

Ground-based Remote Sensing of Total Columnar CO₂, CH₄, and CO using EM27/SUN FTIR spectrometer at a suburban location in India and validation of Sentinel-5p/TROPOMI

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Abstract—Greenhouse gasses (GHGs) play an important role in controlling local air pollution as well as climate change. In the present study, we retrieved column-averaged dry-air (X) mole fractions of carbon dioxide (CO₂), methane (CH₄), and carbon monoxide (CO) using a ground-based EM27/SUN Fourier Transform Infrared Spectrometer (FTIR). The EM27/SUN spectrometers are widely in use in the Collaborative Carbon Column Observing Network (COCCON). The PROFFAST software provided by COCCON has been used to analyze the measured atmospheric solar absorption spectra. In this paper, the diurnal variation and the time series of daily averaged X CO₂, X CH₄, and X CO covering the period December 2020 to May 2021 are analyzed. The maximum values of X CO₂, X CH₄, and X CO are observed to be 420.57 ppm, 1.93 ppm, and 170.40 ppb respectively. Less diurnal (X CO₂ ~0.44 ppm; X CH₄=0.004 ppm, and X CO=4.84 ppb) but clear seasonal changes are observed during the study period. The X CH₄ and X CO from the Sentinel-5Precursor (S5P)/TROPOspheric Monitoring Instrument (TROPOMI) are compared against the EM27/SUN retrievals. The correlation coefficient ‘ r ’ between the EM27/SUN retrieved X CH₄ and X CO against S5P/TROPOMI are 0.75 and 0.94 respectively.

Index Terms— Greenhouse Gases, EM27/SUN, S5P/TROPOMI.

I. INTRODUCTION

Anthropogenic activities entail the release of various trace gasses into the atmosphere, mainly due to industrialization, acceleration of economy, and power generation [1]. Though there are many greenhouse gasses (GHGs), atmospheric carbon dioxide (CO₂) and methane (CH₄) are the leading contributors to anthropogenic global warming [2]. In India, the total emissions rose to 2.52 Gt CO₂ year⁻¹ in 2021 compared to the emissions (2.32 Gt CO₂ year⁻¹) in 2014. Coal fire-based power generation is a major anthropogenic activity contributing about 54 % of the electricity produced in India [3]. Due to its steady growth in the atmosphere and uncertainty of source/sink, anthropogenic CH₄ concentration has attracted the interest of the research community over the last decade [4]. Long-term measurements from the National Oceanic and Atmospheric Administration (NOAA) reveal an annual CH₄ increase of 8 ppb year⁻¹ [5] while the present study site showed 10 ppb year⁻¹ [6]. Carbon monoxide (CO) is an ozone precursor gas that also affects climate due to its role in the formation of OH radicals in the atmosphere [7]. CO is

emitted mostly as a result of incomplete combustion in urban/industrial fossil fuel, biofuel, and biomass combustion. Because atmospheric CO is a serious pollutant, it is vital to investigate the impact of local and transported sources on air quality. Global CO₂ and CH₄ column measurements are especially useful for understanding regional sources and sinks [13], [14]. The satellite-based X CO₂ data with accuracy and precision of 1-2 ppm has the potential to improve understanding of surface fluxes [10]. At present, several GHGs dedicated missions are in orbit namely the Greenhouse Gases Observing Satellite (GOSAT-1 [9], [11], GOSAT-2 [12], the Orbiting Carbon Observatory-2 (OCO-2), and OCO-3 [13] [14] and the Copernicus Sentinel-5 Precursor (S5P)/TROPOspheric Monitoring Instrument (TROPOMI) [15]. As [15] result, combining ground and space-based remote sensing has become a potent tool for studying GHG spatial and temporal variability. The most important source of GHG reference validation data is the ground-based Fourier Transform Infrared (FTIR) from Total Carbon Column Observing Network (TCCON) [16], which has been recently complemented by the portable EM27/SUN FTIR spectrometers from COCCON (Collaborative Carbon Column Observing Network) [17]. The measured near-infrared spectra are used to calculate dry-air column-averaged trace gas abundances.

In the tropics, columnar observations of GHGs are especially important since convection is always prevalent, and as a result, flux patterns are only weakly visible in surface measurements. In the present work, solar spectra collected at NRSC, Shadnagar, a suburban site in India using an EM27/SUN spectrometer are presented and discussed. Shadnagar is a tropical station located at latitude 17.0365° N, longitude 78.1851° S, altitude 540 m a.s.l. The observations were performed from 7th December 2020 to 3rd May 2021.

This study is the first of its kind over the Indian subcontinent. Shadnagar is a tropical station, observations were performed from 7th December 2020 to 3rd May 2021. The present study also includes a comparison of the ground-based retrievals of X CH₄ and X C with the Sentinel-5P/TROPOMI X CH₄ and X CO retrievals during the study period.

II. DATA AND METHODOLOGY

EM27/SUN: The spectrometer has been developed by the Karlsruhe Institute of Technology, Germany, in collaboration with Bruker Optics as described in [17]. It covers a spectral range of 4,000-10,000 cm^{-1} (1.0 μm to 2.5 μm) with spectral resolution of 0.5 cm^{-1} (maximum optical path distance amounts to 1.8 cm). It records near-infrared (NIR) solar spectra using Indium Gallium Arsenide (InGaAs) detectors. While XCO_2 and XCH_4 are covered by the main (shortwave) detector, XCO is covered by the auxiliary (longwave) detector [18]. We have used PROFFAST [19] which is a non-least square spectral algorithm to retrieve the species concentrations. The spectral windows for CO_2 , CH_4 , and CO are 6173-6390 cm^{-1} , 5897-6145 cm^{-1} , and 4233-4290 cm^{-1} respectively.

During the study period, a total of 19494 observations were taken from 7th December 2020 to 3rd May 2021 covering 74 clear sky days from morning 09:20 Indian Standard Time (IST) to till evening 16:45 IST. Pre-processing and trace gas retrievals followed by post-processing are the two processes in the data analysis chain (fig 1). The DC correction, Fast Fourier Transformation including a phase correction scheme and a spectral resampling are performed first for generating spectra from the raw interferograms. The default modulation efficiency (0.983) of an average COCCON spectrometer was used for the trace gas analysis. In the post-processing, the retrieved column values are divided by the dry air column (derived from the co-observed oxygen column) and air mass-independent and air mass-dependent corrections are taken into account for generating the final X_{gas} values. The initial volume files as provided by TCCON (referred to as map-files), which contain the a-priori mixing ratio profiles of the relevant gases and the temperature profiles were used for the trace gas analysis. The column-averaged amount of dry air (X_{air}), which measures the instrument's stability, is found 0.9888 ± 0.008 during the study period. The daily average of X_{air} was in the range of 0.9856 to 0.9924 (Fig 1b), indicating excellent stability of the spectrometer over the observation period.

Sentinel-5p retrieved XCH_4 and XCO : Sentinel-5P/TROPOMI instrument was switched to a better spatial resolution of 5.5km (across-track) X 7 km(along-track) from 6th August 2019. For CO , the processing baseline, product versions, and quality limitations are described in the Product Readme File (PRF) [20]. We have used TROPOMI CO associated with a quality assurance value (qa_value) > 0.5 as recommended in PRF. These criteria separated the retrievals performed under clear sky and clear sky like observations ($\text{qa_value}=1$, Cloud optical thickness < 0.5 and cloud height < 500 m) and cloudy conditions ($\text{qa_value}=0.7$, Cloud optical thickness ≥ 0.5 and cloud height < 5000 m). S5P L2 used in the study provides a total vertically integrated column of CO (TC_{CO}), therefore XCO is calculated by taking the ratio of TC_{CO} and a total vertically integrated column of dry air ($\text{TC}_{\text{dry air}}$).

Since the total column of O_2 and N_2 is not available, $\text{TC}_{\text{dry air}}$ is calculated from the surface pressure (P_s) and vertically integrated water column ($\text{TC}_{\text{H}_2\text{O}}$) from the equation (1).

$$\text{XCO} = \frac{\text{TC}_{\text{CO}}}{\text{TC}_{\text{dry air}}} = \frac{\text{TC}_{\text{CO}}}{P_s / (g \times m_{\text{dry air}}) - \text{TC}_{\text{H}_2\text{O}} \times (m_{\text{H}_2\text{O}} / m_{\text{dry air}})} \quad (1)$$

Where $g=9.82 \text{ m}\cdot\text{s}^{-2}$ is the column-average gravity acceleration and $m_{\text{H}_2\text{O}}=0.01801528 \text{ kg mol}^{-1}$ is the molecular mass of H_2O , $m_{\text{dry air}} = 0.0289644 \text{ kg mol}^{-1}$ is the molecular mass of dry air [21]. The S5P pixels within the radius of 100 km around the observation site are taken for comparison. The ground-based EM27/SUN spectrometer observations are taken within an hour of the satellite overpass. In the same way, the radius of 100 km was taken for the comparison of CH_4 [21], [22]. We have used TROPOMI XCH_4 biased corrected data associated with a quality assurance value ($\text{qa_value} > 0.5$) similar to S5P CO as recommended in CH_4 PRF [23]. The S5P CH_4 is already given in the volume mixing ratio hence compared directly against EM27/SUN retrieved XCH_4 . Details of the data used in the present study are summarized in Table 1.

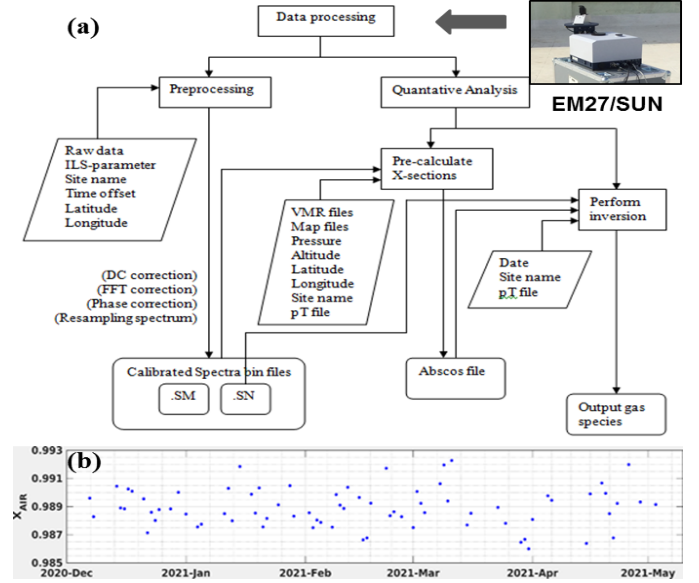


Figure 1. a) EM27/SUN Spectrometer data analysis using PROFFAST. b) X_{air} during the observational period.

Para-meters	Data sources	Resolution	DOI
XCO_2 , XCH_4 , and XCO	EM27/SUN	Point location (17.0365° N, 78.1851° E, 540 m)	NA
XCH_4	S5P L2 version 01.04.00	5.5km x 7km	[3]
XCO	S5P L2 version 01.04.00	5.5km x 7km	[25]

Table 1. Details of XCO_2 , XCH_4 , and XCO used during the study period.

III. RESULTS AND DISCUSSIONS

Diurnal variation: Figure 2 (a) depicts the diurnal variability of XCO_2 , XCH_4 , and XCO over the study period, while 2(b) depicts the deviation (daily mean-value) of XCO_2 and XCO . The variability in XCO_2 is ranged from 0.20 ppm to 1.80 ppm during diurnal time windows, showing that there is a less diurnal effect on XCO_2 . When compared to XCO_2 , however, XCO has a higher diurnal variation. This could mainly be shaped by the local environmental conditions and small-scale industries around the study site. The major sources of CO are the oxidation of CH_4 , industrial activities such as the combustion process, biomass burning, and the oxidation of non-methane hydrocarbons [26]. The major sink of CO is oxidation with hydroxyl radical (OH). Maximum and minimum XCO_2 , XCH_4 , and XCO are summarized in table 2.

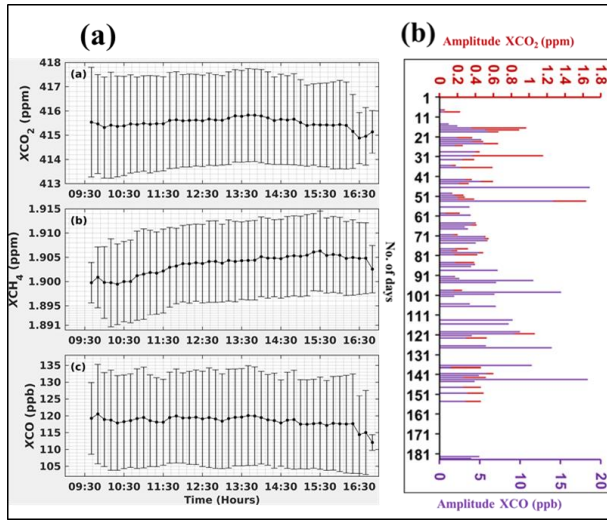


Figure 2. The diurnal variation of XCO_2 , XCH_4 , and XCO respectively during the study period and (b) the diurnal amplitude of XCO_2 and XCO .

Daily and seasonal variation: The time series analysis in Figure 3 shows clear seasonal changes in the XCO_2 , XCH_4 , and XCO (Fig 4b-d) during the study period. The concentrations are higher in the pre-monsoon season (March-April-May) than in winter (December-January-February). In winter, there was less XCO_2 than in the pre-monsoon, which could be due to reduced CO_2 assimilation due to lower temperatures and solar radiation [6]. From winter to pre-monsoon, the XCH_4 shows less seasonality, with a maximum amplitude change of 0.014 ppm. The spatial distribution of surface fluxes and atmospheric transport at the synoptic-scale are substantially responsible for the variability of XCO_2 and XCH_4 . With an amplitude variation of ~ 18 ppb, dry column-averaged CO concentration indicates very considerable seasonality across the study period. This variation could be attributed to CO emissions from fossil fuels, biomass burning, and industrial

production on a daily basis. In comparison to winter, there is a higher surge in fire activity during the pre-monsoon season, which could lead to higher atmospheric concentrations. S5P/TROPOMI XCH_4 and XCO were validated from November 2017 to September 2020 against the TCCON and Infrared Working Group of the Network for the Detection of Atmospheric Composition Change (NDACC-IRWG) observations [21].

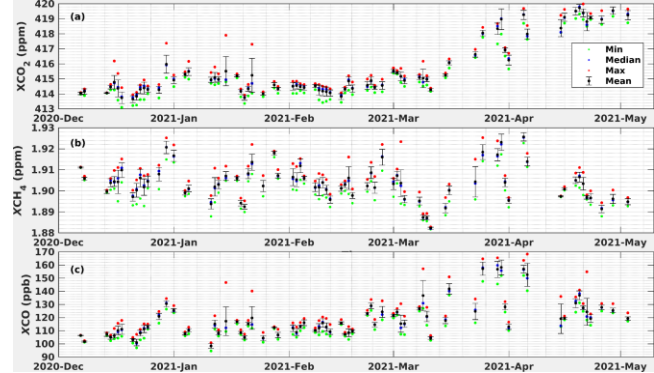


Figure 3. The daily maximum, minimum, median and average values with ± 1 standard deviation of gaseous species, a) XCO_2 , b) XCH_4 and c) XCO during the study period.

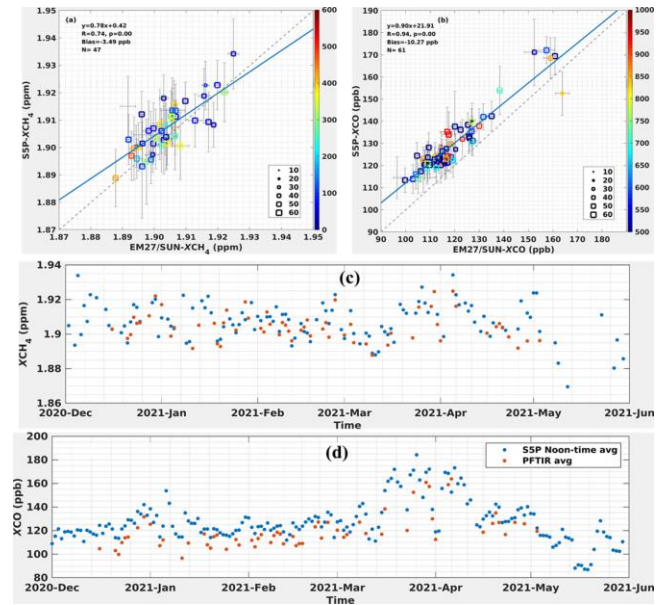


Figure 4. Figure 4. a) XCH_4 and b) XCO comparison between S5p/TROPOMI and EM27/SUN. The color bar and the size of the squares represent the number of S5P total column observations and the number of EM27/SUN ground observations at the same time. The horizontal and vertical bar is one standard deviation of S5P and EM27/SUN observations. c) Time series of XCH_4 between EM27/SUN against TROPOMI XCO d) time series of XCO same as (c).

Validation of Sentinel-5p/TROPOMI: In this work, a comparison of S5P/TROPOMI retrieved XCH_4 and XCO is performed with respect to the EM27/SUN spectrometer retrievals. Further, we compare the atmospheric column averaged dry-air mole fractions of XCH_4 and XCO retrieved from EM27/SUN sensor with S5P/TROPOMI retrievals at Shadnagar, sub-urban region of Telangana, India. During this study period, spatially and temporally collocated XCH_4 and XCO data from S5p/TROPOMI using a 100 km radius from the observational site and ± 1 hour as criteria only considered for the comparison. The number of samples for ground observation, S5P CH_4 and S5P CO range 3-60, 3-686, and 1-967. The resultant comparison is shown in figure 4 (a-d) respectively for XCH_4 and XCO .

Parameter	Mean	$\pm 1 \sigma$	Min	Max	Median
XCO_2 (ppm)	415.57	1.89	413.09	420.57	414.78
XCH_4 (ppm)	1.904	0.008	1.882	1.931	1.902
XCO (ppb)	118.74	13.85	94.62	170.40	114.95

Table 2. Statistics of XCO_2 , XCH_4 , and XCO during the study period.

The average value of XCH_4 for EM27 and S5P were found 1.903 ± 0.01 and 1.907 ± 0.01 . The bias of EM27/SUN w.r.t S5P was observed -3.487 ppb (-0.18 %) with the co-located observations. (EM27-S5P)/S5P).

It is clear that atmospheric XCH_4 and XCO showed a good correlation with correlation coefficients (r) of 0.75 and 0.94 respectively. During the comparison period, TROPOMI retrieved XCH_4 and XCO are following ground-based EM27/SUN spectrometer retrievals (Figure 4c-d). Thus, we report a good agreement between retrievals of atmospheric column-averaged dry-air mole fractions of XCH_4 and XCO from EM27/SUN and S5P/TROPOMI retrievals.

The average of XCO for EM27/SUN and S5P/TROPOMI CO was 119.46 ± 14.57 and 129.72 ± 14.01 with the co-located while the bias was observed -10.26 (-7.91%). The bias and 'r' values are computed under the varied sampling radius as summarized in table 3. The EM27/SUN XCH_4 and XCO exhibit a negative bias of 0.001 ppm and 10.22 ppb with S5P/TROPOMI retrievals respectively.

Parameter	Radius (km)	N	R	Bias (ppb)	EM27 ($\mu \pm 1\sigma$)	S5P ($\mu \pm 1\sigma$)
CH_4	50	45	0.79	-2.58	1.903 ± 0.01	1.905 ± 0.01
CH_4	20	36	0.79	-3.10	1.902 ± 0.01	1.905 ± 0.01
CO	50	61	0.94	-10.22	119.46 ± 14.57	129.68 ± 13.77
CO	20	56	0.94	-9.00	120.12 ± 14.72	129.13 ± 13.88

Table 3. Observed Validation of S5p considering different spatial observations area. N and R show the number of co-located days and correlation coefficient.

IV. CONCLUSION

In this study, we demonstrated first-hand columnar XCO_2 , XCH_4 , and XCO using the EM27/SUN spectrometer data from 7th December 2020 to 3rd May 2021. During this study period, the dry air column was observed to be in the range of (0.9856 -0.9924) indicating the stability of the EM27/SUN spectrometer. The XCO_2 , XCH_4 , and XCO exhibited mild diurnal variation. The retrieved columnar concentrations also observed clear seasonality from winter to pre-monsoon. Further, the present study evaluated Sentinel 5P/TROPOMI retrieved XCH_4 and XCO against EM27/SUN spectrometer retrievals and found very good agreement with 'r' values of 0.74 and 0.94 respectively. The biases between EM27/SUN and S5P/TROPOMI retrieved XCH_4 and XCO are -3.4872 ppb (-0.1829 %) and -10.2670 (-7.9141%) respectively, meeting mission requirements of bias \pm precision of less than $1.5\% \pm 1\%$ for CH_4 and $15\% \pm 10\%$ for CO . Thus, the present study demonstrated the retrieval of XCO_2 , XCH_4 , and XCO while meeting the COCCON observations along with satellite validation.

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