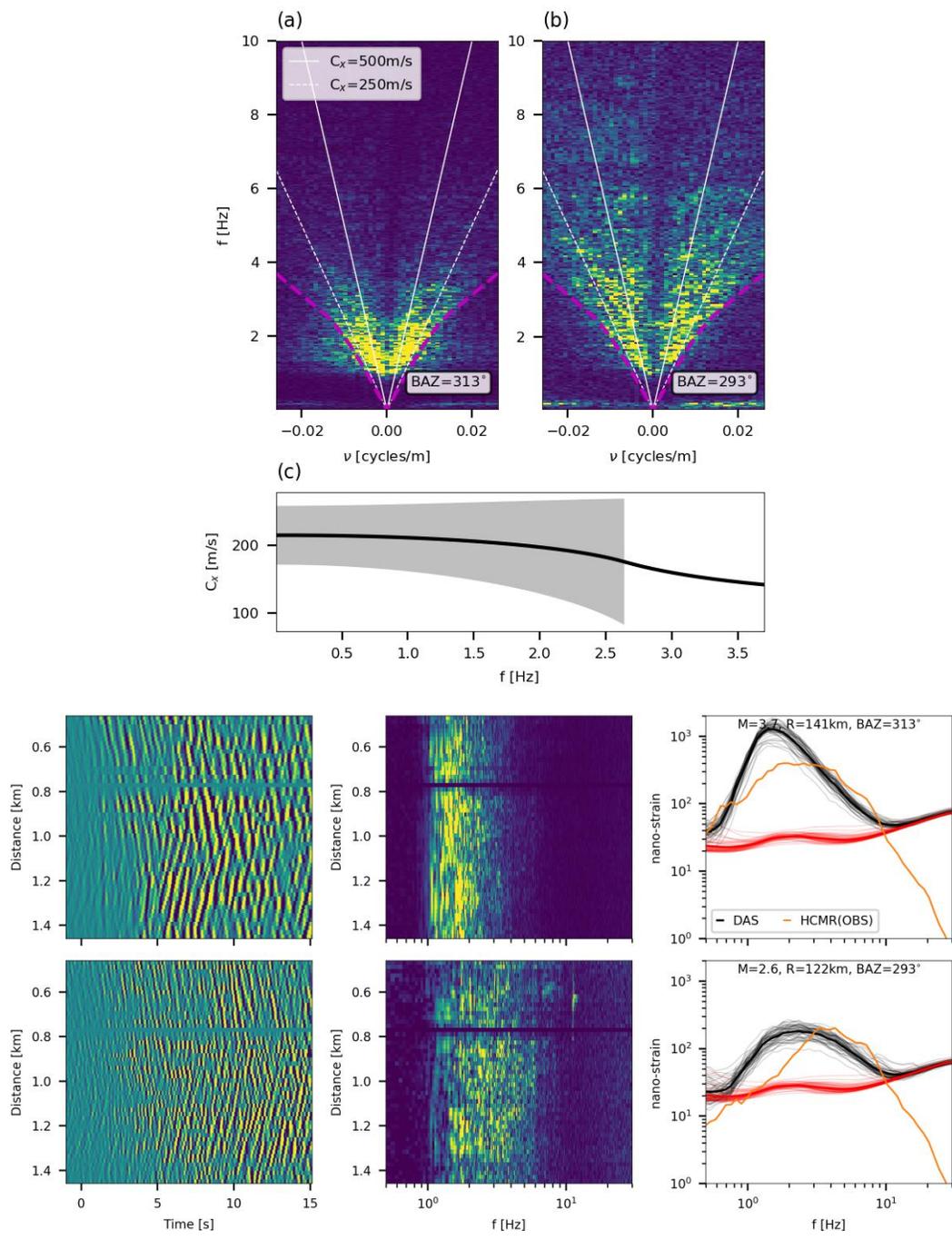
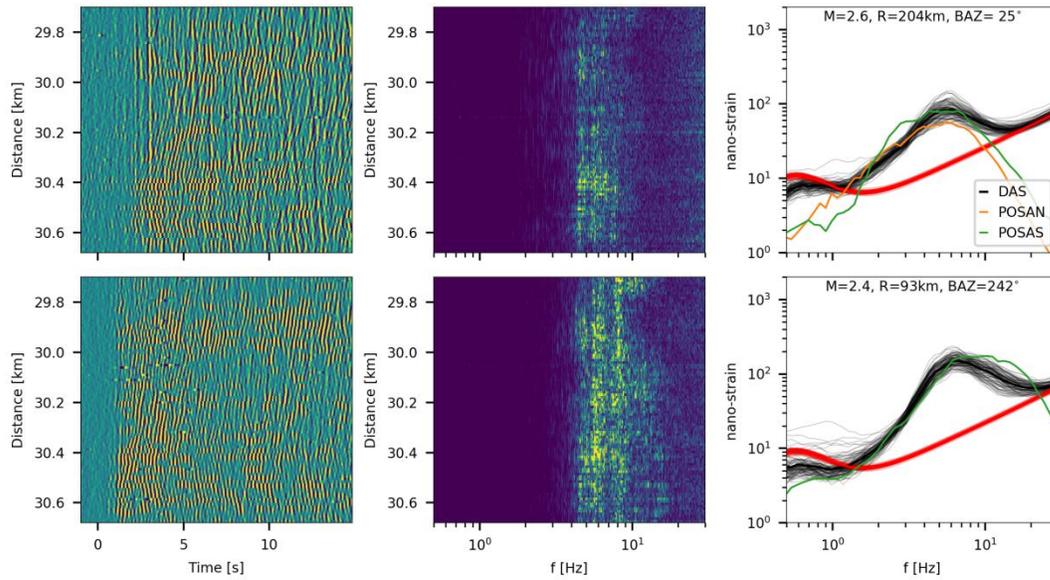
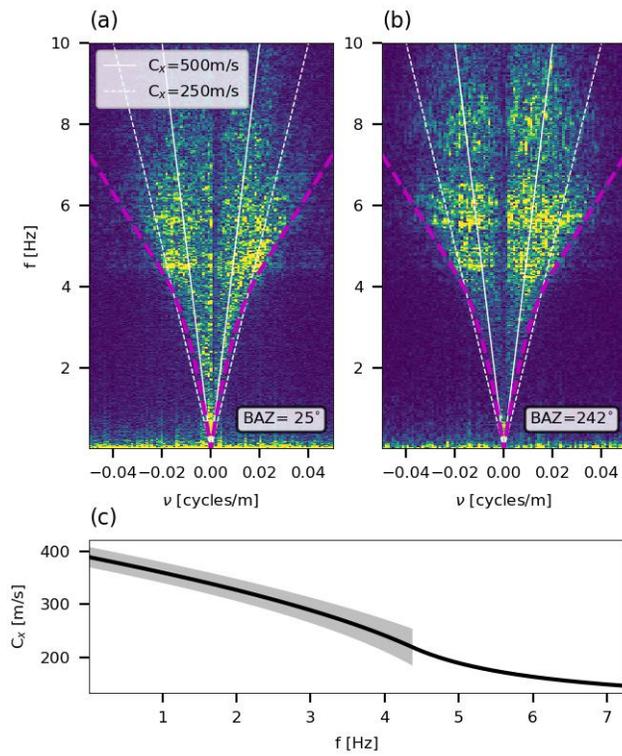


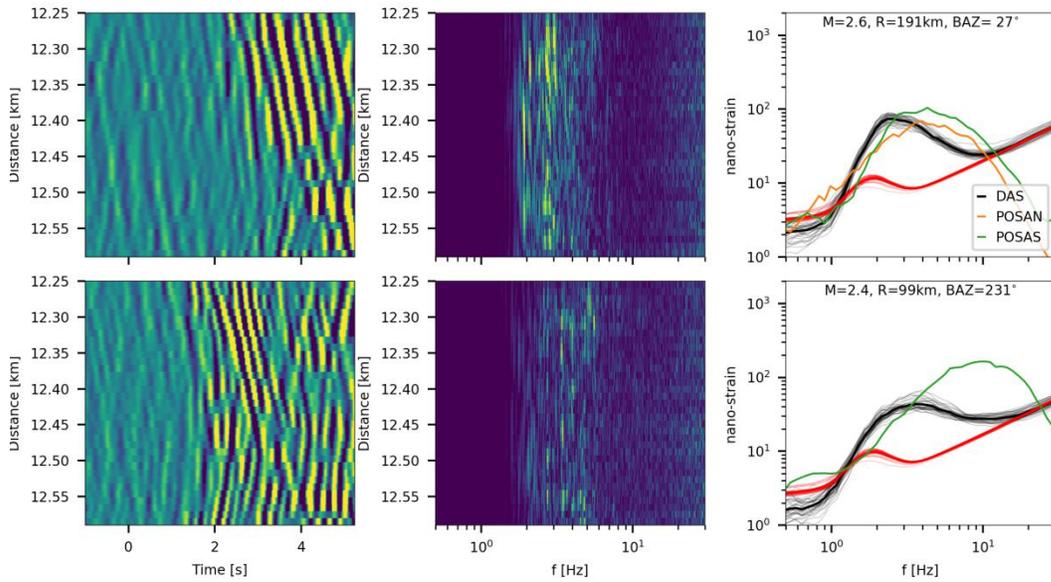
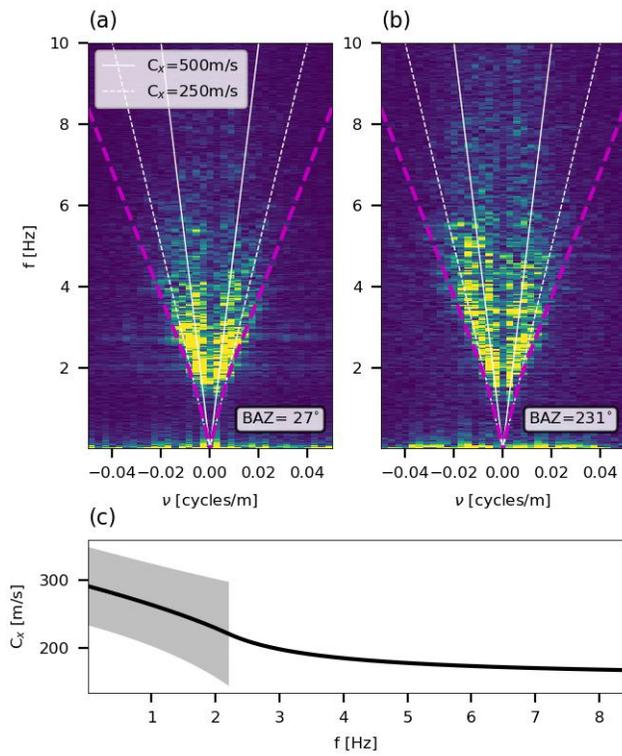
**Figure S5.** As in figures 5 (top) and 6 (bottom) for 2 earthquakes recorded by HCMR between 6 and 6.3 km from the interrogator, at a depth of 160 m.



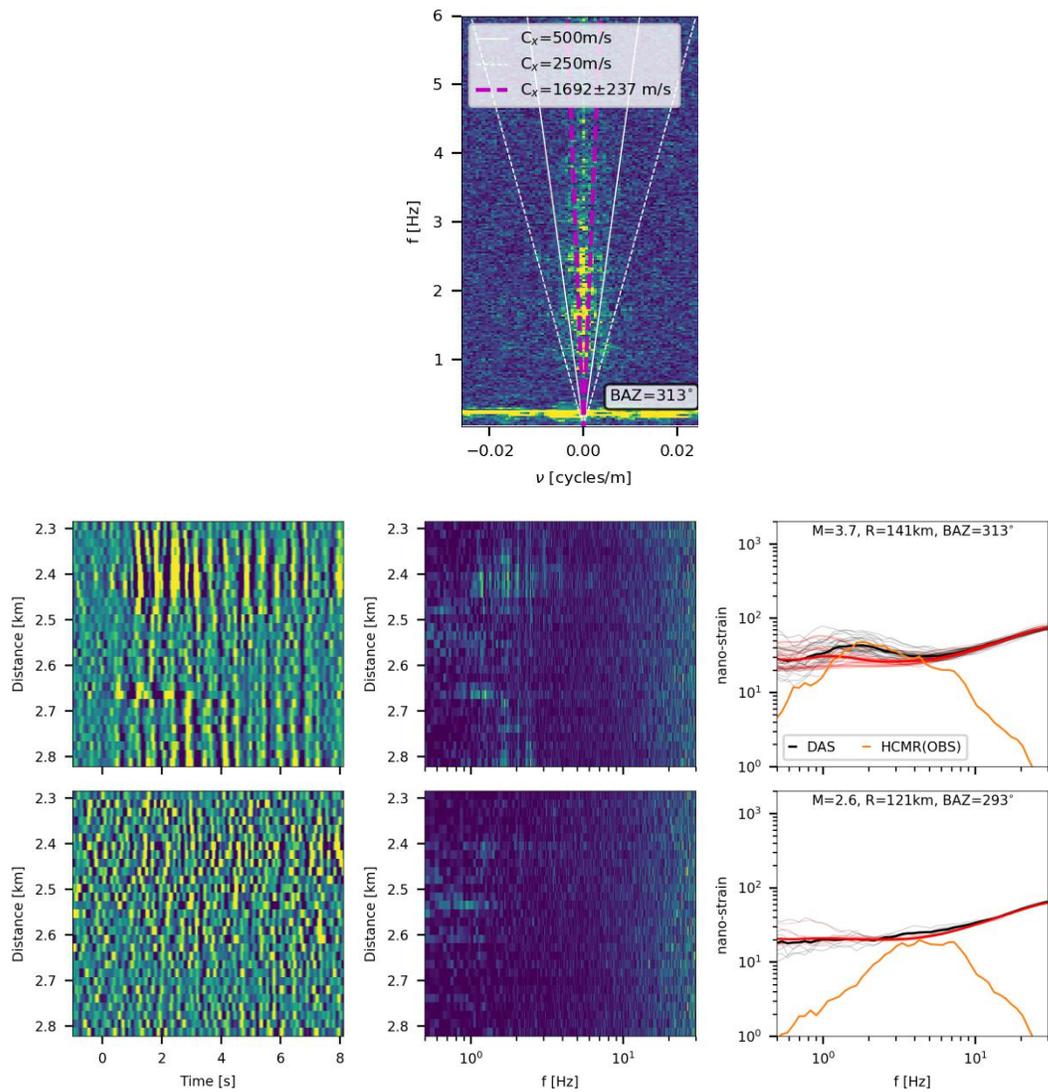
**Figure S6.** As in figures 5 (top) and 6 (bottom) for 2 earthquakes recorded by HCMR between 0.5 and 1.5 km from the interrogator, between 3 and 18 m depth.



**Figure S7.** As in figures 5 (top) and 8 (bottom) for 2 earthquakes recorded by MEUST between 29.7 and 30.7 km from the interrogator, at a depth of 2.35 km.



**Figure S8.** As in figures 5 (top) and 8 (bottom) for 2 earthquakes recorded by MEUST between 12.2 and 12.6 km from the interrogator, at a depth of 550 m.



**Figure S9.** As in figures 5 (top) and 8 (bottom) for 2 earthquakes recorded by HCMR between 2.3 and 2.85 km from the interrogator, between 15 and 60 m depth. The earthquake in top panels is detected while that in the bottom panels is not detected.

**Text S1.**

S-wave ground accelerations were modeled using the omega-squared model (Brune, 1970). This model describes the far-field body-wave radiation in the frequency domain:

$$\ddot{\Omega}(f) = (2\pi f)^2 \frac{\Omega_0}{1 + (\frac{f}{f_0})^2}, \quad (S1)$$

Where  $\ddot{\Omega}(f)$  is acceleration spectra,  $f$  is frequency,  $f_0$  is the source corner frequency, and  $\Omega_0$  is the low frequency displacement value. The parameters  $\Omega_0$  and  $f_0$  are related to the seismic moment,  $M_0$ , and the stress drop,  $\Delta\tau$  (Eshelby, 1957):

$$\Omega_0 = \frac{M_0 U_{\phi\theta} F_s}{4\pi\rho C_s^3 R}, \quad (S2a)$$

and:

$$f_0 = k C_s \left( \frac{16 \Delta\tau}{7 M_0} \right)^{1/3}, \quad (S2b)$$

where  $U_{\phi\theta}$  is the radiation pattern,  $F_s$  is the free-surface correction factor,  $C_s$  is the S-wave velocity,  $R$  is the hypocentral distance,  $\rho$  is the density and  $k$  is a constant. The  $f_0$ - $\Delta\tau$  relation assumes a circular fault, and is a sufficient approximation for many earthquakes, including those analyzed in the manuscript. High frequency attenuation is modeled by multiplying the omega-squared source model (Equation S1) with a decaying exponent:

$$\ddot{\Omega}(f) = (2\pi f)^2 \frac{\Omega_0}{1 + (\frac{f}{f_0})^2} \exp(-\pi\kappa f), \quad (S3)$$

where  $\kappa$  is an attenuation parameter.

The following parameter tuning was set when modeling ground motion accelerations:

Parameter	Value
$\Delta\tau$	4MPa
$\kappa$	0.04 sec
$\rho$	2600 kg/m <sup>3</sup>
$C_s$	3200 m/s
$U_{\phi\theta}$	0.63
$F_s$	2
$k$	0.37