

1 **Dynamic full-field imaging of rupture radiation:**
2 **Material contrast governs source mechanism**

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8 **Abstract**

9 In seismology, the rupture mechanism of an earthquake, a glacier stick-slip and a landslide
 10 is not directly observed, but inferred from surface measurements. In contrast, laboratory
 11 experiments can illuminate near field effects, which reflect the rupture mechanism but are
 12 highly attenuated in the case of real-world surface data. We directly image the elastic
 13 wave-field of a nucleating rupture non-invasively in its near-field with ultrasound speckle
 14 correlation. Our imaging yields the particle velocity of the full shear wave field at the
 15 source location and inside the 3D frictional body. We experimentally show that a strong
 16 bimaterial contrast, as encountered in environmental seismology, yields a unidirectional or
 17 linear force mechanism for pre-rupture microslips and decelerating supershear ruptures. A
 18 weak contrast, characteristic for earthquakes, generates a double-couple source mechanism
 19 for sub-Rayleigh ruptures, sometimes preceded by slow deformation at the interface. This
 20 deformation is reproduced by the near field of a unidirectional force.

21 **Key Points:**

- 22 • Noninvasive elastic near-field laboratory observations reveal source mechanisms of
 23 micro-slips, supershear and sub-Rayleigh ruptures.
- 24 • Strong material contrasts as encountered in glacier stick-slip and landslides cause
 25 single force micro-slips and supershear ruptures.
- 26 • Weak material contrasts lead to a double-couple mechanism, sometimes preceded by
 27 the near field radiation of a slowly rising single force.

28 **Plain Language Summary**

29 Earthquakes, avalanches, icequakes and landslides originate from a common process:
 30 rupture at a material interface. During a rupture, for example when a landslide slips, a
 31 characteristic pattern of seismic waves is created. This pattern differs at the earth's surface
 32 and the rupture interface, which is the source of the seismic waves inside the earth. Usually
 33 scientists only measure the waves arriving at the surface and need to deduce the wave pattern
 34 inside the earth from the surface measurement. We build a laboratory experiment which
 35 enables us to film wave propagation around the rupture surface, as if we had a camera inside
 36 the material. We film waves emitted during and prior to a rupture. For a soft material on
 37 a hard surface, such as encountered in icequakes or landslides, a single force model better
 38 explains the observed wave pattern than the commonly used model of four distributed forces.
 39 The rupture moves faster than shear waves propagate which results in a supershear cone,
 40 the elastic equivalent to the acoustic Mach cone created by supersonic aircrafts. For two
 41 materials of similar hardness, such as encountered in earthquakes, the classic model of four
 42 forces better explains the ruptures, which travel at sub-shear speed.

43 **1 Introduction**

44 For most earthquakes, the longstanding discussion on the appropriate force representa-
 45 tion of the earthquake source has been decided in favor of the double-couple (DC) source.
 46 It is the body force equivalent to slip on a fault and consists of two force couples acting at
 47 the earthquake source point (Pujol, 2003; Aki & Richards, 2009). However, other rupture
 48 observations such as icequakes, landslides, induced seismicity and deep earthquakes are not
 49 always well reproduced by a standard double-couple model. For example, Ben-Zion and
 50 Ampuero (2009) theoretically show that brittle rupture is associated with a non-double-
 51 couple damage related source term. Kwiatek et al. (2011); Kwiatek and Ben-Zion (2013)
 52 discuss the presence of tensile opening during induced seismicity and aftershocks of a Mw

53 1.9 earthquake. In the case of glacial sliding, Ekström et al. (2003) report that a single force
 54 centroid inversion shows a better match than standard moment tensor inversion (Harvard-
 55 CMT). In the laboratory, Lykotrafitis and Rosakis (2006) found indications for wrinkle-like
 56 rupture and tensile opening in a Homalite-on-steel friction experiment.

57 Inversion for earthquake sources is mostly done in the far-field and suffers from ambi-
 58 guity. With the exception of volcanic seismicity, where hypocenters are shallow (Lokmer
 59 & Bean, 2010), the seismic near field suffers strong attenuation and is often concealed by
 60 ambient noise. In contrast to real-world seismic data, laboratory rupture experiments allow
 61 for dense instrumentation and direct imaging of rupture propagation. For example, the first
 62 unambiguous proof of supershear rupture was provided by strain imaging through photo-
 63 elastic experiments of sliding Homalite plates by Rosakis and Coker (1999). Recently, the
 64 group retrieved wave motion displacements of supershear ruptures through digital image
 65 correlation (Rubino et al., 2020, 2022). Latour et al. (2011, 2013) introduced a new direct
 66 rupture imaging method using ultrafast ultrasound imaging that allows for observation of
 67 shear wave radiation during rupture propagation in soft materials: the particle velocity of
 68 a propagating shear wave is retrieved through speckle tracking of subsequent ultrasound
 69 (US) backscatter images. In contrast to photo-elasticity, this method is not restricted to
 70 2D setups. Their results show that during hydrogel-on-sandpaper friction the depinning
 71 of the gel from the sandpaper is well matched by a singular bell shaped (Gaussian) shear
 72 point force. They also directly observed the effects of barriers on rupture propagation on
 73 a hydrogel-glass interface with a granular inter-layer. At first glance hydrogels might seem
 74 counterintuitive as a material choice in rupture experiments. However, they have been ex-
 75 tensively used as geological analogues (van Otterloo & Cruden, 2016). An historic example
 76 is the jelly experiment of Reid (1910) that led to the elastic rebound theory. More recent ex-
 77 amples include a subduction-analogue gelatin setup (Corbi et al., 2011),(Corbi et al., 2017)
 78 and volcanic modeling (Kavanagh et al., 2018).

79 Here we investigate the source mechanism of the failure of a granular asperity in a
 80 laboratory friction experiment using a new setup based on the methodology introduced by
 81 Latour et al. (2011, 2013). Direct imaging of the near field of a propagating rupture allows
 82 us to compare the laboratory rupture to a kinematic rupture simulation using elastodynamic
 83 Green's functions. We compare the case of weak and strong bimaterial contrast and test
 84 single-force and double-couple source models to find the source mechanism depending on
 85 the elastic contrast and type of slip event.

86 **Experimental setup**

87 All results are derived from the dynamic wave field imaging of two experimental sce-
 88 narios: sliding of an asperity along an interface with a strong or a weak bimaterial contrast
 89 (Fig. 1). The strong bimaterial contrast is constituted of a glass - hydrogel (Polyvinyl-alcohol
 90 - PVA) interface (Fig. 1(a)-(b)) and the weak bimaterial contrast by a hydrogel-hydrogel
 91 interface (Fig. 1(d)). Since the hydrogels are homemade and non-standardized, an elasticity
 92 contrast has to be assumed between them. The frictional behaviour is ensured by a sand
 93 patch mimicking an asperity on a smooth surface. The glass plate is moved by a Kollmorgen
 94 stepper motor, which induces the deformation then subsequent sliding of the partly blocked
 95 gel via the frictional contact of the sand asperity. Seismic radiation is emitted upon failure
 96 of frictional contacts due to stick-slip ruptures, and is observed by ultrasonic speckle corre-
 97 lation imaging. The observation plane is centered in the gel, perpendicular to the interface
 98 and reaches from the asperity to the gel surface.

99 The imaging methodology is exemplary shown in Fig. 1(c) with data from the weak
 100 interface experiment. Ultrasound backscatter images show a zone of high reflectivity at the
 101 gel-gel interface at 4 cm depth. It is caused by the sand layer and the presence of air in
 102 between the two hydrogels. Imaging below the interface is feasible, but the speckle quality
 103 is deteriorated due to strong ultrasound backscattering. In both gels, a 1 cm thick layer of

intermediate reflectivity is observed next to the sand. It is likely caused by increased deposition of the backscatter agent (graphite). While graphite changes the ultrasonic impedance, shear wave propagation at the frequencies of interest remains unaffected. The phase correlation of successive ultrasound speckle images allows to resolve the shear wave induced vertical particle displacement between two snapshots, which is the apparent particle velocity $\frac{\partial u_z}{\partial t}$ (Pinton et al., 2005).

The dynamic observation is made possible by the high velocity contrast of the shear and compression waves in the hydrogel: while the compressional ultrasound travels at approximately 1500 m s^{-1} , the rupture induced shear waves propagate at speeds below 10 m s^{-1} . Plane ultrasound pulses at high frame rate allow for the shear wave particle velocity to be temporally well resolved ($\Delta_t = 0.33 \text{ ms}$). The ultrasound frequency (5 MHz) ensures the spatial resolution ($\lambda_{US} = 0.3 \text{ mm}$). Hence, a shear wave propagating at 7 m s^{-1} and 250 Hz is sampled at 25 US-wavelengths per shear wavelength ($\lambda_{shear} = 7 \text{ mm}$). Consequently, the z -component of the entire transverse displacement field, including near-field terms, is observed.

Kinematic modeling of the radiated wavefield

The observed wavefields radiated by the slip events are compared to direct kinematic wavefield modeling of equivalent body-force models (see section S2.6-S2.8) In each case, we compare the single-force and the double-couple solutions. The source moves to simulate propagation of rupture fronts, and its velocity as well as the local source time functions are manually adjusted to obtain a good match to the data.

The displacement $u_{ij}(x, t)$ due to a unidirectional force (UF) in the x_j -direction with a source time function $X_0(t)$ at the source position is the convolution of X_0 with the elastodynamic impulse response (Green's function G_{ij}). It is the superposition of the compression and shear wave far-fields and the elastic near-field:

$$\begin{aligned} u_i(\vec{x}, t) &= X_0 * G_{ij} \\ &= X_0 * G_{ij}^{Near} + X_0 * G_{ij}^{Far-P} + X_0 * G_{ij}^{Far-S} \end{aligned} \quad (1)$$

The full expression is given in Section S2.6 and a thorough derivation can be found in Aki and Richards (2009) Chapter 3-4.

In contrast to the UF-solution in Eq. (1), the Green's function for a double-couple (DC) model can be separated into five physically meaningful terms: Near-field, intermediate S-field and P-field, and far S-field and P-field.

$$G_{DC} = G^{Near} + G^{IP} + G^{IS} + G^{FP} + G^{FS} \quad (2)$$

The full analytical solution for the displacement field of a DC source can be found in Section S2.7.

In the following, we first present the results on the strong and weak contrast bimaterial interface and then discuss their relevance for natural rupture processes.

Strong bimaterial contrast

The wavefield observations for the strong bimaterial contrast (movie S1) reveals two types of slip events at the asperity: strongly localized micro-slips, and moving rupture fronts that propagate along the asperity and cause a global stick-slip behaviour (see fig S6). We analyze one event of each type, representative of the overall observations.

For the microslip event, depicted in Fig. 2(a), a spherical wavefront is radiated from one location on the asperity. No rupture propagation is observed and consequently we model the

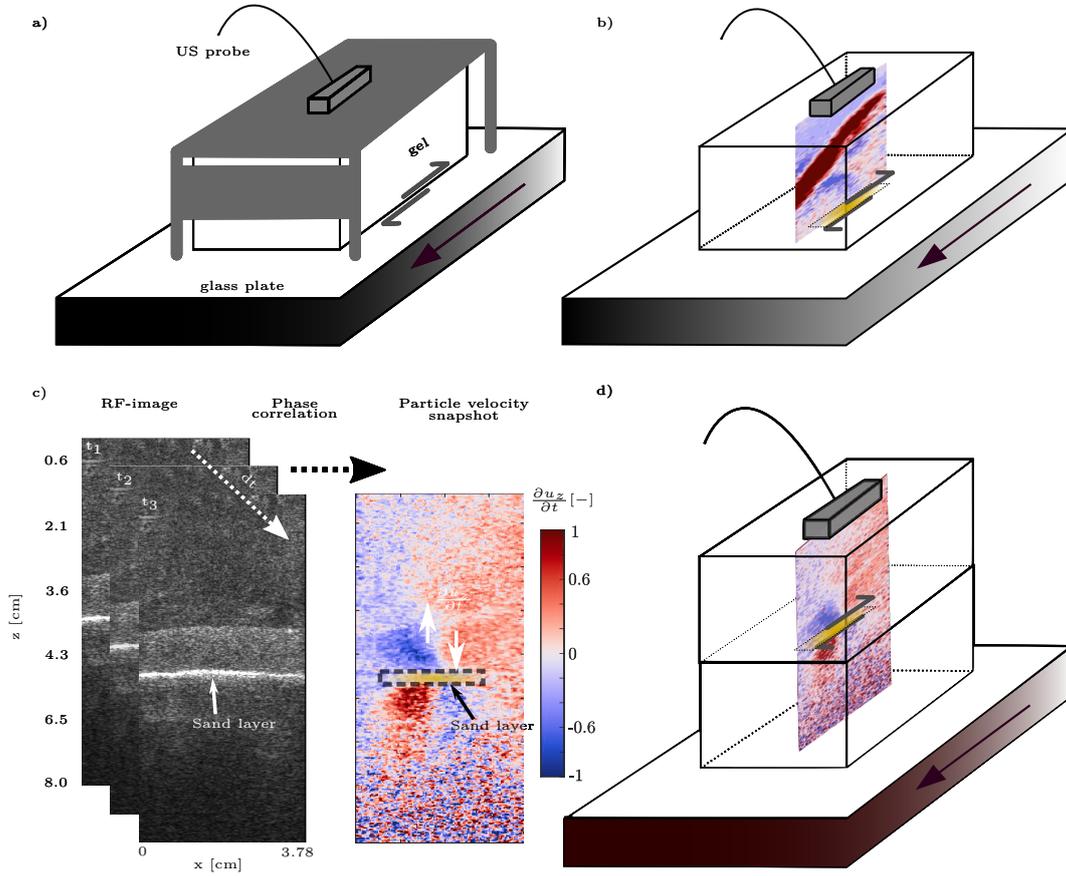


Figure 1. Experimental setup and imaging methodology. (a) Schematic view of the montage. The gel is free at the interface and blocked above. (b) Imaging methodology: Correlation of successive US reflection images results in retrieval of the vertical component of the shear wave’s particle velocity. Blue denotes upwards polarization (negative z) and red denotes downwards polarization (positive z). (b) and (d) Schematic illustration of the imaging plane in the bimaterial setups: (b) A hydrogel - sand asperity - glass interface constitutes the strong bimaterial contrast. (d) A hydrogel - sand asperity - hydrogel constitutes the weak bimaterial contrast. A detailed acquisition and processing workflow is given in Fig. S1.

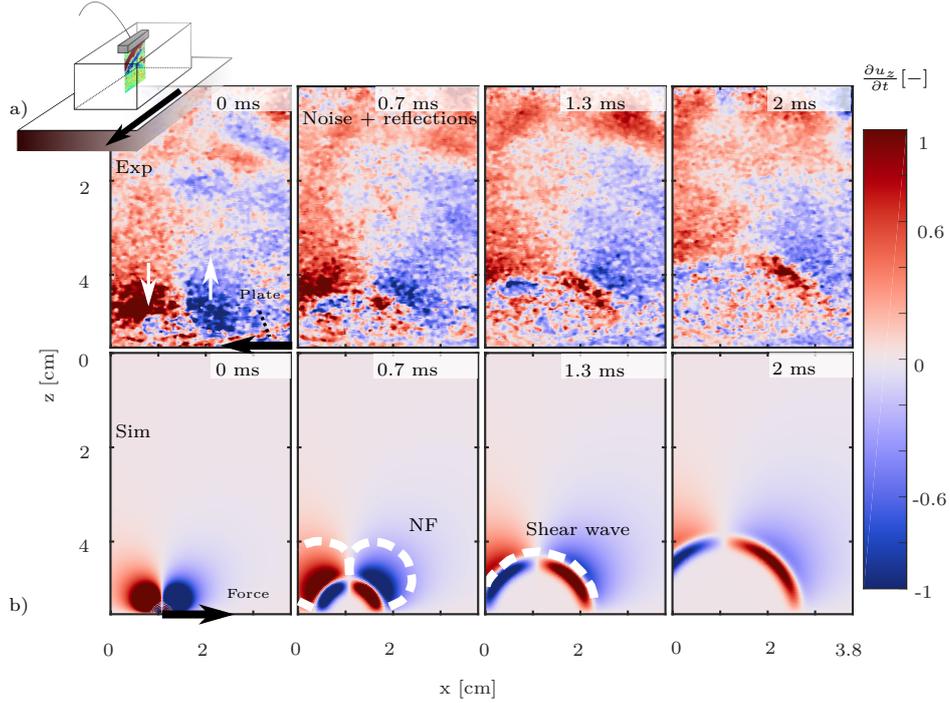


Figure 2. Strong material contrast: Comparison of an experimental micro-slip and simulation. Particle velocity direction is indicated by the white arrows. (a) Particle velocities $\frac{\delta u_z}{\delta t}$ observed by shear wave imaging. The stepper motor drives the plate in negative x-direction. (b) Complete Green's function for displacement u_z of a singular unidirectional shear force in positive x-direction. A median filter was applied to visually highlight the coherent wavefronts. All images are scaled by their extreme values. The near field lobe (NF) and shear wave front are indicated in panels (b).

142 event with a local point source. The experimental radiation pattern is well reproduced by
 143 a unidirectional single force model, as shown in Fig. 2(b) using a ramp shaped source time
 144 function $X_0(t)$. The first lobe represents the near field (NF) lobe of the Green's function
 145 and is quickly attenuated. The second lobe is of opposite polarity and represents the far-
 146 field shear wave. The top 2 cm are artifacts of a previous event (see Fig. S2 event 2-3).
 147 In contrast to the simulation, the experiment undergoes constant charging from the motor.
 148 Thus, noise as well as aseismic displacement due to deformation are present in the snapshots.

149 The gaussian source time function of the force employed by Latour et al. (2011) to model
 150 depinning events of hydrogel on sand paper fails to reproduce the here observed wavefield.
 151 Our ramp shaped source time function with rise time of 0.1ms (see Fig. S10) results in a
 152 better match. The plate displacement deforms the gel and a likely physical explanation is
 153 a localized change from a high- to low-stress state, which we model by a ramp function in
 154 time of a rightward point force (see Fig. S10). Dynamically, this is equivalent a left-
 155 pointing loading force applied to the gel, which drops to zero value, corresponding to a shear
 156 friction drop localized on a micro-asperity. In the granular layer it might correspond to a highly
 157 localized inelastic dislocation or grain micro-slip.

158 The localized event of Fig. 2 precedes a larger event, in which a rupture front traverses
 159 the entire visible interface (see Fig. 3 (a) and Figs. S3-S4 for details). The rupture propa-
 160 gation direction equals the sliding direction of the gel, *i.e.* opposite to the plate movement.
 161 This observation agrees with Dedontney et al. (2011), who found that for bimaterial in-
 162 terfaces, ruptures will preferentially propagate in slip direction of the compliant side. The

163 particle velocity ($\frac{\delta u_z}{\delta t}$) measurements in Fig. 3 (a) are compared to two analytic, kinematic
 164 simulations: a moving unidirectional force (UF) Fig. 3(b), and a moving double-couple (DC)
 165 Fig. 3(c). The simulations result from superposition of point sources along a decelerating
 166 speed profile, which is estimated roughly from the experimental data. Through trial and
 167 error we qualitatively match the near field lobe, supershear- and rupture arrest front. The
 168 source parameters are given in Figs. S10-S14.

169 Key properties of the unidirectional force model, which are also present in the experi-
 170 mental observation, are indicated in Fig. 3 (b). The first phase is an upwards polarized non-
 171 planar lobe with a diffuse front. It corresponds to the near-field (NF) of the right-traveling
 172 and rightwards pointing shear force. A sharp, downwards polarized large amplitude wave
 173 front follows, which is identified as a supershear front in the simulation. It is the result of
 174 a rupture that breaks the asperity faster than the medium's shear wave speed. The front
 175 angle with the x-axis ($\beta=21.8^\circ$) at late observation times in Fig. 3 (a) and the measured
 176 shear wave speed (c_s) of $6.9 \text{ m s}^{-1} \pm 1 \text{ m s}^{-1}$ (see Fig. S15) are used to calculate an average
 177 rupture propagation speed (c_r) of $\approx 18 \text{ m s}^{-1}$: $c_r = \frac{c_s}{\sin(\beta)}$. However, two front angles can be
 178 identified throughout the rupture (see Fig. S17). Furthermore, a time of flight measurement
 179 of the supershear front along the rupture surface (see Fig. S16) suggests a rupture speed
 180 above time resolution on 1 cm and below 12 m s^{-1} afterwards, indicating that the rupture is
 181 decelerating. This justifies the use of a decreasing rupture velocity in the kinematic model.
 182 A low amplitude, downwards polarized wedge is present above the supershear front. It cor-
 183 responds to the imprint of the compressional (P) wave, which propagates at $\approx 1500 \text{ m s}^{-1}$.
 184 Finally, a leftwards propagating and upwards polarized wavefront can be observed in the
 185 last snapshots of Fig. 3 (a). In the simulations it is identified as the rupture arrest front
 186 (RAF), emitted at the asperity border.

187 In comparison, the best moving DC solution(Fig. 3 (c)) exhibits a high wavefield com-
 188 plexity which is absent in the experiment and in the UF force simulation. Furthermore,
 189 the experimental data lack the leading, downwards polarized polarity of the DC simulation.
 190 However, at late times (7 ms), we can observe an upwards polarized front following the su-
 191 pershear front, which has a counterpart in the DC solution (Fig. 3 (c)), but is absent in the
 192 UF simulation Fig. 3 (b). To conclude, we find that the moving unidirectional force better
 193 matches the near field, the supershear front and the rupture arrest front of the experimental
 194 data than the double couple model, but does not capture every detail of the wavefield.

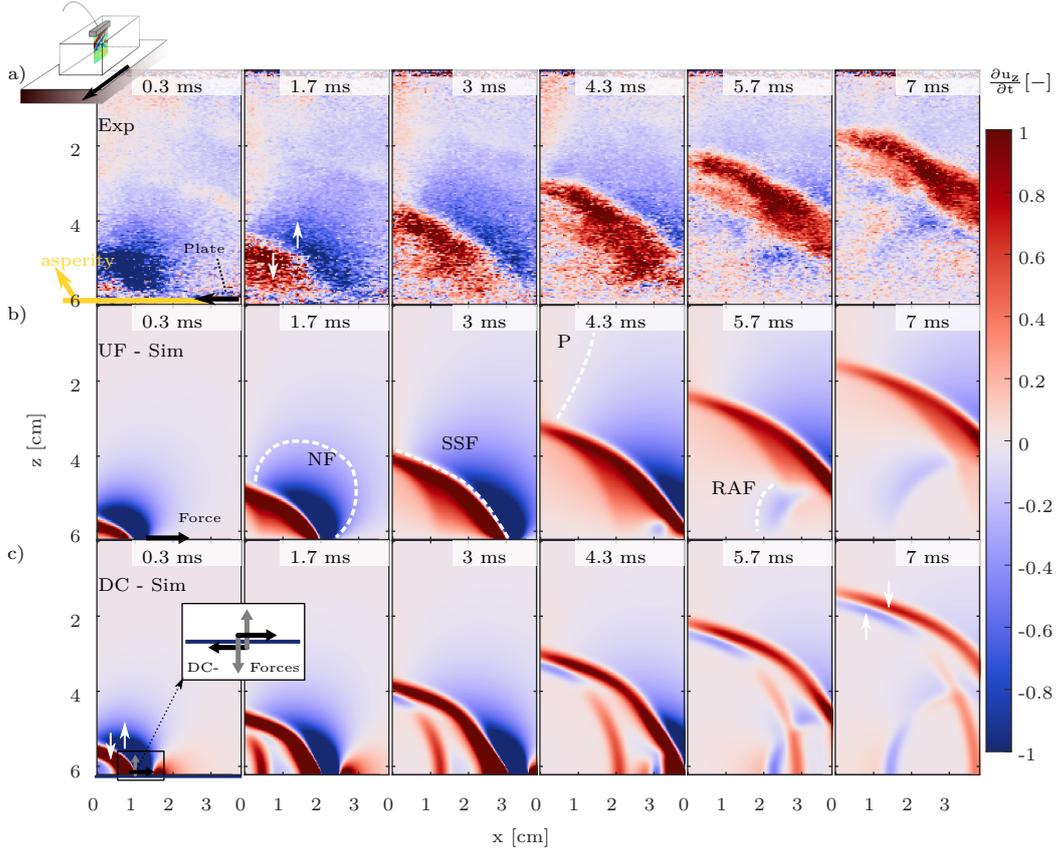


Figure 3. Strong material contrast: Comparison of an experimental supershear rupture and simulation. a) $\frac{\delta u_z}{\delta t}$ as observed by shear wave imaging. The rupture follows the event of Fig. 2. The first snapshot is located 6 ms after the first snapshot of Fig. 2. The motor drives the plate in negative x-direction. b) $\frac{\delta u_z}{\delta t}$ resulting from the superposition of unidirectional shear forces in x-direction. Near field (NF), supershear front (SSF), P-wave imprint (P) and rupture arrest front (RAF) are indicated. c) $\frac{\delta u_z}{\delta t}$ resulting from the superposition of double-couple point sources. The point sources in b) and c) are shifted in time and space, in order to simulate the horizontal advancement of a rupture front (see Supplementary material Section 2.8). All snapshots are normalized with respect to their time-series. The sources are directed in positive x-direction. A higher time-resolution is given in Figs. S4-S5. The source functions and rupture speed profiles can be found in Figs. S11-S12.

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Weak bimaterial contrast

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We observe again two types of slip events on the weak bimaterial contrast interface: propagating ruptures and localized wave radiations (see movie S2).

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A rupture that appears to propagate below shear and Rayleigh wave speed is shown in Fig. 4 (a). Rupture propagation at sub-Rayleigh speed is expected for homogeneous systems (Shlomaï & Fineberg, 2016), but has not been observed by shear wave imaging prior to this observation, which is the first dynamic US observation of a gel-gel rupture (Latour et al., 2011). Fig. 4 (c)-(d) show the corresponding 1D waveforms at specified depth- and time-steps in the upper halfspace. In both displays, the right-traveling front exhibits higher amplitudes than the left-traveling one. This front also exposes a smaller angle to the vertical (inclination difference), indicating a speed difference between the fronts. A straightforward

206 explanation is a right-travelling sub-Rayleigh rupture. A wavefront of continuous polarity
207 throughout both half-spaces exists in the rupture propagation direction.

208 We model the radiation with a double-couple moving to the right at constant sub-
209 Rayleigh velocity (Fig. S13). The simulated wavefield (Fig. 4 (b)-(e)-(f)) reproduces the
210 continuous polarity across the interface. In contrast, the radiation pattern of a unidirectional
211 force exhibits alternating polarities in the two halfspaces (Fig. S9). However, similar to the
212 case of a strong bimaterial contrast, the leading near-field lobe predicted by the double-
213 couple solution is not identified in the experimental data. Fig. 4 a) reveals that a weak
214 upwards polarized zone is present at interface depth, but quickly disappears with depth. This
215 could be an imprint of the near field which gets masked by the continuous deformation of the
216 gel (see movie S2, (Figs. 4 and 5 start at approximately 2396 ms)). Note that the amplitude
217 increase in the rupture direction is reproduced but more pronounced in the simulation than
218 in the experiment. The experiment suffers from shear wave attenuation which is neglected
219 in the kinematic simulation and might mask the amplitude difference between the front in
220 rupture direction and the radiation front in opposite direction. Furthermore, the laboratory
221 rupture might be shorter than the qualitatively simulated rupture of Fig. 4 (b).

222 Situated three milliseconds after Fig. 4, Fig. 5 (a) shows a localized event with a
223 quadripolar radiation pattern (see Fig. S8 for a comprehensive time-series). The radia-
224 tion is qualitatively reproduced as the near-field lobe of a unidirectional point force model,
225 which is shown in Fig. 5(b). The source rise time is several ms long (see Fig. S13). Contrary
226 to the localized event on the strong bimaterial contrast interface, the far-field part of the
227 theoretical force radiation is not observed. Instead, the event is followed by a left-going
228 rupture, shown in Fig. 5 (c) (event 3 in Fig. S8). A similar sequence can be observed at
229 2350 ms in movie S2. One hypothesis is that the long rise-time localized event corresponds
230 to the nucleation process of the subsequent rupture. There appears to be an aseismic lateral
231 displacement of the radiation pattern in the lower half space for the experimental data of
232 unclear origin.

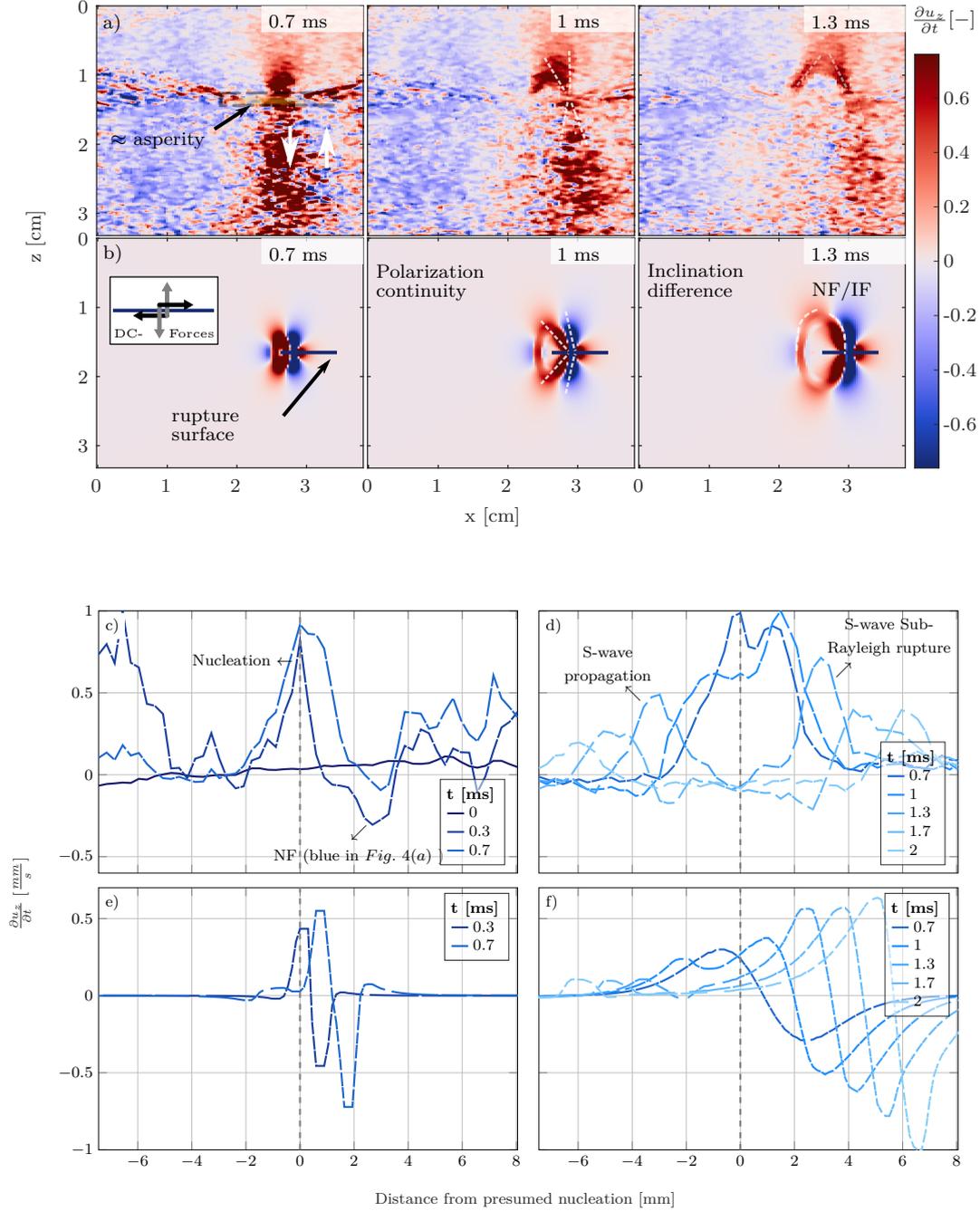


Figure 4. Weak material contrast: Comparison of an experimental sub-Rayleigh rupture and simulation. (a) Experimental particle velocities for a gel-gel rupture. (b) Right-traveling superposition of DC point sources at sub-Rayleigh speed with a ramp source function. The DC-force directions are indicated. The leading near and intermediate fields are indicated as NF/IF. (c)-(f) Spatial Waveforms (x-direction) at fixed depth and time plotted against the distance to the presumed rupture nucleation point. (c) Experimental waveforms during rupture initiation (0 ms - 0.7 ms) at the gel-gel interface. The waveforms are a mean of 27 depthpoints (≈ 1 mm), just above the sand layer, which was identified from the US reflection images. The relative position of the sand layer to the probe varies about 1.5 mm due to gel deformation and sand thickness. (d) Experimental waveforms of 0.7 to 2 ms ≈ 2 mm above the waveforms in (c). (e) - (f) Simulated waveforms corresponding to (c) and (d). (e) is taken 0.2 mm and (f) 3.8 mm above the simulated interface.

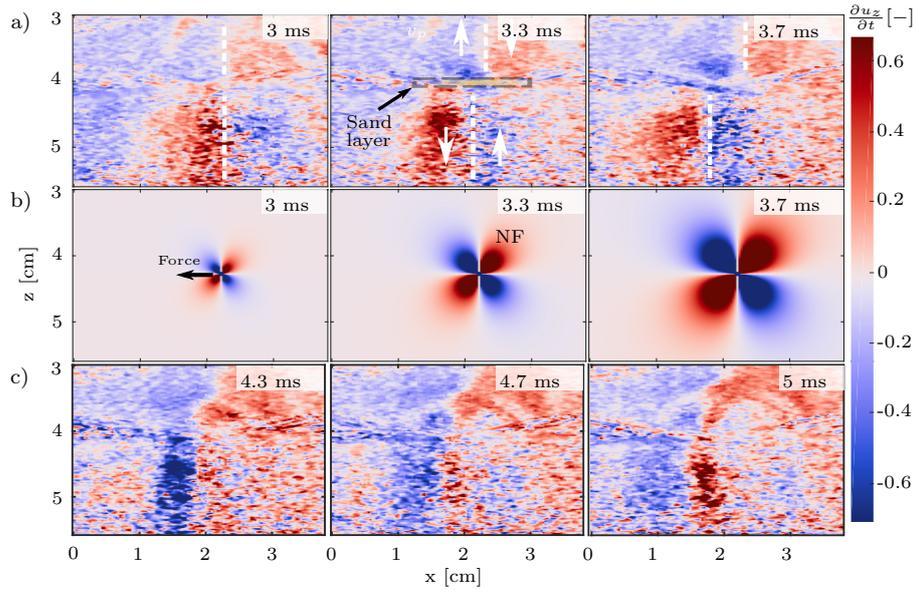


Figure 5. Weak bimaterial contrast: Comparison of an a local event and a UF-simulation. a) The interface is identified by the separation of the upper and lower lobes. Note that imaging quality is deteriorated by US diffraction at the sand, visible in the diagonal artifact in $t = 3\text{ms}$ and the coarse appearance of the displacement field below the frictional interface. The event happens 3 ms after the rupture shown in Fig. 4. b) Green's function simulation of a localized unidirectional shear force in negative x-direction using a 2.16 ms long rise time for the Gaussian. The near field lobe is indicated as NF. c) Consequent time evolution of (a). The local event is followed or transforms into a rupture (see movie S2 $\approx 2400\text{ms}$).

233 **Relevance for natural rupture processes**

234 For the strong bimaterial contrast, we find that microslip events as well as propagating
 235 ruptures radiations are better described by a unique force model than by a double couple
 236 model. This is intuitively understood as due to the strong elastic contrast at the interface:
 237 the unidirectional force corresponds to the relaxation of the gel's loading force when friction
 238 drops at the interface.

239 In nature, strong material contrasts are encountered in environmental seismology, i.e.
 240 for landslides and glacier stick-slip. Both processes exhibit a wet granular layer and a com-
 241 pliant mass sliding on a hard bedrock. Our granular asperity is conceptually comparable
 242 to the "sticky spot" encountered in alpine glacial stick-slip (Umlauf et al., 2021). Unidi-
 243 rectional force source models have been proposed for the 1980 Mt. St. Helens eruption
 244 (Kanamori & Given, 1982) and the 1975 Kalapana, Hawaii, earthquake, where a large land-
 245 slide occurred on Kilauea volcano (Eissler & Kanamori, 1987). In a theoretical analysis
 246 Dahlen (1993) showed that a lower shear wave velocity in the brecciated sliding block of
 247 shallow landslides results in mechanical decoupling of the two fault sides. The decoupling
 248 leads to a single-force rupture source, with the force pointing in the direction of the mass
 249 movement for decelerating sliding (Julian et al., 1998). Ekström et al. (2003) found that for
 250 glacier stick-slip in Greenland, single force inversions perform better than standard moment
 251 tensor inversions. Again, this could be explained by the lower shear wave speed in ice.
 252 Lastly, Trottet et al. (2022) very recently showed rupture propagation at supershear speed
 253 for snow avalanches, another case exposing low shear wave speeds of the sliding mass ($<$
 254 120 m s^{-1}). We confirm through direct experimental observation of the wavefield generation
 255 that unique force mechanisms are relevant for describing slip events between two materials
 256 with strong wave velocity contrasts.

257 In global seismology, the earthquake source corresponds to slip on a planar fault and is
 258 widely modeled by a double couple equivalent body source. Our closest analogue experiment
 259 is the case of the propagating rupture on the asperity at the gel-gel interface. We observe
 260 radiations best described by a moving double couple, which indicates a symmetry in the
 261 strain relaxation process and a coupling between both sides of the fault. However, some
 262 ruptures are preceded by localized events which can be described by the near-field radiation
 263 of a slowly rising unidirectional force, even though the materials are almost symmetric. We
 264 hypothesize that one gel is more deformed than the other during loading. It then begins to
 265 relax slowly as a preparatory process before rupture propagation initiation and both gels
 266 relax the remaining deformation. This non-symmetric process may be possible thanks to
 267 the presence of the sand layer than can locally decouple both sides of the fault through grain
 268 rearrangements. The single force source mechanism may be relevant for slow processes on
 269 natural faults. Shallow thrust faults for example expose an asymmetry in the fault loading,
 270 and fault gouge, damaged layers and fluids can constitute a decoupling mechanism.

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277 **Open Research section**

278 Data archiving is currently underway and will be archived at: <https://zenodo.org>.
 279 Temporary access has been granted via ETH polybox:
 280 <https://polybox.ethz.ch/index.php/s/0r8J638Hs8v1Da0>

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