation of Hessian kernels between the single source-receiver pair and
297 single-source multiple-receiver case. The locations of the perturbations are indicated by
298 the H_c kernels in Figure 3.

299 4.2 Single source-receiver combination

300 We first examine the kernel calculation for a single source-receiver combination based
301 on *Model 1* and *Model 2*. We place a point source at $(x, z)=(100 \text{ km}, -260 \text{ km})$ with the
302 standard Ricker wavelet source-time function of dominant frequency of 0.5 Hz. A sin-
303 gle receiver is placed on the surface of the model at $(x, z)=(600 \text{ km}, 0 \text{ km})$. The simu-
304 lations use $dt = 0.01 \text{ s}$ and run for a total of 10,000 time steps.

305 To see the kernels over the model perturbation, we show here the Fréchet kernels for *Model*
306 *2*, and the Hessian kernels for *Model 1* and *Model 2*. The Fréchet kernels computed for
307 *Model 1* are shown in Figure S16 of the Supporting Information. The Multi-SEM com-
308 putes the Fréchet kernels shared the same solvers with conventional SEM (see Figure 2).
309 The first row of Figure 3 (Part I) shows the Fréchet kernel, the approximate Hessian ker-

310 nel, and the full Hessian kernel. A zoomed-in version around the perturbations is given
311 in the first row of Part II. Detailed descriptions about the kernels are given in the fig-
312 ure caption for Figure 3.

313 For the adjoint field calculations we use traveltimes adjoint sources with waveform win-
314 dow selected for the P phase, and the same procedure can be applied to the full wave-
315 forms. It takes the Multi-SEM method about a total of 31 mins with maximum mem-
316 ory usage of ~ 3.1 GB to simultaneously compute the Fréchet and Hessian kernels on a
317 standard laptop (with 2.3 GHz Dual-Core Intel Core i5 processor and 8GB 2133 MHz
318 LPDDR3 memory). In comparison, the computation of Fréchet kernel alone by the con-
319 ventional SEM and adjoint method takes about 13.5 mins with maximum memory usage
320 of 1.5 GB. Therefore in this case, all the quantities computed by Multi-SEM takes
321 ~ 2.29 times the CPU time and ~ 2.06 times the memory compared to the computation
322 of Fréchet kernels. The storage required for the Multi-SEM is small due to the on-the-
323 fly nature of the calculations, which takes about 1 GB disk space to store the absorb-
324 ing boundary fields, the last-state forward fields as well as the seismograms, while for
325 the wavefield storage method (WSM, see Section 3 of the Supporting Information), it
326 requires about 400 GB disk space to store these fields even without considering the den-
327 sity kernels.

328 **4.3 One source and three receivers**

329 We also show an example with one source and three receivers for the calculation of Hes-
330 sian kernels, where *Model 1* is used as the background model and Hessian kernels are

331 computed with respect to the perturbation in *Model 3*. The source is placed at $(x, z) =$
332 $(150 \text{ km}, -260 \text{ km})$ with the same source time function as in Section 4.2. Three receivers
333 are placed on the top surface of the model horizontally located at 100 km, 200 km, and
334 300 km, respectively. The total number of time steps and time interval are the same as
335 the example in Section 4.2.

336 The second row of Figure 3 (Part I) shows the Fréchet kernel, the approximate Hessian
337 kernels, and the full Hessian kernels computed for P phase on the seismograms. A zoomed-
338 in version of Figure 3 (Part I) around the perturbations is given in Figure 3 (Part II).

339 More detailed descriptions about the Fréchet and Hessian kernels are given in the fig-
340 ure caption. The computational cost for this example is almost the same as for that in
341 section 4.2 since the simulation cost is almost independent of the number of receivers.

342 There is one additional step in the window picking and computation of adjoint source,
343 which is much cheaper than the field calculations. A few selected time steps of the reg-
344 ular wavefields and their perturbations are shown in Figure S3 and Figure S4 in the Sup-
345 porting Information. The adjoint sources computed from the seismograms for \mathbf{m}_1 and
346 \mathbf{m}_2 are also provided there in Figure S5-S7. The key output files for the Multi-SEM pack-
347 age in the forward simulation and in the simultaneous backward and adjoint simulation
348 are presented in Figure S8.

349 **5 Discussions**

350 We found significant differences between the approximate Hessian kernels and the full
351 Hessian kernels for both the one- and multi-receiver case (Figure 3), as also noted in Fichtner

352 and Trampert (2011). Most notably, the amplitudes of the Hessian kernels can be up to
353 100% stronger than those of the approximate Hessian kernels within the red areas, as
354 areas also covered by H_a , $H_b^{(m)}$, and H_c in the full Hessian kernels and usually omitted
355 in the calculation of the approximate Hessian kernels. The greater positive values of the
356 Hessian in the vicinity of the perturbation suggest that the inversion using the Hessian
357 instead of the approximate Hessian will result in better illumination in the region of the
358 model perturbation, in addition to distributing them along the kernel.

359 In the multi-receiver case, we observe a similar higher amplitude in the Hessian kernels
360 near the three model perturbations (Figure 3f) (Part I and II); whereas, for the approx-
361 imate Hessian kernels, the sensitivity has high amplitudes around the middle anomaly
362 only. This again suggests that using the full Hessian kernels in the inversion will focus
363 model perturbations closer to the actual anomalies and the use of full Hessian kernels
364 would provide better resolution for smaller anomalies within the earth model.

365 The Hessian kernels are typically used with the Fréchet kernels for computing the model
366 updating based upon truncated Newton optimization (Nash, 1985; Grippo et al., 1989;
367 Nash & Nocedal, 1991; Nash, 2000), which has demonstrated better results over the L-
368 BFGS based optimization for multi-parameter full-waveform inversion (FWI) in explo-
369 ration seismology (e.g., Métivier et al., 2014, 2017; Pan et al., 2017; Yang et al., 2018;
370 Matharu & Sacchi, 2019). The truncated-Newton FWI, however, is rarely reported based
371 upon the spectral-element and adjoint methods in earthquake seismology due to the com-
372 putational and storage issues. We leave this for further investigation with the on-the-
373 fly Multi-SEM presented.

374 An important question remains as to whether the additional costs of the simultaneous
375 computation of the Fréchet and Hessian kernels at twice the computational cost can be
376 offset by more rapid convergence of the non-linear inversion. As high performance com-
377 puting becomes more accessible and efficient, this may not necessarily be as much of a
378 concern.

379 In addition to the expressions shown here, the approximate Hessian kernels and the full
380 Hessian kernels can be expressed in different model components as given in Section 2 of
381 the Supporting Information. For the anelastic case, the parsimonious storage method
382 (see Komatitsch et al., 2016) can be used which first performs forward simulation with
383 full attenuation to compute predictions to the seismic measurements and construct the
384 proper adjoint sources. The forward field is stored at selected checkpoints and reconstructed
385 back during the adjoint simulation to calculate the kernels for attenuating medium.

386 The ideas of Multi-SEM is not limited to the SEM and it can be also implemented in
387 solvers based on other methods such as finite difference. The Multi-SEM so far is designed
388 to compute Fréchet and Hessian kernels for single event. The Hessian kernels for all events
389 can be summed together as that of the misfit Fréchet kernels (Tromp et al., 2005). The
390 Multi-SEM method computes the Fréchet kernels, the approximate and the full Hessian
391 kernels simultaneously on the fly with only about a 2 fold computational cost when com-
392 pared to the computation for Fréchet kernels alone. The Multi-SEM also supports the
393 simultaneous computation of Fréchet and approximate Hessian kernels as selected func-
394 tion of the Multi-SEM, which is more computationally efficient since only three SEM solvers
395 need to be switched on in the simultaneous backward and adjoint simulation. To fur-

396 ther reduce the computational cost for multiple sources, one may use the source encod-
397 ing techniques (Tromp & Bachmann, 2019).

398 **6 Conclusions**

399 Considering the fast advance in high-performance computing in recent years and the in-
400 creasing demands in high-resolution multi-parameter imaging, we present the Multi-solver
401 spectral-element and adjoint methods (Multi-SEM) for simultaneously computing the
402 Fréchet and the Hessian kernels on the fly. The simultaneous access to Fréchet and Hes-
403 sian kernels may potentially provide better images and convergence properties for FWI
404 iterations than those in gradient-only-based FWI. In contrast to the wavefield storage
405 methods that require saving the wavefields for the duration of the simulation, Multi-SEM
406 constructs the Fréchet and Hessian kernels on the fly. The memory requirement for the
407 Multi-SEM is reasonably small since only a single time step of the wavefields and the
408 integrated kernels are kept in memory. The simultaneous computation by the Multi-SEM
409 requires only about a 2-fold computational time when compared to the computation of
410 Fréchet kernels.

411 The on the fly feature resolves the challenging storage and I/O issues for the Hessian ker-
412 nel calculation, and makes the use of full Hessian possible for multi-parameter full-waveform
413 inversion (FWI) based upon the spectral-element and adjoint methods. It potentially
414 provides a step forward for improving FWI to better image and understand earth struc-
415 ture, particularly in regions characterised by small scale heterogeneities such as subduc-
416 tions zones.

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428 N. Harmon; writing-review and editing, Q. Liu and D. Gajewski; supervision, C. Rychert,
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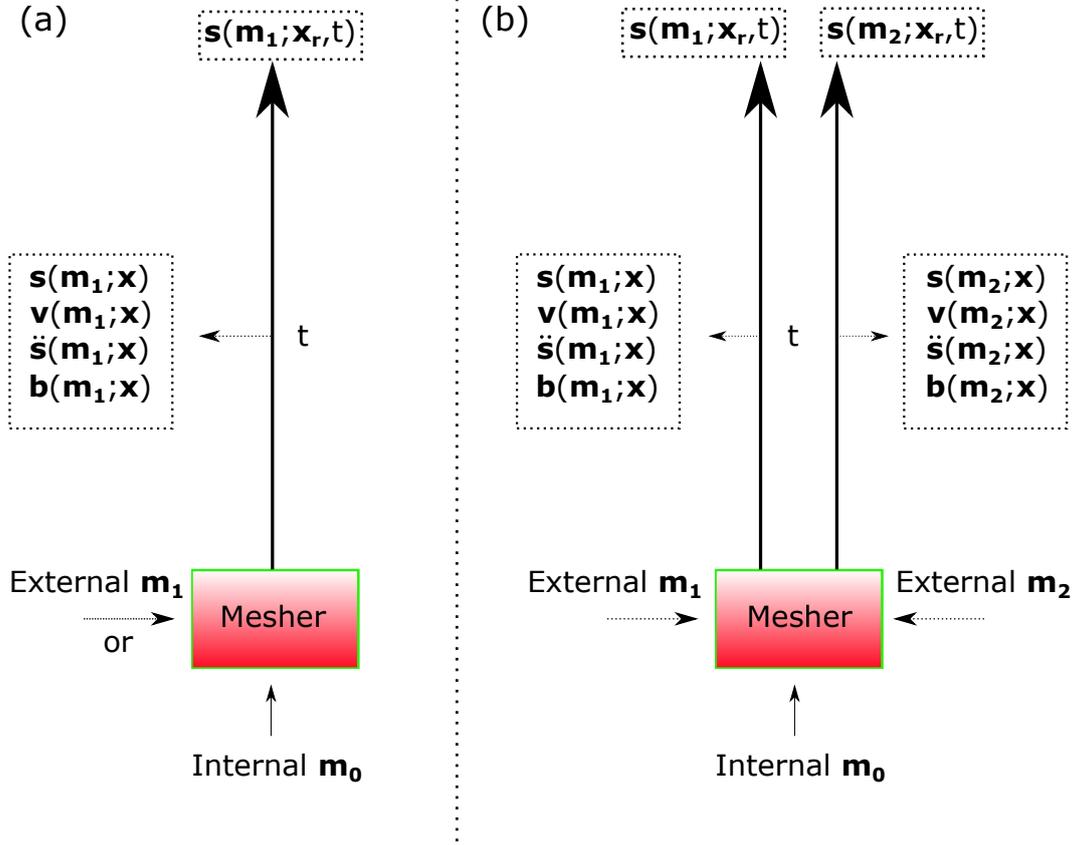


Figure 1. Sketch illustrating the workflow of forward simulation for Conventional SEM vs. Multi-SEM. (a) In Conventional SEM forward simulation, a single model is used and it is set either by the internal mesher (e.g., \mathbf{m}_0) or importing from external file (\mathbf{m}_1) after the mesher is set up. (b) In the Multi-SEM forward simulation, two models (\mathbf{m}_1 and \mathbf{m}_2) are imported into the internal mesher, where $\mathbf{m}_2 = \mathbf{m}_1 + v\delta\mathbf{m}$, and \mathbf{m}_0 will be omitted with models loaded externally.

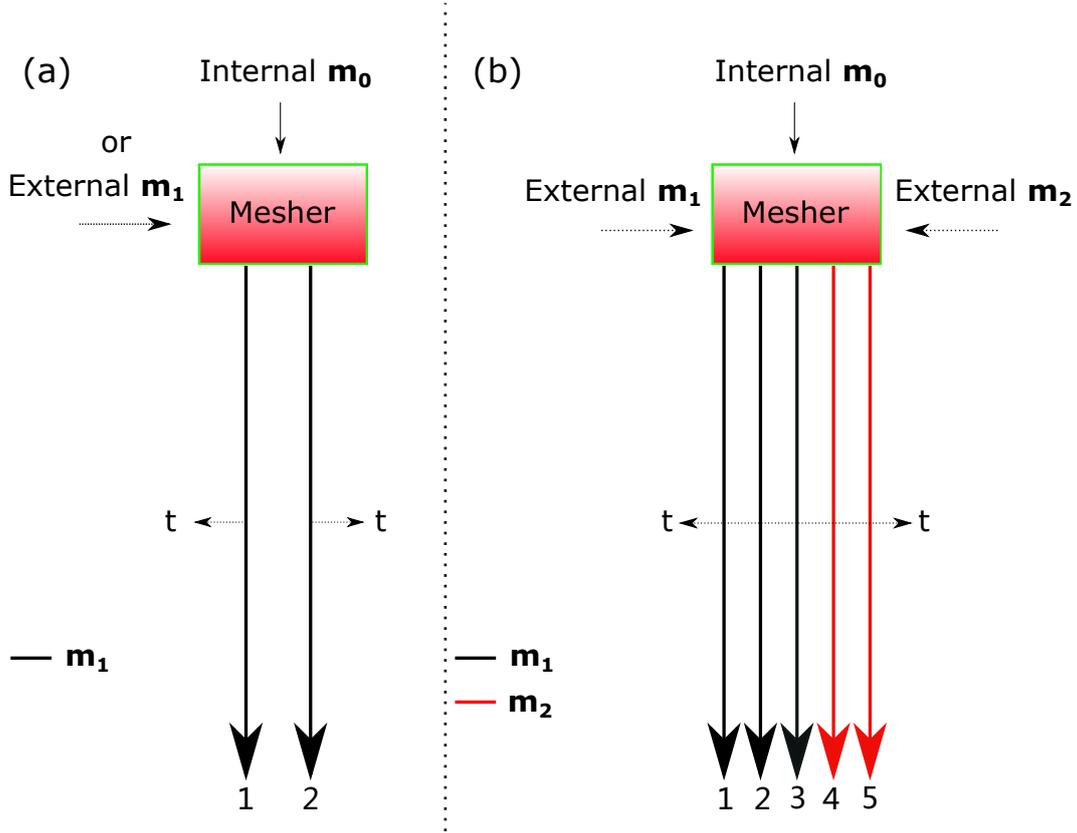


Figure 2. Sketch illustrating the workflows for the simultaneous backward and adjoint simulations for Conventional SEM vs. Multi-SEM. (a) In the simultaneous backward and adjoint simulation of the Conventional SEM, a single model is used. Each arrow represents one solver engine with Arrow 1 indicating the backward simulation (i.e. the reconstruction of the forward field) and Arrow 2 indicating the adjoint simulation which is started from the time-reversed adjoint sources at the receivers. The Fréchet kernel contributions of each time step or an incremental time step are calculated on the fly. (b) In the simultaneous backward and adjoint simulation of the Multi-SEM, Arrows 1, 2, and 3 indicate the solver engines for model \mathbf{m}_1 , where Arrow 1 and 2 performs the same as in (a) and Arrow 3 performs the same as Arrow 2 except with the perturbation of the adjoint source is taken into account. The red Arrows 4 and 5 indicate the computation of the backward and adjoint fields for the perturbed model \mathbf{m}_2 . The calculations of Fréchet kernels (by Arrows 1 and 2), approximate Hessian kernels (by Arrows 1, 2, and 3), and the full Hessian kernels (by Arrows 1, 2, 3, 4, and 5) are simultaneously performed on the fly since the required wavefields are computed for each time step. Some solvers can be switched off for computational efficiency if necessary for instance in the computation of approximate Hessian kernels. The Multi-SEM reduces to Conventional SEM when switched off solvers indicated by Arrows 3, 4, and 5.

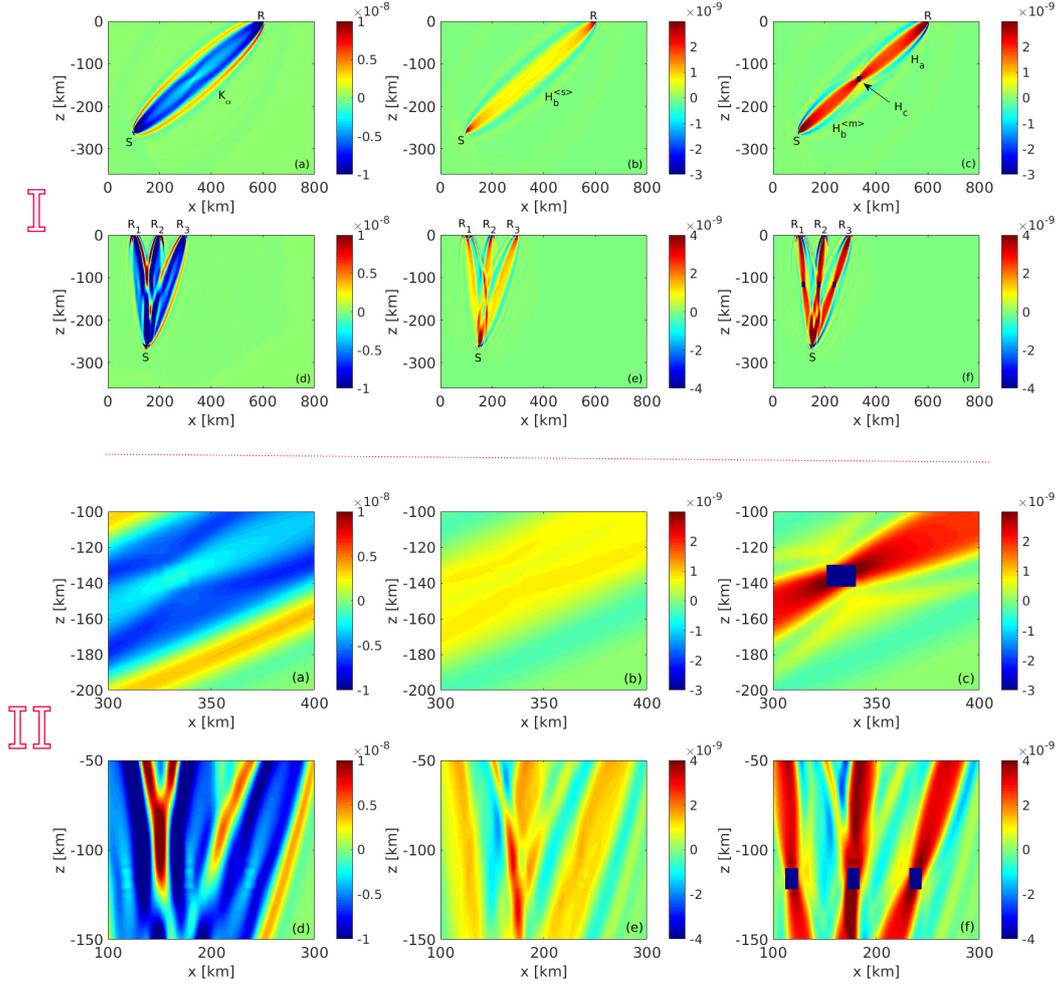


Figure 3. Part I: Fréchet and Hessian kernels computed for *Model 2* (top row) and *Model 3* (bottom row) as discussed in section 4. In the top row we show (a) the Fréchet kernel K_α , (b) the approximate Hessian kernels $H_b^{(s)}$, and (c) the full Hessian kernels for the single source single station case with a single scattering object, where the full Hessian kernels is a summation of H_a , $H_b^{(s)}$, $H_b^{(m)}$ and H_c . The H_c is restricted to the perturbation indicated by the black box in (c) dictated by its expression, eq. (6). Note that the black box here is the H_c Hessian kernels with a negative value of 10^{-9} scale, not the model perturbation although they are located in the same position. The $H_b^{(s)}$ kernel is mostly invisible in (c) except those around the black box due to its relative small amplitude. The H_a and $H_b^{(m)}$ are separated by the black box. Similarly, Panels (d), (e), and (f) in the bottom row show the various kernels for the case of a single source and three stations with three scattering objects. The kernel unit for all sub-figures is $[s m^{-2}]$. A zoomed view of the perturbations within Part I is correspondingly shown in Part II. Significant differences are observed between the approximate and full Hessian kernels.