

**Ammonium sensitivity of biological nitrogen fixation in anaerobic diazotrophs and coastal salt marsh sediments.**

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Key points:

- We identified a porewater ammonium concentration threshold of 9  $\mu\text{M}$  for biological nitrogen fixation inhibition in benthic environments.
- Drawdown of porewater ammonium, timing of enzyme inhibition, and sediment heterogeneity can complicate measurement of nitrogen sensitivity.
- Review of porewater ammonium concentration indicates biological nitrogen fixation is likely limited to the superficial layer of sediments.

## Abstract:

New bioavailable nitrogen (N) from biological nitrogen fixation (BNF) is critical for the N budget and productivity of marine ecosystems. Nitrogen-fixing organisms typically inactivate BNF when less metabolically costly N sources, like ammonium ( $\text{NH}_4^+$ ), are available. Yet, several studies observed BNF in benthic marine sediments linked to anaerobic sulfate-reducing bacteria (SRB) and fermenting firmicutes despite high porewater  $\text{NH}_4^+$  concentrations (10-1,500  $\mu\text{M}$ ), making the importance of and regulating controls on benthic BNF unclear. Here, we evaluate BNF sensitivity to  $\text{NH}_4^+$  in model anaerobic diazotrophs, the sulfate-reducer *Desulfovibrio vulgaris* var. Hildenborough and fermenter *Clostridium pasteurianum* strain W5; in sulfate-reducing sediment enrichment cultures, and in sediment slurry incubations from three Northeastern salt marshes (USA). BNF in sulfate-reducing cultures and sediments is highly sensitive to external  $\text{NH}_4^+$ , with a threshold for BNF inhibition of  $[\text{NH}_4^+] < 2 \mu\text{M}$  in cultures and  $< 9 \mu\text{M}$  in sediment slurries. The prevalence of SRB-like sequences in sediment-derived nitrogenase (*nifH*) genes and transcripts in this and other studies of benthic BNF along with an analysis of benthic  $\text{NH}_4^+$  porewater data suggests a broad applicability of the inhibition thresholds measured here and the confinement of benthic BNF to surficial sediments. The timing of inhibition, fast  $\text{NH}_4^+$  drawdown, and sediment heterogeneity are factors that can complicate studies of benthic BNF sensitivity to  $\text{NH}_4^+$ . We propose a simple theoretical framework based on the affinity of the  $\text{NH}_4^+$  transporter to explain  $\text{NH}_4^+$  control of BNF and improve biogeochemical models of N cycling.

## 44 Introduction

45           The requirement for bioavailable fixed nitrogen (N) is a fundamental constraint for all  
46 life on Earth. Biological nitrogen fixation (BNF), the only biological process capable of  
47 producing newly fixed N, is critical for the function and ecology of the biosphere. Present in a  
48 small subset of taxonomically and metabolically diverse prokaryotes termed diazotrophs, all  
49 BNF is catalyzed by the enzyme nitrogenase in a complex, energy-intensive reaction that  
50 reduces dinitrogen ( $N_2$ ) into ammonia ( $NH_3$ ), which is subsequently incorporated into cell  
51 material. Diazotrophs thus function as natural gatekeepers for ecosystem N input and are  
52 particularly important in oligotrophic systems, such as the surface waters of the tropical and  
53 subtropical ocean (Sohm et al., 2011; Voss et al., 2013; Zehr & Kudela, 2011), where low,  
54 nanomolar levels of dissolved inorganic fixed N species (DIN *i.e.*, ammonium  $NH_4^+$ , nitrite  
55  $NO_2^-$ , nitrate  $NO_3^-$ ) limit primary production (Elser et al., 2007).

56           The distribution of BNF in marine waters was traditionally explained by the well-  
57 documented regulation of nitrogenase synthesis and activity by DIN in model cyanobacterial  
58 and aerobic diazotrophs (*e.g.*, *Trichodesmium* and *Azotobacter*), which downregulate BNF  
59 when ambient DIN concentrations are sufficient to meet cell anabolic demands (*e.g.*,  
60 external  $[NH_4^+] >$  hundreds of nanomolar to tens of micromolar, depending on the species  
61 (Drozd et al., 1972; Hartmann et al., 1986; Holl & Montoya, 2005; Kleiner, 1974; Mulholland  
62 et al., 2001; Ohki et al., 1991; Postgate & Kent, 1984). This downregulation reflects the high  
63 energetic cost of N acquisition through BNF, in comparison to the metabolic cost of  
64 assimilating  $NH_4^+$  or  $NO_3^-$  (Glibert et al., 2016; Großkopf & LaRoche, 2012; Inomura et al.,  
65 2017). Accordingly, marine BNF in hypoxic and anoxic benthic environments, where  
66 porewater  $NH_4^+$  concentrations are typically orders of magnitude higher than in euphotic  
67 zones (see Knapp and reference therein) (Knapp, 2012), has been estimated to account for  
68 only 10% of total marine new N inputs, with benthic BNF largely restricted to highly  
69 productive shallow sedimentary environments like seagrass meadows (Capone, 1988;  
70 Capone & Carpenter, 1982; Zhang et al., 2020).

However, recent studies ranging from open ocean to coastal estuary benthic systems have challenged the above paradigm based on observations of BNF activity and nitrogenase (*nifH*, *nitrogenase reductase*) genes and transcripts in sediments with high porewater  $\text{NH}_4^+$  concentrations (e.g. ~ 50-1500  $\mu\text{M}$ ) (Bertics et al., 2013; Gier et al., 2016; Knapp, 2012). These results, many obtained with closed vessel sediment incubation experiments, have led to the proposal that benthic BNF may be a larger contribution to marine N inputs than previously thought (Fulweiler et al., 2007; Newell, McCarthy, et al., 2016; David T. Welsh, 2000). At a larger scale, increased benthic BNF has been suggested to help resolve an apparent and highly debated imbalance in the modern marine N budget (Codispoti, 2007; Gruber & Galloway, 2008; Zhang et al., 2020). A better understanding of the ecophysiology of benthic diazotrophs, particularly how such microorganisms regulate BNF under fluctuating environmental conditions, is necessary to resolve questions of when and why BNF occurs in benthic settings and to ascertain its importance to the marine N cycle.

Molecular surveys of the nitrogenase gene *nifH* (Bertics et al., 2010; Kapili et al., 2020; Newell, Pritchard, et al., 2016; Zehr et al., 2003) and activity assays designed to target specific microbial groups (e.g., molybdate additions, sulfide production) (Bertics et al., 2010; Gandy & Yoch, 1988; Newell, Pritchard, et al., 2016; D. T. Welsh et al., 1996) have identified sulfate-reducing bacteria (SRB), such as *Desulfovibrio* sp., as the dominant  $\text{N}_2$  fixers in benthic systems (Brown & Jenkins, 2014; Dekas et al., 2018; Gier et al., 2015; Kapili et al., 2020; Newell, McCarthy, et al., 2016). Fermentative firmicute species (e.g. *Clostridium* sp.) have also been consistently observed in surveys of benthic diazotrophy (Gier et al., 2016; Kapili et al., 2020). Such data, which imply a very low sensitivity of BNF to DIN in SRB and Clostridia, have been used to suggest that the physiology and purpose of BNF in benthic organisms could be fundamentally different from that of euphotic zone  $\text{N}_2$  fixers (Knapp, 2012). For example, BNF at high  $\text{NH}_4^+$  (e.g.  $[\text{NH}_4^+] > 50 \mu\text{M}$ ) has been taken as a sign that benthic microorganisms may be using BNF to balance intracellular redox under reducing conditions (Tichi & Tabita, 2000). While this paradox has important implications for our

98 understanding of marine N cycling, the physiologies of N<sub>2</sub> fixation in SRB and other anaerobic  
99 diazotrophs have garnered little attention compared to the aerobic and phototrophic  
100 diazotrophs (Gordon et al., 1981; Postgate & Kent, 1984; Riederer-Henderson & Wilson,  
101 1970; Tubb & Postgate, 1973).

102 Here, we examine N<sub>2</sub> fixation by the model SRB *Desulfovibrio vulgaris* var.  
103 Hildenborough (*DvH*) (Heidelberg et al., 2004) and the model fermenter *Clostridium*  
104 *pasteurianum* strain W5 (*Cp*) (Winogradsky, 1895), two free-living diazotrophic species  
105 common in sediments (Brown & Jenkins, 2014), and in sulfate-reducing sediments from  
106 three salt marshes along the Northeastern coast of the United States. By simultaneously  
107 tracking BNF activity and ambient NH<sub>4</sub><sup>+</sup> concentrations in liquid culture incubations, we find  
108 that BNF by *DvH* only occurs once NH<sub>4</sub><sup>+</sup> concentrations fall below ~ 2 μM and < 20 μM in *Cp*,  
109 where we had lower data resolution. In sediment slurry incubations assayed using two  
110 independent methods (acetylene reduction assay and <sup>15</sup>N tracer), the NH<sub>4</sub><sup>+</sup> threshold for  
111 BNF inhibition is ~ 9 μM. The data indicate benthic BNF is more sensitive to fixed N than  
112 recent studies would suggest and similar to the NH<sub>4</sub><sup>+</sup> sensitivity of BNF by diverse cultured  
113 diazotrophs. Analyses of nitrogen fixation genes and porewater NH<sub>4</sub><sup>+</sup> data suggest the broad  
114 relevance of the ammonium threshold measured. We find that the timing of BNF inhibition by  
115 NH<sub>4</sub><sup>+</sup>, the fast removal of ambient NH<sub>4</sub><sup>+</sup> by active microorganisms, and the heterogeneous  
116 nature of sediments must be considered when investigating the NH<sub>4</sub><sup>+</sup> sensitivity of BNF,  
117 particularly with incubation experiments. To improve mechanistic models of N  
118 biogeochemistry, we propose a simple theoretical framework based on ammonium  
119 transporter kinetics to link ambient porewater NH<sub>4</sub><sup>+</sup> concentrations to BNF activity.

## 120 **Material and Methods**

### 121 *Culture conditions*

122 *Desulfovibrio vulgaris* var. *Hildenborough* (ATCC 29579), (Heidelberg et al., 2004) was  
123 grown anaerobically in minimal diazotrophic media for sulfate reducers using a recipe

modified from Sim et al., 2013 in 27 mL Balch tubes (10 mL media) or 160 mL serum vial (50 mL media) with 20 mm butyl rubber septum stoppers (Bellco glass, Vineland, NJ, USA). Medium modifications include the use of  $0.07\text{g.L}^{-1}$  ascorbate in place of titanium citrate as the reductant and the omission of tungstate from the trace metal mix. Media was prepared using standard anoxic procedures and a glovebox (Coy Laboratories, Grass Lake, MI, USA). Resazurin ( $1\text{mg.L}^{-1}$ ) was used to monitor shifts in the redox potential of the medium during our experiments. Sulfate ( $\text{Na}_2\text{SO}_4$ ) was provided as an electron acceptor at 30 mM. Pyruvate or sodium lactate (30 mM), used as the carbon and electron sources, were the growth limiting nutrients. Sterilization was performed by autoclaving individual tubes for a minimum of 15 min at  $120^\circ\text{C}$ . The final pH of the media was  $7.3 \pm 0.1$ . Cells were grown at  $30^\circ\text{C}$  in the dark, with tubes placed obliquely in an orbital shaker at a constant agitation speed of 130 rpm. *Clostridium pasteurianum* strain W5 was obtained from the DSMZ (German Collection of Microorganisms and Cell Cultures, Braunschweig, Germany). The cells were grown in a diazotrophic media with sucrose as the fermentative substrate as described in Westlake & Wilson, 1959, in 27 mL Balch tubes (10 mL culture) sealed with 20 mm butyl rubber stoppers (Bellco glass, Vineland, NJ, USA) and aluminum seals. Media tubes were sterilized by autoclaving at  $121^\circ\text{C}$  for 45 minutes. Final pH of the media was 6.8. Cells were grown at  $30^\circ\text{C}$  in the dark, with constant orbital shaking at 100 rpm. Culture cell density was followed by turbidimetry at OD 600 nm using a Thermo Fisher Scientific (Waltham, MA, USA) GENESYS spectrophotometer equipped with a tube holder. Anoxic transfer of  $\text{NH}_4^+$  as ammonium chloride was accomplished using syringes and needles from 10- to 100-fold stock solutions prepared in culture media.

#### *Slurry incubations of salt marsh sediments.*

Samples of salt marsh sediments (Table 1) were collected from Barnegat Bay in New Jersey (denoted here as “NJ” saltmarsh, Lat: 40.031319, Long: -74.080609), the Sippewissett Salt Marsh in Massachusetts (denoted here as “MA” saltmarsh, Lat: 41.590533, Long: -70.639711) and from Great Bay in New Hampshire (denoted here as “NH” saltmarsh, Lat:

43.125, Long: -70.868). Sediments were kept at 4°C in the dark until being processed in the laboratory. To determine sediment BNF sensitivity to  $\text{NH}_4^+$ ,  $20 \pm 1\text{g}$  of wet sediment was introduced in a 260 mL serum bottle and acclimated in an anoxic glove bag (2%  $\text{H}_2$  98%  $\text{N}_2$ ) overnight for each incubation. On the day of the experiment, 80 mL of SRB minimal media (described above) containing 30 mM lactate and 30mM  $\text{SO}_4^{2-}$  was added to each sample prior to sealing the bottle. Each experiment had 10 replicates for each condition (NJ,  $n = 2$ ; NH and MA,  $n = 4$ ) where ambient  $\text{NH}_4^+$  concentrations and BNF activity ( $^{15}\text{N}$  and Acetylene Reduction Activity, see methods below) were measured simultaneously. Samples were incubated in the dark in an orbital shaker at 30 °C (ARA) and 26 °C ( $^{15}\text{N}$ ) with constant agitation (100 rpm). Six of the ten sediment samples assayed by  $^{15}\text{N}$  received initial addition of ammonium (~300  $\mu\text{M}$  final concentration) before starting the experiments. After the initial onset of BNF, we added  $\text{NH}_4^+$  (~300  $\mu\text{M}$  final concentration) in all but two samples (ARA, NJ).

#### *Enrichment of sulfate-reducing bacteria*

Small amounts of sediment (~1 g) from the Barnegat Bay salt marsh were introduced into 60 mL serum bottles containing SRB-minimum media (prepared as described above). Initial enrichments were sequentially transferred into fresh DIN-free media at least 6 times before BNF sensitivity experiments. SRBB1 and SRBB2 enrichment cultures were grown under the same conditions as strain *DvH*. Identity and purity of the enrichment consortia (NCBI accessions will be added upon acceptance of the manuscript) were determined using 16S rRNA and *nifH* genes.

#### *Evaluation of $\text{N}_2$ fixation activity.*

To avoid its inhibitory effects, a lower acetylene concentration (~ 2.5 % instead of 10%) than typically used in the traditional Acetylene Reduction Assay (ARA) (Hardy et al., 1968) was employed to assess changes in BNF activity by *DvH* and *Cp* cultures and sulfate-reducing sediments. An experiment was started when culture at the beginning of exponential growth

(OD<sub>600</sub> ~ 0.09 for *DvH* and ~ 0.1 for *Cp*, Supplementary Figs. S3 and S4) by replacing ~ 2.5 % of the total headspace (2.4-2.7%, depending on vessel volume) with 100% acetylene generated anaerobically with N<sub>2</sub> background in a serum bottle from technical grade calcium carbide (Sigma Aldrich, St Louis, MO, USA). At each time point, 1-3 mL of the headspace was exchanged with the same volume of ~2.5% acetylene in N<sub>2</sub> and stored in a 3 mL Exetainer® vial (Labco, Lampeter, UK). Acetylene Reduction (AR) activity was evaluated by measuring ethylene concentration in the headspace at different times using a Shimadzu (Kyoto, Japan) 8A GC-FID as previously described (Bellenger et al., 2011). Results were corrected for the sequential dilution of the headspace during resampling.

Direct nitrogen fixation activity was determined by measuring <sup>15</sup>N incorporation into biomass from <sup>30</sup>N<sub>2</sub> (Montoya et al., 1996). Briefly, the <sup>15</sup>N tracer experiment was started by replacing 20% of the headspace and liquid phase with <sup>30</sup>N<sub>2</sub> gas (98%+ v/v, Cambridge Isotope laboratories, Andover, MA, USA) and media equilibrated with <sup>30</sup>N<sub>2</sub> gas, respectively (Mohr et al., 2010). At each time point, 3-4 mL of sediment slurry or cell culture was sampled and centrifuged at 8000 rpm. Supernatant was filtered and used for NH<sub>4</sub><sup>+</sup> measurement (see below) and solid samples were stored at -20°C until further processing. Solid sediments samples were then dried and weighed, and a known amount was sent for <sup>15</sup>N and <sup>13</sup>C analyses at UC Davis Stable Isotopes Facility (<https://stableisotopefacility.ucdavis.edu>).

#### *Measurement of ammonium concentration.*

At every time point for assessing BNF activity by ARA or the <sup>15</sup>N tracer method, 1-3mL of supernatant was removed from the incubation and filter-sterilized (0.2 µm hydrophilic PTFE or PES syringe filters, Agilent Technologies and MilliporeSigma, respectively). Filtered samples, procedural blanks, and freshly prepared NH<sub>4</sub><sup>+</sup> chloride standards (0.8 µM - 0.1 M) for NH<sub>4</sub><sup>+</sup> analysis were then stored at -20°C until further processing. Ammonium was measured by fluorescence using the o-phthalaldehyde (OPA) method (Holmes et al., 1999). Under our conditions, uncertainty for [NH<sub>4</sub><sup>+</sup>] in the range of 0-10 µM was estimated using

duplicate measurements over 3 different days ( $n=20$ ) and found to be  $1.3\ \mu\text{M}$ . Several filtered supernatant aliquots representing several timepoints sampled from five sediment incubations assessed by ARA (#4, #6, #7, #9, #10) after  $\text{NH}_4^+$  addition exhibited irregular coloration (red-black instead of yellow) during the OPA procedure, even after 8-fold dilution (Supplementary Fig. S2). After careful examination (see Supplementary Material Sup. Discussion), these measurements (highlighted with red stars in Fig. 3 and Supplementary Fig. S5) were disregarded for the purpose of this study. To determine the  $\text{NH}_4^+$  threshold for BNF inhibition (see Fig. 3), several  $\text{NH}_4^+$  concentrations were averaged over two time-points to link BNF rate and  $\text{NH}_4^+$  concentration, without significant influence on the statistical determination of  $[\text{NH}_4^+]_{95\%}$  threshold value estimates.

#### Review of porewater ammonium concentration in sediments.

To estimate the extent to which sediment porewater ammonium concentrations reach the ammonium threshold for BNF inhibition at the global scale, we extracted porewater ammonium concentrations and depth gradients from 26 studies representing 5 continents (Europe, North America, South America, Asia, and Oceania). We reviewed all benthic references included in Knapp et al. 2012, as well as several references reporting BNF at high  $[\text{NH}_4^+]$ . We further identified more recent publications using a keyword search in Google Scholar, searching for “ammonium concentration sediments pore water coastal” and filtering the results to include studies published after 2012 (Supplementary Table S1). Papers were manually inspected, and ammonium concentrations when present were collected from tables or estimated from direct graphic reading when no tabulated values were available ( $\sim 10\%$  or  $5\ \mu\text{M}$  of reading uncertainty, whichever was larger). For each sediment dataset, we derived the general shape of the depth gradient of porewater ammonium (*i.e.*, depth and  $[\text{NH}_4^+]$  values were extracted at initial, break in slope, plateau, and maximal measurement depths). For the value of ammonium concentration at 0 cm (sediment/water interface), the value reported in the study,  $[\text{NH}_4^+]$  of the overlying water, or  $0\ \mu\text{M}$  was used, in this order of preference. We then applied a linear interpolation to estimate the depth at which the

porewater concentration was equal to our threshold (*i.e.* threshold depth). When data were presented as an average value for the whole sample or sediment core, the average porewater  $[\text{NH}_4^+]$  was attributed to the average depth (*i.e.*, half the maximal depth of sampling) before interpolation.

#### *Extraction and analysis of nucleic acid sequence from salt marsh sediments.*

At the end of the ARA and  $^{15}\text{N}$  incubations, 3 mL of slurry was removed from the samples, centrifuged at 5000 g, resuspended in LifeGuard® Soil Preservation solution (Qiagen, Germantown, MD, USA) and stored at  $-20^\circ\text{C}$ . DNA and RNA extraction were conducted using the RNEasy PowerSoil® Total RNA and DNA elution kit (Invitrogen, Carlsbad, CA, USA) following the recommendation of the supplier and extracts were stored at  $-80^\circ\text{C}$ . Nucleic acid purity was assessed using NanoDrop (Thermo Fisher Scientific) and Qubit® (Thermo Fisher Scientific) measurements. RNA samples were converted into cDNA using the SuperScript® IV First Strand kit (Invitrogen) following the recommendation of the supplier. Nested PCR amplification of *nifH* genes from DNA and cDNA were conducted in triplicate according to the modified protocol of Zehr et al., 1998 from Jayakumar et al., 2017. Amplification was verified by gel electrophoresis and amplicons from cDNA were pooled and sent with total DNA extract for *nifH* and 16S rRNA gene sequencing using bTEFAP® and Illumina MiSeq300 at the MRDNA® facility (Shallowater, TX, USA). 16S rRNA genes were sequenced using 515f/806r primer pair targeting the V4 region (Caporaso et al., 2012), and *nifH* genes were sequenced with the nifH1/nifH2 primer pair (Zehr & McReynolds, 1989). The commercial in-house data processing pipeline at MRDNA® was used to process the sequence data and cluster OTUs (97% similarity threshold). OTUs were taxonomically classified against a curated database derived from the Ribosomal Database Project Release 11 (RDPII) for the 16S rRNA gene and the NCBI non-redundant nucleotide sequences database for *nifH*.

We pruned the sequence datasets to retain only OTUs with absolute abundance > 1 000 counts and relative abundance > 0.2%. The top 500 *nifH* gene OTUs and the pruned transcript OTUs (n=55) from ARA and <sup>15</sup>N incubations, as well as *nifH* gene OTUs (n=46) from our two salt marsh enrichments SRBB1 and SRBB2 were aligned against a representative (*i.e.*, all major *nifH* groups represented) subset (n=152) of *nifH* database recently curated in a deep sea sediment study (Kapili et al., 2020). Aligned sequences (see Fig. 4A) were then separated into gene (after pruning, n=59) and transcript datasets to build individual maximum-likelihood phylogenetic trees using FastTree (v2.1.11, GTR+CAT model, 1,000 resamples, Price et al., 2010). Data processing and visualizations were conducted using R Studio (v1.3.1056) and R (v4.0.3) using the *phyloseq* (v.1.32.0) (McMurdie & Holmes, 2013) and *ggtree* (v2.2.4) packages. Hierarchical clustering (*stat* v4.0.3, method Ward.D2) was conducted on *clr*-transformed relative abundance data (Gloor et al., 2017) using the package *compositions* (v2.0-0) (van den Boogaart & Tolosana-Delgado, 2008). Canonical Analysis of Principal coordinate (CAP, 45) were performed on OTUs abundance data using Bray-Curtis dissimilarity, and a Redundancy Analysis (RDA as implemented in *vegan* v2.5-6, Oksanen et al., 2019) was conducted on normalized geochemical data. It is worth noting that we were still able to find *nifH* sequences in the reagent blank control, indicating contamination (Goto et al., 2005; Zehr et al., 2003) or cross-contamination between samples during extraction or sequencing. In particular, two-three sediment slurry samples, characterized by low DNA concentration, purity, and a low number of sequences (< 50 000 counts, Fig. 4B), cluster with our control sequence data, indicating they might be disproportionately affected by contamination (purple box in Fig. 4B top panel and Fig. 5B).

## Results

*Growth rates and yields of Desulfovibrio vulgaris and Clostridium pasteurianum under diazotrophic and ammoniotrophic conditions*

We compared the growth of the two models anaerobic diazotrophs *D. vulgaris* var Hildenborough (strain *DvH*) and *Clostridium pasteurianum* (*Cp*) under fully diazotrophic and  $\text{NH}_4^+$  replete conditions (Fig. 1A and B). For *DvH*, as expected for diazotrophs, including SRB (Lespinat et al., 1987), diazotrophic growth ( $\mu_{\text{N}_2} = 0.046 \pm 0.005 \text{ h}^{-1}$ , SE  $n=3$ ) was much slower ( $\sim 2$ -fold) than  $\text{NH}_4^+$  replete growth ( $\mu_{\text{NH}_4^+} = 0.10 \pm 0.004 \text{ h}^{-1}$ , SE,  $n=3$ ), compatible with the typically higher energetic costs of BNF compared to  $\text{NH}_4^+$  usage for N anabolism (Inomura et al., 2017). Biomass yields ( $\text{OD}_{\text{Max}, \text{NH}_4^+} = 0.82 \pm 0.02$  vs.  $\text{OD}_{\text{Max}, \text{N}_2} = 0.39 \pm 0.01$ , SE  $n=3$ , Fig. 1A) also support this interpretation. Growth at sub-replete initial  $\text{NH}_4^+$  concentrations under our culture conditions (e.g.,  $[\text{NH}_4^+] = 500 \mu\text{M}$ , 30 mM pyruvate, 30 mM sulfate, Fig. 1A) showed clear biphasic growth, with a first exponential growth phase ( $\mu = 0.094 \text{ h}^{-1}$ ), a second lag phase of  $\sim 50 \text{ h}$ , and a second short exponential growth phase ( $\mu \sim 0.007 \text{ h}^{-1}$ ), all being consistent with a transition from ammoniotrophic to diazotrophic growth as medium  $\text{NH}_4^+$  becomes depleted. Accordingly, no BNF activity as measured by ethylene production by ARA was detected during ammoniotrophic growth in similar experiments (data not shown).

*Clostridium pasteurianum* grown fermentatively on sucrose exhibited a shorter lag-phase than *DvH* regardless of N status (5h for *Cp* vs 50h for *DvH*, Fig. 1B). *Clostridium pasteurianum* grew 1.5-fold slower under diazotrophic conditions ( $\mu_{\text{N}_2} = 0.238 \pm 0.006 \text{ h}^{-1}$ , SE  $n=3$ ) (Fig. 1B) than during  $\text{NH}_4^+$  replete growth ( $\mu_{\text{NH}_4^+} = 0.358 \pm 0.002 \text{ h}^{-1}$ , SE  $n=3$ ). While the slower growth and lower biomass yield of diazotrophy relative to ammoniotrophy by *Cp* ( $\text{OD}_{\text{Max}, \text{NH}_4^+} = 1.6 \pm 0.02$  vs.  $\text{OD}_{\text{Max}, \text{N}_2} = 1.3 \pm 0.01$ , SE  $n=3$ , Fig. 1B) demonstrate the higher metabolic cost of N acquisition by BNF during fermentative growth, the differences are not as large as observed for sulfate reducing *DvH*. The faster growth of *Cp* relative to *DvH* corresponded with a very fast and barely measurable transition ( $\sim 1 \text{ h}$  for *Cp* vs 50 h *DvH*) between ammoniotrophic and diazotrophic growth regimes, which was induced by initial growth on sub-replete ammonium levels (Fig. 1B).

*BNF dependency on timing and concentration of ammonium additions to N<sub>2</sub>-fixing pure cultures*

To study the response of BNF to extracellular NH<sub>4</sub><sup>+</sup>, we performed NH<sub>4</sub><sup>+</sup> addition experiments on the slower growing *DvH* acclimated to fully N<sub>2</sub>-fixing conditions (no NH<sub>4</sub><sup>+</sup> addition, Fig. 1C) and tracked BNF using acetylene reduction assays (ARAs) with a 2.5% acetylene headspace concentration (see Material and Methods). This concentration of acetylene decreased but did not prevent the growth of *DvH* or *Cp* (Supplementary Figs. S3A and S4A). BNF generally decreased within the 26 h after the addition of varying concentrations of NH<sub>4</sub><sup>+</sup> (10 – 3,000 μM) (Fig. 1C). However, the response of *DvH* BNF to NH<sub>4</sub><sup>+</sup> addition was highly dependent on both the concentration of added NH<sub>4</sub><sup>+</sup> and the timing of sampling after the addition, as has been found for the DIN sensitivity of other diazotrophs (Drozd et al., 1972; Holl & Montoya, 2005).

To better understand the relative contributions of BNF and external NH<sub>4</sub><sup>+</sup> to fulfilling cellular N growth demands, we calculated the contribution of BNF to new biomass production (Fig. 1D) using the quantity of newly fixed N (derived from ARA activity), the increase in biomass (measured as OD<sub>600</sub>), and the quantity of N required to produce 1 OD equivalent of biomass under N<sub>2</sub>-fixing and NH<sub>4</sub><sup>+</sup>-utilizing conditions (Supplementary Methods S1). Figure 1C shows the reliance on BNF to fulfill N demands depends on the amount of added NH<sub>4</sub><sup>+</sup> and that higher initial spike concentrations result in longer times for the resumption of BNF. At the highest NH<sub>4</sub><sup>+</sup> addition (3,000 μM), the contribution of BNF to cellular N supply dropped to 0 in less than six hours and never resumed, showing that the amount of added NH<sub>4</sub><sup>+</sup> was more than sufficient for the culture to complete growth under our conditions. Lower concentrations of added NH<sub>4</sub><sup>+</sup> (< 300 μM) led to an initial drop in BNF followed by a slow increase back to ~100% of its contribution to N supply, indicating that cells fully assimilated the added NH<sub>4</sub><sup>+</sup> to support their growth. At the first sampling time point (3 h), addition of NH<sub>4</sub><sup>+</sup> > 100 μM showed similarly low residual contributions of BNF to growth (~10-15% of initial BNF activity, Fig. 1C and D) implying that at these concentrations the time to the onset of BNF inhibition (*T<sub>R</sub>*, *i.e.*

the time necessary for the cell to sense external  $\text{NH}_4^+$ , modify its metabolism, and significantly decrease BNF), is independent of  $\text{NH}_4^+$  concentration. By extrapolating the contribution of BNF to the time space, and assuming a drastic shift from nitrogen fixation to  $\text{NH}_4^+$  uptake at  $T_R$ , 100% BNF contribution continues for a maximum of 15% of 3 h = 30 min after addition, then drops to 0%. Hence, in liquid culture of *DvH*,  $T_R < \sim 30$  min. In the first 3 h, residual contributions of BNF to cell growth were higher for  $\text{NH}_4^+$  additions  $< 100 \mu\text{M}$  than when  $> 100 \mu\text{M}$  was added (e.g., 50% at  $30 \mu\text{M}$  vs. 15% at  $100 \mu\text{M}$ ), and we found no inhibition at the lowest level of added  $\text{NH}_4^+$  ( $10 \mu\text{M}$ ), suggesting that  $T_R$  might be concentration-dependent at low external  $\text{NH}_4^+$  concentrations ( $10 - 30 \mu\text{M}$ ).

#### *Ammonium sensitivity of BNF in batch cultures of sulfate-reducing and fermenting bacteria*

To determine the ambient  $\text{NH}_4^+$  concentrations associated with changes in BNF activity, we simultaneously measured BNF activity and medium  $[\text{NH}_4^+]$  in  $\text{NH}_4^+$  addition experiments with *DvH*, (Fig. 2A and Supplementary Fig. S3), *Cp* (Fig. 2B and Supplementary Fig. S4), and two sulfate-reducing microbial enrichment consortia SRBB1 and SRBB2 dominated by *Desulfovibrio desulfuricans* sp. (Fig. 2C and Supplementary Fig. S5). Biological N fixation rates in all culture incubations decreased after  $\text{NH}_4^+$  was added to cultures growing in fully diazotrophic conditions (Fig. 2A-C, Supplementary Fig. S3-S5). We did not find any evidence of BNF activity in *DvH* and SRBBs enrichment when measured ambient  $\text{NH}_4^+$  concentrations were  $> 2 \mu\text{M}$  (Fig. 2A and C, Supplementary Fig. S3 and S5). Consistent with our previous experiments with *DvH* (Fig. 1D), interpolations of BNF rates before and after  $\text{NH}_4^+$  addition for *DvH*, SRBB2, and *Cp* (dashed lines, Fig. 2A-C), showed noticeable inhibition of BNF in all cultures within 0.5 - 3 h post-addition. After ammonium addition, bacterial uptake of  $\text{NH}_4^+$  for growth led to the removal of  $> 400 \mu\text{M}$   $\text{NH}_4^+$  in less than 30 h for *DvH* (max uptake rate  $\sim 500 \mu\text{M}_{\text{NH}_4^+} \cdot \text{OD}^{-1} \cdot \text{h}^{-1}$ ) and in less than 4 h for *Cp* (see Supplementary Figs. S3-S5). Full removal of  $\text{NH}_4^+$  coincides with the recommencement of BNF (e.g., resumption of ethylene production in Fig. 2A-C).

## *Ammonium sensitivity of BNF in sulfate-reducing salt marsh sediments*

To evaluate the sensitivity of BNF by benthic organisms to  $\text{NH}_4^+$  under conditions that are more representative of natural environments, we performed incubation experiments on slurries of salt marsh sediments collected from Barnegat Bay (NJ), Sippewissett Bay (MA), and Great Bay (NH) (Fig. 2D-F, Supplementary Figs. S6 and S7). We evaluated BNF using both ARA and  $^{15}\text{N}$  tracer methods to account for possible artifacts associated with the use of acetylene to estimate BNF (Fulweiler et al., 2015; Payne & Grant, 1982). For both ARA and  $^{15}\text{N}_2$  tracer incubations (10 incubations per method), we observed the onset of BNF only after the levels of dissolved  $\text{NH}_4^+$  initially present in slurries decreased drastically (to  $[\text{NH}_4^+] < 20 \mu\text{M}$ ) (Fig. 2D-F, Supplementary Figs. S6 and S7). In the  $^{15}\text{N}_2$  tracer incubations, five of ten samples (Supplementary Fig. S7 #3, #5, #6, #7, and #9) showed a decrease in  $^{15}\text{N}$  incorporation into organic matter before  $\text{NH}_4^+$  addition, and two of them completely stopped incorporation of  $^{15}\text{N}$  (Supplementary Fig. S7 #3 and #9). In parallel,  $\text{NH}_4^+$  concentrations slightly increased from  $< 8 \mu\text{M}$  to levels around 10-18  $\mu\text{M}$  in several samples prior to  $\text{NH}_4^+$  addition (Supplementary Fig. S7 #3, #4, #5, and #9), suggesting the presence of some endogenic  $\text{NH}_4^+$  production within our experiments.

Similar to our results with cultured diazotrophs,  $\text{NH}_4^+$  additions (e.g., target  $[\text{NH}_4^+] \sim 300 \mu\text{M}$ , indicated with arrow, Fig. 2E and F) to actively fixing samples coincided with a fast decrease in ethylene production to a complete stop in less than  $\sim 3 \text{ h}$  in all our samples ( $n=20$ ). For four of the eight ARA-assayed sediment samples (two each from NH and MA salt marshes) spiked with  $\text{NH}_4^+$ , BNF responded normally to  $\text{NH}_4^+$ , *i.e.* BNF rates resumed to pre-addition values once ambient  $\text{NH}_4^+$  was drawn down to near background levels ( $[\text{NH}_4^+] < 10 \mu\text{M}$ , Fig. 2E, Supplementary Fig. S6), likely due to assimilatory or dissimilatory biological activities. The post-addition results from the other four of eight ARA samples with  $\text{NH}_4^+$  added (two each from NH and MA salt marshes) were excluded because of dubious OPA measurements of  $[\text{NH}_4^+]$  (see Methods and Supplementary Discussion). Importantly, none of the 10 sediment incubations assayed using the  $^{15}\text{N}$  tracer method, which directly reflects

BNF activity, showed any evidence of a resumption of BNF activity after  $\text{NH}_4^+$  addition, and post-addition  $\text{NH}_4^+$  concentrations never decreased below 50  $\mu\text{M}$  (Fig. 2F and Supplementary Fig. S7). We note that the general decrease in  $\delta^{15}\text{N}$  of OM in these samples after  $\text{NH}_4^+$  addition is consistent with the assimilation of the unlabeled  $\text{NH}_4^+$  added to the incubation, which could also impact the timing and sensitivity of the method to detect BNF.

#### *Determination of $\text{NH}_4^+$ threshold for benthic BNF*

Fast removal of ammonium from media due to high metabolic activity (e.g., growth) of the samples, uncertainties on the ethylene concentration in headspace due to sequential dilution during sampling, and variations in the timing for the BNF inhibition response after  $\text{NH}_4^+$  addition all preclude a direct measurement of the exact ammonium concentration at which BNF stops. To obtain a best estimate of the threshold for ammonium inhibition of BNF for each type of experiment, we used the 95<sup>th</sup> percentile of all measured  $\text{NH}_4^+$  concentrations where BNF was present (before  $\text{NH}_4^+$  addition) or resumed (after  $\text{NH}_4^+$  addition) as a robust estimate for the threshold  $\text{NH}_4^+$  concentration that induce changes in BNF activity (Fig. 3, number of datapoints for cultures in liquid media:  $n=10$ , ARA sediment slurries:  $n=52$ , and  $^{15}\text{N}$  tracer sediment slurries:  $n=20$ ). For each sample, significant BNF activity corresponds to BNF activity > 5% of the maximal BNF activity. Results from *Cp* were not included in this analysis as we could not verify the complete inhibition followed by resumption of BNF in any of our replicates (Fig. 2B and Supplementary Fig. S4). This is likely due to the removal of  $\text{NH}_4^+$  faster being than the response time for complete BNF inhibition by the fast-growing *Cp* cultures, as well as by our limited sampling frequency.

Threshold values of  $[\text{NH}_4^+]$  are < 2  $\mu\text{M}$  (average  $\pm$  SD of  $0.7 \pm 0.7 \mu\text{M}$ ) for all sulfate-reducing liquid media cultures (*DvH* and *SRBB1* and *SRBB2*, Fig. 3A), consistent with values obtained for other cultured diazotrophs (Dekaezemacker & Bonnet, 2011; Sweet & Burris, 1981) and oligotrophic pelagic BNF (Knapp, 2012). We note that the actual threshold value is likely lower than our reported estimate of 2  $\mu\text{M}$  because of the precision of the  $\text{NH}_4^+$

measurement in our conditions (reproducibility<sub>0-10 $\mu$ M</sub>  $\sim$  1  $\mu$ M). While the fast growth, fast removal of ammonium, and our insufficient sampling frequency does not allow for an accurate determination of  $\text{NH}_4^+$  threshold in *Cp*, we estimate that its value to be anywhere between 0 and 20  $\mu$ M, as evidenced by BNF slow-down  $\sim$ 3 h post-ammonium addition and its recommencement  $\sim$ 10 hr post-addition (Fig. 2B, ammonium at  $t \sim$  8h vs 14 hr). Threshold values of  $\sim$  7  $\mu$ M and  $\sim$  11  $\mu$ M were found for BNF in slurries assessed by ARA and  $^{15}\text{N}$  tracer methods, respectively (Fig. 3B and C). This value is in good agreement with similar experiments performed with deep-sea sediments ( $< 25 \mu\text{M}$ ) (Dekas et al., 2018). All these direct measurements of  $\text{NH}_4^+$  thresholds are consistent with calculations of the external  $[\text{NH}_4^+]$  at which BNF resumes in experiments with *DvH* (Fig. 1C and D) that account for the drawdown of external  $\text{NH}_4^+$  contributing to biomass growth ( $[\text{NH}_4^+]_{\text{calculated}} < 23 \mu\text{M}$ , Supplementary Methods S2 and Fig. S8).

#### *Diazotrophic community in enrichment cultures and slurry experiments*

In both SRBB1 and SRBB2, *nifH* taxonomic diversity was dominated by several operational taxonomic units (OTUs) closely related to the SRB *Desulfovibrio desulfuricans*. SRBB1 showed the additional presence of an OTU closely related to the genus *Clostridium* (Supplementary Fig. S1A). Other OTUs (e.g., *Desulfovibrio fructosovorans*, *Azotobacter vinelandii*, and *Pelobacter carbinolicus*) were found in low abundance ( $<1\%$  of total *nifH* genes). The 16S rRNA gene diversity confirmed the dominance of *Desulfovibrio* and *Clostridium* species in the enrichments (Supplementary Fig. S1B).

To determine the diazotrophic community that developed in our sediment incubations and compare them to those in other benthic environments, we sequenced *nifH*, which encodes a component of the Mo-nitrogenase and is the most common phylogenetic marker used in studies of diazotroph diversity (Fig. 4). Analysis of *nifH* gene diversity in DNA extracted from sediments at the end of incubation (5 days) revealed the dominance of sequences related to those from delta-proteobacteria, particularly sulfate-reducing genera

(closely related to *Desulfovibrio desulfuricans*, *D. salexigens*, *Desulfatibacillum alkenivorans*, *Desulfuromonas acetoxidans*), as well as other anaerobic bacteria (e.g. *Pelobacter carbinolicus* and *Marichromatium purpuratum*). These diazotrophic taxa have also been identified in other salt marshes (Burns et al., 2002; Steppe & Paerl, 2005), and in estuarine (Burns et al., 2002; Newell, Pritchard, et al., 2016), and deep-sea sediments (Gier et al., 2016; Kapili et al., 2020). Examination of *nifH* transcript diversity at the end of incubations showed similar results suggesting that the putative diazotrophs identified based on *nifH* gene diversity are likely active during our incubation (Fig. 4A and Supplementary Fig. S9).

#### *Relationship between nitrogenase activity and sediment biogeochemical characteristics*

The richness of *nifH* transcripts (expressed as observed number of OTUs, Fig. 5A) was similar at the end of incubations, even in treatments exhibiting high final concentrations of  $\text{NH}_4^+$  ( $> 100 \mu\text{M}$ ) and no  $^{15}\text{N}_2$  incorporation or acetylene reduction activity in the previous 20-30 h (Fig. 2, symbol gradient shade in Fig. 5A, Supplementary Fig. S6 and S7). Biological N fixation activity at the end of incubation was only detected for  $\text{NH}_4^+$  concentrations less than  $\sim 20 \mu\text{M}$ . This strongly supports a post-transcriptional inhibition of nitrogenase activity by  $\text{NH}_4^+$  and slow turnover of *nifH* transcripts in our incubations.

Hierarchical clustering of our samples and Canonical Analysis of Principal Coordinates on the OTU abundance data (Hellinger transformation) indicates that geographical origin (NJ, MA, and NH) rather than the type of incubation (i.e., BNF assayed by ARA or  $^{15}\text{N}$ ) best explains sediment *nifH* composition (see colored boxes in top panel Fig. 4B, Fig. 5B, and Supplementary Fig. S10). We identified several OTUs showing significant differences in relative abundance between ARA and  $^{15}\text{N}$  incubation conditions based on 16S rRNA gene diversity, but not for *nifH* gene and transcript abundance (Supplementary Fig. S10C). Net  $\text{NH}_4^+$  removal rates (as the balance between all processes that remove  $\text{NH}_4^+$  minus the potential release of  $\text{NH}_4^+$  from sediments) are correlated with initial (pre-addition)  $\text{NH}_4^+$

concentrations, and maximum BNF rates are positively correlated with sediment C:N ratios (Fig. 5C).

#### *Literature review of benthic porewater ammonium concentrations*

To estimate the extent to which BNF can occur in benthic sediments (Fig. 6A), we combine all data by transforming the ARA- and  $^{15}\text{N}$ -based results so that they are relative to the maximum in each replicate  $((V_{\text{BNF,Max}} - V_{\text{BNF}}) / V_{\text{BNF,Max}})$ . It allows us to derive a single estimate of  $\text{NH}_4^+$  threshold for BNF in sediments of  $[\text{NH}_4^+]_{\text{threshold}} = 9 \mu\text{M}$  (Fig. 6B,  $n=72$ ). We then analyzed porewater ammonium concentration data from 26 studies including 151 sites and more than 300 individual samples representing a diversity of benthic environments (Fig. 6C and D, SI Table S1). We estimate the fraction of global sediments where BNF might occur by deriving the depth ('threshold depth', see Fig. 6A) at which published porewater  $[\text{NH}_4^+]$  equaled the threshold value of BNF inhibition of  $\text{NH}_4^+ = 9 \mu\text{M}$  (Fig. 6C and D). The median depth at which porewater ammonium reaches the threshold value is 0.3 cm (Fig. 6C, average 0.65 cm, range 0 – 25 cm,  $n=151$ ). In this review, ~ 87 % of the values collected (based on individual cores,  $n=334$ ) exceeded the  $[\text{NH}_4^+]_{\text{threshold}}$  at depths deeper than 1 cm below surface, and 95 % of values exceeded the threshold at depths of ~ 3 cm below surface. Porewater  $[\text{NH}_4^+]$  was lower than the threshold at any measured depth in only ~ 0.1% of all individual cores (3 of 334). Results separated along the different benthic zones, *i.e.*, tidal (< 4 m,  $n=64$ ), sub-littoral (between 4 and 200 m,  $n=47$ ), and bathyal (> 200 m,  $n=40$ ), are available in Supplementary Fig. S12.

## **Discussion**

#### *Metabolic cost of BNF in anaerobic diazotrophs*

The metabolic cost of using BNF versus fixed N sources like ammonium to support growth can yield insights on ecophysiology of microorganisms in nature. The substantively lower

growth rates and yields observed for diazotrophic versus ammoniotrophic growth of the sulfate-reducer *DvH* (2-fold growth rate and yield) and the fermenter *Cp* (1.5-fold growth rate, 1.2-fold yield) indicate a significant metabolic cost of BNF as an N source in anaerobic bacteria. The cost of BNF in diazotrophs that have metabolisms involving  $O_2$  is even more substantial given the additional energetic costs of protecting nitrogenase and other  $O_2$ -sensitive BNF-related enzymes from oxidative damage (Großkopf & LaRoche, 2012; Inomura et al., 2017). Collectively, the physiological data for anaerobic diazotrophs presented here, as for other types of diazotrophs, support the paradigm of prevalent BNF in environments with limiting amounts of fixed N.

#### *Ammonium threshold for BNF in diazotrophic bacteria and its relation to cellular N metabolism*

Our results show that BNF activity in sulfate-reducing bacteria and sulfate-reducing sediments is sensitive to fixed nitrogen (as  $NH_4^+$ ) at micromolar concentrations. This conclusion does not appear to support the use of BNF as an important mechanism to balance intracellular redox in ammonium-rich benthic systems, as previously hypothesized (Tichi & Tabita, 2000). The onset of inhibition of BNF activity occurs for  $[NH_4^+]_{\text{threshold}} < 2 \mu M$  in liquid cultures of SRB; a slightly higher threshold of  $[NH_4^+]_{\text{threshold}} < 11 \mu M$  was found in slurry incubations. This is in good agreement with existing knowledge of the biochemistry and metabolism of nitrogen fixers (Dekaezemacker & Bonnet, 2011; Hartmann et al., 1986; Holl & Montoya, 2005; Kleiner, 1974; Ohki et al., 1991; Postgate & Kent, 1984). Our threshold estimates are also consistent with early studies in salt-marsh sediments ( $[NH_4^+]_{\text{threshold}} < 33 \mu M$ ) (Carpenter et al., 1978; Patriquin & Keddy, 1978), a recent studies of BNF in deep-sea sediments ( $[NH_4^+]_{\text{threshold}} < 25 \mu M$ ) (Dekas et al., 2018), and pure culture studies with *Desulfovibrio gigas* ( $[NH_4^+]_{\text{threshold}} \sim 10 \mu M$ ) (Kessler et al., 2001; Postgate & Kent, 1984), *Rhodospirillum rubrum* ( $[NH_4^+]_{\text{threshold}} \sim 3-5 \mu M$ ) (Sweet & Burris, 1981), *Klebsiella oxytoca* ( $[NH_4^+]_{\text{threshold}} < 10 \mu M$ ) (Schreiber et al., 2016), *Methanococcus maripaludis* ( $[NH_4^+]_{\text{threshold}} < 25 \mu M$ ) (Kessler et al., 2001), *Azotobacter vinelandii*

516 ( $[\text{NH}_4^+]_{\text{threshold}} < 10 \text{ } \mu\text{M}$ ) (Kleiner, 1974), and *Trichodesmium* ( $[\text{NH}_4^+]_{\text{threshold}} < 10 \text{ } \mu\text{M}$ )  
517 (Mulholland et al., 2001), and *Crocospaera watsonii* ( $[\text{NH}_4^+]_{\text{threshold}} < 1 \text{ } \mu\text{M}$ )  
518 (Dekaezemacker & Bonnet, 2011). While the ammonium threshold of BNF inhibition could  
519 not be precisely estimated for the fermenter *Clostridium pasteurianum* ( $[\text{NH}_4^+]_{\text{threshold}} < 20$   
520  $\mu\text{M}$ ) due to rapid ammonium draw-down and low sampling frequency, the presence of  
521 *Clostridia* relatives in the isolate SRBB1 (Supplementary Figs. S1 and S4) and Clostridia-like  
522 *nifH* sequences in sediment slurries suggest a similarly high sensitivity.

523 Biological N fixation and  $\text{NH}_4^+$  uptake are complementary N sources to cells. Nitrogenase  
524 activity and gene regulation is controlled by intracellular  $\text{NH}_4^+$  through the GS-GOGAT  
525 systems. In most diazotrophs, low  $\text{NH}_4^+$  supply to the cell leads to a decrease in intracellular  
526 glutamine concentrations, turning on nitrogenase gene transcription and/or nitrogenase  
527 protein activity through the PII signal-transduction cascade (Dixon & Kahn, 2004).  
528 Conversely, sufficient  $\text{NH}_4^+$  supply to the cell decreases BNF, either at the pre- or post-  
529 transcriptional level. Cellular N and C metabolisms are coupled via the balance between  $\alpha$ -  
530 ketoglutarate, a metabolite of the TCA cycle, and glutamine (Huergo & Dixon, 2015).  
531 Extracellular N sources are acquired through the use of specific membrane transporters (e.g.,  
532 AmtB for ammonium (Kleiner, 1985; Zheng et al., 2004)). Most other fixed-N sources, such  
533 as  $\text{NO}_3^-$ , are first taken up and then intracellularly converted into ammonia by specific  
534 enzymes (e.g., nitrate reductase) before being used anabolically. Assuming a Michaelis  
535 Menten mechanism for  $\text{NH}_4^+$  transport (Kuzyakov & Xu, 2013), we expect the affinity  
536 constant of the  $\text{NH}_4^+$  transporter to play a critical role in cellular sensing of and ability to grow  
537 using external  $\text{NH}_4^+$  (Javelle et al., 2004), thereby modulating internal N status and BNF  
538 down-regulation (Fig. 6A). For example, at  $\text{NH}_4^+$  concentrations  $< K_m$  of the transporter, less  
539 than 50% of  $\text{NH}_4^+$  uptake activity can take place, reducing intracellular N concentrations and  
540 inducing a metabolic cascade that ultimately leads to the onset of BNF activity. This  
541 simplified framework would constrain the switch between ammoniotrophy and BNF or their  
542 co-occurrence to  $\text{NH}_4^+$  concentrations close to transporter  $K_m$  values (Schreiber et al., 2016).

In reviewing the literature, we found that the  $K_m$  values for AmtB, the most common  $\text{NH}_4^+$  transport protein, in a variety of nitrogen fixers, range in value from approximately 1 to 20  $\mu\text{M}$  (Kleiner, 1985). These values are in good accordance with our experimental data and prior literature which show BNF inhibition thresholds of 2-10  $\mu\text{M}$   $\text{NH}_4^+$ . Given these estimates, BNF activity in environment with high N loadings would not be unexpected when residual dissolved  $\text{NH}_4^+$  concentrations are  $< 20 \mu\text{M}$  (Foster & Fulweiler, 2014).

#### *Timing of inhibition and regulation of nitrogenase in anaerobic diazotrophs*

We found that  $T_R$ , the time to the onset of BNF inhibition after  $\text{NH}_4^+$  addition, was concentration-independent at  $[\text{NH}_4^+] > 100 \mu\text{M}$ , and concentration-dependent at  $[\text{NH}_4^+] < 30 \mu\text{M}$ , in accordance with a Michaelis-Menten mechanism for  $\text{NH}_4^+$  sensing. Within this framework, the lowest addition of  $\text{NH}_4^+$  in this study (10  $\mu\text{M}$ ), which led to the longest delay ( $> 3 \text{ h}$ ) in BNF inhibition (see Fig. 1D), can be explained by a sub-maximal uptake of  $\text{NH}_4^+$  leading to the delay in cellular metabolic response. The passive diffusion of ammonia through cell membranes is not likely to be important as the pH of media was  $\sim 7$  and rate of removal was faster in the faster growing bacteria (i.e., *Cp*) than in *DvH*, indicating an active process.

The relatively short-time response of BNF to  $\text{NH}_4^+$  additions in both pure culture and slurry incubations (0.5 - 3h, as estimated from BNF rate interpolations, Fig. 2) and mass balance calculations in Fig. 1D) and the fast resumption of BNF activity upon depletion of the external  $\text{NH}_4^+$  pool indicates that the control of  $\text{NH}_4^+$  over nitrogenase activity is most likely post-transcriptional. This is also supported by the presence of multiple *nifH* mRNA from SRB clades within incubations with  $[\text{NH}_4^+] > 100 \mu\text{M}$  and no BNF activity. Boyd and co-workers have suggested that post-transcriptional regulation arose during the evolutionary transition from anaerobic to aerobic BNF as an early strategy of anaerobes for optimizing BNF activity in fluctuating environments (Boyd et al., 2015). Pure culture studies on anaerobic organisms (Heiniger & Harwood, 2015; Kessler et al., 2001; Masepohl et al., 2002) and field studies in

salt marsh sediments using methionine sulfoxide, an inhibitor of glutamine synthase, also support this interpretation (Gandy & Yoch, 1988; Yoch & Whiting, 1986). The results of this study, with frequent measurement of ammonium and BNF activity over the incubation period, and the likely post-transcriptional nature of  $\text{NH}_4^+$  inhibition of nitrogenase in anaerobic systems, reiterates the need for higher time resolution flux measurements for better *in situ* BNF estimates in natural environments.

Interestingly, there is a drastic difference between the long lag-phase (~ 20-50h) at the initial onset of diazotrophy both in *DvH* cultures (without initial  $\text{NH}_4^+$  and during the transition between ammoniotrophy and diazotrophy) and in slurry incubations, and the rather rapid inhibition and resumption of BNF following  $\text{NH}_4^+$  removal in actively fixing samples (< 3h) (Fig. 1C and Fig. 2). This suggests that SRB organisms living in environments with fluctuating N concentrations close to the  $\text{NH}_4^+$  threshold value might maintain an intracellular nitrogenase pool that can help them quickly resume BNF activity should there be local  $\text{NH}_4^+$  depletion. Interestingly, BNF in *Cp* cultures remained active for ~ 3 hr after  $\text{NH}_4^+$  addition while  $\text{NH}_4^+$  in media was quickly removed. Consistent with this observation, Yoch and Whiting, 1986, found that undisturbed sediments amended with mannose, a preferred substrate of fermenting bacteria, took longer to inhibit BNF than when lactate, a substrate favoring SRBs, was used as the amended C source. In addition, and in opposition to our findings with SRB, only scarce *nifH* transcripts attributed to *Clostridia* were found at the end of our incubation experiments compared to their general presence in our samples (Fig. 4 and Supplementary Fig. S10), indicating that transcriptional regulation of the *nif* genes could be more prominent in this strain. The data suggest that even though the ammonium threshold between SRB and firmicutes is likely similar, the sensing mechanism or inhibition pathway could slightly differ (Gandy & Yoch, 1988; Gordon et al., 1981; Postgate & Kent, 1984; Tubb & Postgate, 1973). How these differences in regulatory pathways influence the adaptation and competitiveness of these organisms in benthic systems remains to be investigated.

597 Consistent with previous studies of saltmarsh, estuarine, and carbonate mud sediments  
598 (reviewed by Capone, 1988; see also Welsh et al., 1996), we find strong and consistent  
599 inverse relationships between ambient ammonium and BNF activity. Our estimate of  $\text{NH}_4^+$   
600 threshold for BNF in sediments of  $[\text{NH}_4^+]_{\text{threshold}} = 9 \mu\text{M}$  (Fig. 6B,  $n=72$ ). This value is lower  
601 than any previously reported value for BNF inhibition in sediments, and is consistent with a  
602 very early assessment of *in-situ* BNF sensitivity to porewater  $\text{NH}_4^+$  in *Spartina* marshes  
603 ( $[\text{NH}_4^+]_{\text{threshold}}$  between 3 and 33  $\mu\text{M}$ ) (Patriquin & Keddy, 1978) and recent estimates in deep-  
604 sea sediments ( $[\text{NH}_4^+]_{\text{threshold}} < 25 \mu\text{M}$ ) (Dekas et al., 2018).

605 The potential diazotrophic genera identified in our sediment slurry experiments overlap well  
606 with those found in other benthic systems, from intertidal salt marshes (Burns et al., 2002;  
607 Steppe & Paerl, 2005), and estuaries (0-20m water level) (Burns et al., 2002; Newell,  
608 Pritchard, et al., 2016) to deep oceanic floor sediment (100-1000m) (Gier et al., 2015; Kapili  
609 et al., 2020), with the dominance of *Desulfovibrionaceae*, *Desulfobulbaceae*, as well as some  
610 Firmicutes species (see Fig. 4). We found evidence in several of our samples for the  
611 presence and transcriptional activity of *Pelobacter carolinicus* and *Desulfuromonas*  
612 *acetoxidans*, which have been identified in a recent study on the diazotrophic community of  
613 deep sea sediments using  $^{15}\text{N}$ -SIP-RNA to probe the identity of the dominant diazotrophs  
614 (Kapili et al., 2020). The observation of similar microbiome composition indicates that our  
615 threshold estimates likely represent the general response of SRBs and possibly of other  
616 types of anaerobic N fixers present in benthic systems (e.g., potential fermenters in SRBB1  
617 diazotroph enrichment, Supplementary Fig. S1).

618 *In situ* porewater  $[\text{NH}_4^+]$  in most benthic sediments (Knapp, 2012) is usually higher than our  
619 observed threshold ( $> 10\text{-}20 \mu\text{M}$ ) and increases with depth (Metzger et al., 2019) (Fig. 6A).  
620 Our review of porewater ammonium data from 26 studies, including 151 sites and more than  
621 300 individual samples indicates that the median depth at which our threshold value was

reached is 0.3 cm below sediment surface (Fig. 6C, based on site, average 0.65 cm, range 0 – 25cm, n=151). Overall, in our review, ~ 87% of the values collected (based on individual cores, n=334) were over our  $[\text{NH}_4^+]_{\text{threshold}}$  at a depth greater than 1 cm, and 95% at depth of ~ 3 cm. This analysis suggests that BNF rates in sediments are likely to be greatest at the surface and decrease rapidly to zero. Diffusion of  $\text{O}_2$  from the sediment-water interface into deeper layers could further reduce the opportunity for BNF to occur.

The previously reported relatively high ammonium values associated with BNF inhibition, like those reviewed in Capone 1988 (~200-300  $\mu\text{M}$ ) and Knapp et al 2012 (50-2,000  $\mu\text{M}$ ), were likely due to long incubation times (> 24 hr), the fast depletion of porewater ammonium, and the reporting of *in situ* porewater  $[\text{NH}_4^+]$  rather than measurement of ambient  $[\text{NH}_4^+]$  during incubation. Consistent with our results, many studies that track the timeline of N fixation show very little if any activity in the first ~ 24 - 48 hr of incubation in saltmarsh sediments when roots were excluded (Patriquin & Keddy, 1978; Yoch & Whiting, 1986). In contrast, *in situ* incubations of vegetated saltmarshes sediments (Yoch & Whiting, 1986) demonstrate rapid onset of BNF and the highest activity was found with surface-washed living roots and rhizome samples, and with C addition, indicating a predominant role of root-associated microbes and available C (*i.e.*, carbon substrate amendment, root exudates).

Thus, our results, which are compatible with our knowledge of the regulation of diazotrophy (Dixon & Kahn, 2004), suggest that the extent of natural BNF activity in benthic sediments may be mostly limited to N-limited superficial regions of the sediments (top 1 cm, Fig. 6) and highly productive environments like seagrass meadows and *Spartina* marshes. The observation of BNF associated with live vegetation roots in *Spartina* marsh with high  $[\text{NH}_4^+]$  (Yoch & Whiting, 1986) is likely explained by local zones around and within plant roots (Kuzyakov & Xu, 2013) in which  $\text{NH}_4^+$  is depleted by plant and microbial uptake to concentrations below our threshold estimates.

Our findings do not contradict the traditional view of a relatively small contribution of benthic systems to total marine BNF compared to pelagic new N sources. Early studies have suggested benthic contributions to global marine BNF of ~ 15 Tg N (Capone, 1988; Capone & Carpenter, 1982), ~10% of total BNF (Zhang et al., 2020). Most of this activity was then attributed to shallow benthic environments, particularly vegetated and coral reef ecosystems. Based on the recent findings of Dekas et al. 2017 and consistent with our analyses of ammonium thresholds (Supplementary Fig.S12), it is possible that unvegetated sediments from the deep sea could constitute a larger source of new N than previously thought. More precise accounting of benthic BNF however requires additional measurements in these and other benthic systems that account for several methodological challenges in measuring BNF, which we discuss below.

#### *Current methodologies for BNF investigations in sediments*

Our results show that the timing of inhibition is a critical factor to consider when assessing the sensitivity of BNF to ammonium as a fixed N source. Measurement of BNF activity before the onset of the inhibition response would lead to the false conclusion that BNF in a tested sample is not or little sensitive to ammonium. Conversely, it is particularly critical to measure and report the evolution of  $[\text{NH}_4^+]$  over time along or at the end of incubations, as typical incubation times, which range from 8 hours to several days, are sufficient for pore water  $[\text{NH}_4^+] > 100 \mu\text{M}$  to become depleted to background  $\text{NH}_4^+$  levels (below threshold for BNF activation). This precaution is particularly important in closed and semi-closed vessel experiments with high activity samples (e.g., experiments with carbon addition, high C:N sediments, and pure cultures), where ammonium-supported growth would drastically increase the biomass and lead to higher BNF rates after fixed N addition.

Our measurements show that the apparent  $\text{NH}_4^+$  concentration required to stop BNF in slurry experiments ( $[\text{NH}_4^+] < 11 \mu\text{M}$ , Fig. 3B and C) is roughly five times higher than in liquid media

experiments (*DvH* and SRBB enrichments,  $[\text{NH}_4^+] < 2 \mu\text{M}$ , Fig. 3A). It is probable that our measurement of the  $\text{NH}_4^+$  concentration threshold for BNF is higher in the sediment slurries due to higher spatial and temporal heterogeneity. Small biogeochemically distinct zones, or microniches, could originate from slow nutrient diffusion to hotspots of highly active biomass in heterogeneous sediments and lead to imbalanced macronutrient ratios (C or P excess relative to N). Indeed, our results show an association between maximal BNF rate and high C:N (Fig. 5C), and reports on the occurrence of BNF at high fixed N have been correlated with high C:N ratios of sediments (Hou et al., 2018). We particularly expect microniche formation to be favored at high cell densities when the requirement for N in a large population of active cells could outpace the diffusion of porewater  $\text{NH}_4^+$  through the solid phase. These conditions would prevail at the end of our incubations (~4-5 d) when the addition of  $\text{NH}_4^+$  and of 30 mM lactate, a readily available source of carbon to SRBs, resulted in a high density of biomass. In addition, the heterogeneity of organic matter in our samples is suggested by the high variability of sample C and N composition (Table 1), and even though agitation was constant during our experiment, we observed macroscopic clumping of slurry material and vegetation debris throughout the incubation period. The threshold  $\text{NH}_4^+$  concentration values reported here were obtained for a wide range of sediment chemical and isotopic composition (*i.e.*, %C, %N, C:N,  $\delta^{13}\text{C}$ , and  $\delta^{15}\text{N}$ , see Table 1), suggesting that our threshold value should be general and apply to many benthic environments.

It is important to note both here and in other studies, the presence of *nifH* genes and transcripts do not provide definitive proof of BNF activity in anaerobic settings, as we found evidence consistent with post-transcriptional BNF regulation in benthic SRBs. In addition, prolonged incubation (5 days for ARA or  $^{15}\text{N}_2$  uptake experiments) along with addition of carbon and sulfur (lactate + sulfate) could have drastically altered the biogeochemical conditions within our microcosms (buildup of  $\text{H}_2\text{S}$ , removal of nutrients) and selected for diazotroph species different from those important in natural microbial assemblages. These long incubation periods could also have led to other artifacts that directly interfered with our

study, such as the rapid drawdown of high *in situ*  $\text{NH}_4^+$  and abnormal  $\text{NH}_4^+$  measurements with the OPA method (see Methods and Supplementary Fig. S2 and Supplementary Discussion), which strongly affect interpretations on the sensitivity of BNF activity to  $\text{NH}_4^+$  (Supplementary Fig. S6 #4, #6, #9 & #10). In this regards, the use of *in situ* measurements (Yoch & Whiting, 1986) and of several existing (ARA,  $^{15}\text{N}$ , MIMS) and newly developed methodological approaches (Aoki & McGlathery, 2019) could help produce a more accurate view of N cycling in sediments.

The use of high concentration of acetylene (~25% v/v) has been shown to influence the microbiome (16S mRNA) of estuary sediment samples during short-term incubations (~ 7 h) (Fulweiler et al., 2015). Indeed, we were able to identify changes in microbial composition based on 16S rRNA genes at 2.5% acetylene over 5 days of incubation (Supplementary Fig. S10C). However, there was no significant differences in the *nifH* gene or transcript diversity (Fig. 5B and Supplementary Fig. S10A and B) or in the response of BNF to  $\text{NH}_4^+$  addition between ARA and  $^{15}\text{N}_2$  tracer incubations (Fig. 3), indicating that the use of a sub-saturating concentration of acetylene (2.5%v/v) did not interfere with our study. Indeed, estimates of *DvH* growth inhibition under these low acetylene conditions ( $\mu_{\text{C}_2\text{H}_2} = 0.01 \text{ h}^{-1}$  vs  $\mu_{\text{ctrl}} = 0.016 \text{ h}^{-1}$ , 30-40% inhibition, Supplementary Fig. S3) is consistent with the flux of electron diverted from  $\text{N}_2$  reduction to produce ethylene during ARA, as evaluated by nitrogenase acetylene saturation in the presence of dinitrogen ( $K_{\text{mC}_2\text{H}_2} \sim 0.5\text{-}4 \text{ kPa}$ , ~40-60% saturation) (Davis & Wang, 1980). This is in agreement with a previous report of acetylene inhibition in sulfate-reducing bacteria (Payne & Grant, 1982), where an acetylene concentration of 5-10% only partially inhibited growth rate and yield. In addition, the presence and activity of molybdenum-independent nitrogenases in sediment (McRose et al., 2017; Zhang et al., 2016) could also influence BNF rate determinations and the overall N input budget in benthic areas when the ARA method is used without  $^{15}\text{N}$  calibration (Bellenger et al., 2020). Indeed, in our experiments, the average R ratio (i.e.,  $\text{C}_2\text{H}_2 : \text{N}_2$  ratio of activity), calculated with the highest BNF rate reached for each method was 2.3 (range 0.16 – 5.5, n=10), possibly

indicative of the presence of alternative nitrogenase. Finally, in a set of unsuccessful experiments containing fresh salt marsh sediment from Barnegat Bay (n=6), overlying water, and addition of lactate, we did not observe any BNF activity (as AR and  $^{15}\text{N}$  incorporation) over the course of one week and one month, respectively, even though *nifH* genes were presumably present in these sediments before our incubation (Fig. 4B). These results indicate that other factors, such as water quality, (micro-) nutrient availability, and the nature of the dissolved organic matter also likely influence the *in-situ* activity.

## Conclusion

In this study, we demonstrate that various sulfate-reducing bacteria, and very likely fermenting clostridia, found in coastal sediments exhibit high sensitivity to  $\text{NH}_4^+$  addition, with a  $[\text{NH}_4^+]$  threshold for BNF inhibition below  $10\ \mu\text{M}$ . To mechanistically link BNF inhibition and external  $\text{NH}_4^+$  concentrations, we propose a simple framework of cellular N metabolism based on the affinity constant of the  $\text{NH}_4^+$  transporter. The measured threshold is likely to apply to several benthic systems as the different diazotrophic genera identified here are similar to those described in sediments ranging from the deep sea to coasts. Analysis of published porewater data suggests that the vast majority of  $[\text{NH}_4^+]$  reported for sediments sampled at 1 cm below the sediment water interface exceed the threshold for BNF inhibition. The data are consistent with the existing framework for a relatively small contribution of benthic systems to marine new N inputs. Apparent discrepancies between our results and previous reports of low  $\text{NH}_4^+$  sensitivity in benthic environments could be explained by porewater  $\text{NH}_4^+$  concentrations close to the threshold value, fast removal of porewater  $\text{NH}_4^+$ , sediment heterogeneity, and by considering the timing necessary after addition of  $\text{NH}_4^+$  for inhibition to be recorded (e.g. 30 min – 3 h as found here). Our research provides new mechanistic insights on the biogeochemistry of nitrogen input into marine ecosystems that can help improve N cycling models and guide future measurement studies of benthic BNF.

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## Availability statement

All the data used for this study are currently available in Supplementary Materials (Supplementary Materials \_Dataset). Upon manuscript acceptance, the dataset files will be deposited to Figshare, all molecular datasets will be deposited on NCBI database, and the R Markdown code and associated files to reproduce all figures will be deposited at GitHub.

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1051

1052 **Figure Caption:**

1053 **Figure 1: Effect of ammonium on growth and biological nitrogen fixation (BNF) in**  
1054 ***Desulfovibrio vulgaris* var. Hildenborough (DvH, Panels A, C and D) and *Clostridium***  
1055 ***pasteurianum* (Cp, Panel B).** (A) Representative growth curves of DvH for initial [NH<sub>4</sub><sup>+</sup>] at  
1056 the background concentration in diazotrophic media (< 10 μM), 500 μM, and 3,000 μM.  
1057 Growth was supported by 30 mM pyruvate and sulfate. Error bars are standard deviations  
1058 from one experiment (n=3). Growth rates are calculated from independent replicates (n=3)  
1059 (B) Representative growth curves of Cp for initial [NH<sub>4</sub><sup>+</sup>] at the background concentration in  
1060 diazotrophic media (< 10 μM), 500 μM, and 3,000 μM. Growth was supported by  
1061 fermentation of sucrose. Error bars are standard deviations from one experiment (n=3).  
1062 Growth rates are calculated from independent replicates (n=3)(C) Relative BNF activity of  
1063 DvH was measured as the Acetylene Reduction Rate at 3h, 6h and 26h after addition of  
1064 NH<sub>4</sub><sup>+</sup> (to concentrations of 10, 30, 100, 300, and 3,000 μM) compared to BNF activity before  
1065 addition to DvH grown under N<sub>2</sub>-fixing conditions. Symbols represent individual samples

from a single experiment. (D) Estimated contribution of BNF to total N supply after ammonium additions. Error bars represent standard errors of the mean.

**Figure 2: Effect of ammonium concentration on biological nitrogen fixation (BNF) activity in incubations of sulfate-reducing and fermenting microbes in culture and in sediments.** (A-C) Medium  $[\text{NH}_4^+]$  (blue circles) and BNF activity (as accumulation of headspace ethylene, green squares for  $\text{NH}_4^+$ -amended) before and after addition of  $\text{NH}_4^+$  to nitrogen limited, diazotrophic cultures of *Desulfovibrio vulgaris* var. Hildenborough (DvH) (A), *Clostridium pasteurianum* (Cp) (B), and Barnegat Bay salt marsh enrichment strain SRBB2 (C) at OD ~ 0.1 (Supplementary Figs. S3-S5). Open symbols show BNF as ethylene (square) and dissolved  $[\text{NH}_4^+]$  (circle) in control experiments without addition of  $\text{NH}_4^+$ . (C-E) Dissolved  $[\text{NH}_4^+]$  (blue circles) and  $\text{N}_2$  fixed equivalent (green squares, headspace ethylene for ARA and  $^{15}\text{N}$  enrichment of particulate organic matter for  $^{15}\text{N}$  tracer) over time in salt marsh sediment slurry incubations using ARA (C-D) and  $^{15}\text{N}$  tracer methods (E). Additions are indicated with arrows. Only one representative replicate is shown to illustrate results for reasons of clarity. All incubation data (*Desulfovibrio vulgaris* n=6, SRBB1 n=2, SRBB2 n=2, salt marsh slurry n=10 for each method) are found in Supplementary Figures S3-S6 and summarized in Figure 3. Dashed lines illustrate interpolation of the BNF rate before and after  $\text{NH}_4^+$  addition to estimate  $T_R$ , the time from ammonium addition to BNF inhibition.

**Figure 2: Effect of ammonium concentration on biological nitrogen fixation (BNF) activity in incubations of sulfate-reducing microbes in culture and in sediments.** (A and B) Medium  $[\text{NH}_4^+]$  (blue circles) and BNF activity (as accumulation of headspace ethylene, green squares for  $\text{NH}_4^+$ -amended) before and after addition of  $\text{NH}_4^+$  to nitrogen limited, diazotrophic cultures of *Desulfovibrio vulgaris* var. Hildenborough (DvH) (A) and Barnegat Bay salt marsh enrichment strain SRBB2 at OD<sub>600</sub> ~ 0.1 (Supplementary Fig. S3 and S4). Open square symbols show control experiments without addition of ammonium. (C and D) Dissolved  $[\text{NH}_4^+]$  (blue circles) and  $\text{N}_2$  fixed equivalent (green squares, headspace ethylene for ARA and  $^{15}\text{N}$  enrichment of particulate organic matter for  $^{15}\text{N}$  tracer) over time in

salt marsh sediment slurry incubations using ARA (C) and  $^{15}\text{N}$  tracer methods (D). Additions are indicated with arrows. Only one representative replicate is shown to illustrate results for reasons of clarity. All incubation data (*Desulfovibrio vulgaris* n=6, SRBB1 n=2, SRBB2 n=2, salt marsh slurry n=10 for each method) are found in Supplementary Figures S3-S6 and summarized in Figure 3. Dashed lines illustrate interpolation of the BNF rate before and after  $\text{NH}_4^+$  addition to estimate  $T_R$ , the time from ammonium addition to BNF inhibition.

**Figure 3: Determination of threshold ammonium concentration (95<sup>th</sup> percentile) for inhibition of biological nitrogen fixation (BNF) in liquid cultures and sediment slurries.**

Combined results from incubations of *Desulfovibrio vulgaris* var. Hildenborough and SRBB1&2 consortia enriched from NJ salt marshes (A), ARA (B), and  $^{15}\text{N}$  tracer (C) slurry incubations. BNF rates are the slope between time points from Figure 2. Blue boxes indicate the range of ammonium concentrations at which BNF was detected; black lines indicate average ammonium concentration, and dotted lines the 95-percentile (*i.e.* threshold values). The symbols distinguish data for BNF onset before (closed symbols) and BNF resumption after (open symbols)  $\text{NH}_4^+$  addition. Panel C shows no post-addition data as BNF activity never resumed in  $^{15}\text{N}$  tracer incubations of sediment (See Supplementary Fig. S6 and Fig. 2D). In panel B, post addition data for the 4 of 10 ARA incubations showing abnormal OPA measurements (See Supplementary Discussion) were highlighted with red stars.

**Figure 4: Microbial composition of diazotroph community in salt marsh sediment incubations.** (A) Phylogeny of the top 500 *nifH* genes OTUs (n=497) and the most abundant *nifH* transcripts OTUs (n=55, relative abundance>0.2%, total count >1000) from salt marsh slurry incubations relative to a representative subset (n=152) of a recently curated *nifH* library (Kapilli et al. 2020). Information on *nifH* genes for *Desulfovibrio vulgaris* v. Hildenborough (DvH), and SRBB1 and SRBB2 enrichments consortia (n=46) are also shown. (B) Heatmap of *nifH* genes OTU relative abundance (log-scale) at the end of incubation with class-level information for the 500 first OTUs (bottom bar plot, > 80 % total abundance) and detailed phylogeny and closest relatives for the most abundant sequences (left panel, OTUs

with relative abundance >0.2% & total count >1000, n=59). Hierarchical clustering of individual sediment samples based on *nifH* composition (top left of panel) shows the relationship between sample *nifH* genes composition, incubation condition (ARA vs.  $^{15}\text{N}$ ), and geographical origin of the sediments. Purple box highlights low DNA samples with *nifH* composition similar to that of control extraction samples (reagent blank). Initial condition represents the samples before incubation.

**Figure 5: Relationships between nitrogenase activity, microbial diversity, and sediment biogeochemical characteristics in sediment incubation.** (A) Changes in richness (number of unique OTUs) in nitrogen fixer *nifH* transcripts at the end of incubation as a function of final ammonium concentrations (n=10). Symbol shade is proportional to the final BNF activity. (B) Discrimination analysis of nitrogen fixer communities (CAP analysis on Bray-Curtis distance). (C) Discriminant analysis of sediment biogeochemical activity and characteristics (redundancy analysis). In B and C, samples using acetylene reduction assay and  $^{15}\text{N}$  tracer are highlighted in blue and green, respectively (shaded ovals represent 95% confidence ellipses) and axes show the percent of variance explained and the statistical significance for each explanatory axis as estimated using permutation Anova. Variables Bold are statistically significant in the model. In panel B, results were unchanged when we control for the origin of the sediment (NJ, NH, and MA).

**Figure 6: Theoretical model for ammonium threshold for biological nitrogen fixation in benthic environments.** (A) Overview of the Michaelis-Menten framework for  $\text{NH}_4^+$  sensitivity applied to N-fixers in benthic environment, showing the effect of increasing pore water ammonium with depth on the transition from diazotrophy to ammoniotrophy. (B) Compilation of experimental data of model sulfate reducers *D. vulgaris* H. (ARA, N = 6 + 3 controls), SRBB enrichments (ARA, N = 4 + 2 controls), and three salt marshes of the Northeastern US (both ARA and  $^{15}\text{N}$ , N=20) for  $[\text{NH}_4^+] < 100 \mu\text{M}$ . The red dashed vertical

1146 line shows the 95-percentile  $[\text{NH}_4^+]$  (*i.e.* threshold value), and the dashed horizontal line  
1147 indicated where BNF is 95% inhibited by ambient ammonium. (C) Frequency distribution of  
1148 threshold depth (*i.e.*, depth at which  $[\text{NH}_4^+]_{\text{threshold}}$  was reached) derived from literature review  
1149 of 26 study in several benthic environments (151 sites, 334 replicates, Supplementary Fig.  
1150 S12 and Table S1). The red line represents cumulative frequencies, and the dashed vertical  
1151 line represent the median value. (D) Geographic location and threshold depth (color scale) of  
1152 reviewed study sites separate by benthic area (shape). Replicate numbers are proportional  
1153 to point size.

Figure 1.

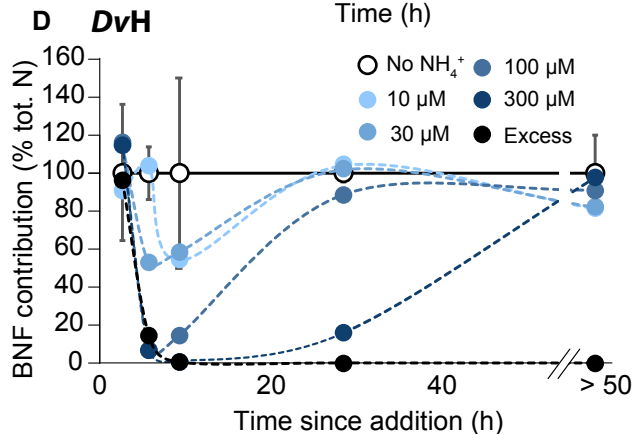
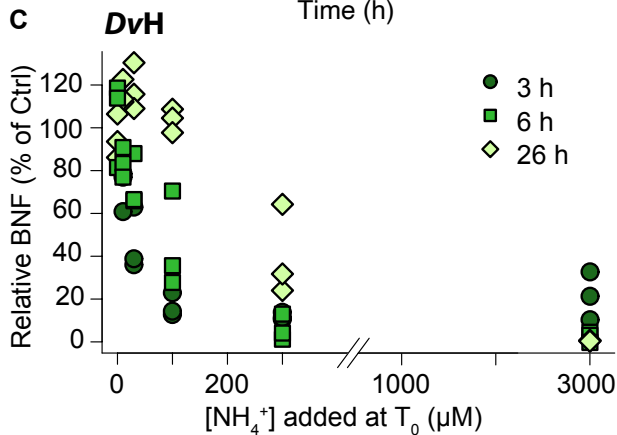
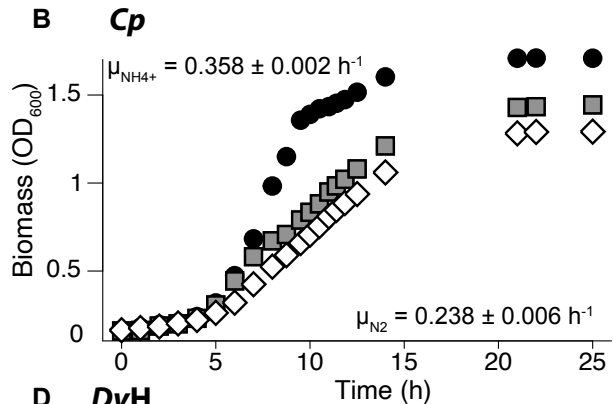
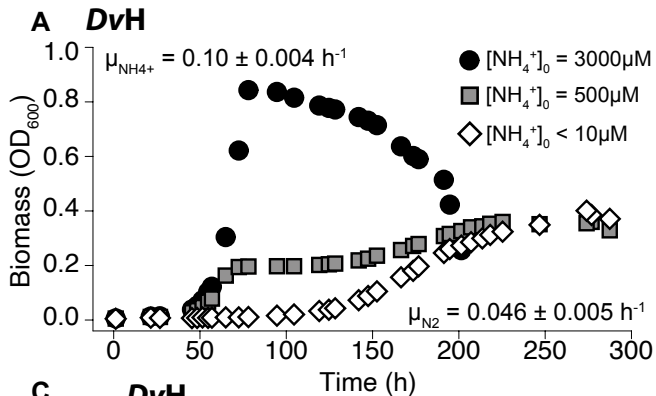


Figure 2.

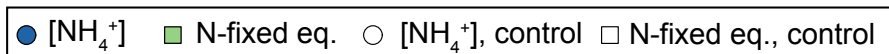
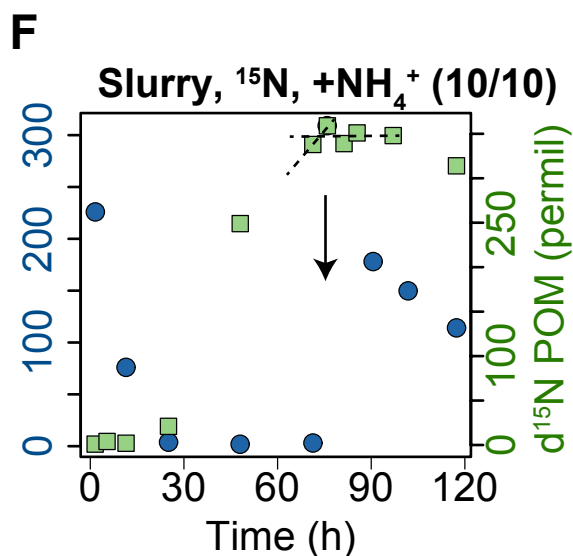
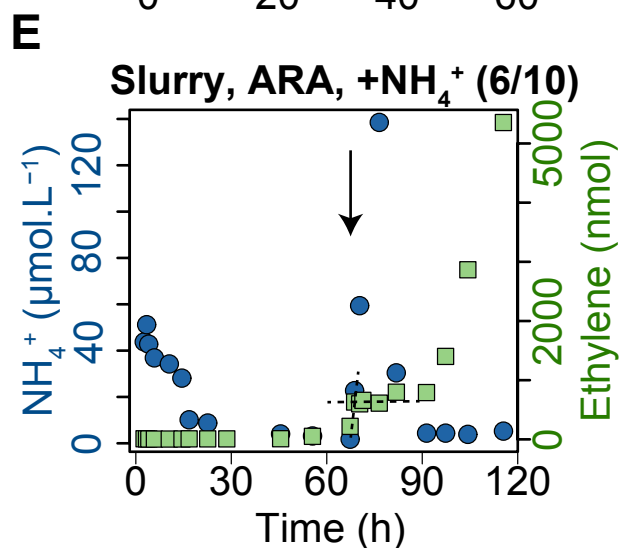
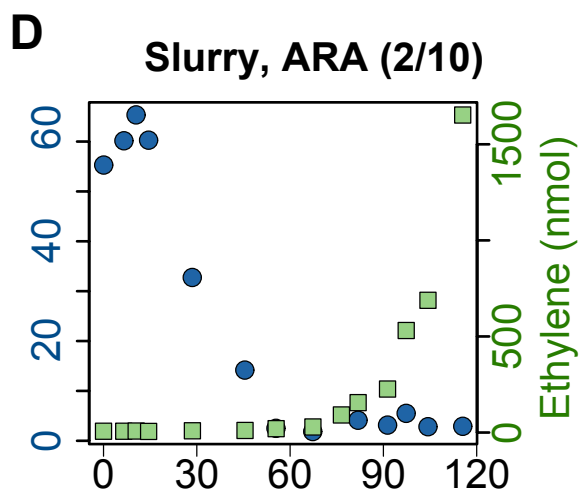
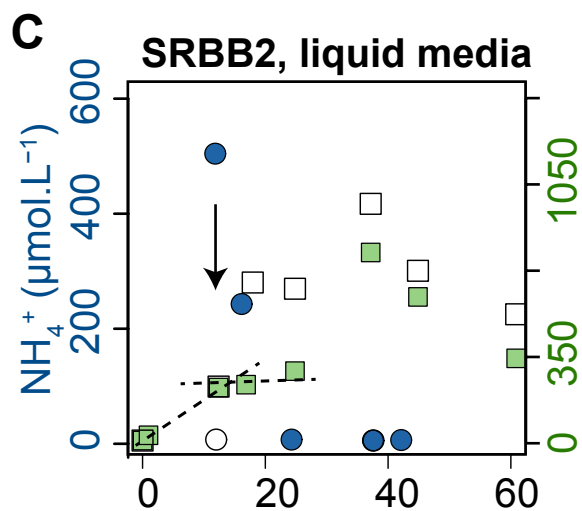
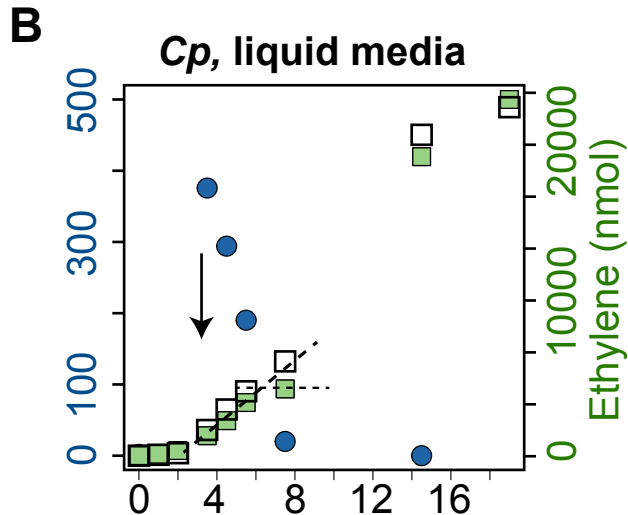
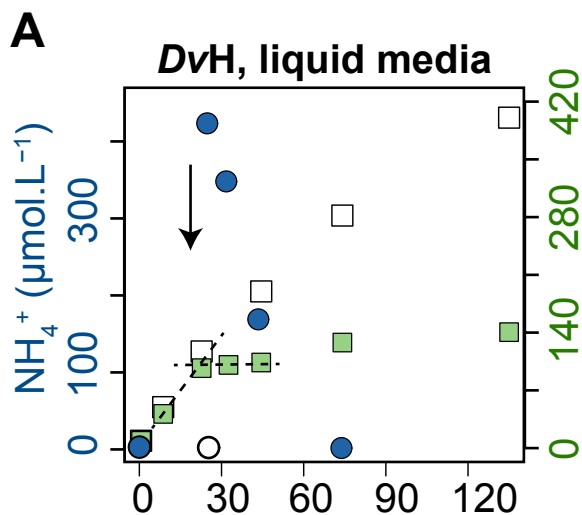


Figure 3.

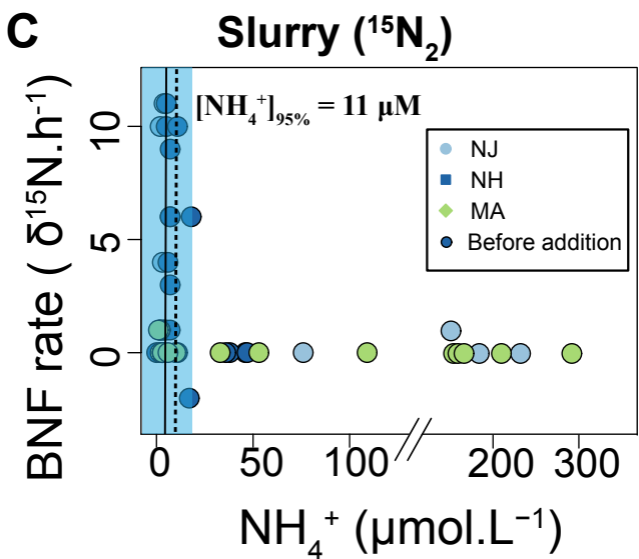
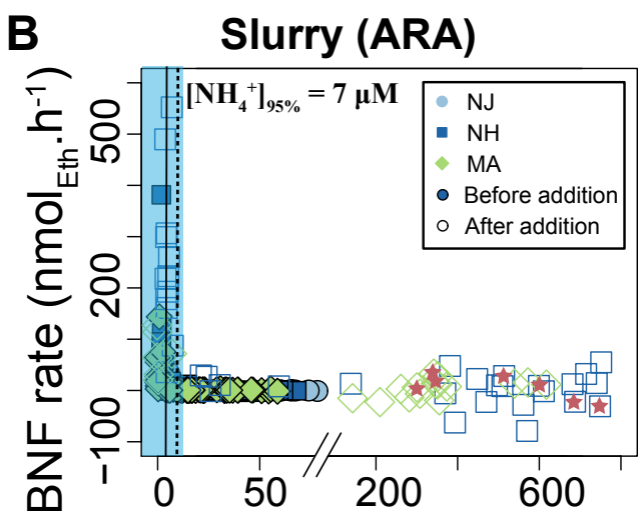
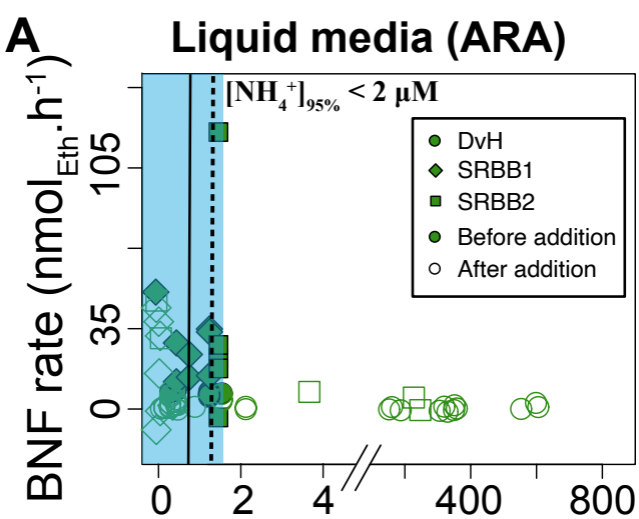


Figure 4.

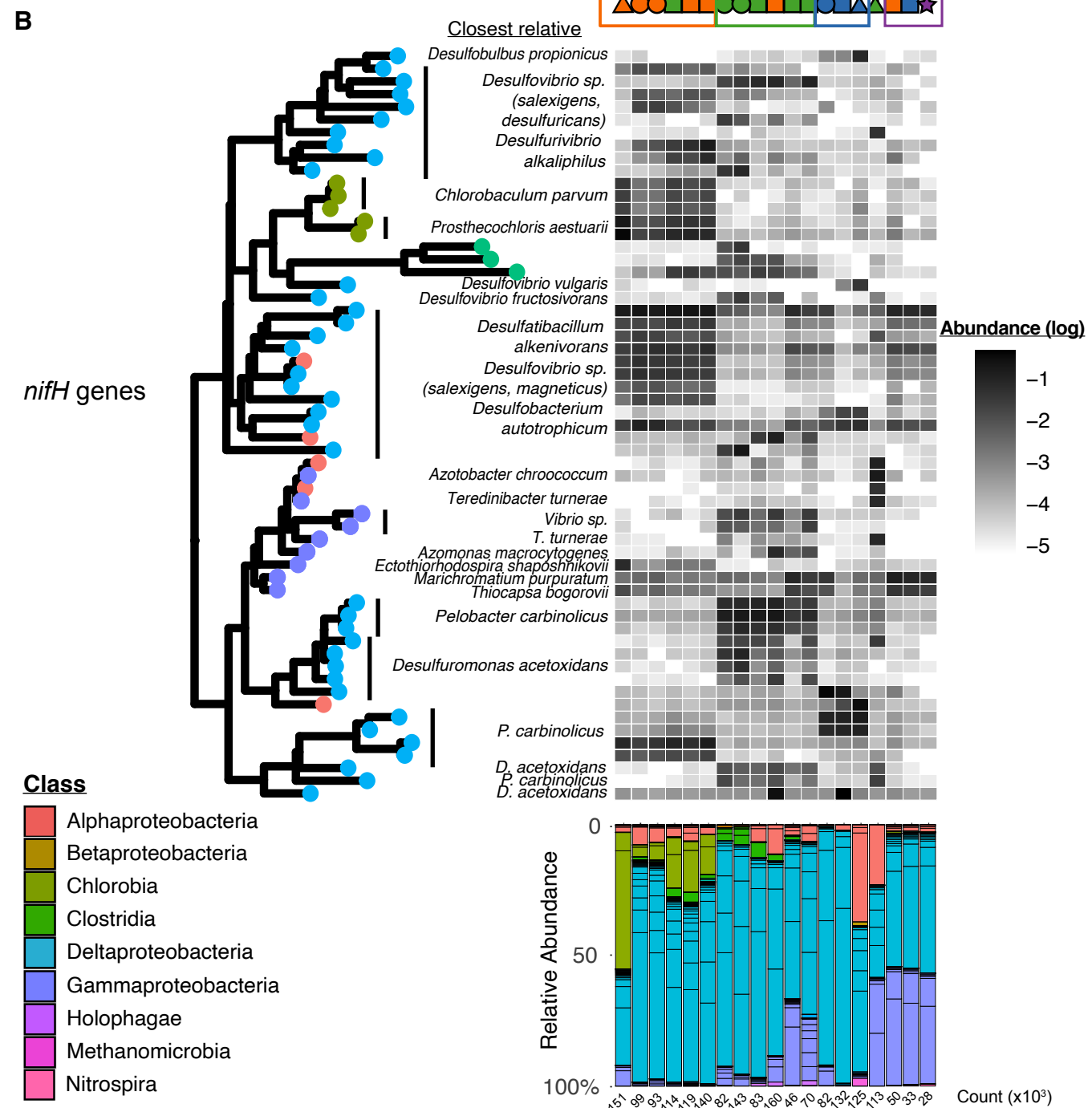
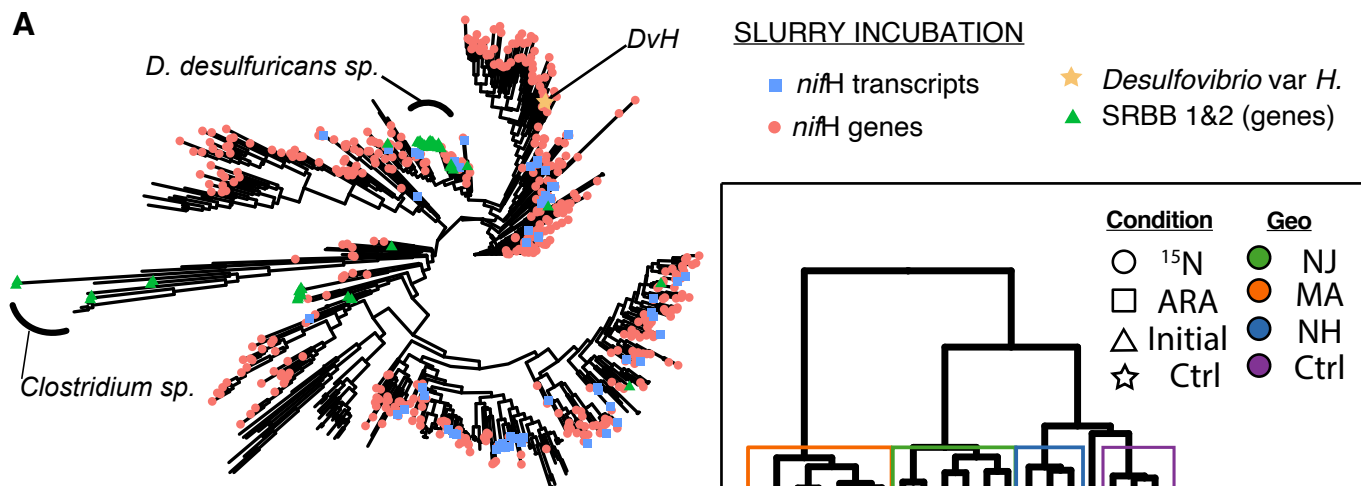


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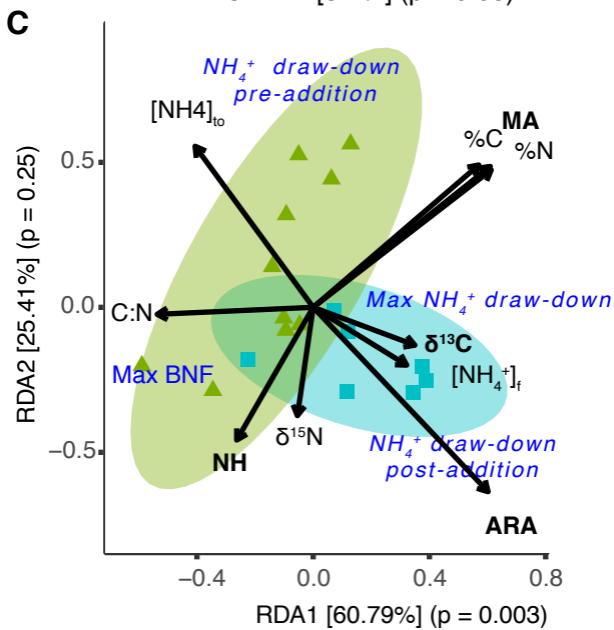
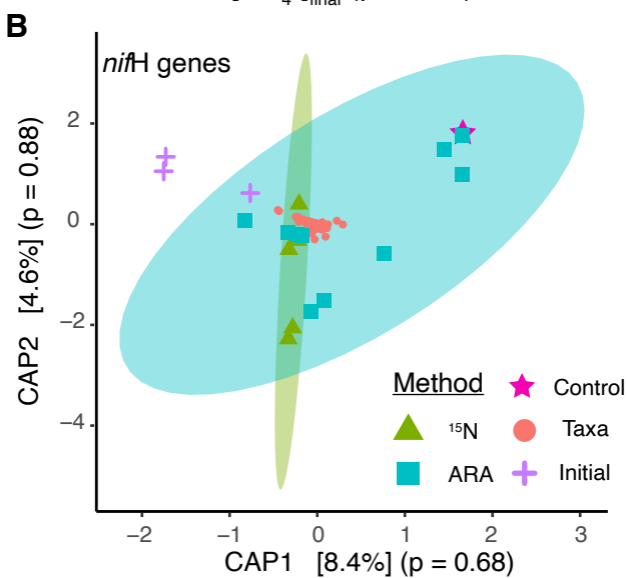
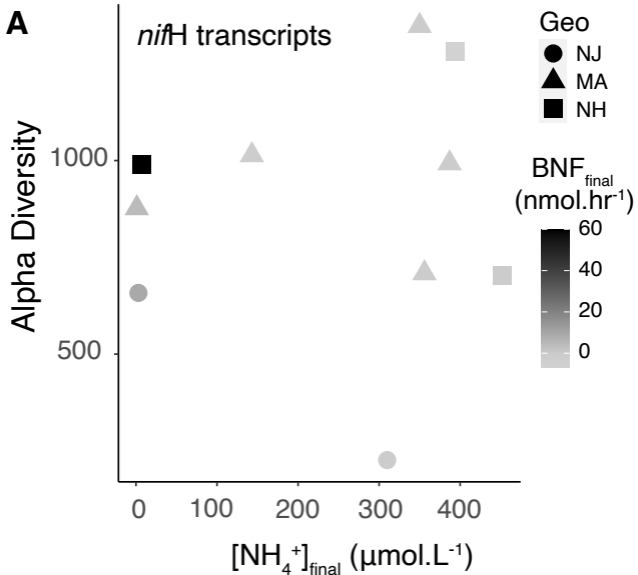
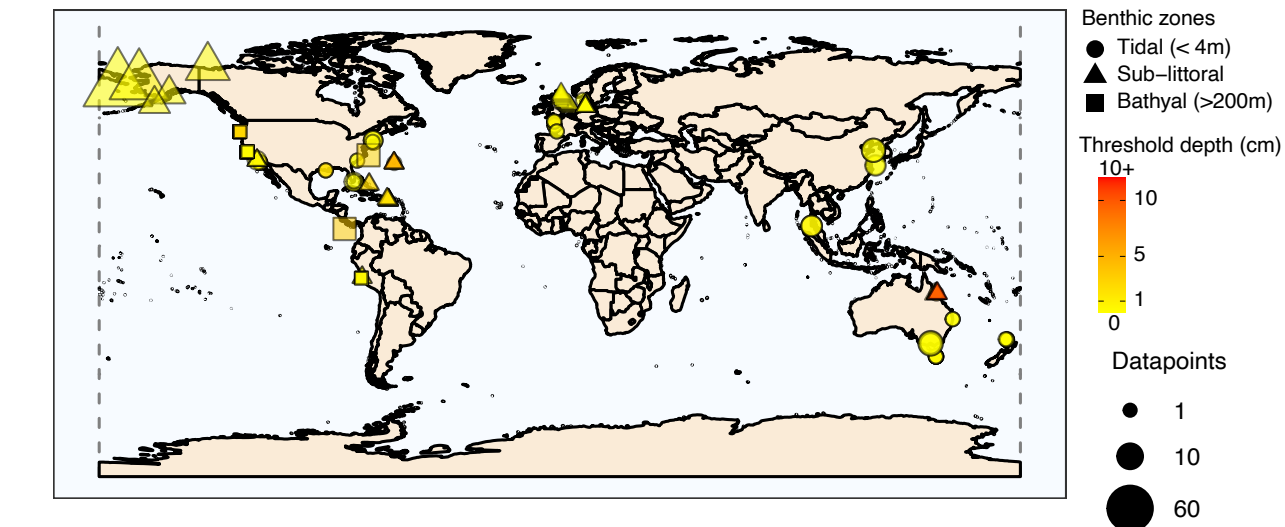
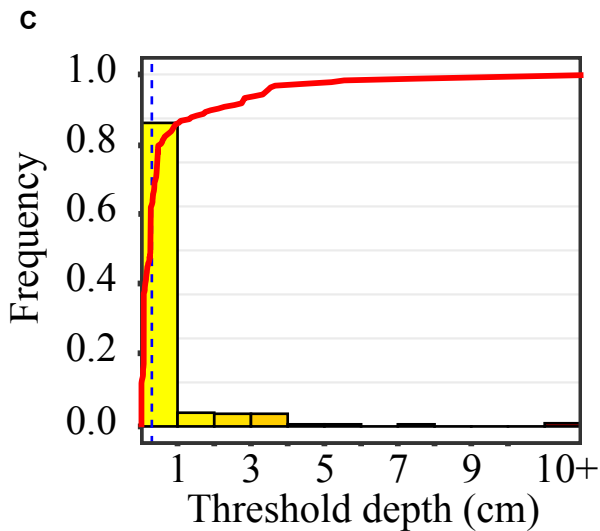
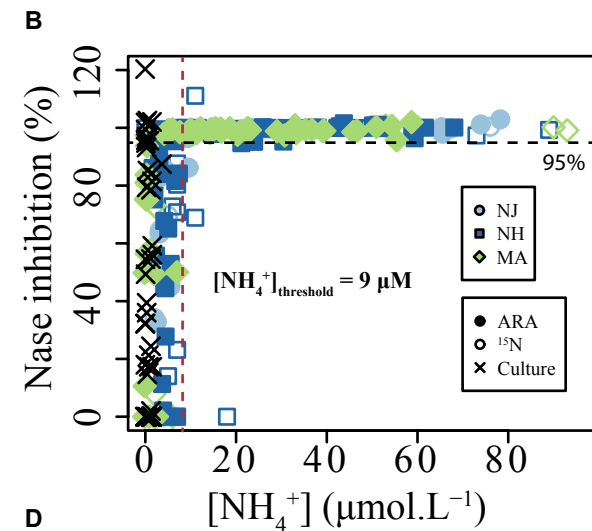
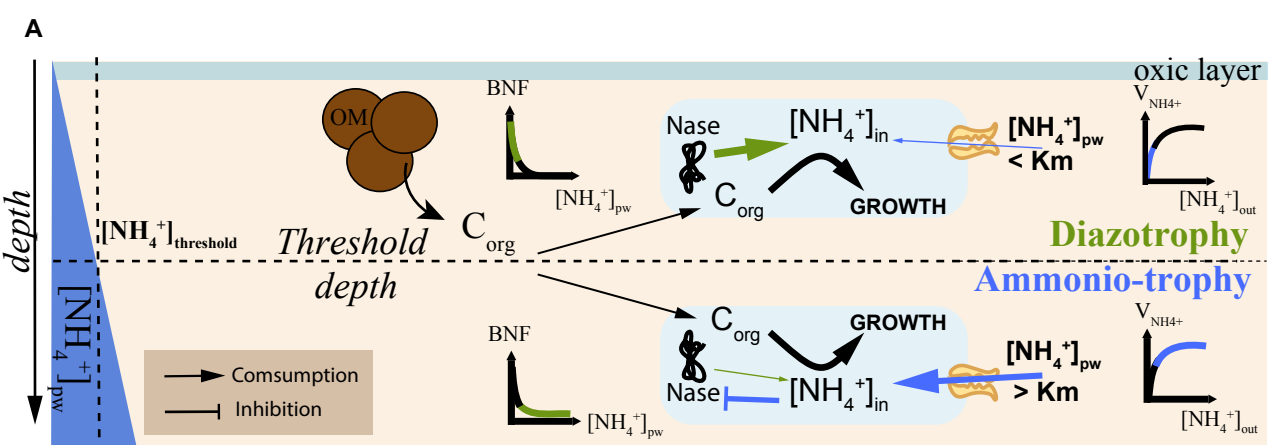


Figure 6.



**Table 1 : Sediments carbon and nitrogen compositions**

	%C <sup>1</sup>	SD* <sub>intra</sub> (n=10)	%N	SD <sub>intra</sub> (n=10)	C:N <sup>2</sup>	SD <sub>intra</sub> (n=10)	δ <sup>13</sup> C	SD <sub>intra</sub> (n=10)	δ <sup>15</sup> N	SD <sub>intra</sub> (n=10)
<b>Barnaget Bay (NJ)</b>	6.7	3.4	0.30	0.18	28.3	7.0	-27.7	0.4	2.2	0.9
<i>SEM** (n=2)</i>	<i>0.3</i>	<i>0.1</i>	<i>0.01</i>	<i>0.01</i>	<i>0.4</i>	<i>2.6</i>	<i>0.2</i>	<i>0.1</i>	<i>0.2</i>	<i>0.6</i>
<b>Great Bay (NH)</b>	5.4	0.9	0.43	0.07	14.7	0.5	-15.5	0.2	5.1	0.4
<i>SEM (n=4)</i>	<i>0.2</i>	<i>0.1</i>	<i>0.01</i>	<i>0.01</i>	<i>0.1</i>	<i>0.1</i>	<i>0.1</i>	<i>0.1</i>	<i>0.1</i>	<i>0.2</i>
<b>Sippewissett Marsh (MA)</b>	22.2	6.9	1.74	0.53	14.9	0.3	-18.1	0.4	3.2	0.4
<i>SEM (n=4)</i>	<i>1.1</i>	<i>0.5</i>	<i>0.08</i>	<i>0.05</i>	<i>0.2</i>	<i>0.0</i>	<i>0.1</i>	<i>0.2</i>	<i>0.1</i>	<i>0.1</i>

<sup>1</sup>. %C and %N are on a mass basis relative to oven-dried sediments ( $\text{g}_{\text{element}} \cdot \text{g}_{\text{ODW}}^{-1}$ ). <sup>2</sup>. C:N are molar ratio.

\*SD<sub>intra</sub> represents sample heterogeneity, as the SD of 10 replicates taken from the same sample over multiple timepoints and averaged per sediment geography.

\*\*SEM represents the standard error of the mean for each sediment geography.