

1 **Tile drainage causes flashy streamflow response in Ohio watersheds**

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7

8 **Abstract**

9 Artificial subsurface (tile) drainage is used to increase trafficability and crop yield in much of the
10 Midwest due to soils with naturally poor drainage. Tile drainage has been researched extensively
11 at the field scale, but knowledge gaps remain on how tile drainage influences the streamflow
12 response at the watershed scale. The purpose of this study is to analyze the effect of tile drainage
13 on the streamflow response for 59 Ohio watersheds with varying percentages of tile drainage and
14 explore patterns between the Western Lake Erie Bloom Severity Index to streamflow response in
15 heavily tile-drained watersheds. Daily streamflow was downloaded from 2010-2019 and used to
16 calculate mean annual peak daily runoff, mean annual runoff ratio, the percent of observations
17 in which daily runoff exceeded mean annual runoff ($T_{Q_{mean}}$), baseflow versus stormflow
18 percentages, and the streamflow recession constant. Heavily-drained watersheds (> 40 % of
19 watershed area) consistently reported flashier streamflow behavior compared to watersheds with
20 low percentages of tile drainage (< 15% of watershed area) as indicated by significantly lower
21 baseflow percentages, $T_{Q_{mean}}$, and streamflow recession constants. The mean baseflow percent for
22 watersheds with high percentages of tile drainage was 20.9 % compared to 40.3 % for
23 watersheds with low percentages of tile drainage. These results are in contrast to similar research
24 regionally indicating greater baseflow proportions and less flashy hydrographs (higher $T_{Q_{mean}}$) for
25 heavily-drained watersheds. Stormflow runoff metrics in heavily-drained watersheds were
26 significantly positively correlated to western Lake Erie algal bloom severity. Given the recent
27 trend in more frequent large rain events and warmer temperatures in the Midwest, increased
28 harmful algal bloom severity will continue to be an ecological and economic problem for the

29 region if management efforts are not addressed at the source. Management practices that reduce
30 the streamflow response time to storm events, such as buffer strips, wetland restoration, or
31 drainage water management, are likely to improve the aquatic health conditions of downstream
32 communities by limiting the transport of nutrients following storm events.

33 **Keywords:** tile drainage, agriculture, baseflow, recession analysis, intensively managed
34 landscapes

35 1. INTRODUCTION

36 Artificial subsurface (tile) drainage is required for increased crop yield in much of the cropland
37 in the Midwestern U.S. ('Midwest') due to soils with naturally poor drainage capabilities. Tile
38 drains increase soil drainage by removing excess subsurface water that can inhibit plant growth,
39 resulting in lower water tables that increase the trafficability of heavy machinery to operate in
40 farm fields. Tile drains began to be installed in the Midwest during the late 20th century with the
41 initial goal of strategically draining wet areas of farm fields that were susceptible to ponding, but
42 installations are now common throughout the entire field to lower the water table (Blann *et al.*,
43 2009). Drainage pipes are typically installed between 0.6 – 1.2 m below the surface
44 approximately 10-30 m apart, depending on site-specific soils, crop type, and cost (Skaggs and
45 van Schilfhaarde, 1999). Infiltrated water is captured underground by perforated drainage pipes
46 and routed away from the field into adjacent ditches and streams.

47 According to the U.S. Department of Agriculture (USDA) National Agriculture Statistics
48 Service (NASS) 2017 Census of Agriculture, 225,024 km² of cropland are estimated to have tile
49 drainage with the vast majority occurring in the Midwest (USDA NASS, 2019). The amount of
50 land with tile drainage in the U.S. increased by 28,484 km² (14.5%) between 2012 and 2017
51 (USDA NASS, 2014), with the largest increases occurring in the Midwest. Recent changes to
52 precipitation patterns that generate more frequent large rain events in the Midwest (Williams and
53 King, 2020) may partly explain the growing adoption of tile drainage. Further, as heavy rainfall
54 events are projected to increase in frequency into the future due to climatic change, we can
55 expect an expansion of land under tile drainage globally (Gordon *et al.*, 2017). Understanding
56 how tile drainage impacts the hydrologic response of downstream waterways, and subsequent
57 transport of nutrients, is critical for the development holistic management plans that improve
58 downstream aquatic life and help communities assess flood risks.

59 However, the streamflow response, and subsequent export of nutrients, from farm fields
60 under tile drainage is complicated to ascertain and predict due to compounding environmental,
61 management, and site-specific soil conditions (Hanrahan *et al.*, 2020). For example, Boland-
62 Brien *et al.* (2014) reviewed several field studies (< 10 ha) and suggested that tile drainage can
63 cause peak streamflow to decrease when water tables are close to the surface due to clayey soils
64 with low permeability or during high rainfall events. In contrast, peak flows may increase on
65 fields with deeper water tables with drier climates or more permeable soils (Boland-Brien *et al.*,
66 2014). Such changes in peak flows across large scales could have impacts on timing and
67 magnitudes of flood peaks for downstream communities. In addition, management practices that
68 target particular flow pathways (e.g. reducing surface runoff or reducing tile outlet discharge)
69 could have adverse effects on other nutrient transport mechanisms, and thus have unintended
70 impacts to nutrient loads not initially targeted (Smith *et al.*, 2015). Studies have consistently
71 showed that water exiting tile drains contribute significant amounts of nutrients (e.g. nitrogen
72 and phosphorus) to downstream waterbodies. In Illinois, riverine nitrate flux from tile-drained
73 land was over twice the value compared to non-tile drained land despite higher net nitrogen
74 inputs on non-tile drained land (McIsaac and Hu, 2004). Tile drainage exported 80% of stream
75 nitrogen load, despite only contributing 15-43% of the streamflow in a 122 km² watershed in
76 northeast Iowa (Arenas Amado *et al.*, 2017). In a headwater watershed in Ohio (<4 km²), tile
77 drainage accounted for 47% of total discharge, 48% of dissolved phosphorus, and 40% of total
78 phosphorus (King *et al.*, 2015).

79 Tile drains have been shown to reduce mean groundwater travel times, which is
80 problematic for example when considering the transport of nitrogen which tends to have higher
81 concentrations in groundwater compared to surface runoff (Schilling *et al.*, 2012). A modeling
82 study on a 74.3 km² watershed in north-central Iowa revealed that mean groundwater travel times
83 are more than 150 times faster than those that existed prior to settlement, resulting in the
84 majority of groundwater (>98%) bypassing perennial riparian buffers (Schilling *et al.*, 2015),
85 which drastically reduces the effectiveness of installing stream buffers to reduce nitrogen
86 concentrations (Schilling *et al.*, 2015). A study in western Indiana compared the residence time
87 of baseflow in agricultural and adjacent undisturbed forested watersheds using multiple isotopic
88 tracers (specifically CFC, SF₆, ³⁶Cl, and ³H) and suggested that baseflow in the agricultural
89 watershed with tile drainage was controlled by a large contribution of tile drainage and/or soil

90 water with short residence times (Frisbee *et al.*, 2017). In contrast, Frisbee *et al.* (2017)
91 concluded that baseflow in the adjacent, undisturbed forested catchments was supported by
92 groundwater with much older residence times (at least 40 years old). Baseflow comprised of
93 large contributions from tile drainage is problematic for the aquatic health of waterways due to
94 the often high concentrations of nutrients measured in tile drainage.

95 Baseflow proportions can be used to assess hydrologic impacts of land use and
96 conservation practices and have been found to be strongly correlated to legacy nutrient
97 concentrations; thus, baseflow estimations provide a first approximation of stream vulnerability
98 to legacy nutrients (Tesoriero *et al.*, 2013). Tile drainage was demonstrated to increase the
99 proportion of baseflow to receiving streams in Iowa (Schilling and Libra, 2003; Schilling and
100 Helmers, 2008; Boland-Brien *et al.*, 2014), but a gap remains understanding the relationship
101 between tile drainage and baseflow in other regions, particularly in regions with different soil
102 and precipitation characteristics, such as Ohio. Baseflow proportions are generally thought to
103 increase in larger or flatter watersheds as groundwater tends to be the main contributor to
104 streamflow. According to Boland-Brien *et al.* (2014), while watersheds in Iowa with large
105 proportions of tile drainage tended to have larger baseflow proportions compared to non-tiled
106 watersheds, the variability of baseflow percentage with watershed size was much lower for
107 watersheds with large proportions of tile drainage compared to non-tiled watersheds which
108 exhibited an increase in baseflow proportion with watershed size. Boland-Brien *et al.* (2014)
109 found that tile drainage had a similar homogenizing effect on all flow regimes, where heavily
110 tile-drained watersheds showed little to no variability in streamflow response across a range of
111 drainage areas compared to watersheds with a smaller proportion of tile drainage that exhibited
112 larger variability in streamflow response when considering various streamflow metrics across a
113 range of watershed sizes, which is expected for natural systems.

114 In addition to baseflow assessments, hydrograph recession analysis has proven to be a
115 helpful mathematical exercise that estimates the potential change in the storage-discharge
116 relationship for a particular watershed. Recession analysis can be used to evaluate storm
117 responses and thus infer storage properties and mean residence times (e.g. Troch *et al.*, 2013).
118 For example, Schilling and Helmers (2008) found the master recession curves for tile-drained
119 watersheds in Iowa to be more linear compared to less-tiled watersheds that showed a non-linear

120 recession, typical of natural systems where hydraulic conductivity decreases with depth. They
121 suggested that downstream hydrograph recession may be controlled by longer recession times
122 from tiled regions, but also found inconsistent recession coefficients between tiled and non-tiled
123 regions and advocated for additional research in this field. Boland-Brien et al. (2014) also
124 performed streamflow recession analysis on watersheds with varying percentages of tile drainage
125 across Iowa and concluded that tiled regions were less flashy compared to non-tiled regions
126 based on master recession curve analysis.

127 Clearly, tile drainage can have confounding impacts on hydrological response depending
128 on scale and the combination of physical and climatic characteristics considered. Given an
129 emphasis in the literature on tile drainage impacts to streamflow response in Iowa (e.g., Schilling
130 and Helmers, 2008; Boland-Brien et al., 2014; Schilling et al., 2015; Arenas Amado et al., 2017),
131 we wondered how tile drainage impacts hydrological response under other landscapes and
132 climatic conditions? As such, the goal of this study is to assess the impact of tile drainage on the
133 streamflow response of Ohio watersheds with varying percentages of tile drainage. The shallow,
134 poorly-drained soils of Ohio provide an excellent contrast to those in Iowa, which tend to be
135 deeper and coarser, thus have different drainage tendencies. We used an automated baseflow
136 separation technique combined with hydrograph recession analysis to determine if the effects of
137 tile drainage on the storage-discharge relationship are evident at the watershed scale and
138 postulate the consequences for downstream nutrient transport. To this latter aspect, phosphorous
139 loads from March to July have recently been identified as a major driver of the severity of HABs
140 in the western Lake Erie basin (Baker *et al.*, 2019), which is where the majority of tile drainage
141 occurs in Ohio. Therefore, we focused on this critical time period in order to isolate the effects of
142 tile drainage from heavily-drained watersheds in the western Lake Erie basin on hydrograph
143 partitioning that could be exacerbating HAB severity by creating a quicker hydrologic
144 connection between agricultural fields and adjacent streams.

145

146 **2. MATERIALS AND METHODS**

147 **2.1 Data**

148 Daily mean streamflow for each study watershed was downloaded from 2010 – 2019 for 59
149 United States Geologic Survey (USGS) stream gaging stations in Ohio using the R package
150 ‘dataRetrieval’ (De Cicco and Hirsch, 2014). The station ID for each stream gage is included as
151 supplementary material. Streamflow was converted to area-weighted runoff (‘runoff’) using the
152 total watershed area and daily time interval. The time period of data considered was selected to
153 match with the responses from the recent county-level tile drainage census data used to generate
154 AgTile-US (Valayamkunnath *et al.*, 2020). Monthly PRISM precipitation data from the same
155 period (i.e. 2010 – 2019) was aggregated to watershed boundaries to determine mean monthly
156 and annual precipitation for each study watershed (PRISM Climate Group, 2019).

157 The 59 study watersheds were selected based on streamflow record and limited
158 hydrological modifications using the following criterion: (1) had at least eight years of complete
159 data from 2010 - 2019, with each year having at least 90% daily streamflow records available,
160 (2) had less than 6 major dams, (3) were located at least 5 miles downstream of dams, (4) had
161 less than 25% developed land, (5) had at least 25% agricultural land, and (6) had area less than
162 2,000 km² (Falcone, 2011). The watershed size limitation was suggested as a threshold in which
163 the effects of tile drainage were likely to become less apparent due to channelization and in-
164 stream attenuation (Boland-Brien *et al.*, 2014). These 59 watersheds were split into three roughly
165 equal groups with increasing proportions of tile drainage to evaluate the mean streamflow
166 response for watersheds with low (< 15 % area), medium (15% - 40% area), and high (> 40%
167 area) amounts of tile drainage.

168 Watershed characteristics and boundaries were obtained from the GAGES-II dataset
169 (Falcone, 2011). For each of the 59 watersheds in Ohio, a 30-m resolution tile drainage map
170 (AgTile-US) was aggregated to calculate the percent of each watershed under tile drainage
171 (Valayamkunnath *et al.*, 2020). This dataset was generated using soil drainage information,
172 topographic slope, and county-level tile drainage census data for the most-likely tile-drained area
173 of the contiguous United States. Accuracy across the Midwest ranges from 82.7% to 93.6%
174 (Valayamkunnath *et al.*, 2020). The raster dataset is available in binary format, where 1 indicates
175 tile-drained land and 0 indicates undrained land. For each watershed, the percent of tile drainage
176 was calculated by summing the total amount of tile-drained area divided by the total watershed
177 area.

178

179 **2.2 Runoff metrics**

180 To evaluate the effects of tile drainage on streamflow response, we calculated several of the
181 runoff metrics suggested by Boland-Brien et al. (2014) including runoff ratio, mean annual peak
182 runoff, and the percent of time daily runoff exceeded mean annual runoff ($T_{Q_{mean}}$). Runoff ratios
183 were calculated by dividing annual runoff by annual precipitation from PRISM and multiplying
184 the result by 100 to have ratios expressed as a percent. To evaluate the impact of tile drainage on
185 peak runoff conditions, we calculated mean annual peak daily runoff for each watershed
186 considered. The final metric considered the percent of time daily runoff exceeded mean annual
187 runoff, $T_{Q_{mean}}$, which measures the flashiness of the hydrograph (Konrad and Booth, 2002). As
188 such, a low value corresponds to a flashier response and a high value suggests a more dampened
189 hydrograph. Differences in runoff metrics were compared among the three drainage categories
190 using the Tukey test and Pearson's correlation coefficient. All significant results are considered
191 when $p < 0.05$.

192 Daily baseflow was calculated from the total daily runoff hydrograph using the R
193 package 'lfstat' (Koffler *et al.*, 2016) following methodology from Tallaksen and van Lanen
194 (2004) and WMO (2008). This procedure was developed for rainfall regimes with a typical
195 runoff response in hours or days and partitions the hydrograph into delayed and quick
196 components by identifying turning points of runoff minima for each non-overlapping five-day
197 period. Turning points are joined by straight lines to obtain the baseflow hydrograph. Daily
198 stormflow was subsequently calculated by subtracting daily baseflow from daily total runoff. A
199 baseflow index (BFI) was then calculated by dividing baseflow by total runoff, expressed as a
200 percent.

201 All runoff metrics were calculated from daily records and summarized to mean annual
202 and monthly values to assess potential seasonality effects. Further, a time period of particular
203 interest for the study area is from March – July, for which runoff (and nutrient loads) have been
204 shown to be critical for determining HAB severity in the western Lake Erie Basin (Baker *et al.*,
205 2019). For this reason, we calculated runoff metrics by averaging daily values for these months
206 in watersheds with high (> 40% area) amounts of tile drainage. In addition, we calculated the day
207 of calendar year in which 50% of annual runoff occurs to evaluate the effect that annual

208 streamflow timing had on bloom severity. Runoff metrics were compared to the Western Lake
 209 Erie Bloom Severity Index, calculated by the United States National Oceanic and Atmospheric
 210 Administration based on algal bloom biomass, to evaluate relationships between streamflow
 211 response and HAB severity.

212 To assess how water is stored and released following storm events, we performed
 213 hydrograph recession analyses for each of the watersheds considered in this study. The
 214 calculation of the recession constant required selecting an analytical expression to fit to the
 215 recession curve, determining the typical recession period, and optimizing the recession
 216 parameters (WMO, 2008). We used the R package ‘lfstat’ to determine recession rates (Koffler
 217 *et al.*, 2016). The recession curve was modelled using an exponential equation assuming a single
 218 linear reservoir where storage is proportional to outflow:

$$219 \quad Q_t = Q_o e^{\left(\frac{-t}{C}\right)} \quad (1)$$

220 where Q_t is total runoff at time t ; Q_o is total runoff at the beginning of the recession period ($t=0$),
 221 and C is the recession constant [time], which is the number of days needed for runoff to decrease
 222 one log cycle. The recession curve plots as a straight line with slope $-1/C$ on a semi-logarithmic
 223 plot of t versus $\ln Q_t$. Both master recession curve (MRC) and individual recession segments
 224 (IRS) methods require criteria for selecting recession segments and the period of discharge to
 225 disregard following peak runoff to avoid selecting times of rapid response following a rainfall
 226 event that were not caused by groundwater discharge. For both analyses, a minimum segment
 227 length of five days was chosen and recession segments began at least two days after peak flood
 228 discharge and after runoff was below a Q25 threshold (i.e. the highest 25% of runoff following
 229 peak flood discharge was omitted). The MRC method constructs a single mean recession curve,
 230 while in the IRS method a recession model is fit to each segment and the recession constant C , is
 231 determined as the mean value of individual recession segments.

232

233 **3. Results**

234 **3.1 Watershed characteristics**

235 According to the AgTile-US dataset (Valayamkunnath *et al.*, 2020), mean areal coverage of tile
236 drainage for the watersheds analyzed in this study was 27.8% and ranged from 0.5% to 61.0%.
237 Watersheds were split into three roughly equal groups to compare the mean streamflow
238 response: low (< 15% tile drained, n = 18), medium (15% - 40% tile drained, n = 24), and high
239 (>40% tile drained, n = 17) (Figure 1). Watersheds in the medium and high drainage categories
240 were located primarily in northwestern Ohio, while watersheds in the low drainage category
241 were spread out throughout the state. Mean watershed size was similar for all three drainage
242 categories (Table 1). There was a significant positive relationship with agricultural land and tile
243 drainage (Pearson's $r = 0.86$, $p < 0.001$, Figure 2a, Table 1) and a significant negative
244 relationship with mean watershed slope and tile drainage (Pearson's $r = -0.77$, Figure 2b, Table
245 1). Tile drainage was significantly positively correlated to clay content (Pearson's $r = 0.54$,
246 Figure 2c, Table 1) and significantly negatively correlated to the depth of the seasonally high
247 water table provided in the GAGES-II dataset (Pearson's $r = -0.71$, Figure 2d, Table 1; Falcone,
248 2011). [Insert Figure 1] [Insert Figure 2] [Insert Table 1]

249 **3.2 Precipitation**

250 Mean annual precipitation (PRISM Climate Group, 2019) for the 59 watersheds over the ten-year
251 period (2010-2019) was 1109 mm and ranged from 945 mm in 2010 to 1465 mm in 2011. Mean
252 annual precipitation was significantly greater for the low drainage category (1160 mm) compared
253 to the medium (1099) or high drainage (1067 mm) categories (Table 2). On average across all 59
254 watersheds, spring and summer months (April – September) were wetter than fall and winter
255 months (October – March). The five-month period from March – July contributed 50% of annual
256 precipitation in high drainage watersheds.

257 For high drainage watersheds June had the most precipitation (133 mm), followed by
258 May (123 mm) and July (111 mm), while January had the least precipitation (60 mm), followed
259 by February (66 mm) and December (70 mm). In medium drainage watersheds June had the
260 most precipitation (132 mm) followed by May (115 mm) and July (111 mm), while January had
261 the least precipitation (63 mm), followed by February (72 mm) and October (79 mm). For low
262 drainage watersheds June had the most precipitation (140 mm), followed by July (113 mm) and
263 May (111 mm), while January had the least precipitation (71 mm), followed by February (80
264 mm) and November (83 mm). [Insert Table 2]

265 3.3 Runoff metrics

266 Mean annual runoff for the 59 watersheds from 2010 – 2019 was 435 mm and ranged from 298
267 mm in 2012 to 717 mm in 2011. Mean annual runoff was significantly greater for the low
268 drainage category (469 mm) compared to the medium (429 mm) and high drainage (407 mm)
269 categories (Table 2). On average, across all 59 watersheds winter and spring months (January –
270 June) produced more runoff compared to summer or fall months (July – December). The five-
271 month period from March – July contributed 57% of annual runoff in high drainage watersheds.

272 For high drainage watersheds March had the most runoff (59 mm), followed by April (56
273 mm) and June (44 mm), while August had the least runoff (6 mm), followed by September (7
274 mm) and October (12 mm). In medium drainage watersheds March had the most runoff (64 mm)
275 followed by April (59 mm) and February (49 mm), while September had the least runoff (8 mm),
276 followed by August (9 mm) and October (13 mm). For low drainage watersheds March had the
277 most runoff (71 mm), followed by April (65 mm) and February (57 mm), while August and
278 September had the least runoff (13 mm), followed by October (16 mm).

279 Mean annual runoff ratio for the 59 study watersheds from 2010 – 2019 was 39.1% and
280 ranged from 33.6% in 2010 to 47.8% in 2018. There was a significant positive relationship
281 between mean annual runoff ratio and mean annual precipitation among all 59 watersheds
282 (Pearson's $r = 0.48$). Despite significantly greater mean annual precipitation and runoff for the
283 low drainage category, mean annual runoff ratio was the same among all drainage categories
284 (Table 2). Mean annual runoff ratio was not significantly correlated to tile drainage (Figure 3a)
285 or watershed area (Figure 4a) for any of the drainage categories considered. Peak daily runoff
286 was similar among all drainage categories and not significantly correlated to tile drainage (Figure
287 3b; Table 2). Peak daily runoff was significantly negatively correlated to watershed area for the
288 medium (Pearson's $r = -0.67$) and high (Pearson's $r = -0.76$) drainage categories, but not for the
289 low drainage category (Figure 4b). [Insert Figure 3] [Insert Figure 4]

290 The percent of time in which mean daily streamflow was greater than mean annual
291 streamflow ($T_{Q_{\text{mean}}}$) was significantly negatively correlated to tile drainage (Pearson's $r = -0.57$,
292 Figure 3c, Table 2). A lower $T_{Q_{\text{mean}}}$ value implies a flashier hydrograph response for the high
293 drainage category watersheds. There was a significant positive relationship between watershed
294 area and $T_{Q_{\text{mean}}}$ for the medium (Pearson's $r = 0.47$) and high (Pearson's $r = 0.51$) drainage

295 categories, but not for the low drainage category (Figure 4c). The mean annual baseflow index
296 (BFI) was significantly negatively correlated to tile drainage (Pearson's $r = -0.58$, Figure 3d,
297 Table 2). The mean annual BFI for the high drainage category was 20.9% compared to 40.3% for
298 the low drainage category. Conversely, watersheds with a high percentage of tile drainage had
299 significantly higher stormflow proportions compared to watersheds with low to medium
300 percentages of tile drainage. There was no significant relationship between watershed area and
301 BFI for any of the drainage categories (Figure 4d).

302 Both MRC and IRS techniques for hydrograph recession analysis revealed a significant
303 negative correlation between recession constants and tile drainage (Pearson's $r = -0.45$, Figs. 5a;
304 Pearson's $r = -0.46$, Figure 5c, Table 2). There was no significant relationship with watershed
305 area and recession constant using either MRC or IRS methods for any of the drainage categories
306 (Figs. 5b & 5d). Both MRC and IRS recession constants were significantly positively correlated
307 to annual BFI (Pearson's $r = 0.89$), $T_{Q_{\text{mean}}}$ (Pearson's $r = 0.70$), and average soil permeability
308 (Pearson's $r = 0.79$) (Falcone, 2011). These relationships suggest a flashier hydrograph response
309 from watersheds with higher percentages of tile drainage and poorer drainage capabilities. [Insert
310 Figure 5]

311 It should be noted that the March – July BFI was similar to the annual BFI and was
312 significantly lower for the high drainage category watersheds (Table 2). In addition, the amount
313 of March – July stormflow as a percentage of total annual runoff was significantly positively
314 correlated to tile drainage (Pearson's $r = 0.58$, Figure 6a). The amount of annual runoff from
315 March – July stormflow approached 50% for watersheds with high percentages of tile drainage,
316 while the percent of total annual runoff from March – July stormflow in watersheds with low
317 percentages of tile drainage was around 30%. [Insert Figure 6]

318 We compared runoff metrics from the high drainage category watersheds, which
319 predominantly drain into western Lake Erie, to the Western Lake Erie Bloom Severity Index and
320 found mean March – July total stormflow (mm) to be the best predictor of bloom severity for all
321 of the runoff metrics (Pearson's $r = 0.90$, Figure 6b). The March – July stormflow runoff ratio
322 (i.e. the ratio of total stormflow to total precipitation during March – July) was also highly
323 positively correlated to bloom severity (Pearson's $r = 0.87$, Figure 6c), unlike the March – July
324 baseflow runoff ratio that did not show any correlation (Figure 6c). Another runoff metric that

325 was highly correlated to the bloom severity index was the mean day of year in which 50% of
326 annual runoff occurred (Pearson's $r = 0.89$, Figure 6d). Recent years with the highest bloom
327 severity index (>10) observed 50% of annual streamflow in June, while years with less severe
328 blooms saw 50% of annual streamflow occurring much earlier in the year.

329

330 **4. Discussion**

331 **4.1 Comparison with other studies across the Midwestern U.S.**

332 Our results on the streamflow response of watersheds with varying percentages of tile
333 drainage in Ohio are markedly different from previous studies conducted in Iowa watersheds.
334 We showed a significant negative relationship between tile drainage percent and mean annual
335 baseflow index (BFI) (Figure 3d) and a significant positive relationship between tile drainage
336 percentage and Mar-Jul total stormflow (Figure 6a) for 59 watersheds in Ohio. This is in contrast
337 to extensive research performed with Iowa watersheds that showed an increase in baseflow
338 proportions with tile drainage percentage (Schilling and Libra, 2003; Schilling and Helmers,
339 2008; Boland-Brien et al., 2014). This should not come as a surprise since previous work showed
340 a linear relationship between rainfall and tile drainage in which 12.6% of rainfall was recovered
341 in tile drainage in Iowa, compared to 22.2% in Ohio (Logan *et al.*, 1980). According to 30-year
342 climate normal, the watersheds used our study have significantly greater mean annual
343 precipitation (979 mm) compared to Iowa watersheds (869 mm) analyzed by Boland-Brien et al.
344 (2014) (Falcone, 2011). In the Midwest, Ohio and Iowa roughly represent two end-members in
345 terms of the meteorological and physical characteristics of watersheds with high percentages of
346 tile drainage; thus, it is fair to assume that tile drainage could result in greater baseflow or
347 stormflow proportions, depending on site-specific meteorological and physical conditions.

348 When our results are compared to the work from Boland-Brien et al. (2014) - who
349 calculated similar runoff metrics - it is clear that large percentages of tile drainage can cause a
350 notably different hydrologic response at the watershed scale in terms of baseflow and stormflow
351 proportions and the general flashiness behavior. Boland-Brien et al. (2014) reported a mean BFI
352 of 67% for the Iowa watersheds considered with a high degree of tile drainage ($>50\%$),
353 compared to 22% reported for the high drainage category ($>40\%$) in our Ohio study. The mean
354 annual runoff ratio was notably higher for watersheds analyzed in our study (39%) compared to

355 those by Boland-Brien et al. (2014) (28%). In addition, our results suggest $T_{Q_{\text{mean}}}$ and the
356 recession constants indicate flashier streamflow behavior in watersheds with high amounts of tile
357 drainage compared to the Iowa watersheds that showed the opposite trend. Given low drainage
358 category watersheds had significantly greater mean annual precipitation and runoff (Table 2), we
359 would usually expect to observe a significantly greater mean annual runoff ratio and peak daily
360 runoff for the low drainage category. However, there were no significant difference between
361 mean annual runoff ratio or peak daily runoff for any of the drainage categories, suggesting
362 medium and high drainage category watersheds had greater mean annual runoff ratios and peak
363 daily runoff than expected. All of these results suggest an increasing percentage of tile drainage
364 leads to flashier watersheds in Ohio.

365 Of course, the watersheds analyzed by Boland-Brien et al. (2014) were substantially
366 larger (average area of 1,666 km²) compared to the ones presented in this study (average area of
367 605 km²), which likely partially explains the larger observed BFI in Iowa watersheds. This
368 difference, however, does not explain the opposite trend observed between the relationship of
369 percent tile drainage and runoff metrics. Despite similar soil textures (i.e. sand, silt, clay
370 percentages) between our watersheds and the ones presented in Boland-Brien et al. (2014), the
371 Ohio watersheds showed significantly greater soil bulk density (1.54 g/cm³) compared to the
372 Iowa watersheds (1.44 g/cm³) (Falcone, 2011). The lower bulk density values observed in Iowa
373 favor faster infiltration rates compared to Ohio, which likely results in greater groundwater
374 recharge and smaller proportions of stormflow in Iowa. In fact, the Ohio watersheds analyzed in
375 this study had a significantly greater percent of soils in hydrologic group C (62%), characterized
376 by moderately fine or fine texture, slow soil infiltration rates with layers impeding the downward
377 movement of water (Falcone, 2011). In contrast the Iowa watersheds had a significantly lower
378 percent of soils in hydrologic group C (16%) and were dominated by soils in hydrologic group
379 B, characterized by moderately deep, coarse, well drained soils with moderate infiltration rates.
380 Another substantial difference between the two areas is the depth to seasonally high water table,
381 which was significantly smaller for the Ohio watersheds (which averaged 0.80 m) compared to
382 the Iowa watersheds (which averaged 1.23 m) (Falcone, 2011). Tile drains installed in Ohio
383 fields with slow soil infiltration rates and shallow water tables creates a more direct response to
384 rainfall events observed in tile drainage outlets compared to installations in fields with moderate
385 infiltration rates and deeper water tables since soil bulk density typically increases with depth.

386 Another major difference in hydrologic response of Ohio and Iowa watersheds to varying
387 percentages of tile drainage was the homogenization of all runoff metrics with high percentages
388 of tile drainage reported by Boland-Brien et al. (2014). While our study watersheds with high
389 percentages of tile drainage did not show a relationship with drainage area for mean annual
390 runoff ratio or mean annual BFI, we found significant correlations between drainage area and
391 $T_{Q_{mean}}$ (Figure 4c) and peak daily runoff (Figure 4d) for the medium and high tile drainage
392 category watersheds, but not for the low drainage category watersheds. In contrast, drainage area
393 was not correlated to any of the runoff metrics for the low drainage category (Figure 4). As
394 mentioned before, larger watersheds typically show higher percentages of baseflow and a more
395 attenuated streamflow response as groundwater contributions increase (Price, 2011). However,
396 the influence of geological conditions on streamflow response will be most apparent during dry
397 conditions when baseflow contributions are high (Cross, 1949). Since the low drainage category
398 watersheds are more dispersedly located throughout Ohio (Figure 1), it is possible the geological
399 conditions are more variable for these watersheds compared to the medium or high drainage
400 categories, which are predominantly located in northwest Ohio and likely have more similar
401 geological conditions.

402 **4.2 Implications for nutrient transport**

403 The agricultural economic benefits of tile drainage are accompanied with environmental and
404 economic costs associated with impaired water quality. Water exiting tile drain outlets transport
405 agricultural pollutants (e.g. nitrogen, phosphorous, pesticides) downstream which can
406 accumulate leading to hypoxic zones and harmful algal blooms (HABs), with detrimental effects
407 to human and aquatic systems (Diaz, 2001). Harmful algal blooms are not unique to Ohio and
408 have become a global problem in recent decades (Ho *et al.*, 2019). The environmental
409 consequences of HABs are difficult to remediate and can negatively impact tourism, recreation,
410 property values, wildlife, and commercial fishing. In August of 2014, elevated microcystin toxin
411 levels associated with a HAB resulted in 400,000 residents left without drinking water. In Lake
412 Erie, the world's largest walleye fishery, summer-long HABs can result in \$5.6 million in lost
413 fishing expenditures alone (Wolf *et al.*, 2017).

414 Tile drainage is thought to reduce surface runoff, therefore improve soil stability and
415 limit the amount of erosion and particulate nutrient concentrations exporting via surface runoff.

416 While nutrient concentrations measured in tile drainage are often low during low discharge
417 periods, elevated nutrient concentrations have been measured during high discharge periods,
418 proving that tile drains can act as effective conduits for nutrient export from agricultural fields
419 (Dils and Heathwaite, 1999). Numerous studies have showed a strong surface connection to tile
420 drainage through macropores and other preferential flow paths (Stamm et al., 1998; Smith et al.,
421 2015; Williams et al., 2016; Macrae et al., 2019), and thus potential to transport nutrients applied
422 to the soil surface. In addition, recent research suggests storm events can accelerate the
423 subsurface transport of particulate and dissolved nutrient species (Jiang *et al.*, 2021).

424 The results reported in this study suggest that Ohio watersheds with large percentages of
425 tile drainage could be exacerbating the problem with downstream nutrient transport due to
426 increases in total stormflow amounts and proportions (Figure 3d; Figure 6a). In fact, recent HAB
427 severity observed in the western Lake Erie basin was significantly correlated to March-July
428 stormflow amounts (Figure 6b). It should be noted that one of the strongest correlations of
429 watershed attributes from the GAGES-II dataset with tile drainage percentage were estimates of
430 applied nitrogen (Pearson's $r = 0.79$) and phosphorus (Pearson's $r = 0.70$) from agricultural
431 censuses (Falcone, 2011). This should not be surprising given the strong correlation between
432 agriculture and tile drainage (Figure 2a) but emphasizes the role that watersheds with high
433 percentages of tile drainage, and higher percentages of stormflow, play in the downstream
434 transport of nutrients.

435 Direct HAB remediation is costly and involves either physical, chemical, or biological
436 control measures, but will not help mitigate future severe HABs. If left uncontrolled, HABs in
437 Lake Erie are estimated to cost Canada alone \$5.3 billion over the next 30 years (Smith *et al.*,
438 2019), thus targeting conservation efforts at the source could prove to be cost-effective. A
439 combination of both nutrient and water management practices are probably needed to improve
440 downstream aquatic conditions (Hanrahan *et al.*, 2019). In Ohio, soil test phosphorus
441 concentrations were found to be linearly related to dissolved concentration loads in tile-drained
442 fields, thus soil test phosphorus can be a good screening method to identify fields at risk for
443 greater phosphorus loss (Duncan *et al.*, 2017). Limiting fertilizer application prior to spring
444 storm events or incorporating fertilizer into the soil structure could help to reduce the
445 downstream transport of nutrients from tile-drained fields (Williams *et al.*, 2016). The strong

446 positive correlation between the timing of 50% of annual streamflow and HAB severity (Figure
447 6d) supports an earlier application of fertilizer to avoid excess nutrient transport during large
448 late-spring storms which could be contributing to more severe HABs when water temperatures
449 are greater.

450 Conservation practices that decrease the hydrologic response time to storm events in
451 Ohio watersheds could benefit the aquatic health of downstream communities (e.g. buffer strips,
452 wetland restoration). Restoring 5-10% of the 4,000 km² Great Black Swamp in the Maumee
453 River basin could reduce phosphorus loading by 18-37% (Mitsch, 2017). Another technique that
454 could decrease the hydrologic response time and thus greatly reduce the export of nutrient loads
455 from agricultural fields is drainage water management, which has been shown to significantly
456 reduce annual tile drainage discharge and subsequent nutrient loads (e.g. Williams et al., 2015).
457 Through drainage water management, tile drainage outlets can be manipulated at the edge of
458 field to reduce discharge during winter fallow periods and times in which field accessibility is
459 not imperative.

460 **4.3 Limitations and future research needs**

461 One of the main limitations to our analyses was accurately selecting appropriate watersheds to
462 compare the hydrologic response. The medium and high drainage category watersheds analyzed
463 in this study are primarily located in northwest Ohio, while the low drainage category watersheds
464 are scattered more throughout the state. Thus, the low drainage category watersheds have more
465 variable soil properties, land cover, and precipitation patterns compared to the medium and high
466 drainage category watersheds. In addition, some of the low drainage category watersheds have
467 much greater mean slope (>4 %) and forest cover, thus the processes leading to the observed
468 streamflow response in these low drainage category watersheds are likely quite different
469 compared to the medium or high drainage category watersheds or the remaining low drainage
470 category watersheds with lower mean watershed slope (< 4%). We performed the same analyses
471 after removing the steepest watersheds (> 4% mean watershed slope, n = 13), which tended to be
472 located in eastern and southern Ohio and none of the results changed, suggesting that our results
473 and interpretations presented are robust across a range of tile drainage percentage for Ohio
474 watersheds.

475 Another limitation for this study was relying on the modeled tile drainage dataset
476 (Valayamkunnath *et al.*, 2020) for accurate identification of land drained by subsurface tiles.
477 While recent advancements using thermal infrared sensors deployed with drones have provided
478 adequate representation of tile delineation at agricultural fields (Allred *et al.*, 2018), it is
479 currently unrealistic to obtain this information at the scale of the watersheds analyzed in this
480 study. In Ohio, the total land area in the AgTile-US dataset is within 0.22 % of the total tile
481 drained area reported in the USDA Census of Agriculture. However, neither of these datasets are
482 able to provide information on whether drainage water management is implemented. For this
483 reason, we assumed that drainage water management did not contribute substantially to the tile
484 drained land or that drainage water management is uniformly practiced throughout the study
485 watersheds, thus would not impact any particular watershed or drainage category.

486 Baseflow is a fairly ambiguous term but is generally thought to be representative of the
487 water that sustains streamflow in between storms. In contrast, stormflow (i.e. quickflow, Hewlett
488 and Hibbert, 1967) is a term used to represent the remaining streamflow not accounted for in
489 baseflow. While mathematical baseflow separation techniques have been used since the early
490 20th century, more recently, chemical and isotopic mass-balance methods have become a popular
491 alternative to mathematical approaches and are generally considered to be more physically-based
492 due to incorporating chemical and/or isotopic information (Schilling and Helmers, 2008;
493 Tesoriero *et al.*, 2013; Frisbee *et al.*, 2017; Schilling *et al.*, 2019). However, mathematical
494 approaches continue to be used widespread due to fewer data requirements, with only stream
495 discharge being needed to perform baseflow separation (Schilling and Helmers, 2008; Boland-
496 Brien *et al.*, 2014; Schilling and Jones, 2019). Since the calculations for baseflow and stormflow
497 used in this study are strictly based on the shape of the hydrograph, mathematical derivations of
498 these terms cannot differentiate the geographic sources or ages and residences times of these two
499 hydrograph sources. For example, under dry conditions tile drainage is likely composed of
500 primarily baseflow derived from relatively older groundwater, whereas during wet storm
501 conditions tile drainage could be comprised from a mixture of older groundwater and younger
502 rainfall event water. Thus, the water discharging from tile drainage cannot be assumed to be
503 entirely baseflow or stormflow. Additional research utilizing unique tracer signatures would be
504 valuable for assessing the relative age of stream water and discharge from tile drainage outlets
505 and downstream rivers and lakes.

506

507

508 5. Conclusion

509 This study analyzed the effect of tile drainage on various runoff metrics for 59 Ohio watersheds.
510 We used a recently developed 30-m resolution tile drainage dataset to calculate the percentage of
511 tile drainage in each watershed. Our results indicate that high percentages of tile drainage (> 40%
512 of watershed area) result in significantly greater percentages of stormflow and a flashier
513 hydrograph response in general, which contrasts with similar studies conducted in Iowa that
514 showed increases in baseflow percentages and less flashy hydrographs for heavily tiled
515 watersheds. Using baseflow and recession analysis, watersheds with high percentages of tile
516 drainage consistently reported flashier behavior compared to watersheds with low percentages of
517 tile drainage. The total amount of March – July stormflow and the stormflow proportion during
518 this time was significantly positively correlated to western Lake Erie harmful algal bloom
519 severity during the study period (2010-2019).

520 Increases in stormflow proportions, or the fast-varying portion of the hydrograph, are
521 problematic for the downstream transport of nutrients and could be linked to exacerbated
522 harmful algal bloom severity in Lake Erie observed in recent years. Given the recent trend in
523 more frequent large rain events and warmer temperatures in the Midwest, increased harmful
524 algal bloom severity will continue to be an ecological and economic problem for the region if
525 management efforts are not addressed at the source. Management practices that reduce the
526 hydrologic response time to storm events, such as buffer strips, wetland restoration, or drainage
527 water management, are likely to improve downstream aquatic health conditions by limiting the
528 transport of nutrients after storm events.

529

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534

535 **DATA AVAILABILITY**

536 All data sources (streamflow, precipitation, watershed characteristics, tile drainage) are available
537 publically. Code is available from the corresponding author upon reasonable request.

538

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686

687 **TABLES**

688 Table 1: Mean watershed characteristics for the three drainage categories. Area, Agricultural
 689 land, slope, clay, and depth to seasonally high water table from the GAGES-II dataset (Falcone,
 690 2011). Unique letters represent significant differences ($p < 0.05$) using the Tukey Test.

Drainage Category	Number	Area (km ²)	Tile Drainage (%)	Agricultural Land (%)	Slope (%)	Clay (%)	Water Table Depth (m)
Low	18	493 a	4.9 a	42.2 a	5.3 a	27.4 a	1.08 a
Medium	24	655 a	28.5 b	72.6 b	1.3 b	27.9 a	0.80 b
High	17	653 a	51.2 c	82.2 c	0.4 b	33.7 b	0.52 c

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698 Table 2: Mean annual precipitation (PRISM Climate Group, 2019), area-weighted runoff, runoff
 699 ratio, peak daily runoff (Peak Q), proportion of time mean daily streamflow is greater than mean
 700 annual streamflow ($T_{Q_{mean}}$), mean annual baseflow index (BFI), March – July BFI, and recession
 701 constants from MRC and IRS methods for the three drainage categories during all study years
 702 (2010-2019). Unique letters represent significant differences ($p < 0.05$) using the Tukey Test.

Drainage Category	Precip. (mm)	Runoff (mm)	Runoff Ratio (%)	Peak Q (mm/day)	$T_{Q_{mean}}$ (%)	BFI (%)	Mar-Jul BFI (%)	MRC	IRS
Low	1160 a	469 a	40.3 a	17.7 a	27.9 a	40.3 a	41.1 a	6.3 a	7.6 a
Medium	1099 b	429 b	38.9 a	18.3 a	25.7 a	35.6 a	35.8 a	6.5 a	7.5 a
High	1067 b	407 b	38.1 a	19.6 a	22.3 b	20.9 b	22.2 b	3.5 b	4.2 b

703

704 FIGURE LEGENDS

705 Figure 1: Location and drainage category of 59 watersheds used in this study.

706

707 Figure 2: Watershed tile drainage (%) versus agricultural land (a), mean watershed slope (b),
 708 average value of soil clay content (c), and depth to seasonally high water table (d) (Falcone,
 709 2011).

710

711 Figure 3: Mean annual runoff ratio (%) (a), mean annual peak daily runoff (b), percent of
 712 observations daily runoff exceeds mean annual runoff ($T_{Q_{mean}}$) (c), and mean annual baseflow
 713 index (BFI) (d) versus watershed tile drainage (%).

714

715 Figure 4: Mean annual runoff ratio (%) (a), mean annual peak daily runoff (b), percent of
 716 observations daily runoff exceeds mean annual runoff ($T_{Q_{mean}}$) (c), and mean annual baseflow
 717 index (BFI) (d) versus watershed drainage area colored by drainage category.

718

719 Figure 5: Recession (MRC) constant versus watershed tile drainage (a) and watershed area (b)
 720 colored by drainage category. Recession (IRS) constant versus watershed tile drainage (c) and
 721 watershed area (d) colored by drainage category.

722

723 Figure 6: March – July stormflow as a percentage of total annual runoff versus tile drainage (a).

724 March – July mean total stormflow (mm) for the high drainage category watersheds vs Western

725 Lake Erie Bloom Severity Index (b). March – July mean stormflow (blue) and baseflow (red)

726 runoff ratio for the high drainage category watersheds vs Western Lake Erie Bloom Severity

727 Index (c). Day of calendar year (DOY) when 50% of annual streamflow occurs for the high

728 drainage category watersheds vs Western Lake Erie Bloom Severity Index (d).

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