

# The Composition of $\sim 96$ keV $W^+$ in Saturn's Magnetosphere

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## Key Points:

- On average, for  $\sim 96$  keV ions, fractional abundances at  $L = 7-16$  are  $O^+/W^+ \sim 0.53$ ,  $OH^+/W^+ \sim 0.22$ ,  $H_2O^+/W^+ \sim 0.22$ , and  $H_3O^+/W^+ \sim 0.03$
- The partial number densities  $O^+$ ,  $OH^+$ ,  $H_2O^+$  and  $H_3O^+$ , peak at dipole  $L \sim 9.5$  and the ratios  $O^+/W^+$ ,  $OH^+/W^+$ , and  $H_2O^+/W^+$  vary little with  $L$
- At  $L < 13$ ,  $H_3O^+/W^+$  decreases significantly with increasing  $L$ , whereas at  $L > 13$   $H_3O^+/W^+$  varies little with  $L$

## Abstract

The plumes of Enceladus produce a cloud of neutral  $H_2O$  molecules and, via dissociation,  $OH$  and  $O$ . These neutrals are ionized by charge exchange, solar UV, and electron impacts, producing the thermal water group ions  $W^+$  ( $O^+$ ,  $OH^+$ ,  $H_2O^+$ , and  $H_3O^+$ ) which become energized in Saturn's magnetosphere. We first separate the components of energetic ( $\sim 96$  keV)  $W^+$  using Cassini Charge-Energy-Mass Spectrometer (CHEMS) data from 78 near equatorial main ring current passes (dipole  $L = 7-16$ ,  $\pm 10^\circ$  in latitude) in 2004-2010. We find  $\sim 53\%$   $O^+$ ,  $\sim 22\%$   $OH^+$ ,  $\sim 22\%$   $H_2O^+$ , and  $\sim 3\%$   $H_3O^+$  when averaged over  $L = 7-16$ , resulting in a mean water group mass of 16.7 amu. At  $7 < L < 21$ , we find abundance ratios for  $O^+/W^+$ ,  $OH^+/W^+$ , and  $H_2O^+/W^+$  that vary little with  $L$ . However, while  $H_3O^+/W^+$  is nearly constant at  $L > 13$ ,  $H_3O^+/W^+$  tends to increase persistently at  $L < \sim 10$ . The large  $O^+$  abundance qualitatively agrees with the broad atomic  $O$  cloud observed by Cassini and predicted by some models. Our observation of  $H_2O^+/W^+ > \sim 20\%$  out to  $L \sim 21$  suggests that neutral  $H_2O$  spreads throughout the magnetosphere rather than being confined to a narrow  $H_2O$  torus centered on Enceladus' orbit.

## Introduction

[1] The discovery of OH in Saturn's magnetosphere by the Hubble Space Telescope (HST) indicated that Saturn's magnetosphere contains a significant source of neutral  $H_2O$  (Shemansky et al., 1993). Using Voyager and HST observations, Jurac and Richardson (2005) modeled the plasma and neutral particles in Saturn's magnetosphere and predicted that the majority of the  $H_2O$  was coming from Enceladus at a rate of  $\sim 10^{28}$   $H_2O/s$ , a small moon at  $\sim 4 R_S$  (the Saturn radius; please note that near the equator, dipole  $L = 1$  is about  $1 R_S$ ). The discovery of water plumes in the south-pole region of Enceladus by the Cassini mission confirmed their predictions (Waite et al., 2006; Hansen et al., 2006; Porco et al., 2006). Typical estimates of the plume source strength range from  $\sim 1-4 \times 10^{28}$   $H_2O/s$  (Burger et al., 2007; Jia et al., 2010; Dong et al., 2011), but some studies suggest that the source strength varies significantly with time and can be as low as  $\sim 0.6-0.7 \times 10^{28}$   $H_2O/s$  (Saur et al., 2008; Smith et al., 2010). Cassini measurements of the ionized  $H_2O$  products ( $O^+$ ,  $OH^+$ ,  $H_2O^+$ , and  $H_3O^+$ , collectively called  $W^+$ ) in the plume indicated that  $H_3O^+$  was an important component of the Enceladus plume's thermal ionosphere (Cravens et al., 2009), whereas, at energetic, that is, suprathermal energies,  $O^+$  dominates, followed by  $OH^+$  and  $H_2O^+$  (Krimigis et al., 2005). Sergis et al. (2010) found that (a) the contribution of the energetic particles to the total particle pressure becomes significant at  $>9 R_S$  and progressively overtakes the thermal plasma beyond  $12 R_S$ , and (b) Saturn's ring current at  $\sim 8-18 R_S$  is intense and variable, being primarily inertial, determined by thermal particles, at  $<8.5 R_S$  but increasingly pressure gradient-driven by suprathermal particles in its maximum region (8 to  $12 R_S$ ) and certainly farther out.

[2] Charge exchange is especially important for both the ions and the neutral particles in Saturn's inner magnetosphere because of the high neutral gas densities (Mauk et al., 2009). Johnson et al.

(2006) concluded that charge exchange reactions between ions and the neutral particles emitted by Enceladus produce a broad OH cloud and leave a narrow  $H_2O$  torus centered on Enceladus' orbit (a region roughly collocated with a torus of dust and ice grains called the E-ring, see e.g., Showalter et al., 1991; and Cassidy & Johnson, 2010). Although this model reproduced the OH torus observed by HST, it, along with other models that were based on data prior to the Cassini mission (e.g. Richardson et al., 1998; Jurac and Richardson, 2005), had to be revised. Re-analysis of the OH observations by HST revealed that the OH cloud was broader than initially thought (Melin et al., 2009). Additionally, the Cassini Ultraviolet Imaging Spectrograph (UVIS) observed an O cloud twice as broad as the OH cloud (Melin et al., 2009; Persoon et al., 2009; 2013). Further, Cassini observations of the plasma electrons (Schippers et al., 2008) and ions (Wilson et al., 2008; McAndrews et al., 2009; Thomsen et al., 2010; Sittler et al., 2008; Holmberg et al., 2012; 2014; 2017; Livi et al., 2014) greatly improved the understanding of the plasma environment in Saturn's magnetosphere, so the dissociation and ionization rates of the neutrals had to be revised.

[3] Interactions of the neutral particles in Saturn's neutral cloud with plasma and solar UV photons, such as charge exchange, electron impact ionization and photoionization, produce  $O^+$ ,  $OH^+$ ,  $H_2O^+$ , and  $H_3O^+$  (Ip, 1997; 2000; Johnson et al., 2006; Smith et al., 2010; Cassidy & Johnson, 2010; Fleshman et al., 2010a). The combination of these four species is typically called the water group  $W^+$ .  $H_3O^+$  is produced largely by charge exchange reactions in the Enceladus torus (Cravens et al., 2009; Fleshman et al., 2010) and was observed to be a prominent  $W^+$  component at thermal energies during Cassini's Saturn Orbit Insertion (SOI) (Young et al., 2005; Sittler et al., 2008) and during Enceladus flybys (Tokar et al., 2006, 2009; Cravens et al., 2009). However, the overall importance of thermal  $OH^+$  and  $H_2O^+$ , and a diminished importance of

thermal  $H_3O^+$ , except possibly near Enceladus and at  $<4.75 R_S$ , has recently been demonstrated in a preliminary study by Wilson et al. (2015).

[4] Electron impact dissociation and photo-dissociation of  $H_2O$  and  $OH$  are common sources of  $H$ ,  $OH$ , and  $O$  in Saturn's magnetosphere (e.g., Fleshman et al., 2010a,b), efficiently spreading these neutral species throughout the magnetosphere (e.g. Smith et al., 2010a,b). Outside of  $\sim 8 R_S$ , the  $O$  cloud predicted by Smith et al. (2010) was not in agreement with the UVIS observations (Melin et al., 2009). Assuming the UVIS results are correct, the discrepancy may be the result of neglecting the effect of neutral-neutral collisions in the model, which is another effective mechanism for spreading neutral particles (Farmer, 2009). Cassidy & Johnson (2010) include the effects of neutral-neutral collisions in their model, and their estimates of the  $O$  density are in better agreement with the results from Melin et al. (2009). However, the  $O$  density from Cassidy & Johnson (2010) is from a factor of  $\sim 2$ -4 less than that estimated by Melin et al. (2009) at  $\sim 10$  to  $20 R_S$ . Including the collisions between neutrals significantly increased the  $H_2O$  density in Saturn's middle and outer magnetosphere. Models that ignore neutral-neutral collisions (e.g., Johnson et al., 2006) predict a narrow  $H_2O$  torus centered on Enceladus with a negligible  $H_2O$  density outside of  $6 R_S$ , while including the neutral-neutral collisions results in an  $H_2O$  cloud with approximately the same density as  $OH$  outside of  $7 R_S$  (Cassidy & Johnson, 2010).

[5] Ion observations from many Cassini orbits revealed that  $W^+$  is the most abundant ion species in Saturn's equatorial plane from thermal (Thomsen et al., 2010) to suprathermal energies (DiFabio et al., 2011). Initial Cassini Plasma Spectrometer (CAPS) observations from SOI indicated that inside of  $10 R_S$ , the thermal water group consists mostly of molecular ions with an

average mass of approximately 18 amu (Figure 16 in Sittler et al., 2008; see also Persoon et al., 2009). Subsequent analysis of CAPS ion data from two Cassini orbits in 2011 by Wilson et al. (2015) found a similar value,  $\sim 17.6$  amu.

[6] The results from our initial analysis of the Cassini Charge-Energy-Mass Spectrometer (CHEMS) data from two years of Cassini orbits at Saturn was reported by Mauk et al. (2009) in their Table 11.3, where 127-145 keV  $W^+$  was estimated to consist of 46%  $O^+$ , 24%  $OH^+$ , 25%  $H_2O^+$ , and 5%  $H_3O^+$  (see also Difabio, 2012). In the present study, we use the data from CHEMS to separate  $\sim 96$  keV  $W^+$  into  $O^+$ ,  $OH^+$ ,  $H_2O^+$ , and  $H_3O^+$ . The data are now more clearly understood, the fitting procedure used to determine the composition of  $W^+$  has been refined since the analysis reported in Mauk et al. (2009), and the data interval extended, so the results presented here provide a better estimate of the suprathermal  $W^+$  composition. First, we determine the average  $W^+$  composition in the equatorial ring current region (at  $L = 7-16$  and  $\pm 10^\circ$  Latitude) from 2004-348 to 2010-356, in order to complement DiFabio et al. (2011), the first Cassini study of the long-term variations of the major suprathermal ions in Saturn's magnetospheric equatorial ring current region. Then, extending the radial range to  $L=21$ , we present the first examination of suprathermal  $W^+$  and its compositional component variations with  $L$ , which we then compare to the thermal ion composition measurements from CAPS reported by Wilson et al. (2015) and model results, including those of Fleshman et al (2013) which provide background in the Wilson et al. analysis.

#### CHEMS Instrument

[7] CHEMS uses energy per charge ( $E/Q$ ) selection followed by time-of-flight (TOF) versus energy measurements to identify positive ions with  $E/Q = 3-220$  keV/e. CHEMS has an overall

field of view of  $159^\circ \times 4^\circ$  with the large angle divided into three independent  $53^\circ$  sections called telescopes. With the measurement of  $E/Q$ , TOF, and kinetic energy with a solid-state detector (SSD) located at the end of the TOF path, the mass, energy and charge state can be determined. The SSD has an electronic threshold of 26 keV. When nuclear defect is included, energy measurements can only be made for protons above 27 keV and oxygen above 47 keV. Herein, we focus on the narrow energy channel at  $\sim 96$  keV (specifically  $95.9 \pm 2.9$  keV, Vandegriff et al., 2018), which has the lowest  $W^+$  instrumental background, in order to achieve the most effective  $H_2O^+$  and  $H_3O^+$  separation. Below that energy, only the  $E/Q$  and mass per charge ( $M/Q$ ) of an ion are determined. See Krimigis et al. (2004) for a more detailed description of the CHEMS instrument. The entire Cassini/MIMI/CHEMS data set and Cassini ephemeris data are publicly available in the Planetary Data System and can be found at <http://pds.nasa.gov>.

## Method of Data Analysis

[8] We combine the CHEMS data from 78 of the 82 equatorial ring current passes ( $L = 7$ -16; Latitude =  $-10^\circ$  to  $10^\circ$ ) used earlier by Difabio et al. (2011) to study long term variations of the major suprathermal ions at Saturn. We focus our efforts on the main ring current where the highest, most important suprathermal ion fluxes are present (Sergis et al., 2009) in order to separate  $W^+$  into four components:  $O^+$ ,  $OH^+$ ,  $H_2O^+$ , and  $H_3O^+$ . We focus on the  $\sim 96$  keV/e  $E/Q$  channel of telescope 1 data where all four  $W^+$  components are best resolved. Suprathermal ions at Saturn, at the energies considered here, do not rigidly corotate with Saturn outside of  $\sim 5$ -6  $R_S$  (Richardson et al., 1998; Kane et al., 2008), and the observed corotation speeds inside  $\sim 6 R_S$  are  $< \sim 50$  km/s, small compared to the suprathermal ion speeds. Convection speeds, as estimated in the above studies, are up to  $\sim 50$ -100 km/s lower than rigid corotation speeds in the equatorial ring current out to 15-18  $R_S$ , so only using telescope 1 does not affect our results significantly.

We only use the CHEMS data that includes an energy measurement from the SSD because of its lower background.

[9] A plot of the SSD energy channel, Essd channel, versus the time-of-flight channel, TOF channel, for all 78 ring current passes combined is shown in Figure 1a. Separating the  $W^+$  components was found to be easier using the raw channel values. (DiFabio, 2012), so we do not convert Essd channel and TOF channel to Mass (amu) and M/Q (amu/e). We make an Essd channel cut selecting the data between the horizontal red dashed lines shown in Figure 1a and create a TOF channel histogram of the data (see Figure 1b). All four major components of the  $W^+$  group can be identified, although the  $H_3O^+$  appears as a shoulder. These data are from CHEMS telescope 1 which has the best heavy ion resolution after signal to noise ratio (DiFabio, 2012). We fit these peaks to estimate the abundances of these four species.

[10] We use the  $\sim 96$  keV  $O^{++}$  ( $E/Q \sim 48$  keV/e) TOF distribution from a different energy step (see Vandegriff et al., 2018) as a guide to choosing appropriate fitting parameters for the  $W^+$  species, which overlap. The pre-launch calibration data for  $\sim 96$  keV  $O^+$  and  $H_2O^+$  are not used because the calibration data distributions tend to be narrower than those in the in flight data due to the mono-energetic and unidirectional nature of the calibration beam. Since the  $O^{++}$  ions at  $E/Q = \sim 48$  keV/e have the same energy as  $\sim 96$  keV  $O^+$ , the  $\sim 48$  keV/e  $O^{++}$  TOF distribution is nearly identical to that of  $\sim 96$  keV  $O^+$ . We can easily fit the  $\sim 96$  keV  $O^{++}$  since there is no overlap from other ion species and use the result in fitting the  $O^+$ .

[11] A TOF histogram of the  $\sim 96$  keV  $O^{++}$  distribution for all 78 equatorial ring current passes is shown in Figure 2. The  $\sim 96$  keV  $O^{++}$  distribution is asymmetric, so a simple Gaussian, such as



that used to obtain our preliminary CHEMS data analysis results reported in Mauk et al. (2009), is not the best fitting function. An asymmetric Gaussian fits the central peak of the distribution, where its centroid (the peak's center of symmetry) is at the TOF channel  $\tau_c$ , which we call the peak's "center". At TOF channels higher than the centroid, the counts decline more slowly than expected for a Gaussian distribution as a result of straggling in the carbon foil (Northcliffe & Schilling, 1970; Gloeckler & Hsieh, 1970), so we must add a small additional term to the Gaussian on the right side of the peak. We refer to this higher-TOF section as the "tail". The functional form we use to fit the count distribution in TOF channel,  $\tau$ , is

$$f(\tau) = A \left[ e^{-\frac{(\tau-\tau_c)^2}{2\sigma_1^2}} P(-\tau + \tau_c) + \left( e^{-\frac{(\tau-\tau_c)^2}{2\sigma_2^2}} + \frac{(\tau - \tau_c)^2}{100 + [(\tau - \tau_c)^2]} \right) P(\tau - \tau_c) \right] \quad (1)$$

This function maximizes at the centroid  $\tau = \tau_c$  with an amplitude A. The sigmoid function  $P(x) = 1/(1+e^{-x})$  is an approximation to the Heaviside function which is  $\sim 0$  for  $x < 0$  and  $\sim 1$  for  $x > 0$ . The parameters  $\sigma_1$  and  $\sigma_2$  are the Gaussian standard deviations, which we call the "widths", for the  $\tau < \tau_c$  and  $\tau \geq \tau_c$  sides of the distribution, respectively. The parameter  $\alpha$  controls the rate of decline in the tail, in that the function declines as  $\sim ((\tau - \tau_c)^2)^{(1-\alpha)}$  for  $\tau \gg \tau_c$ .

[12] Gaussian fits performed by von Steiger et al. (2000) on ion species distributions from the Ulysses Solar Wind Composition Spectrometer, a functionally similar type of instrument, showed (in their Plate A2) that the width to center ratios of the TOF distributions vary slowly, so that differences between species spaced closely in TOF channel, such as O<sup>+</sup> and H<sub>3</sub>O<sup>+</sup>, vary by  $\lesssim 3$ -4% percent. Therefore, we use the  $\sigma_1/\tau_c$  and  $\sigma_2/\tau_c$  ratios as fitting parameters rather than  $\sigma_1$

and  $\sigma_2$ . The form of our fitting function is similar to the kappa distribution used by Allegrini et al. (2006) to fit the energy distribution of 1-50 keV ions traveling through thin, carbon foils.

[13] Using the Levenberg-Marquardt fitting procedure (Press et al., 1989), we use the function given in Equation 1 to fit the  $\sim 96$  keV  $O^{++}$ . The best fit (reduced  $\chi^2 = 0.73$ ) is obtained with  $\alpha = 1.81$  and is shown in Figure 2. Then, when fitting the  $\sim 96$  keV  $O^+$  distribution, we set its  $\sigma_1/\tau_c$  ratio equal to that of the  $\sigma_1/\tau_c$  ratio from the  $\sim 96$  keV  $O^{++}$  fit. Because the widths of the molecular ion distributions are not necessarily identical to those of atomic ions, we let the  $\sigma_1/\tau_c$  and  $\sigma_2/\tau_c$  ratios of  $OH^+$ ,  $H_2O^+$ , and  $H_3O^+$  differ from that of  $O^+$ , but force them to be identical for these molecules, because they are dominated by their heaviest atom, O.

[14] Although the CHEMS pre-launch calibration data show that the  $W^+$  tail counts grow with increasing  $M/Q$ , we were unable to determine how much the  $\alpha$  parameter would vary among the four species from the in-flight data. Thus, we assumed that the  $\alpha$  parameter is same for all four species. This assumption increases the uncertainty in our estimates of the ion abundances, particularly that of  $H_3O^+$ . In order to determine the sensitivity of our results to the form of the fitting function, we performed three different types of fits on the  $W^+$  distribution with these conditions on  $\alpha$ : type 1: asymmetric Gaussian ( $\alpha = \infty$ ), type 2: requiring  $\alpha$  to equal the value from the  $O^{++}$  fit ( $\alpha = 1.81$ ), and, type 3: allowing  $\alpha$  to vary. (We include this discussion although, in the results of our analysis, variations of the distributions' tails tended to be of minor importance.) For the first two types, we require the  $\sigma_2/\tau_c$  value for  $O^+$  to equal the results from the  $O^{++}$  fits. Since the value of the  $\alpha$  parameter affects the value of  $\sigma_2$ , we let the  $O^+$   $\sigma_2/\tau_c$  and  $\alpha$  vary for the third fit. The centers  $\tau_c$  for all four species and the  $\sigma_1/\tau_c$  ratio for the molecular ions

are only allowed to vary in the asymmetric Gaussian fit (type 1). In the type 2 and 3 fits, the centers and molecular ion  $\sigma_1/\tau_c$  ratio are forced to equal the values from the type 1 fit. We let the amplitudes  $A$  and  $\sigma_2/\tau_c$  ratio of the molecular ions vary for all three types of fits.

[15] Figure 3 shows our best fit (reduced  $\chi^2 = 4.0$ ) of the entire  $W^+$  distribution, which is found using the type 3 fit and with  $\alpha = 1.90$ . Numerically integrating each peak from the type 3 fit provides our best estimate of the  $O^+$ ,  $OH^+$ ,  $H_2O^+$ , and  $H_3O^+$  counts. We then convert the counts into partial number density and calculate the fractional abundance relative to  $W^+$ . As discussed in the next section, the type 1 and 2 fits are used to determine the upper and lower limits of the  $H_3O^+$  abundance.

[16] After obtaining the fits in the main ring current ( $L = 7-16$ ; Latitude =  $-10^\circ$  to  $10^\circ$ ), we divide the Cassini passes by  $L$  shell and look for variations with  $L$  in the  $L = 7-21$  range. When fitting the data by  $L$  shell, the statistics are worse, and we only let the amplitudes of the peaks vary, forcing the center  $\tau_c$ ,  $\sigma_1$ ,  $\sigma_2$ , and  $\alpha$  of each species peak to be the same as the results from the  $L = 7-16$  fit. We use a  $\Delta L = 1$  for  $L = 7$  to  $11$  where there are sufficient counts, but we must combine  $L$  shells beyond this distance where the counting statistics decline, using  $\Delta L = 2$  for  $L = 11$  to  $15$ , and  $\Delta L = 3$  for  $L = 15$  to  $21$ .

## Results and Discussion

[17] Results and comparison with earlier studies. Table 1 lists the counts from the best fit to the  $W^+$  distribution plotted in Figure 3 (i.e., the type 3 fit with  $\alpha = 1.90$ ) along with the partial number densities and the relative abundances of the  $\sim 96$  keV  $W^+$ . The  $W^+$  group has an average mass of 16.7 amu and consists of  $\sim 53\%$   $O^+$ ,  $\sim 22\%$   $OH^+$ ,  $\sim 22\%$   $H_2O^+$ , and  $\sim 3\%$   $H_3O^+$ . Since the

number of counts in the distribution tails contain only a small percentage of the total counts, our estimates of the  $O^+$ ,  $OH^+$ , and  $H_2O^+$  abundances do not vary significantly among the three types of fits. The quoted uncertainties for these three species reflect the statistical uncertainties from our type 3 fits. Because  $H_3O^+$  is much less abundant and the tails from the other three species make up a significant fraction of the counts at the location of the  $H_3O^+$  peak, our estimate of the  $H_3O^+$  counts strongly depends on the  $\alpha$  parameter. This can possibly introduce a larger systematic error and the quoted  $H_3O^+$  uncertainties reflect the range of values from our three different fits. The asymmetric Gaussian fit with no tail provides an upper limit on the  $H_3O^+$  abundance, while forcing  $\alpha = 1.81$ , which overestimates the tail counts, provides a lower limit on the  $H_3O^+$  abundance.

[18] The partial number densities and fractional abundances of the four  $\sim 96$  keV  $W^+$  components are plotted versus  $L$  in Figures 4a and 4b, respectively. Figure 4a shows that all four  $W^+$  components increase similarly at  $L \sim 8-10$ , peak at  $L \sim 9.5$ , and decrease at  $L < \sim 9$ , but not quite similarly. Figure 4b shows that (a) fractional  $O^+$ ,  $OH^+$ , and  $H_2O^+$  vary little from  $L = 5-21$ , with each species' variations being statistically insignificant, (b) fractional  $H_3O^+$  is approximately constant at  $L > \sim 13$ , and (c) the enhancement of fractional  $H_3O^+$  probably starts at  $L \sim 13$ , but certainly only increases monotonically inside of  $L \sim 11$ , increasing by almost a factor of  $\sim 3$  above fractional  $H_3O^+$  levels at  $L > \sim 13$ . Paranicas et al. (2008) noted that  $\sim 5-100$  keV  $H^+$  and  $O^+$  (products of the magnetospheric processes CRAND and , respectively) populate Saturn's inner magnetosphere episodically, being energized by injections, and concluded that they are then primarily lost to charge exchange in the Enceladus torus as they drift inward. We agree that this interpretation is consistent with our  $W^+$  radial variations in Figure 4a. The  $W^+$  components may originate from any object in Saturn's magnetosphere before their dispersal and subsequent

inward transport from  $L \sim 15$ -20, but most are currently presumed to originate in Saturn's main rings and the Enceladus neutral torus. On their way inward, the ions' partial number density at a given energy increases, peaking at  $L \sim 8$ -10, inward of which losses increase as they encounter the neutral torus. Paranicas et al. (2008) demonstrated that energetic/suprathermal  $O^+$  flux decreases persistently in to  $L \sim 5.5$ -6.5 (depending on magnetospheric conditions), where it then becomes undetectable (see Figure S3 in the Supporting Information). They note that "...The shape of the flux drop off inward to the planet has different behavior on each orbit we have surveyed and is not related to the locations of the moons." They argued that the  $W^+$  components are transformed by the neutral gas cloud in this region which converts the inward drifting ions to energetic neutral atoms via charge exchange. Kollmann et al. (2011) show that  $\sim 46$  keV  $H^+$  and  $\sim 91$  keV  $e^-$  display differential intensity peaks at  $\sim 9 < L < \sim 10.5$  and  $L \sim 8$ , respectively (see Figure S4 in the Supporting Information), similar to the energetic protons and oxygen in Paranicas et al., with fluxes decreasing in to  $L \sim 5$ -6. Christon et al. (2014) show that 83-167 keV  $W^+$ , a portion of its straggling tail  $W^+_{tail}$  at  $\sim 22$ -26 amu/e,  $O_2^+$ , and  $^{28}M^+$  all display peaks at  $L \sim 9$  and decreases in to  $L \sim 6$  (see Figure S5 in the Supporting Information). The fractional  $W^+_{tail}$  varies little outside of  $L \sim 7$ , similar to this study's fractional  $O^+$ ,  $OH^+$ , and  $H_2O^+$ , and like the fractional  $H_3O^+$  component, the fractional  $O_2^+$  and  $^{28}M^+$  abundances vary insignificantly at  $L \geq 13$ , but vary inside  $L \sim 13$ . However, rather than increasing as fractional  $H_3O^+$  does, fractional  $O_2^+$  and  $^{28}M^+$  levels decrease markedly inside  $L \sim 13$ , much further than the fractional  $H_3O^+$ . Note though, in this study, while the variation is not statistically significant, fractional  $H_3O^+$  is disrupted with respect to its behavior beyond  $L \sim 15$  and starts increasing inside  $L = 13$ . Interestingly, the magnetic pressure of Saturn's inner, "core", dipolar magnetic field is rather well ordered out to  $\sim 10 R_S$ , with disturbed ordering persisting somewhat out to  $\sim 13 R_S$ , beyond which low level disturbed magnetic pressure extends to larger radial distances (Sergis et al.,

2010), suggesting that the  $L < 10$  region may be characteristically differentiated from the  $L > 13$  magnetosphere. It is apparent that the  $H_3O^+$  is not depleted as much as the other three  $W^+$  components inside of  $L \sim 10$ , suggesting some as yet unidentified, differentiated interaction(s) possibly resulting from this magnetospheric magnetic field transition near  $L \sim 10$  occur near the neutral torus for these ions.

[19] The energetic  $O^+$ ,  $OH^+$ , and  $H_2O^+$  fractional abundances in Figure 4b show no significant variation with  $L$  from 7 to 21; they do not deviate far from their average values throughout the main ring current. The dashed box encompasses the entire radial and abundance range of the four thermal ion fractional abundances from Wilson et al. (2012). Suffice it to say, most of the thermal ion radial variation is evident at  $L < 8$ . Clear depiction of both data sets, not necessary here, is presented in Figure S2 in the Supporting Information. We expect few, if any, correlated variations or processes between the energetic and thermal- $W^+$ -components as the thermal  $W^+$  components propagate outward from their inner magnetospheric source region, and the suprathermal  $W^+$  components diffuse inward from their outer magnetospheric acceleration region(s). The suprathermal  $W^+$  components have much wider variation than the thermal components, with  $O^+$  being somewhat more abundant and  $H_3O^+$  being much less abundant than either  $OH^+$  or  $H_2O^+$ .

[20] The suprathermal  $H_3O^+$  fractional abundance has an average value of  $\sim 0.017$  at  $13 < L < 21$  and shows no measurable radial dependence. However, at  $L < 10$ , the  $H_3O^+$  partial density peak and relative decrease are not as substantial as that of the other  $W^+$  components, resulting in its fractional abundance increasing to  $H_3O^+/W^+ \sim 0.048$  - apparently due to a weaker interaction with the nearer-Saturn environment. One possible cause of this more moderate decrease of  $H_3O^+$  at

lower L might be that the  $H_3O^+$  charge exchange lifetime is longer than those of  $O^+$ ,  $OH^+$ , and  $H_2O^+$ , or, alternatively, that  $H_3O^+$  interactions with the negatively charged E-ring grains inside the orbit of Rhea ( $\sim 8.74$  Rs) are weaker than those of  $O^+$ ,  $OH^+$ , or  $H_2O^+$ . The fractional abundances of suprathermal ions with longer charge exchange lifetimes (smaller cross sections), such as possibly  $H_3O^+$ , would be expected to increase with respect to other ions whose cross sections are larger, as losses become more important going inward. Unfortunately, although cross sections for 10's of eV-energy  $W^+$  ions are available (e.g., Lishawa et al., 1990), to our knowledge, the cross sections for  $\sim 10$ -500 keV water group ions colliding with Saturn's most important thermal neutral cloud atoms and molecules (O, OH, and  $H_2O$ ) are not all available in the literature. As a result, we can only offer qualitative analysis of the  $W^+$  component charge exchange lifetimes below using available lower energy charge exchange cross sections as a guide. An alternative, independent, and concurrent process contributing to the difference of the  $H_3O^+$  radial dependence might result from varying strengths of interactions between  $H_3O^+$  and the other  $W^+$  components with the charged E-ring grains. E-ring grains are observed to vary in size, from several nanometers up to tens of micrometers and have radially varying charged components, in which negatively charged grains dominate inside the orbit of Rhea, transitioning to both positive and negative within an Rs or so around the orbit of Rhea, and then become preponderantly positive with increasing distance from Saturn outside the orbit of Rhea, as reported, for example, by Wahlund et al. (2005), Sittler et al. (2006), Kempf et al. (2008), Hsu et al. (2013), and Ye et al. (2017). We address both possible scenarios below.

[21] In Figure 4b we compare our fractional  $W^+$  abundance results at  $\sim 96$  keV to the range of the preliminary results for the thermal plasma fractional abundances reported by Wilson et al. (2015) in their Figure 3. Both studies analyze near-equatorial data, with a radial overlap at  $L = 7$ -10; that

is, this study uses  $L = 7-21$  and Wilson et al. (2015) used  $R_S = 4.75-10$ . Considering the small  $7 \leq L \leq 10$  overlap region, the fractional abundance values are somewhat comparable, although the range of fractional suprathermal ion abundances is much greater than that of the thermal ions. Ratio ranges from the two studies and model results from Figures 3 in Fleshman et al. (2013) and Wilson et al. (2015) are listed in Table 2. To date, we are unaware of any models characterizing the energetic/suprathermal water products' distribution, propagation, and/or interactions in Saturn's magnetosphere. The observed  $OH^+$  and  $H_2O^+$  fractional abundances are roughly comparable in the two energy ranges, although the thermal population values are somewhat larger. Major differences between the thermal and energetic population  $O^+$  and  $H_3O^+$  fractional abundances are evident: the  $O^+$  fractional abundance is much larger for the suprathermal ions and the  $H_3O^+$  fractional abundance is much smaller for the suprathermal ions. The thermal ion fractional abundances are all between  $\sim 0.18$  and  $\sim 0.30$ , which is about twice the range of the  $H_2O^+$  and  $OH^+$  suprathermal uncertainties. Understanding that acceleration to suprathermal energies is not instantaneous, one might expect for extensive dissociation of the  $W^+$  molecular species to occur, producing relatively more atomic  $O^+$ . In summation, the suprathermal fractional abundances show little radial variation except for the  $H_3O^+$  increase by a factor of  $\sim 3$  from  $L \sim 13$  to  $L \sim 7$ . Of note, the thermal  $H_3O^+$  fractional abundance also shows a large increase with decreasing  $L$ . However, at thermal energies, the fractional abundances of the other species also show radial variations over the  $L = 7-10$  range with the  $O^+$  and  $OH^+$  increasing and the  $H_2O^+$  decreasing with increasing  $L$ .

[22] Origin of the suprathermal ions. The majority of Saturn's water group ions are produced by the dissociation and ionization of the neutral  $H_2O$  emitted from the Enceladus plumes (Burger et al., 2007; Smith et al., 2010; Cassidy & Johnson, 2010). Since the ions observed here are much



more energetic than the thermal ions, we must consider the acceleration and transport processes when comparing our results to the thermal ion observations and neutral cloud models. The most likely acceleration processes are injections (Mauk et al., 2005; Hill et al., 2005; Burch et al., 2005; Paranicas et al., 2007; Rymer et al., 2009) and inward radial diffusion (Van Allen et al., 1980; Armstrong et al., 1983) conserving the first and second adiabatic invariants. Injection events are a result of the centrifugal interchange instability where narrow “fingers” of hot plasma are injected inward and replace broad, outward flowing cold plasma. Injection events can be grouped into two categories: low energy and high energy (Mitchell, Carbary, et al., 2009). Low energy events originate at L = 9-11 (Rymer et al., 2009), while high energy events originate in the magnetotail (Mitchell, Krimigis, et al., 2009). Acceleration by inward radial diffusion also requires the thermal ions to start in the outer magnetosphere, so the accelerated ions studied here likely reflect the plasma ion composition in the outer magnetosphere, a region not addressed by Wilson et al. (2015).

[23] The composition of the thermal ions in the outer magnetosphere depends on a number of factors. One is the ionization rate of the neutral particles in the outer magnetosphere. Although possible differences among the neutral species’ ionization and dissociation rates prevent us from directly comparing the ion composition to the neutral densities, our results do provide constraints. If a large percentage of the thermal ions in the outer magnetosphere is produced by the ionization of neutrals in this region, then the suprathermal ion composition will closely reflect the neutral particles that were initially ionized in the outer magnetosphere.

[24] Because the ionization rates are much higher in the inner magnetosphere, we must consider the possibility that thermal ions produced in the inner magnetosphere are then transported to the

outer magnetosphere where they are accelerated. As discussed above, broad outward flows of cold plasma are consistent with the properties of the centrifugal interchange instability (Hill et al., 2005; Rymer et al., 2008). Electron and ion observations suggest that this outward transport process is not adiabatic and does not cool the plasma (Rymer et al., 2007, 2008; Wilson et al., 2008). Therefore, the outward transport of thermal ions followed by acceleration in the outer magnetosphere can produce ions at suprathermal energies. Observations by Cassini UVIS (Melin et al., 2009) and models of the neutral cloud (Smith et al., 2010; Cassidy & Johnson 2010) indicate that the neutral O cloud is much broader than the neutral OH and  $H_2O$  clouds. Outside of  $6 R_S$ , the neutral O density is larger than that of OH and  $H_2O$ , and the neutral O density is about an order of magnitude greater than the other water group neutrals near  $\sim 20 R_S$  (Cassidy & Johnson, 2010), whereas, near the orbit of Rhea,  $\sim 8.75 R_S$ , the neutral O to OH (and, likely  $H_2O$ ) density ratio is modeled to be from  $\sim 2$  to  $\sim 3.5$  (by Smith et al., 2010 and Cassidy & Johnson, 2010, respectively). Since the ionization rates of the neutral particles throughout the magnetosphere are estimated to be approximately the same (Cassidy & Johnson, 2010), we would have expected the suprathermal  $O^+/OH^+$  and  $O^+/H_2O^+$  ratios to be larger than our observed value of  $\sim 2.4$ , because their presumed source is closer to  $\sim 20 R_S$  than  $\sim 8.75 R_S$ .

[25] It is possible that a large percentage of the ions in the outer magnetosphere is produced inside of  $10 R_S$ , where the difference between densities of neutral O and of neutral molecules is not as large, as noted above, and the ions are then transported to the outer magnetosphere. However, Wilson et al. (2015) find that  $O^+$  is generally the least abundant water-group ion at thermal energies, not the most abundant water-group ion, as at suprathermal energies. Wilson et al. (2015) reported that the mean mass of thermal water group ions has a mean of 17.56 amu in the radial region of  $4.75$  to  $10 R_S$ , and decreases with distance from Saturn. On the other hand,

the mean mass of suprathermal ions, 16.7 amu, reflects the large percentage of suprathermal energy  $O^+$ ,  $\sim 2.5$  times greater than  $OH^+$  or  $H_2O^+$  throughout  $7 \leq L \leq 16$ . Because of the larger fractional abundance of molecular ions inside of  $10 R_S$ , the outward transport of these ions would result in a lower  $O^+/W^+$  ratio relative to the neutral  $O/W$  ratio. To perform quantitative analysis, better constraints on the ion transport rates, production rates, and loss rates are needed. However, our observation that  $W^+$  consists mostly of  $O^+$  is qualitatively consistent with the broad atomic O cloud observed by Cassini UVIS (Melin et al., 2009) and predicted by some models (Smith et al., 2010; Cassidy & Johnson, 2010; Dialynas et al., 2013). Our observation of nearly equal amounts of suprathermal  $OH^+$  and  $H_2O^+$  suggests that neutral  $H_2O$  is spread to Saturn's middle or outer magnetosphere rather than confined to a narrow ( $\sim 2-3 R_S$ ) torus centered on Enceladus' orbit. This result supports the conclusions of the Cassidy & Johnson (2010) model, which predicts approximately equal neutral OH and  $H_2O$  densities spread outside of  $7 R_S$ . If neutral  $H_2O$  were confined to a narrow torus near  $4 R_S$  as suggested by other models (e.g. Johnson et al., 2006; Smith et al., 2010), we would expect very low abundances of suprathermal  $H_2O^+$  compared to  $O^+$  and  $OH^+$ .

[26] Our results do not rule out the possibility that most of the  $H_2O^+$  is produced near Enceladus' orbit and then transported to the outer magnetosphere where it is then accelerated. However, this scenario seems unlikely, because we would expect the suprathermal  $H_2O^+$  fractional abundance to be similar to that of the suprathermal  $H_3O^+$ . Thermal  $H_3O^+$  is the most abundant ion species in the Enceladus torus where the majority of it is produced via charge exchange reactions (Sittler et al., 2008; Cravens et al., 2009; Fleshman et al., 2010a). As the  $H_3O^+$  is transported outward, its fractional abundance drops significantly due to its short recombination lifetime in the Enceladus torus (Ip, 2000; Sittler et al., 2008). Because the recombination lifetime of  $H_2O^+$  is even shorter

than the  $H_3O^+$  lifetime in the Enceladus torus (Fleshman et al., 2010b), we would expect a similar decrease in the thermal  $H_2O^+$  fractional abundance. Therefore, the production of  $H_2O^+$  in the inner magnetosphere followed by outward transport and acceleration in the outer magnetosphere would probably not result in our observed suprathermal  $H_2O^+/H_3O^+$  ratio of  $\sim 7.8$ . A more detailed model is necessary before this possibility can be conclusively ruled out.

[27] L dependence at  $L < \sim 9$ . We now qualitatively discuss the processes that might cause the large losses of suprathermal ions inside of  $L \sim 9$ . This is a qualitative discussion because, to our knowledge, relevant functional forms and/or laboratory measurements for the water group suprathermal ions of interest relevant to their interactions with the thermal ions and neutral particles present in the E-ring (or Enceladus torus) and discussed herein do not exist or are not available in the literature. In this discussion, we will also present some measurements at low center of mass, CM, collision speeds because, they are all that exists. Charge exchange and proton-transfer cross sections likely vary among the species studied herein and are probably responsible for most of the observed changes/differences in the relative abundances of the  $W^+$  constituents. For  $W^+$  constituents, fractional differences appear to be seen only for  $H_3O^+$ , but note that important differences with respect to  $W^+$  exist for heavier important molecules at Saturn (see Figure S5 in the supporting information).

[28] Charge exchange with neutral gas. Consider the charge exchange reaction involving an ion  $A^+$  colliding with the neutral B ( $A^+ + B \rightarrow A + B^+ + \Delta E$ ), where  $\Delta E$  is the difference between the ionization energies of A and B. For resonant processes (i.e.  $\Delta E = 0$ ), the charge exchange cross section decreases with energy and behaves like  $\sigma = (C-D*\ln(v))^2$ , where C and D are constants and  $v$  is the relative velocity (Dalgarno, 1958). At low velocities, cross-sections of non-resonant

collisions are smaller than those of resonant collisions and increase with increasing velocity (Rapp and Francis, 1962). Non-resonant cross sections maximize at a velocity,  $v_{\max}$ , that can be approximated by the Massey adiabatic hypothesis (Massey, 1949):

$$v_{\max} = \frac{v_o |\Delta E|}{h} \quad (2)$$

where  $v_o = 7 \times 10^8$  cm and  $h$  is Planck's constant. Above this maximum velocity, the non-resonant cross section converges with the cross section of the resonant collision. The collision cross section maximizes at higher velocities as the difference in ionization energies increase.

[29] Although the above relation is only meant for slower collisions, in the absence of related laboratory results and in order to get a sense of any differences that may exist between H<sub>3</sub>O<sup>+</sup> and the other W<sup>+</sup> components, we apply Equation 2 to collisions of the water group ions with atomic O, which is ~6 times more plentiful than OH at ~10 R<sub>S</sub> as shown by Melin et al. (2009). The ionization energies of O, OH, and H<sub>2</sub>O are 13.6 eV, 13.0 eV, and 12.6 eV (Weast, 1968; Wiedmann et al., 1992), respectively, while the ionization energy of H<sub>3</sub>O is much less at 5.5 eV (Melin et al., 2005). Using these ionization energies, the collision cross sections of OH<sup>+</sup> on O and H<sub>2</sub>O<sup>+</sup> on O are estimated to peak at  $v_{\max} \sim 1\text{--}2 \times 10^7$  cm/s, while the H<sub>3</sub>O<sup>+</sup> on O cross section is estimated to peak near at  $v_{\max} \sim 1.4 \times 10^8$  cm/s, an order of magnitude larger. Because the velocities of ~96 keV OH<sup>+</sup> and H<sub>2</sub>O<sup>+</sup> are much greater than the velocity at which the cross section maximizes, we might expect the charge exchange cross sections of these species to be approximately the same. The cross section of ~96 keV H<sub>3</sub>O<sup>+</sup>, however, will not have reached its maximum value and will be less than the cross sections of the other three ion species. Therefore, ~96 keV H<sub>3</sub>O<sup>+</sup> will have a longer lifetime than the other ~96 keV water group ion species. As a result, the fractional abundance of H<sub>3</sub>O<sup>+</sup> should be expected to increase inward, while the fractional abundance of O<sup>+</sup>, OH<sup>+</sup>, and H<sub>2</sub>O<sup>+</sup> would not show significant variations. The result of

this argument is qualitatively consistent with and provides a reasonable framework for understanding why suprathermal  $H_3O^+/W^+$  increases with decreasing  $L$ , and has a very different explanation than that for thermal  $H_3O^+/W^+$  increasing with decreasing  $L$ . Hopefully, in future analysis related to these particles in this process, the charge exchange cross sections of these energetic ions will be more fully documented and other reactions, such as dissociation and interactions with charged dust should also be included and modeled for these suprathermal ions in Saturn's magnetosphere.

[30] Grain interactions. Grains and dust permeate Saturn's magnetosphere, as demonstrated in Figure 5, where we show three distributions of 0.1-3.0  $\mu$  grains that Hornyi et al. (2008) modeled from a continuous scan of grain densities derived from data measured during Cassini's ring-plane crossing on Orbit 7. Their study provides one of the few snapshots of the magnetosphere's grain population. Three different size ranges demonstrate the spatial distribution of grains relative to our overlaid  $W^+$  measurements in the upper panels. A large measurable portion of E-ring grains have either positive or negative charge (e.g., Kempf et al., 2006; Ye et al., 2014). Kempf et al. (2006) found that grains bigger than 2–3  $\mu$ m detected inside Rhea's orbit ( $L \sim 8.75$ ) were always negatively charged, whereas grains outside Rhea's orbit were all positively charged. Dong et al. (2015) determined that the size distribution of nanometer-sized grains obtained by CAPS ( $\sim 0.001 \mu$ m, herein called dust) suggests that  $\sim 1$ -10 nm sized grains dominate both charge density and number density, at least near Enceladus. Little information is available on total grain population characteristics (from nanometer to micrometer sizes) and their properties throughout Saturn's magnetosphere, except near Enceladus' orbit. Nevertheless, we estimate below whether ion interactions with Saturn's charged grain/dust environment should be considered in addition to Saturn's neutral gas populations, even when exclusion of the grains

might provide acceptable explanations of phenomena in Saturn's magnetosphere without such consideration. The grain interactions we consider here are a possible alternative reason for the difference between the radial profiles of  $H_3O^+$  and the other  $W^+$  components.

[31] The water-group ions should be attracted to any negatively charged particles, grains or ions, in the dusty plasma through electrostatic attraction, and repulsed by positively charged particles. The  $W^+$  components' electric dipole moments ( $\partial$ ) for  $OH^+$ ,  $H_2O^+$ , and  $H_3O^+$  have different values: 2.32, 2.40, and 1.44 D (debye), respectively (see e.g., González-Alfonso et al., 1983; Vogelius et al., 2004; and Botschwina et al., 1983). For comparison, neutral molecules can have an induced electric dipole moment, two are (up to)  $\partial_{OH} = 1.65$  D for OH, Werner et al., 1985, and  $\partial_{H_2O} = 1.85$  D for  $H_2O$ , Dyke & Muentert, 1973), so the  $H_3O^+$  dipole moment is  $\sim 40\%$  smaller than those of  $OH^+$  and  $H_2O^+$  and comparable to induced neutral molecule dipole moments. Given that,  $H_3O^+$  will be less attracted to the prevalent negatively charged E-ring grains near and inside of Rhea's orbit than the other  $W^+$  components.  $H_3O^+$  would therefore have a longer mean free path in the E-ring planetward of Rhea, if not for some distance outside of  $L \sim 9$ . Some portion of the negatively charged E-ring grains will capture  $W^+$  components, more likely the  $O^+$ ,  $OH^+$ , and  $H_2O^+$  than the  $H_3O^+$ . This weakness of the  $H_3O^+$  dipole moment is a possible, as yet unexplored, interaction difference with the E-ring material, in addition to the aforementioned possible cross section difference between  $H_3O^+$  and the other  $W^+$  components. The difference in dipole moments would result in fewer collisions with and less attraction between charged grains and  $H_3O^+$  than with the other  $W^+$  components.

[32] E-ring grain measurements in two different size ranges by Cassini instruments were: (a)  $\geq 9$   $\mu m$  by design, Cosmic Dust Analyzer, CDA (Srama et al., 2004; 2011; Kempf et al., 2008; and

>0.5 by impact identification algorithm, Kempf et al., 2004); (b)  $\sim 2$ -10  $\mu\text{m}$  from spacecraft charging, Radio Wave and Plasma Wave Science (RPWS) investigation (Gurnett et al., 2004; Ye et al., 2014; 2016); and (c)  $\sim 0.6$ -3 nm by CAPS, from large energy-per-charge measurements (e.g., Hill et al., 2012; Dong et al., 2015). Only CDA was designed to measure the grains and dust, but the measurements have been compared and appear to be complementary (Meier et al., 2014; Dong et al., 2015). Complementary overlays of data from Kempf et al. (2006) and Ye et al. (2014) in the bottom panel of Figure 5 show that grains >0.5 microns detected inside Rhea's orbit,  $\sim 8.75$  Rs, are mostly negatively charged and grains outside Rhea's orbit were positively charged, with the typical potentials being about -2 V and +3 V, respectively (Figure 1 of Kempf et al., 2006). However, the Ye et al. (2014) data clearly show that some negatively charged grains persist beyond Rhea's orbit. Dust, detected in and near the Enceladus plume with both negative and positive net charges (Hill et al., 2012), likely follows a similar, but more likely complex large scale pattern, than the grains (e.g., Engelhardt et al., 2015), but, to our knowledge, this has not yet been fully investigated. In the Appendix, we estimate the ratio of mean free paths for  $\text{H}_2\text{O}^+$  near the orbits Enceladus and Rhea (a) in O and OH gas from the models of Cassidy and Johnson (2010), and (b) in 1 and 2 nm dust at expected densities there, using dust (CAPS) and grain (CDA and RPWS) measurements at Enceladus' orbit and values at Rhea's orbit, using CDA's measured radial grain variation and complementary CDA and RPWS data near Rhea's orbit combined with extrapolated dust estimates assuming the same radial variation as grains. The range of estimates for the ratios of  $\text{H}_2\text{O}^+$  mean free paths in neutral O and OH gas to  $\text{H}_2\text{O}^+$  mean free paths in uncharged grains and dust are neither orders of magnitude apart nor widely divergent from unity,  $\approx 1$ , where the mean free path estimates are equal; the estimates range instead from ratios of  $\sim 2$  to  $\sim 0.2$ , for dust sizes at 1 and 2 nm, respectively, values slightly below and slightly above those characteristic of the most likely nanograin size (Dong et al., 2015).



Therefore, depending on the true gas and grain densities and grain sizes in the E-ring near the orbit of Rhea, in addition to charge exchange interactions, consideration of ion-grain/dust interactions through focused modeling, might be necessary in order to fully understand the radial variations of suprathermal  $W^+$  component partial densities.

#### Concluding Remarks

[33] The fractional abundances of suprathermal  $OH^+$  and  $H_2O^+$  are comparable and both consistently near  $\approx 0.2$  for  $7 < L < 21$ , compared to (a) thermal  $OH^+$  clearly increasing from  $\sim 0.23$  to  $\sim 0.3$  and (b) thermal  $H_2O^+$  decreasing somewhat from  $\sim 0.31$  to  $\sim 0.26$  from 5 to 10  $R_S$ . Over  $7 < L < 21$ , the fractional abundance of suprathermal  $O^+$  is much larger than the other  $W^+$  components and the fractional abundance of suprathermal  $H_3O^+$  is much smaller than the other  $W^+$  components. Suprathermal  $O^+$  ( $H_3O^+$ ) fractional abundance is larger (much smaller) than thermal  $O^+$  ( $H_3O^+$ ) fractional abundance in the  $\sim 7$ -10  $R_S$  radial range. The suprathermal  $O^+$ ,  $OH^+$ , and  $H_2O^+$  fractional abundances show no significant variation with  $L$  over the range  $L = 7$ -21. The suprathermal  $H_3O^+$  fractional abundance, with an average value of 0.017, shows no dependence on  $L$  at  $L = 13$ -21. Inside of  $L = 13$ , however, the suprathermal  $H_3O^+$  fractional abundance increases substantially to 0.048 at  $L = 7$ -8, a factor of 2.8. The radial variations of suprathermal  $W^+$  components are distinctly different from available thermal energy  $W^+$  components radial variations. A full examination of the interesting  $H_3O^+$  variation inside  $L \sim 9$  is hampered by both the lack of charge exchange cross sections for suprathermal water-group ions on thermal  $O$  and  $H_2O$  and the lack of more comprehensive information.

## Appendix 1. Scattering of a water ion in the near-Rhea E-ring

In simple grazing, elastic collisions, the mean free path of an ion traveling in a neutral gas (or grain field) of interaction distances  $d$  is determined, in part, by the ratio of the distance traveled ( $vt$ ) to the product of the interaction volume ( $\pi d^2 vt$ ) times the number of target particles per unit volume ( $Nt$ ).

We assume that:

(a)  $H_2O^+$  is the representative ion for the water group;

(b) O or OH are the most likely neutral gas targets (Cassidy & Johnson, 2010) with ion radius,  $r_N$ , and density  $N_N$ ;

(c) The incoming ion (I) and a target (T) neutral particle will interact if they touch, that is, if the distance between their centers is  $d \leq (r_I + r_T)$ . The interaction area, or cross section, is  $\pi(r_I + r_T)^2$ ; and (d) In estimating the varying scenarios, the grain population is characterized by their minimum grain radius,  $r_G$ , with  $r_G \gg r_N$ , and density  $N_G$ .

A general form of the MFP, given in Chemistry LibreTexts. 4.12: The Frequency of Collisions between Unlike Gas Molecules (<https://chem.libretexts.org>, 2019), is:

$$MFP = vt / (\pi(r_I + r_T)^2 vt N_T), \text{ where,}$$

$v$  = speed of the ion through the medium,

$t$  = time for the ion to travel the distance  $vt$ ,

$r_I$  = the radius of the ion,

$r_T$  = the radius of the target population's particles,  $r_T$ , which, for the

neutral gas are  $r_{N,O}$ , for O atoms, or  $r_{N,H_2O}$ , for  $H_2O$  molecules, the most populous E-ring neutrals (see the reference list below), and the second case, E- ring grains, characterized above, and

$N_T$  = the target particle population's density (the same subscripts apply).

For our purpose of obtaining order of magnitude estimates for ions in different media, we simply need to examine  $\Gamma_{N:G}$ , the ratio of the ions' mean free paths in the neutral gas (N) to grain (G)

populations, which reduces to:

$$\Gamma_{N:G} = \frac{MFP_N}{MFP_G} = \frac{vt / (\pi(r_I + r_N)^2 vt N_N)}{vt / (\pi(r_I + r_G)^2 vt N_G)} \sim \frac{(r_I + r_G)^2 N_G}{(r_I + r_N)^2 N_N} \quad (A1).$$

If  $\Gamma_{N:G} \approx 1$ , the ion's MFP in the neutral gas is about the same as it is in the grains. It is easier for the ion to pass through (a) the neutral gas if  $\Gamma_{N:G}$  is  $\gg 1$ , or, vice versa, (b) the grains if  $\Gamma_{N:G}$  is  $\ll 1$ . We calculate values for a range of estimated situations below, as detailed information for the grains throughout Saturn's magnetosphere is not complete in order to see if the grains should even be considered as possibly affecting the suprathermal ion populations.

Neutral Gas. For the neutral gas, we use the models of Cassidy & Johnson (2010) and Smith et al. (2010), which is accepted as being a reasonable representation of the neutral gas populations in Saturn's magnetosphere derived from Cassini measurements. Their model shows that while the O and OH number densities,  $N_O$  and  $N_{OH}$ , respectively, are nearly equal (to within  $\sim 25\%$ ) at Enceladus' orbit,  $N_O$  is  $\sim 3.5$  times higher than  $N_{OH}$  at Rhea's orbit. As they are the two most populous neutral species, we perform calculations for both O and OH, because, OH, with  $\sim 22\%$  of the plasma density near Rhea, can still produce measurable effects in that region.

Grains and Dust. Saturn's magnetosphere is filled with an icy, dusty plasma, and is the closest environment to Earth similar in any way to vast expanses of the interstellar medium. The radial variation of CDA data collected from 2004-200 to 2008-259 for nanometer grains and the fit to it in Figure 11 of Srama et al. (2011) is a clear, comprehensive, long-term display of magnetospheric-scale E-ring grain data. Srama et al. (2011) do not state whether data within several to tens of Enceladus radii, near the moon and its plume, were included in this long-term data set or not, but the display, a data collection intended to represent an overall average perspective of the E-ring grains, suggests that near-Enceladus data were not included. To our knowledge, no study E-ring data at/near the orbit of Enceladus, characterizes the azimuthal

variation of the grain population in a manner useful in our study. In their Figure 11, Srama et al. (2011) fit the radial variation of CDA densities with a power law, given as  $n(r) = 20 (r - 2.8)^{-4.6}$ , in the range  $\sim 4.5$ - $19 R_S$ , for data obtained from 3 to  $19 R_S$ . The  $n(r)$  fit is very close to the density averages plotted from  $\sim 5$  to  $\sim 15 R_S$  at intervals of  $1$ - $3 R_S$ . The mean density point plotted at  $< 5 R_S$  is far below their  $n(r)$  fit, consistent with their statement that, on account of dead time limitations, the CDA counting rates saturated at times near Enceladus' orbit,  $3.95 R_S$ , so that the derived number density from those saturated counting rates (as shown) at  $< \sim 4.5 R_S$  are a lower limit. The fall-off with increasing radial distance is similar to other reports (see e.g., Ye et al., 2016), so we take their fit in this radial range as a reliable representation of average-Enceladus orbit data which does not include the near-moon plume flux and, therefore, representative of the overall average E-ring population from  $\sim 3.95$  -  $15 R_S$  at locations not directly influenced by the Enceladus plume. Therefore, for our purposes, we can estimate the average grain density near the orbit of Rhea,  $8.75 R_S$ , from their  $n(r)$  fit in determining  $\Gamma_{N:G}$  near Rhea's orbit from measurements typically made near the orbit of Enceladus. Their measurements are for micrometer sized grains, whereas most of the grains are at nanometer sizes, at least near Enceladus. The values of  $\Gamma_{N:G}$  we obtain (see Tables S1-S3 in the Supporting Information) range within a factor of  $\sim 3$ - $4$  about 1 for characteristic grain sizes of 1 and 2 nm. This is not an exhaustive assessment and only elastic collisions were considered, as well as much of the necessary, relevant information particular to the grains and the gas is not yet well established, nor will be in the near future. Nevertheless, these estimates are not orders of magnitude higher or lower than unity, they are clustered around  $\sim 1$ , the value at which the  $H_2O^+$  mean free path in the dust/grain populations is approximately equal to that in the neutral gas populations. Finally, this exercise does suggest that, when the extant grain population sizes and gas and grain densities in Saturn's magnetosphere are known: that is, measured, confirmed, and characterized, the effect

666 of grains on Saturn's suprathermal ions and various other observables can be established and  
667 might be found to be of significance.

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## References

- Allegrini, F., McComas, D. J., Young, D. T., Berthelier, J. -J., Covinhes, J., Illiano, J. -M., et al. (2006). Energy loss of 1-50 keV H, He, C, N, O, Ne, and Ar ions transmitted through thin carbon foils. *Reviews of Scientific Instruments*, 77(4). <https://doi.org/10.1063/1.2185490>
- Armstrong, T. P., Paonessa, M. T., Bell II, E. V., & Krimigis, S. M. (1983). Voyager observations of Saturnian ion and electron phase space densities. *Journal of Geophysical Research*, 88, 8893-8904. <https://doi.org/10.1029/JA088iA11p08893>
- Botschwina, P., Rosmus, P., & Reinsch, E. A. (1983). Spectroscopic properties of the hydroxonium ion calculated from scep cepa wave functions. *Chemical Physics Letters*, 102, 299-306. [https://doi.org/10.1016/0009-2614\(83\)87045-6](https://doi.org/10.1016/0009-2614(83)87045-6)
- Branscomb, L. M., (1966). Photodetachment cross section, electron affinity, and structure of the negative hydroxyl ion. *Physical Review*, 148, 1, 11-18. <https://doi.org/10.1103/PhysRev.148.11>
- Burch, J. L., Goldstein, J., Hill, T. W., Young, D. T., Crary, F. J., Coates, A. J., et al. (2005). Properties of local plasma injections in Saturn's magnetosphere. *Geophysical Research Letters*, 32, L14S02. <https://doi.org/10.1029/2005GL022611>
- Burger, M. H., Sittler Jr., E. C., Johnson, R. E., Smith, H. T., Tucker, O. J., & Shematovich, V. I. (2007). Understanding the escape of water from Enceladus. *Journal of Geophysical Research*, 112, A06219. <https://doi.org/10.1029/2006JA012086>
- Cassidy, T. A., & Johnson, R. E. (2010). Collisional spreading of Enceladus' neutral cloud. *Icarus*, 209, 696-703. <https://doi.org/10.1016/j.icarus.2010.04.010>
- Christon, S. P., Hamilton, D. C., Mitchell, D. G., DiFabio, R. D., & Krimigis, S. M. (2014). Suprathermal magnetospheric minor ions heavier than water at Saturn: Discovery of <sup>28</sup>M<sup>+</sup> seasonal variations. *Journal of Geophysical Research*, 119, 7, pp. 5662-5673.

<https://doi.org/10.1002/2014JA020010>

Cravens, T. E., McNutt Jr., R. L., Waite Jr., J. H., Robertson, I. P., Luhmann, J. G., Kasprzak, W., & Ip, W. -H. (2009). Plume ionosphere of Enceladus as seen by Cassini ion and neutral mass spectrometer. *Geophysical Research Letters*, 36, L08106. <https://doi.org/10.1029/2009GL037811>

Dalgarno, A. (1958). The mobilities of ions in their parent gases. *Philosophical Transactions Royal Society London, Series A.*, 250, 426. <https://doi.org/10.1098/rsta.1958.0003>

Dialynas, K., Brandt, P. C., Krimigis, S. M., Mitchell, D. G., Hamilton, D. C., Krupp, N. et al. (2013). The extended Saturnian neutral cloud as revealed by global ENA simulations using Cassini/MIMI measurements. *Journal of Geophysical Research*, 118, 3027–3041. <https://doi.org/10.1002/jgra.50295>

DiFabio, R. D. (2012). Spatial and temporal variations of the suprathermal (3-220 keV/e) ion composition in Saturn's equatorial magnetosphere (Doctoral dissertation). Retrieved from Digital Repository at the University of Maryland, University of Maryland, College Park, Maryland, USA. <https://drum.lib.umd.edu/handle/1903/12700>

DiFabio, R. D., Hamilton, D. C., Krimigis, S. M., & Mitchell, D. G. (2011). Long term time variations of the suprathermal ions in Saturn's magnetosphere. *Geophysical Research Letters*, 38, L18103. <https://doi.org/10.1029/2011GL048841>

Dong, Y., Hill, T. W., Teolis, B. D., Magee, B. A., & Waite, J. H. (2011). The water vapor plumes of Enceladus. *Journal of Geophysical Research*, 116, A10204. <https://doi.org/10.1029/2011JA016693>

Dong, Y., & Hill, T. W. (2013). A study of the charged ice grains in the Enceladus plume with a composite size distribution. American Geophysical Union Fall Meeting, abstract, P32B-04. 2013AGUFM.P32B..04D.

- 717 Dong, Y., Hill, T. W., & Ye, S. -Y. (2015). Characteristics of ice grains in the Enceladus plume  
718 from Cassini observations. *Journal of Geophysical Research*, *120*, 915–937.  
719 <http://dx.doi.org/10.1002/2014JA020288>
- 720 Dyke, T. R., & Muentner, J. S. (1973). Electric dipole moments of low J states of H<sub>2</sub>O and D<sub>2</sub>O.  
721 *Journal of Chemical Physics*, *59*, 3125-3127. <https://doi.org/10.1063/1.1680453>
- 722 Engelhardt, I. A. D.; Wahlund, J. -E.; Andrews, D. J.; Eriksson, A. I.; Ye, S.; Kurth, W. S.; et al.  
723 (2015). Plasma regions, charged dust and field-aligned currents near Enceladus. *Planetary and*  
724 *Space Science*, *117*, 453-469. <https://doi.org/10.1016/j.pss.2015.09.010>
- 725 Farmer, A. J. (2009). Saturn in hot water: Viscous evolution of the Enceladus torus. *Icarus*, *202*,  
726 280-286. <https://doi.org/10.1016/j.icarus.2009.02.031>
- 727 Fleshman, B. L., Delamere, P. A., & Bagenal, F. (2010a). Modeling the Enceladus plume-plasma  
728 interaction. *Geophysical Research Letters*, *37*, L03202.  
729 <https://doi.org/10.1029/2009GL041613>.
- 730 Fleshman, B. L., Delamere, P. A., & Bagenal, F. (2010b). A sensitivity study of the Enceladus  
731 torus. *Journal of Geophysical Research*, *115*, E04007. <https://doi.org/10.1029/2009JE003372>
- 732 Fleshman, B. L., Delamere, P. A., Bagenal, F., and Cassidy, T. (2013). A 1-D model of physical  
733 chemistry in Saturn's inner magnetosphere. *Journal of Geophysical Research*, *118*, 1567–  
734 1581. <https://doi.org/doi:10.1002/jgre.20106>
- 735 González-Alfonso, E., Fischer, J., Bruderer, S., Müller, H. S. P., Graciá-Carpio, J., Sturm, E., et  
736 al. (2013). Excited OH<sup>+</sup>, H<sub>2</sub>O<sup>+</sup>, and H<sub>3</sub>O<sup>+</sup> in NGC 4418 and Arp 220. *Astronomy &*  
737 *Astrophysics*, *550*, A25, 1-23. <https://doi.org/10.1051/0004-6361/201220466>
- 738 Hansen, C. J., Esposito, L., Stewart, A. I. F., Colwell, J., Hendrix, A., Pryor, W., Shemansky, D.,  
739 & West, R. (2006). Enceladus' water vapor plume. *Science*, *311*, 1422-1425.  
740 <https://doi.org/10.1126/science.1121254>



- 741 Hill, T. W., Rymer, A. M., Burch, J. L., Crary, F. J., Young, D. T., Thomsen, M. F., et al. (2005).  
742 Evidence for rotationally driven plasma transport in Saturn's magnetosphere. *Geophysical*  
743 *Research Letters*, 32, L14S10. <https://doi.org/10.1029/2005GL022620>
- 744 Hill, T. W., Thomsen, M. F., Tokar, R. L., Coates, A. J., Lewis, G. R., Young, D. T., et al.  
745 (2012). Charged nanograins in the Enceladus plume. *Journal of Geophysical Research*, 117,  
746 A05209. <https://doi.org/doi:10.1029/2011JA017218>
- 747 Holmberg, M. K. G., Wahlund, J. -E., & Morooka, M. W. (2014). Dayside/nightside asymmetry  
748 of ion densities and velocities in Saturn's inner magnetosphere. *Geophysical Research Letters*,  
749 41, 3717–3723. <https://doi.org/10.1016/10.1002/2014GL060229>
- 750 Holmberg, M. K. G., Wahlund, J. -E., Morooka, M. W., & Persoon, A. M. (2012). Ion densities  
751 and velocities in the inner plasma torus of Saturn. *Planetary and Space Science*, 73, 1, 151-  
752 160. <https://doi.org/10.1016/j.pss.2012.09.016>
- 753 Holmberg, M. K. G., Shebanits, O., Wahlund, J.-E., Morooka, M., Vigren, E., André, N., et al.  
754 (2017). Density structures, dynamics, and seasonal and solar cycle modulations of Saturn's  
755 inner plasma disk. *Journal of Geophysical Research*, 122, 12,258-12,273.  
756 <https://doi.org/10.1002/2017JA024311>
- 757 Hood, L. L. (1989). Radial diffusion and losses of energetic protons in the 5 to 12 R<sub>S</sub> region of  
758 Saturn's magnetosphere. *Journal of Geophysical Research*, 94, A7, 8721-8730.  
759 <https://doi.org/10.1029/JA094iA07p08721>
- 760 Horányi, M., Juhász, A., & Morfill, G. E. (2008). Large-scale structure of Saturn's E-ring.  
761 *Geophysical Research Letters*, 35, L04203. <https://doi.org/10.1029/2007GL032726>
- 762 Ip, W.-H. (1997). On the neutral cloud distribution in the Saturnian magnetosphere. *Icarus*, 126,  
763 42-57. <https://doi.org/10.1006/icar.1996.5618>
- 764 Ip, W. -H. (2000). Thermal plasma composition in Saturn's magnetosphere. *Planetary & Space*

- 765 *Science.*, 48, 755-783. [https://doi.org/10.1016/S0032-0633\(00\)00036-2](https://doi.org/10.1016/S0032-0633(00)00036-2)
- 766 Jia, Y. -D., Russell, C. T., Khurana, K. K., Ma, Y. J., Kurth, W., & Gombosi, T. I. (2010).  
767 Interaction of Saturn's magnetosphere and its moon: 3. Time variation of the Enceladus plume.  
768 *Journal of Geophysical Research*, 115, A12243. <https://doi.org/10.1029/2010JA015534>
- 769 Johnson, R. E., Smith, H. T., Tucker, O. J., Liu, M., Burger, M. H., Sittler, et al. (2006). The  
770 Enceladus and OH tori at Saturn. *Astrophysical Journal*, 644, L137.  
771 <https://doi.org/10.1086/505750>
- 772 Jurac, S., & Richardson, J. D. (2005). A self-consistent model of plasma and neutrals at Saturn:  
773 Neutral cloud morphology. *Journal of Geophysical Research*, 110, A09220.  
774 <https://doi.org/10.1029/2004JA0106635>
- 775 Kempf, S., Beckmann, U. Moragas-Klostermeyer, G., Postberg, F., Srama, R., et al. (2008). The  
776 E ring in the vicinity of Enceladus I. Spatial distribution and properties of the ring particles.  
777 *Icarus*, 193, 420–437. <https://doi.org/10.1016/j.icarus.2007.06.027>
- 778 Kempf, S., Beckmann U., Srama, R., Horanyi, M., Auer, S., & Grun, E. (2006). The electrostatic  
779 potential of E ring particles. *Planetary and Space Science*, 54, 999–1006.  
780 <https://doi.org/10.1016/j.pss.2006.05.012>
- 781 Kempf, S., Srama, R., Altobelli, N., Auer, S., Tschernjawski, V., Bradley, et al. (2004). Cassini  
782 between Earth and asteroid belt: first in-situ charge measurements of interplanetary grains.  
783 *Icarus*, 171, 2, 317-335. <https://doi.org/10.1016/j.icarus.2004.05.017>
- 784 Kollmann, P., E. Roussos, C. Paranicas, N. Krupp, C. M. Jackman, E. Kirsch, and K. •H.  
785 Glassmeier (2011). Energetic particle phase space densities at Saturn: Cassini observations and  
786 interpretations. *Journal of Geophysical Research*, 116, A05222.  
787 <https://doi.org/10.1029/2010JA016221>
- 788 Krimigis, S. M., Mitchell, D. G., Hamilton, D. C., Liv, S., Dandouras, J., Jaskulek, et al. (2004).

- 789 Magnetospheric Imaging Instrument (MIMI) on the Cassini Mission to Saturn/Titan. *Space*  
790 *Science Reviews*, 114, 233-329. <https://doi.org/10.1007/s11214-004-1410-8>.
- 791 Krimigis, S. M., Mitchell, D. G., Hamilton, D. C., Krupp, N., Livi, S., Roelof, E. C., et al.  
792 (2005). Dynamics of Saturn's magnetosphere from MIMI during Cassini's orbital insertion.  
793 *Science*, 307, 5713, 1270-1273. <https://doi.org/10.1126/science.1105978>
- 794 Lishawa, C. R., Dressler, R. A., Gardner, J. A., Salter, R. H., & Murad, E. (1990). Cross sections  
795 and product kinetic energy analysis of H<sub>2</sub>O<sup>+</sup>-H<sub>2</sub>O collisions at suprathermal energies. *Journal*  
796 *of Chemical Physics*, 93, 3196-3206; <https://doi.org/10.1063/1.458852>
- 797 Livi, R., Goldstein J., Burch, J. L., Crary, F., Rymer, A. M., Mitchell, D. G., & Persoon, A. M.  
798 (2014). Multi-instrument analysis of plasma parameters in Saturn's equatorial, inner  
799 magnetosphere using corrections for spacecraft potential and penetrating background radiation.  
800 *Journal of Geophysical Research*, 119, 3683-3707. <https://doi.org/10.1002/2013JA019616>
- 801 Marcus, Y. (2012). Volumes of aqueous hydrogen and hydroxide ions at 0 to 200 °C. *Journal of*  
802 *Chemical Physics*, 137, 15, 154501-254501-5. <https://doi.org/10.1063/1.4758071>
- 803 Massey, H. S. W. (1949). Collisions between atoms and molecules at ordinary temperatures.  
804 *Rep. Progr. Phys.*, 12, 248. <https://doi.org/10.1088/0034-4885/12/1/311>
- 805 Mauk, B. H., Saur, J., Mitchell, D. G., Roelof, E. C., Brandt, P. C., Armstrong, T. P., et al.  
806 (2005). Energetic particle injections in Saturn's magnetosphere. *Geophysical Research Letters*,  
807 32, L14S05. <https://doi.org/10.1029/2005GL022485>
- 808 Mauk, B. H., Hamilton, D. C., Hill, T. W., Hospodarsky, G. B., Johnson, R. E., Paranicas, C., et  
809 al. (2009). Fundamental plasma processes in Saturn's magnetosphere. In M. K. Dougherty, L.  
810 W. Esposito, & S. M. Krimigis (Eds.), *Saturn from Cassini-Huygens* (pp. 281-331). B.V:  
811 Springer. [https://doi.org/10.1007/978-1-4020-9217-6\\_11](https://doi.org/10.1007/978-1-4020-9217-6_11)
- 812 McAndrews, H. J., Thomsen, M. F., Arridge, C. S., Jackman, C. M., Wilson, R. J., Henderson,

- 813 M. G., & Dougherty, M. K. (2009). Plasma in Saturn's nightside magnetosphere and the  
814 implications for global circulation. *Planetary and Space Science*, 57 (14-15), 1714.  
815 <https://doi.org/10.1016/j.pss.2009.03.003>
- 816 Meier, P., Kriegel, H., Motschmann, U., Schmidt, J., Spahn, F., Hill, T. W., et al. (2014). A  
817 model of the spatial and size distribution of Enceladus' dust plume. *Planetary and Space*  
818 *Science*, 104, 216–233. <http://dx.doi.org/10.1016/j.pss.2014.09.016>
- 819 Melin, J., Ortiz, J. V., Martín, I., Velasco A. M., & Lavín, C. (2005). Ground and excited states  
820 of the Rydberg radial H3O: Electron propagator and quantum defect analysis. *Journal of*  
821 *Chemistry and Physics*, 122, 23, 234317-234317-7. <https://doi.org/10.1063/1.1926286>
- 822 Melin, H., Shemansky, D. E., & Liu, X. (2009). The distribution of atomic hydrogen and oxygen  
823 in the magnetosphere of Saturn. *Planetary and Space Science*, 57, 1743-1753.  
824 <https://doi.org/10.1016/j.pss.2009.04.014>
- 825 Mitchell, D. G., Carbary, J. F., Cowley, S. W. H., Hill, T. W., & Zarka, P. (2009a). *The*  
826 *Dynamics of Saturn's magnetosphere*. In M. K. Dougherty, L. W. Esposito, & S. M. Krimigis  
827 (Eds.), *Saturn from Cassini-Huygens* (pp. 257-279). B.V: Springer.  
828 [https://doi.org/10.1007/978-1-4020-9217-6\\_10](https://doi.org/10.1007/978-1-4020-9217-6_10)
- 829 Mitchell, D. G., Krimigis, S. M., Paranicas, C., Brandt, P. C., Carbary, J. F., Roelof, E. C., et  
830 al.(2009b). Recurrent energization of plasma in the midnight-to-dawn quadrant of Saturn's  
831 magnetosphere and its relationship to auroral UV and radio emissions. *Planetary and Space*  
832 *Science*, 57, 1732. <https://doi.org/10.1016/j.pss.2009.04.002>
- 833 Nemeth, Z., Szego, K., Foldy, L., Kivelson, M. G., Jia, X., Ramer, K. M., et al. (2015). The  
834 latitudinal structure of the nightside outer magnetosphere of Saturn as revealed by velocity  
835 moments of thermal ions, *Annales Geophysicae*, 33, 1195–1202.  
836 <https://doi.org/10.5194/angeo-33-1195-2015>

- 837 Northcliffe, L. C., & Schilling, R. F. (1970). Range and stopping-power tables for heavy ions.  
838 Atomic Data and Nuclear Data Tables, 7, 3-4, 233-463. <https://doi.org/10.1016/S0092->  
839 640X(70)80016-X
- 840 Paranicas, C., Mitchell, D. G., Krimigis, S. M., Hamilton, D. C., Roussos, E., Krupp, N., et al.  
841 (2008). Sources and losses of energetic protons in Saturn's magnetosphere. *Icarus* 197, 2, 519-  
842 525. <https://doi.org/10.1016/j.icarus.2008.05.011>
- 843 Paranicas, C., Mitchell, D. G., Roelof, E. C., Mauk, B. H., Krimigis, S. M., Brandt, P. C., et al.  
844 (2007). Energetic electrons injected into Saturn's neutral cloud. *Geophysical Research Letters*,  
845 34, L02109. <https://doi.org/10.1029/2006GL028676>
- 846 Persoon, A. M., Gurnett, D. A., Leisner, J. S., Kurth, W. S., Groene, J. B., & Faden, J. B. (2013).  
847 The plasma density distribution in the inner region of Saturn's magnetosphere, *Journal of*  
848 *Geophysical Research*, 118, 2970–2974, <https://doi.org/10.1002/jgra.50182>
- 849 Persoon, A. M., Gurnett, D. A., Santolik, O., Kurth, W. S., Faden, J. B., Groene, J. B., et al.  
850 (2009). A diffusive equilibrium model for the plasma density in Saturn's magnetosphere,  
851 *Journal of Geophysical Research*, 114, A04211, <https://doi.org/10.1029/2008JA013912>
- 852 Porco, C. C., Helfenstein, P., Thomas, P. C., Ingersoll, A. P., Wisdom, J., West, R., et al. (2006).  
853 Cassini observes the active south pole of Enceladus, *Science*, 311, 1393.  
854 <https://doi.org/10.1126/science.1123013>
- 855 Press, W. H., Flannery, B. P., Teukolsky, S. A., & Vetterling, W. T. (1989). *Numerical Recipes:*  
856 *The Art of Scientific Computing*, Cambridge University Press, New York
- 857 Rapp, D., & W. E. Francis (1962). Charge exchange between gaseous ions and atoms. *Journal*  
858 *Chemical Physics*, 37, 2631. <https://doi.org/10.1063/1.1733066>
- 859 Richardson, J. D. (1998). Thermal plasma and neutral gas in Saturn's magnetosphere. *Reviews of*  
860 *Geophysics*, 36, 501-524. <https://doi.org/10.1029/98RG01691>

- 861 Richardson, J. D., Eviatar, A., McGrath, M. A., & Vasyliunas V. M. (1998). OH in Saturn's  
862 magnetosphere: Observations and implications. *Journal of Geophysical Research*, 103, 20245.  
863 <https://doi.org/10.1029/98JE01127>
- 864 Rymer, A. M., Mauk, B. H., Hill, T. W., Paranicas, C., André, N., Sittler Jr., E. C., et al. (2007).  
865 Electron sources in Saturn's magnetosphere. *Journal of Geophysical Research*, 112, A02201.  
866 <https://doi.org/10.1029/2006JA012017>
- 867 Rymer, A. M., Mauk, B. H., Hill, T. W., Paranicas, C., Mitchell, D. G., Coates, A. J., & Young,  
868 D. T. (2008). Electron circulation in Saturn's magnetosphere. *Journal of Geophysical*  
869 *Research*, 113, A01201. <https://doi.org/10.1029/2007JA012589>
- 870 Rymer, A. M., Mauk, B. H., Hill, T. W., André, N., Mitchell, D. G., Paranicas, C., et al. (2009).  
871 Cassini evidence for rapid interchange transport at Saturn. *Planetary and Space Science*, 57,  
872 1779. <https://doi.org/10.1016/j.pss.2009.04.010>
- 873 Saur, J., N., Schilling, Neubauer, F. M., Strobel, D. F., Simon, S., Dougherty, M. K., et al.  
874 (2008). Evidence for temporal variability of Enceladus' gas jets: Modeling of Cassini  
875 observations. *Geophysical Research Letters*, 35, L20105.  
876 <https://doi.org/10.1029/2008GL035811>
- 877 Schippers, P., Blanc, M., André, N., Dandouras, I., Lewis, G. R., Gilbert, L. K., et al. (2008).  
878 Multi-instrument analysis of electron populations in Saturn's magnetosphere. *Journal of*  
879 *Geophysical Research*, 113, A07208. <https://doi.org/10.1029/2008JA013098>
- 880 Sergis, N., Krimigis, S. M., Mitchell, D. G., Hamilton, D. C., Krupp, N., Mauk, B. M., et al.  
881 (2007). Ring current at Saturn: Energetic particle pressure in Saturn's equatorial  
882 magnetosphere measured with Cassini/MIMI. *Geophysical Research Letters*, 34, L09102.  
883 <https://doi.org/10.1029/2006GL029223>
- 884 Sergis, N., Krimigis, S. M., Mitchell, D. G., Hamilton, D. C., Krupp, N., Mauk, B. M., et al.

- 885 (2009). Energetic particle pressure in Saturn's magnetosphere measured with the  
886 Magnetospheric Imaging Instrument on Cassini. *Journal of Geophysical Research*, 114,  
887 A02214. <https://doi.org/10.1029/2008JA013774>
- 888 Shemansky, D. E., Matheson, P., Hall, D. T., Hu, H. -Y., & Tripp, T. M. (1993). Detection of the  
889 hydroxyl radical in the Saturn magnetosphere. *Nature*, 363, 329-331.  
890 <https://doi.org/10.1038/363329a0>
- 891 Sergis, N., Krimigis, S. M., Roelof, E. C., Arridge, C. S., Rymer, A. M., Mitchell, D. G. et al.  
892 (2010). Particle pressure, inertial force, and ring current density profiles in the magnetosphere  
893 of Saturn, based on Cassini measurements. *Geophysical Research Letters*, 37, L02102.  
894 <https://doi.org/10.1029/2009GL041920>
- 895 Showalter, M. R., Cuzzi, J. N., Larson, S. M. (1991). Structure and particle properties of Saturn's  
896 E Ring. *Icarus*, 94, 2, 451-473. [https://doi.org/10.1016/0019-1035\(91\)90241-K](https://doi.org/10.1016/0019-1035(91)90241-K)
- 897 Sittler Jr., E. C., Andre, N., Blanc, M., Burger, M., Johnson, R. E., Coates, A., et al. (2008). Ion  
898 and neutral sources and sinks within Saturn's inner magnetosphere: Cassini results. *Planetary*  
899 *and Space Science*, 56, 3. <https://doi.org/10.1016/j.pss.2007.06.006>
- 900 Slater, J. C., (1964). Atomic Radii in Crystals. *Journal of Chemical Physics*, 41, 10, 3199-3204.  
901 <https://org/doi/10.1063/1.1725697>
- 902 Smith, H. T., Johnson, R. E., Perry, Mitchell, M. E., D. G., McNutt, R. L., & Young, D. T.  
903 (2010). Enceladus plume variability and the neutral gas densities in Saturn's magnetosphere.  
904 *Journal of Geophysical Research* 115, A10252. <https://doi.org/10.1029/2009JA015184>
- 905 Srama, R., Ahrens, T. J., Altobelli, N., Auer, S., Bradley, J. G., Burton, M., et al. (2004). The  
906 Cassini Cosmic Dust Analyzer. *Space Science Reviews*, 114, 1-4, 465-518.  
907 <https://doi.org/10.1007/s11214-004-1435-z>

- 908 Srama, R., Kempf, S., Moragas-Klostermeyer, G., Altobelli, N., Auer, S., Beckmann, U., et al.  
909 (2011). The cosmic dust analyser onboard Cassini: ten years of discoveries. *CEAS Space*  
910 *Journal*, 2, 1-4, 3-16. <https://doi.org/10.1007/s12567-011-0014-x>
- 911 Thomsen, M. F., Reisenfeld, D. B., Delapp, D. M., Tokar, R. L., Young, D. T., Crary, F. J., et al.  
912 (2010). Survey of ion plasma parameters in Saturn's magnetosphere. *Journal of Geophysical*  
913 *Research*, 115, A10220. <https://doi.org/10.1029/2010JA015267>
- 914 Tokar, R. L., Johnson, R. E., Hill, T. W., Pontius, D. H., Kurth, W. S., Crary, F. J., et al. (2006).  
915 The interaction of the atmosphere of Enceladus with Saturn's plasma. *Science*, 311, 1409.  
916 <https://doi.org/10.1126/science.1121061>
- 917 Tokar, R. L., Johnson, R. E., Thomsen, M. F., Wilson, R. J., Young, D. T., Crary, et al. (2009).  
918 Cassini detection of Enceladus' cold water-group plume ionosphere. *Geophysical Research*  
919 *Letters*, 36, L13203. <https://doi.org/10.1029/2009GL038923>
- 920 Van Allen, J. A., Randall, B. A., & Thomsen, M. F. (1980). Sources and sinks of energetic  
921 electrons and protons in Saturn's magnetosphere. *Journal of Geophysical Research*, 85, 5679-  
922 5694. <https://doi.org/10.1029/JA085iA11p05679>
- 923 Vandegriff, J., DiFabio, R., Hamilton, D., Kusterer, M., Manweiler, J., Mitchell, D., et al. (2018).  
924 Cassini/MIMI Instrument Data User Guide. NASA, Washington, D. C. Available at  
925 <http://pds.nasa.gov>
- 926 von Steiger, R., Schwadron, N. A., Fisk, L. A., Geiss, J., Gloeckler, G., Hefti, et al. (2000).  
927 Composition of quasi-stationary solar wind flows from Ulysses/Solar Wind Ion Composition  
928 Spectrometer. *Journal of Geophysical Research*, 105, 27,217-27,238.  
929 <https://doi.org/10.1029/1999JA000358>
- 930 Waite Jr., J. H., Combi, M. R., Ip, W. -H., Cravens, T. E., McNutt Jr., R. L., Kasprzak, W., et  
931 al.(2006). Cassini ion and neutral mass spectrometer: Enceladus plume composition and



- structure. *Science*, 311, 1419-1422. <https://doi.org/10.1126/science.1121290>
- Walt, M. (1971). The radial diffusion of trapped particles induced by fluctuating magnetospheric fields. *Space Science Reviews*, 12, 4, 446-485. <https://doi.org/10.1007/BF00171975>
- Weast, R. C. (Ed.) (1968). *CRC Handbook of Chemistry and Physics*, 49th Ed., p. E-71, Cleveland.
- Weis, B., Carter, S., Rosmus, P., Werner, H.-J., & Knowles, P. J. (1989). A theoretical rotationally resolved infrared spectrum for H<sub>2</sub>O<sup>+</sup> (X 28 1). *Journal of Chemical Physics*, 91, 2818-2833. <https://doi.org/10.1063/1.456951>
- Werner, H.-J., Rosmus, P., & Reinsch, E.-A. (1983). Molecular properties from MCSCF-SCEP wave functions. I. Accurate dipole moment functions of OH, OH<sup>-</sup>, and OH<sup>+</sup>. *Journal of Chemical Physics*, 79, 905-916. <https://doi.org/10.1063/1.445867>
- Wiedmann, R. T., Tonkyn, R. G., White, M. G., Kwanghsi, W., & McKoy, V. (1992). Rotationally resolved threshold photoelectron spectra of OH and OD. *Journal Chemical Physics*, 97, 768-772. <https://doi.org/10.1063/1.463179>
- Wilson, R. J., Bagenal, F., Cassidy, T., Fleshman, B. L., & Crary, F. (2015), The relative proportions of water group ions in Saturn's inner magnetosphere: A preliminary study, *Journal Geophysical Research, Space Physics*, 120, 6624–6632, <https://doi.org/10.1002/2014JA020557>
- Wilson, R. J., Tokar, M. G., Henderson, R. L., Hill, T. W, Thomsen M. F., & Pontius, D. H. (2008). Cassini plasma spectrometer thermal ion measurements in Saturn's inner magnetosphere. *Journal of Geophysical Research*, 113, A12218. <https://doi.org/10.1029/2008JA013486>
- Ye, S.-Y., Gurnett, D. A., Kurth, W. S., Averkamp, T. F., Kempf, S., Hsu, H.-W., Srama, R., & Grün, E. (2014). Properties of dust particles near Saturn inferred from voltage pulses induced by dust impacts on Cassini spacecraft. *Journal of Geophysical Research*, 119, 6294–6312.

956 <https://doi.org/10.1002/2014JA020024>

957 Young, D. T., Berthelier, J.-J., Blanc, M., Burch, J. L., Bolton, S., Coates, A. J., et al. (2005).

958 Composition and dynamics of plasma in Saturn's magnetosphere. *Science*, 307, 1262-1266.

959 <https://doi.org/10.1126/science.1106151>

960 Ye, S.-Y., D. A. Gurnett, W. S. Kurth, T. F. Averkamp, S. Kempf, H.-W. Hsu, R. Srama, and E.

961 Grün (2014), Properties of dust particles near Saturn inferred from voltage pulses induced by

962 dust impacts on Cassini spacecraft, *J. Geophys. Res. Space Physics*, 119, 6294–6312,

963 doi:10.1002/2014JA020024

964 Ye, S.-Y., W. S. Kurth, G. B. Hospodarsky, T. F. Averkamp, and D. A. Gurnett (2016). Dust

965 detection in space using the monopole and dipole electric field antennas, *J. Geophys. Res. Space*

966 *Physics*, 121, 11,964–11,972. doi:10.1002/2016JA023266.

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## 976 Figure Captions

977

978 Figure 1. (a) Essd channel, the Solid State Detector (SSD) energy channel number, versus TOF  
979 channel, the time-of-flight channel number of the  $\sim 96$  keV  $W^+$ . (b) The  $\sim 96$  keV  $W^+$  TOF  
980 channel count histogram for the data in the Essd channel range indicated by the horizontal red  
981 dashed lines in (a). Data from 78 equatorial ring current passes from 2004 to 2010 are summed.  
982 These data are from CHEMS telescope 1 which has the best heavy ion resolution (DiFabio,  
983 2012).

984 Figure 2.  $\sim 96$  keV  $O^{++}$  ( $E/Q = 48$  keV/e) TOF Channel count distribution for all 78 equatorial  
985 ring current passes from 2004 to 2010 ( $L = 7-16$ , Latitude =  $-10^\circ$  to  $10^\circ$ ) combined. The solid red  
986 line represents the best fit to the distribution using Equation 1 (See text). We use an asymmetric  
987 Gaussian to fit the counts near the center of the distribution and add a small correction to the  
988 Gaussian on the right side to fit the counts in the “tail.”

989 Figure 3. Best fit of the  $\sim 96$  keV  $W^+$  distribution. The fits to each individual species are shown  
990 along with the sum of the four fits (red). The data are from CHEMS telescope 1.

991 Figure 4. (a) The partial number density of energetic, 96 keV, water group components versus  $L$ .  
992 (b) The fractional abundance of 96 keV water group components versus  $L$ . The error bars of  $O^+$ ,  
993  $OH^+$ , and  $H_2O^+$  represent the statistical uncertainties from the best fit. The error bars of  $H_3O^+$   
994 represent the spread of values from our three different fits (see text).

995 Figure 5. (top 3 panels) The water-group fractional abundances plotted over model E-ring grain  
996 (Hornáyi et al., 2008) and OH density (Jurac et al., 2005) information. (bottom panel) Grain  
997 charge information from Kempf et al. (2008) and Ye et al. (2014).

998 Equations 1 (top) and 2 (bottom)

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$$f(\tau) = A \left[ e^{-\frac{(\tau - \tau_c)^2}{2\sigma_1^2}} P(-\tau + \tau_c) + \left( e^{-\frac{(\tau - \tau_c)^2}{2\sigma_2^2}} + \frac{(\tau - \tau_c)^2}{100 + [(\tau - \tau_c)^2]} \right) P(\tau - \tau_c) \right] \quad (1)$$

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$$v_{\max} = \frac{v_o |\Delta E|}{h} \quad (2)$$

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Table 1: Partial number density and relative abundance of the four major ~96 keV W<sup>+</sup> components in Saturn's equatorial ring current region (L = 7-16)<sup>a</sup>.

Species	Counts	Partial Number Density (x10 <sup>-6</sup> cm <sup>-3</sup> )	Fraction of W <sup>+</sup>
O <sup>+</sup>	14400	23.4 ± 0.7	0.53 ± 0.02
OH <sup>+</sup>	5743	9.6 ± 0.4	0.22 ± 0.01
H <sub>2</sub> O <sup>+</sup>	5607	9.7 ± 0.4	0.22 ± 0.01
H <sub>3</sub> O <sup>+</sup>	690	1.2 +0.4, -0.48 <sup>b</sup>	0.028 +0.01, -0.012 <sup>b</sup>

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a The abundances are determined using the results from our best fit ( $\alpha = 1.90$ ). The uncertainties of O<sup>+</sup>, OH<sup>+</sup>, and H<sub>2</sub>O<sup>+</sup> reflect the statistical uncertainties from our fits.

b The H<sub>3</sub>O<sup>+</sup> uncertainty reflects the range of values from our three fits (see text).

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1022 Table 2: Component to Total\* W<sup>+</sup> Fractional Abundance Ratio Comparisons

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W <sup>+</sup> Analysis	O <sup>+</sup> /W <sup>+</sup>	OH <sup>+</sup> /W <sup>+</sup>	H <sub>2</sub> O <sup>+</sup> /W <sup>+</sup>	H <sub>3</sub> O <sup>+</sup> /W <sup>+</sup>
This Study (Suprathermal Ions) (~ 7 ≤ L ≤ ~16)	0.52-0.55	0.20-0.22	0.22-0.22	0.047-0.023
Wilson et al. (2015) (Thermal Ions) (~ 7 ≤ L ≤ ~10)	0.20-0.24	0.28-0.30	0.29-0.28	0.23-0.18
Fleshman et al. (2013) (Model) (~ 7 ≤ L ≤ ~10)	0.24-0.33	0.19-0.22	0.29-0.26	0.28-0.19

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1027 \* [W<sup>+</sup>] = [O<sup>+</sup>] + [OH<sup>+</sup>] + [H<sub>2</sub>O<sup>+</sup>] + [H<sub>3</sub>O<sup>+</sup>], the W<sup>+</sup> population is the sum of the O<sup>+</sup>, OH<sup>+</sup>, H<sub>2</sub>O<sup>+</sup>,  
 1028 and H<sub>3</sub>O<sup>+</sup> populations

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Figure 1.

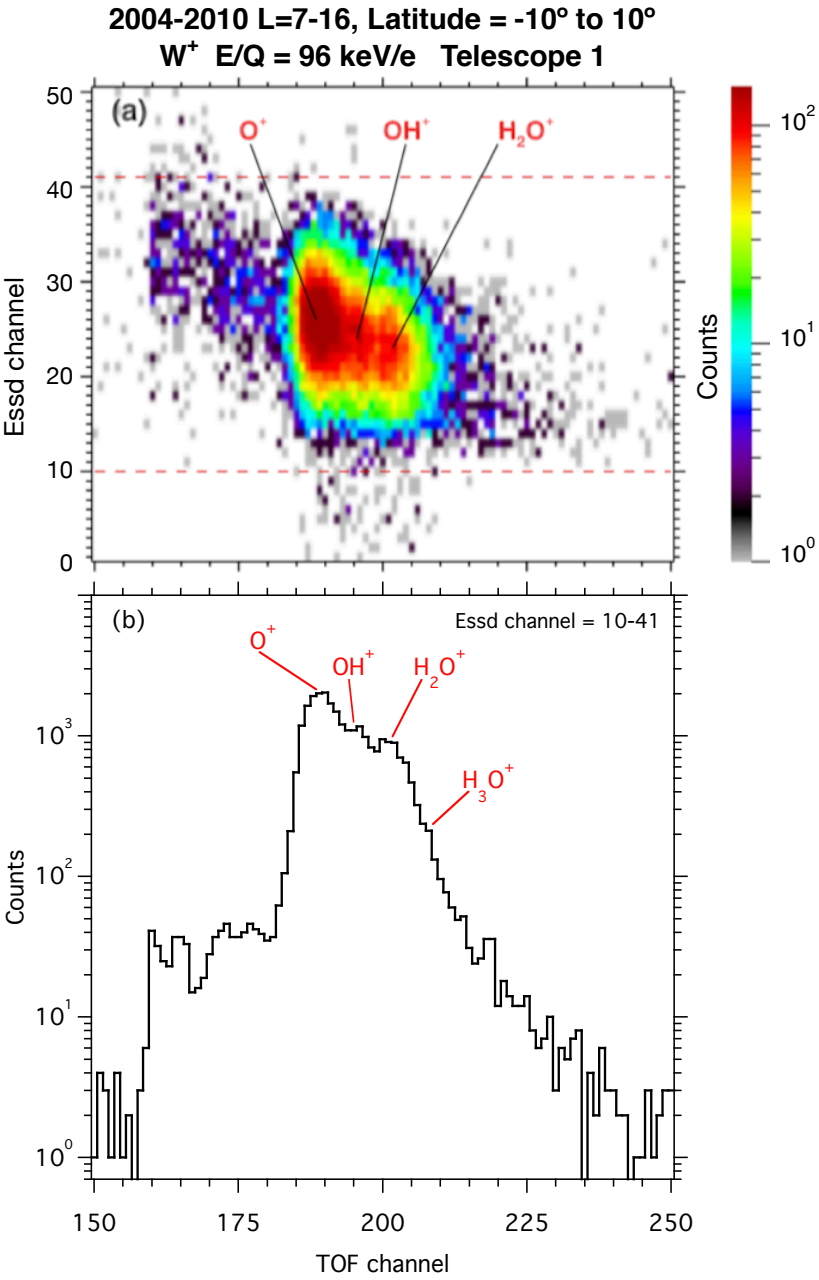




Figure 2.

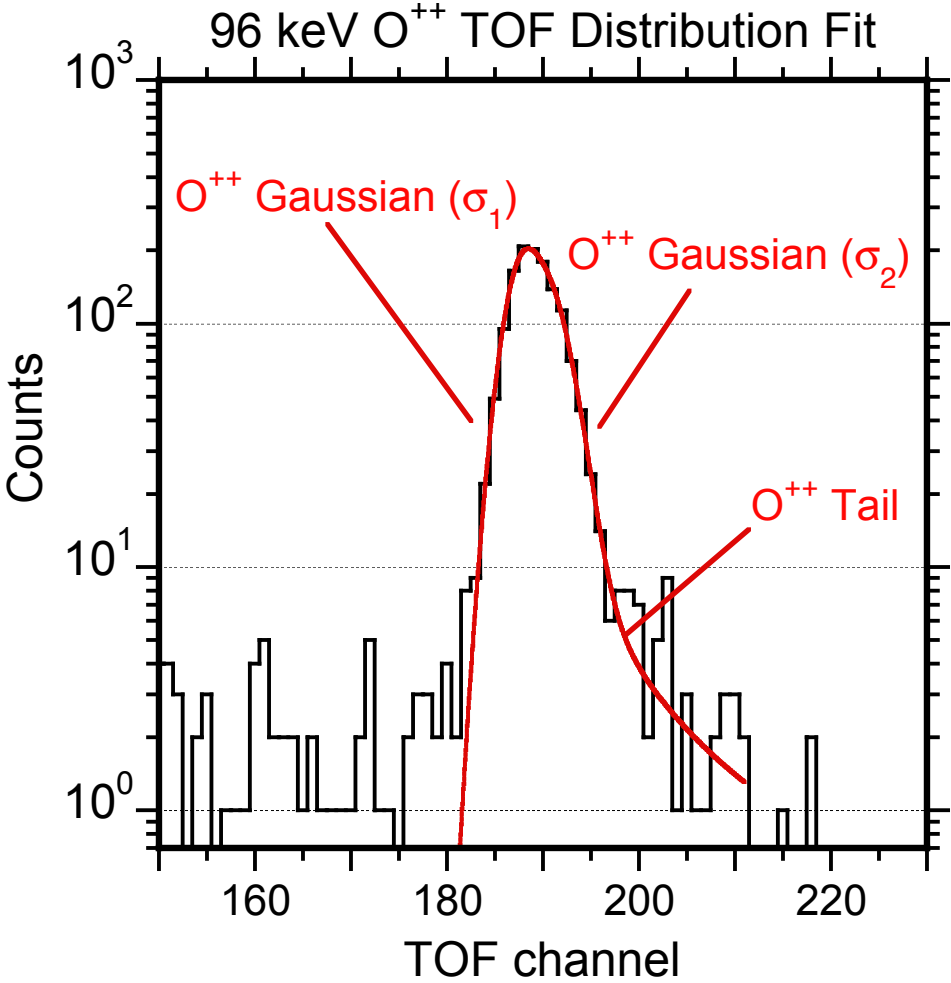


Figure 3.

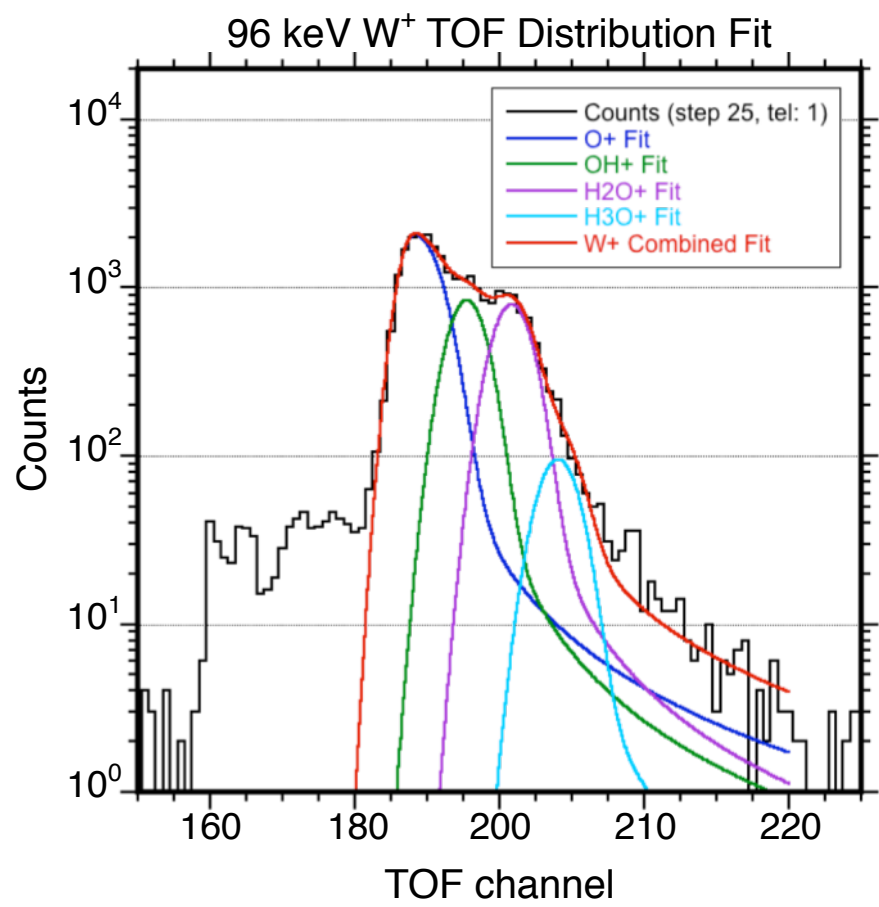


Figure 4.

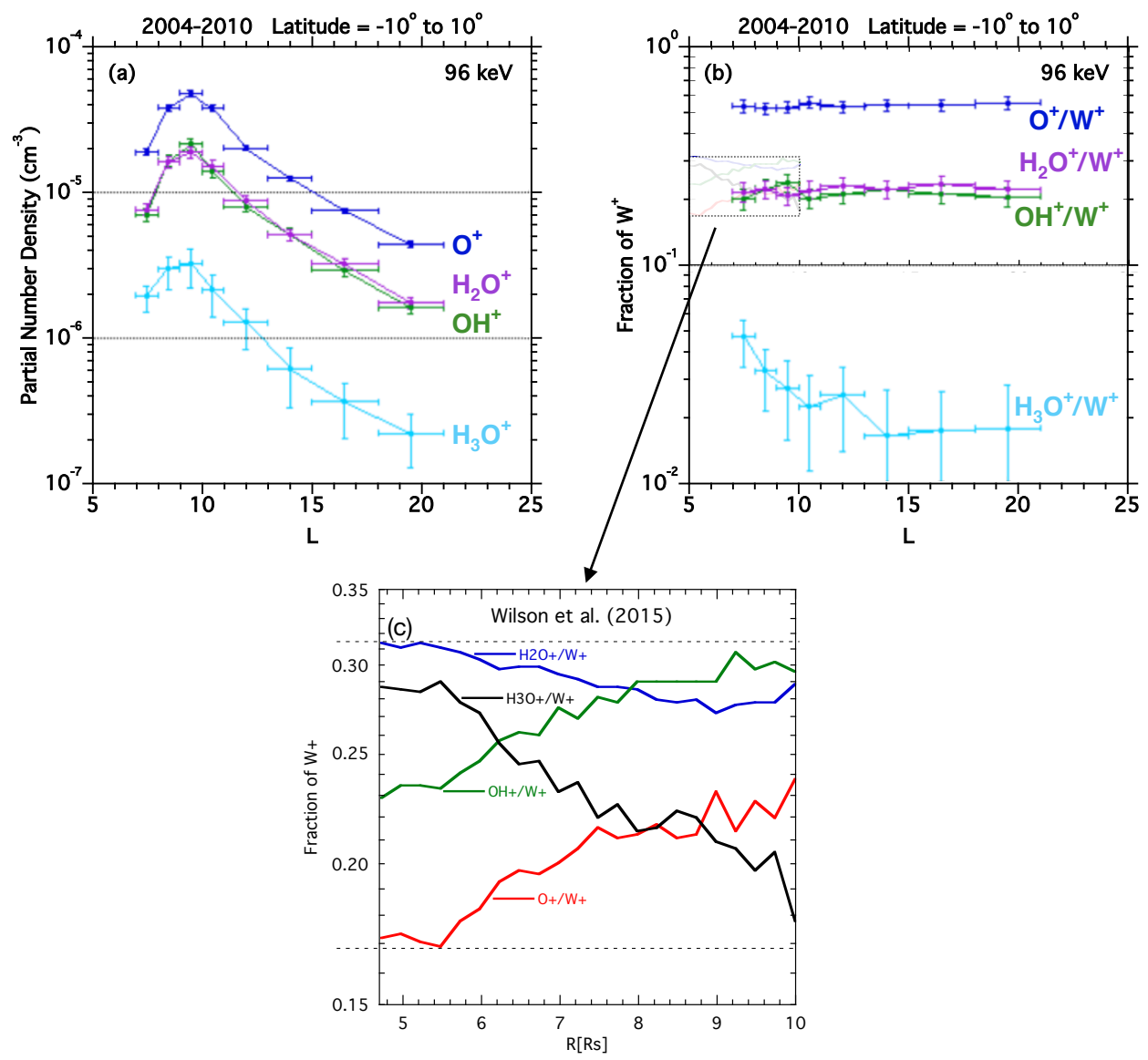


Figure 5.

