

Denitrification losses in response to N fertiliser rates - a synthesis of high temporal resolution N₂O, in-situ ¹⁵N₂O and ¹⁵N₂ measurements and fertiliser ¹⁵N recoveries in intensive sugarcane systems

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Abstract

Denitrification is a key process in the global nitrogen (N) cycle, causing both nitrous oxide (N₂O) and dinitrogen (N₂) emissions. However, estimates of seasonal denitrification losses (N₂O+N₂) are scarce, reflecting methodological difficulties in measuring soil-borne N₂ emissions against the high atmospheric N₂ background and challenges regarding their spatio-temporal upscaling. This study investigated N₂O+N₂ losses in response to N fertiliser rates (0, 100, 150, 200 and 250 kg N ha⁻¹) on two intensively managed tropical sugarcane farms in Australia, by combining automated N₂O monitoring, in-situ N₂ and N₂O measurements using the ¹⁵N gas flux method and fertiliser ¹⁵N recoveries at harvest. Dynamic changes in the N₂O/(N₂O+N₂) ratio (< 0.01 to 0.768) were explained by fitting generalised additive mixed models (GAMMs) with soil factors to upscale high temporal-resolution N₂O data to daily N₂ emissions over the season. Cumulative N₂O+N₂ losses ranged from 12 to 87 kg N ha⁻¹, increasing non-linearly with increasing N fertiliser rates. Emissions of N₂O+N₂ accounted for 31–78% of fertiliser ¹⁵N losses and were dominated by environmentally benign N₂ emissions. The contribution of denitrification to N fertiliser loss decreased with increasing N rates, suggesting increasing significance of other N loss pathways including leaching and runoff at higher N rates. This study delivers a blueprint approach to extrapolate denitrification measurements at both temporal and spatial scales, which can be applied in fertilised agroecosystems. Robust estimates of denitrification losses determined using this method will help to improve cropping system modelling approaches, advancing our understanding of the N cycle across scales.

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4 **intensive sugarcane systems**
5

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16
17 **Key Points:**

- 18 • A novel method to estimate N₂O+N₂ losses by combining high-frequency N₂O data, the in-situ
19 ¹⁵N gas flux method and fertiliser ¹⁵N recoveries
- 20 • Denitrification losses in sugarcane systems were 12–87 kg N ha⁻¹ mostly as N₂ (> 94%) and
21 increased non-linearly with increasing N rates
- 22 • Denitrification accounted for 31–78% of N fertiliser losses while the proportion of reactive N
23 losses increased with increasing N rates

24 **Abstract**

25 Denitrification is a key process in the global nitrogen (N) cycle, causing both nitrous oxide (N₂O)
26 and dinitrogen (N₂) emissions. However, estimates of seasonal denitrification losses (N₂O+N₂) are scarce,
27 reflecting methodological difficulties in measuring soil-borne N₂ emissions against the high atmospheric
28 N₂ background and challenges regarding their spatio-temporal upscaling. This study investigated N₂O+N₂
29 losses in response to N fertiliser rates (0, 100, 150, 200 and 250 kg N ha⁻¹) on two intensively managed
30 tropical sugarcane farms in Australia, by combining automated N₂O monitoring, in-situ N₂ and N₂O
31 measurements using the ¹⁵N gas flux method and fertiliser ¹⁵N recoveries at harvest. Dynamic changes in
32 the N₂O/(N₂O+N₂) ratio (< 0.01 to 0.768) were explained by fitting generalised additive mixed models
33 (GAMMs) with soil factors to upscale high temporal-resolution N₂O data to daily N₂ emissions over the
34 season. Cumulative N₂O+N₂ losses ranged from 12 to 87 kg N ha⁻¹, increasing non-linearly with increasing
35 N fertiliser rates. Emissions of N₂O+N₂ accounted for 31–78% of fertiliser ¹⁵N losses and were dominated
36 by environmentally benign N₂ emissions. The contribution of denitrification to N fertiliser loss decreased
37 with increasing N rates, suggesting increasing significance of other N loss pathways including leaching and
38 runoff at higher N rates. This study delivers a blueprint approach to extrapolate denitrification
39 measurements at both temporal and spatial scales, which can be applied in fertilised agroecosystems.
40 Robust estimates of denitrification losses determined using this method will help to improve cropping
41 system modelling approaches, advancing our understanding of the N cycle across scales.

42

43 **Plain Language Summary**

44 Denitrification is a key soil process in the global nitrogen (N) cycle. Denitrification produces a potent
45 greenhouse gas, nitrous oxide (N₂O), but also turns reactive N into environmentally benign dinitrogen (N₂).
46 The response of these N losses to N fertiliser inputs is critical to reducing environmental impacts while
47 maintaining crop productivity in agriculture. However, difficulties in measuring and upscaling N₂ emissions
48 at the farm scale hinder estimation of denitrification losses, leaving denitrification as a major uncertainty
49 for N budgets. This study quantified denitrification losses in response to N fertiliser rates on sugarcane
50 farms in Australia, by combining automated greenhouse gas monitoring systems, N isotope techniques and
51 statistical models. This unique approach demonstrated denitrification as a major N loss pathway, increasing
52 nonlinearly with increasing N rates. Fertiliser N budgets showed that environmentally harmful N losses
53 increased more than proportionally with N inputs. These findings emphasise that excessive N fertiliser use
54 leads to agronomic inefficiency with severe adverse effects on the surrounding ecosystems such as the Great
55 Barrier Reef. The novel approach presented here will advance our understanding of N cycling across scales
56 and thus aid in reducing the environmental footprint of global agricultural production.

57

58 **1 Introduction**

59 Denitrification is a key process in the global nitrogen (N) cycle, reducing nitrate (NO_3^-) to gaseous
60 N emissions in the form of nitrous oxide (N_2O) and dinitrogen (N_2). Emissions of N_2O contribute to climate
61 change, as N_2O is a long-lived atmospheric trace gas with a global warming potential 273 times higher than
62 that of carbon dioxide (CO_2) over a 100-year period (IPCC, 2021) and the largest remaining threat to the
63 stratospheric ozone layer (Portmann et al., 2012; Ravishankara et al., 2009). Emissions of N_2 , while
64 environmentally benign, still represent a loss of N from the system, with potential detrimental effects on
65 crop growth and productivity in agricultural systems. Despite a growing body of denitrification research
66 delivering both N_2O and N_2 data from different agroecosystems, the ratio between reactive N_2O and N_2
67 remains a major uncertainty for N budgets across scales (Friedl et al., 2020a; Scheer et al., 2020). Growing
68 evidence of non-linear responses of N_2O emissions to N fertiliser rates (Shcherbak et al., 2014; Takeda et
69 al., 2021a) together with increasing fertiliser ^{15}N loss with increasing N rates (Rowlings et al., 2022;
70 Schwenke & Haigh, 2016; Takeda et al., 2021b) in intensive cropping systems suggests excessive N inputs
71 promote denitrification losses and lead to inefficiency of N use and adverse environmental impacts.
72 Constraining the response of denitrification losses to N fertiliser rates is therefore critical for sustainable N
73 management strategies to reduce N losses while maintaining crop productivity.

74 Yet, measuring N_2 emissions from the soil against the high atmospheric N_2 background remains
75 challenging (Friedl et al., 2020a; Groffman et al., 2006), reflected in the small number of studies quantifying
76 both N_2O and N_2 in the field. The Helium/Oxygen atmosphere method (He/ O_2 method) (Butterbach-Bahl
77 et al., 2002; Scholefield et al., 1997) and the ^{15}N gas flux method (Mosier & Schimel, 1993) are considered
78 suitable for the direct quantification of N_2 and N_2O from soils. For the He/ O_2 method, soil cores are
79 incubated in the laboratory and the headspace atmosphere inside the closed incubation system is replaced
80 with a He/ O_2 mixture to measure soil-borne N_2 emissions. Field-scale seasonal/annual N_2 emissions can be
81 estimated by repeated short laboratory measurements of soil cores, which are returned to the field after
82 incubation. Uncertainty in the cumulative emissions with this approach however remains high due to
83 disturbance of the soil, as in-situ measurements are not possible with this method (Chen et al., 2019; Zistl-
84 Schlingmann et al., 2019). The ^{15}N gas flux method is the only method to measure N_2 emissions under both
85 laboratory and field conditions. The method requires highly enriched ^{15}N fertiliser to be applied to a
86 designated plot. Gas samples are taken using the static chamber method and analysed for their different
87 isotopologues of N_2 and N_2O via isotope ratio mass spectrometry (IRMS) (Friedl et al., 2020a). As a result,
88 evaluation of denitrification losses under field conditions is scarce and mostly limited to measurement
89 periods of less than a month (Baily et al., 2012; Buchen et al., 2016; Friedl et al., 2017; Warner et al., 2019;

90 Weier et al., 1998), as the sensitivity of this method declines in response to the decrease of the ^{15}N
91 enrichment in the soil NO_3^- pool. Due to the shortcomings of available direct measurement methods,
92 estimates of cumulative denitrification losses over the crop growing season require upscaling approaches
93 accounting for the highly dynamic response of denitrification to its drivers.

94 Denitrification losses have been estimated by applying the average ratio between N_2O and N_2
95 emissions measured for a short period under laboratory conditions to N_2O emissions measured over the
96 crop growing season under field conditions (Scheer et al., 2009). Burchill et al. (2016) measured the $\text{N}_2:\text{N}_2\text{O}$
97 ratio bimonthly in the field and interpolated the ratio linearly between sampling events to apply to more
98 frequent N_2O measurements. However, the ratio between N_2O and N_2 is highly variable and changes rapidly
99 in a non-linear fashion depending on interactions between environmental drivers of denitrification such as
100 soil water content (Friedl et al., 2016), temperature (Bizimana et al., 2021), C availability (Qin et al., 2017)
101 and N substrate availability (Chen et al., 2019; Warner et al., 2019), leading to considerable bias and large
102 uncertainty in N_2 estimation if a fixed ratio is used. Wang et al. (2020) correlated the $\text{N}_2\text{O}/(\text{N}_2\text{O}+\text{N}_2)$ ratio
103 measured under laboratory conditions to multiple soil factors and applied the ratio to field-measured N_2O
104 to estimate field-scale seasonal N_2 emissions. These approaches account for the dynamic response of the
105 $\text{N}_2:\text{N}_2\text{O}$ ratio to key drivers. However, the absence of plants may bias the measured ratios, as plant-soil-
106 microbe interactions are known to both affect magnitude and partitioning of N_2 and N_2O emissions (Henry
107 et al., 2008; Malique et al., 2019). Furthermore, inevitable disturbance of soil through sampling is also of
108 concern, while the lack of in-situ measurements hinders the direct validation of the $\text{N}_2:\text{N}_2\text{O}$ ratio calculated
109 as a function of key drivers. These shortcomings denote a high uncertainty of field-scale seasonal N_2
110 estimates using current approaches and demand a refined method that allows for robust estimates of N_2 and
111 N_2O emissions. Critically, accounting for the dynamic responses of the ratio between N_2O and N_2 to soil
112 factors needs to occur under field conditions in the presence of plants. Such estimates are urgently needed
113 to constrain N budgets in different agroecosystems and to refine N fertiliser management strategies for both
114 agronomic and environmental benefits.

115 The aim of this study was to estimate seasonal denitrification losses ($\text{N}_2\text{O}+\text{N}_2$) in response to N
116 fertiliser rates in intensively managed tropical sugarcane (*Saccharum* spp.) systems in Australia, by
117 combining high temporal resolution N_2O measurements with automated greenhouse gas (GHG) monitoring
118 systems, in-situ measurements of $\text{N}_2\text{O}/(\text{N}_2\text{O}+\text{N}_2)$ ratio with the ^{15}N gas flux method and fertiliser ^{15}N
119 recoveries. The dynamic changes in the $\text{N}_2\text{O}/(\text{N}_2\text{O}+\text{N}_2)$ ratio observed in the field were explained by fitting
120 generalised additive mixed models (GAMMs) with soil temperature, water-filled pore space (WFPS), soil
121 mineral N contents and CO_2 emissions, enabling spatio-temporal upscaling of high temporal frequency N_2O
122 measurements to N_2 emissions. Fertiliser-derived $\text{N}_2\text{O}+\text{N}_2$ losses were further calculated and compared

123 with fertiliser ^{15}N loss, corroborating the estimates of $\text{N}_2\text{O}+\text{N}_2$ at the cumulative scale and differentiating
124 fertiliser ^{15}N loss pathways. Establishing the response of $\text{N}_2\text{O}+\text{N}_2$ losses as well as their proportion of
125 fertiliser ^{15}N loss to N fertiliser application rates with this innovative approach will refine N budget
126 estimates across scales and allow evaluation of N fertiliser management strategies accounting for N losses
127 from agroecosystems.

128

129 **2 Materials and Methods**

130 In this study, in-situ measurements of N_2O and N_2 emissions from two sugarcane systems were
131 combined with previously reported high temporal resolution measurements of N_2O (Takeda et al., 2022;
132 Takeda et al., 2021a) and recovery of ^{15}N -labelled fertiliser in the plant, soil and N_2O (Takeda et al., 2022;
133 Takeda et al., 2021b) presented in the previous studies to quantify seasonal N_2O and N_2 losses.

134

135 **2.1 Study site**

136 The field experiments were conducted on commercial sugarcane farms in Burdekin, QLD (19° 37'
137 4'' S, 147° 20' 4'' E) from October 2018 to August 2019 and in Mackay, QLD (21° 14' 4'' S, 149° 04' 6''
138 E) from October 2019 to August 2020, described in details in Takeda et al. (2022). The climate is tropical
139 in both Burdekin and Mackay. The soil is classified as Brown Dermosol and Brown Kandosol in the
140 Australian Soil Classification (Isbell, 2016), or Luvisol and Fluvisol in the World Reference Base (WRB)
141 Classification (IUSS Working Group, 2014), at the Burdekin and Mackay sites, respectively. Sugarcane
142 varieties Q240 and Q208 were planted in 2015 and 2016 and the crop was the 3rd ratoon during the
143 experiment at the Burdekin and Mackay sites, respectively. Irrigation was applied by furrow irrigation at
144 the Burdekin site and overhead sprinkler at the Mackay site. Sugarcane is burnt before harvest to remove
145 the leaves at the Burdekin site, leaving little trash (crop residues) on the ground. 'Green cane trash
146 blanketing (GCTB)', a practice where the cane is harvested green and the trash is spread over the ground,
147 is practised at the Mackay site. Selected soil physical and chemical parameters are shown in Table 1.

148

149 **Table 1** Soil properties at 0-0.2 m depth at the Burdekin and Mackay sites

Variable	Burdekin	Mackay
BD (g cm ⁻³)	1.3	1.1
pH (H ₂ O)	6.92	4.13
Total C (%)	1.60	1.35
Total N (%)	0.08	0.09
Clay (%)	35.4	22.2
Silt (%)	26.0	15.9
Sand (%)	38.7	61.9
Mineral N (kg N ha ⁻¹)	37.0	31.8

150

151 2.2 Experimental design

152 A detailed description of the experimental design and setup at the Burdekin and Mackay sites can
 153 be found in Takeda et al. (2021a) and Takeda et al. (2022), respectively. Briefly, treatments at the Burdekin
 154 site were arranged in a randomised strip design with four plots across two strips for each N treatment. The
 155 experiment at the Mackay site had a completely randomised block design with three replicates per
 156 treatment, accompanied by an unfertilised control (0N) plot with three subplots. Fertiliser N rate treatments
 157 included 0N, 150 kg N ha⁻¹ (150N), 200 kg N ha⁻¹ (200N) and 250 kg N ha⁻¹ (250N), plus 100 kg N ha⁻¹
 158 (100N) at the Mackay site only. The recommended N application rate was based on the district yield
 159 potential and soil C content as outlined in the SIX EASY STEPS protocol of the Australian sugar industry
 160 (Schroeder et al., 2010) and was 150N at the Mackay site and 200N at the Burdekin site. Urea was applied
 161 by banding the fertiliser 10 cm deep and 30 cm from the bed centre on both sides of the cane row at the
 162 Burdekin site and by stool splitting 10 cm deep at the bed centre of the cane row at the Mackay site. For
 163 the ¹⁵N recovery in the soil and the plant, a 2.0 m section was excluded from the application of unlabelled
 164 N fertiliser in each plot and ¹⁵N enriched urea fertiliser (5 atom%) in solution was manually applied at the
 165 corresponding rate, matching the N fertiliser placement at the respective site.

166

167 2.3 Measurement of N₂O emissions using an automated chamber system

168 Soil-borne N₂O and CO₂ emissions were measured at a high temporal resolution using an automated
 169 chamber system (Grace et al., 2020) from 17 October 2018 to 15 August 2019 at the Burdekin site and from
 170 3 October 2019 to 24 August 2020 at the Mackay site. Details of the automated chamber system are given
 171 in Supporting Information S1.1. Manual gas sampling was conducted for the control plots of the Mackay
 172 site by the static closed chamber method (Friedl et al., 2017), detailed in Supporting Information S1.2. The

173 placement of the chambers accounted for N fertiliser placement and irrigation practice at each site: At the
 174 Burdekin site, chambers were installed covering the area from (a) the fertiliser band to the centre of the bed
 175 (bed chamber) and (b) the fertiliser band the centre of the furrow (furrow chamber). At the Mackay site,
 176 bed chambers (a) were placed at the centre of the bed (i.e., on the fertiliser band) and furrow chamber
 177 measurements (b) were substituted with those from the control plots. Daily N₂O and CO₂ emissions were
 178 calculated by averaging the measured hourly fluxes over a 24-h period from each chamber and multiplying
 179 by 24. Missing daily N₂O and CO₂ emissions between measurements were imputed by linear interpolation.

180

181 2.4 ¹⁵N-labelled N₂ and N₂O sampling and analysis in the micro plots

182 The application of highly enriched ¹⁵N urea fertiliser enabled us to quantify N₂ and N₂O emissions
 183 and their respective ratio, as well as the contribution of N fertiliser to N₂ and N₂O emissions. Micro plots
 184 were established alongside the main plots with N fertiliser rates of 150, 200 and 250 kg N ha⁻¹ at the
 185 Burdekin site and with 100, 150, 200 and 250 kg N ha⁻¹ at the Mackay site. The micro plots were arranged
 186 in a completely randomised block design with four replicates. A steel base (0.22 m × 0.22 m at the Burdekin
 187 site and 0.2 m × 0.4 m at the Mackay site) was installed in each micro plot and ¹⁵N enriched urea fertiliser
 188 (70 atom%) was applied inside the base at the corresponding rates. Gas sampling was conducted with static
 189 closed chambers at the Burdekin site from November 2018 to February 2019 and with semi-automated
 190 chambers at the Mackay site from October 2019 to January 2020 (Takeda et al. (2022), Supporting
 191 Information S1.3). The gas samples were analysed for the concentration of N₂O and CO₂ using a Shimadzu
 192 GC-2014 Gas Chromatograph (Shimadzu, Kyoto, Japan) and for different isotopologues of N₂ and N₂O
 193 using an Isotope Ratio Mass Spectrometer (IRMS) (20–22 Sercon Limited, UK).

194

195 2.5 The ¹⁵N gas flux method

196 The ¹⁵N enrichment of the soil NO₃⁻ pool undergoing denitrification (*a_p*) and the fraction of N₂ and
 197 N₂O emitted from this pool (*f_p*) were calculated following the equations outlined by Spott et al. (2006) and
 198 given in the Supporting Information S1.4. Multiplying the headspace concentrations of N₂ by the respective
 199 *f_p* value gave N₂ emitted via denitrification, with fluxes expressed in g N₂-N emitted ha⁻¹ d⁻¹. The precision
 200 of the IRMS for N₂ based on the standard deviation of atmospheric air samples (n = 18) at 95% confidence
 201 intervals was 4.4 × 10⁻⁷ and 6.0 × 10⁻⁷ for ²⁹R (²⁹N₂/²⁸N₂) and ³⁰R (³⁰N₂/²⁸N₂), respectively. The
 202 corresponding method detection limit ranged from 0.005 g N₂-N ha⁻¹ d⁻¹ with *a_p* assumed at 50 atom % to
 203 0.014 g N₂-N ha⁻¹ d⁻¹ with *a_p* assumed at 20 atom %. For each gas sample, the product ratio RN₂O was
 204 calculated as N₂O/(N₂O+N₂).

205

206 2.6 Plant and soil sampling and analyses

207 Plant and soil samples were taken from each of the 2.0 m sections prior to harvest (on 27–28 August
208 2019 at the Burdekin site and 25–26 August 2020 at the Mackay site). The procedure of plant and soil
209 sampling and analyses are detailed in Takeda et al. (2021b) and Takeda et al. (2022) as well as Supporting
210 Information S1.5. Briefly, aboveground sugarcane biomass, trash on the ground, two green leaves at the 3rd
211 node from the section and the adjacent row and remaining stools and major roots of sugarcane were
212 harvested. Soil samples were taken at three to four points between the bed and furrow centres using a soil
213 corer and a post-hole driver down to 1.0 m. The dried plant and soil samples were then finely ground and
214 analysed for N and ¹⁵N content via IRMS analysis (20–22 Sercon Limited, UK).

215

216 2.7 ¹⁵N calculations

217 Fertiliser ¹⁵N recovered in the plant, soil, N₂O and N₂ emissions were then calculated by ¹⁵N mass
218 balance (Friedl et al., 2017; Rowlings et al., 2016; Takeda et al., 2022) using equations detailed in the
219 Supporting Information S1.6. Overall fertiliser ¹⁵N loss was calculated by the difference between the N
220 applied and fertiliser ¹⁵N recovered in the soil and plant. The contribution of soil-derived N to plant N
221 uptake, N₂O and N₂ emissions was calculated by the difference between total N and fertiliser ¹⁵N recovered
222 in each N pool. This contribution of soil-derived N includes residue fertiliser N from the previous seasons,
223 N in the crop residue and other sources such as N deposition or fixation.

224

225 2.8 Auxiliary measurements

226 For soil NH₄⁺ and NO₃⁻ measurements, soil samples (0–20 cm depth) were taken in each plot one
227 day after fertilisation, every 3–7 days for the first three months and monthly thereafter. At each sampling
228 event, soils were taken from the bed near the fertiliser band at the Burdekin site where N fertiliser was
229 applied on both sides of the bed while from both bed and furrow at the Mackay site where N fertiliser was
230 applied at the centre of the bed. Soil NH₄⁺ and NO₃⁻ were extracted by adding 100 mL of 2 M KCl to 20 g
231 of air-dried soil and shaking the solution for one hour, followed by NH₄⁺ and NO₃⁻ content measurements
232 using a Gallery™ Discrete Analyzer (Thermo Fisher Scientific, USA). Volumetric soil water content was
233 measured at 10 cm depth every 30 minutes using a field-calibrated FDR soil moisture probe (EnviroSCAN,
234 Sentek, Australia) and then averaged per day. Then, WFPS was calculated from the volumetric soil water
235 content using the measured bulk density assumed constant during the season. Soil temperature in the surface
236 soil layer (0–10 cm) was measured every five minutes using a PT100 probe (IMKO, Germany) and then
237 averaged per day.

238

239 2.9 Upscaling N₂ emissions and statistical analysis

240 Statistical analyses and graphical presentations in this study were conducted using R statistical
241 software version 3.5.2 (R Core Team, 2018) with a significant level set at $P < 0.05$. Gap-filling of missing
242 daily measurements of N₂O and CO₂ emissions and soil NH₄⁺ and NO₃⁻ contents was conducted with linear
243 interpolation using “imputeTS” package (Moritz & Bartz-Beielstein, 2017).

244 Emissions of N₂ at the plot scale were calculated by fitting a statistical model trained with RN₂O
245 observed in the micro plots and applying the predicted RN₂O to high-frequency measurements of N₂O
246 emissions in the main plots. First, daily RN₂O measured in the micro plots at both sites were modelled per
247 N rate using the following predictors: (i) soil temperature and WFPS measured at each site, (ii) soil NH₄⁺
248 and NO₃⁻ contents measured near the band at the corresponding rate in the main plots, (iii) CO₂ emissions
249 measured in the micro plots and (iv) site as a factor. Then, daily RN₂O in the main plots were predicted per
250 plot for each bed and furrow position for the whole crop growing season using soil temperature, WFPS,
251 soil NH₄⁺ and NO₃⁻ contents and daily CO₂ emissions measured in the main plots. Daily N₂ emissions were
252 calculated per plot for each bed and furrow position for the whole crop growing season as the product of
253 predicted RN₂O and daily N₂O emissions measured in the main plots. Finally, N₂ emissions were upscaled
254 to the plot scale by the area ratio bed:furrow = 1:1 at the Burdekin site and bed:furrow = 1:2 at the Mackay
255 site. Cumulative N₂ emissions were calculated by the sum of daily upscaled N₂ emissions for each plot over
256 the whole crop growing season.

257 Modelling of RN₂O and gap-filling of *Ndff*N₂ were conducted by fitting generalised additive mixed
258 models (GAMMs), using a package “mgcv” (Wood, 2011) and detailed in Supporting Information S1.7.
259 Briefly, GAMMs can quantify non-linear relationships without specifying the functional forms (De Rosa
260 et al., 2020; Dorich et al., 2020), which were used to analyse RN₂O in response to soil variables and *Ndff*
261 N₂ in response to days after fertilisation (DAF) and N rates. Furthermore, GAMMs allow the use of (i) the
262 beta family suitable to model proportions ranging from 0 to 1 and (ii) random factors to handle repeated
263 measurements.

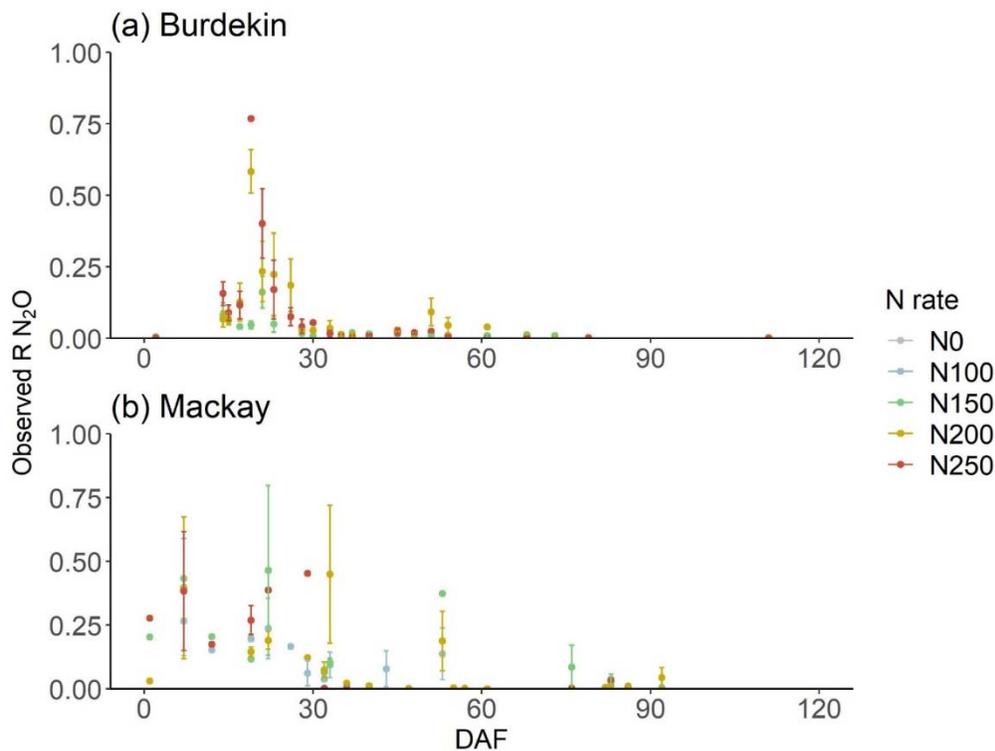
264 Effects of the sites, N fertiliser treatments and bed/furrow positions on RN₂O and N₂ emissions as
265 well as fertiliser-derived N₂O+N₂ in the proportion of the N fertiliser applied and the N fertiliser lost were
266 examined by fitting generalised linear (mixed) models, using packages “lme4” (Bates et al., 2015) and
267 “mgcv” (Wood, 2011). The beta family was specified for RN₂O and the proportions of fertiliser-derived
268 N₂O+N₂ and the gamma family for N₂ together with chamber/plot as a random factor in the case of daily
269 variables. To establish the response of cumulative N₂O+N₂ losses to N rates, (generalised) linear models
270 were fitted for each site.

271

272 **3 Results**273 3.1 Daily RN_2O and N_2 emissions

274 Daily RN_2O observed ranged from < 0.01 to 0.768 (Fig. 1) during ~ 120 DAF of the measurement
 275 period, peaking at values > 0.25 within 30 DAF at the Burdekin and within 60 DAF at the Mackay site. For
 276 the remainder of the measurement period, RN_2O stayed below 0.1. The range of observed RN_2O averaged
 277 for each N rate was 0.030–0.092 at the Burdekin site, smaller than 0.082–0.189 at the Mackay site (Table
 278 2). Overall, the observed daily RN_2O correlated positively with the N fertiliser rates (Table 2).

279



280

281 **Figure 1** Observed RN_2O near the band in the micro plots over the measurement period at N rates of 100,
 282 150, 200 and 250 kg N ha^{-1} at the Burdekin (a) and Mackay (b) sites. Points and error bars indicate mean
 283 values and standard errors

284

285

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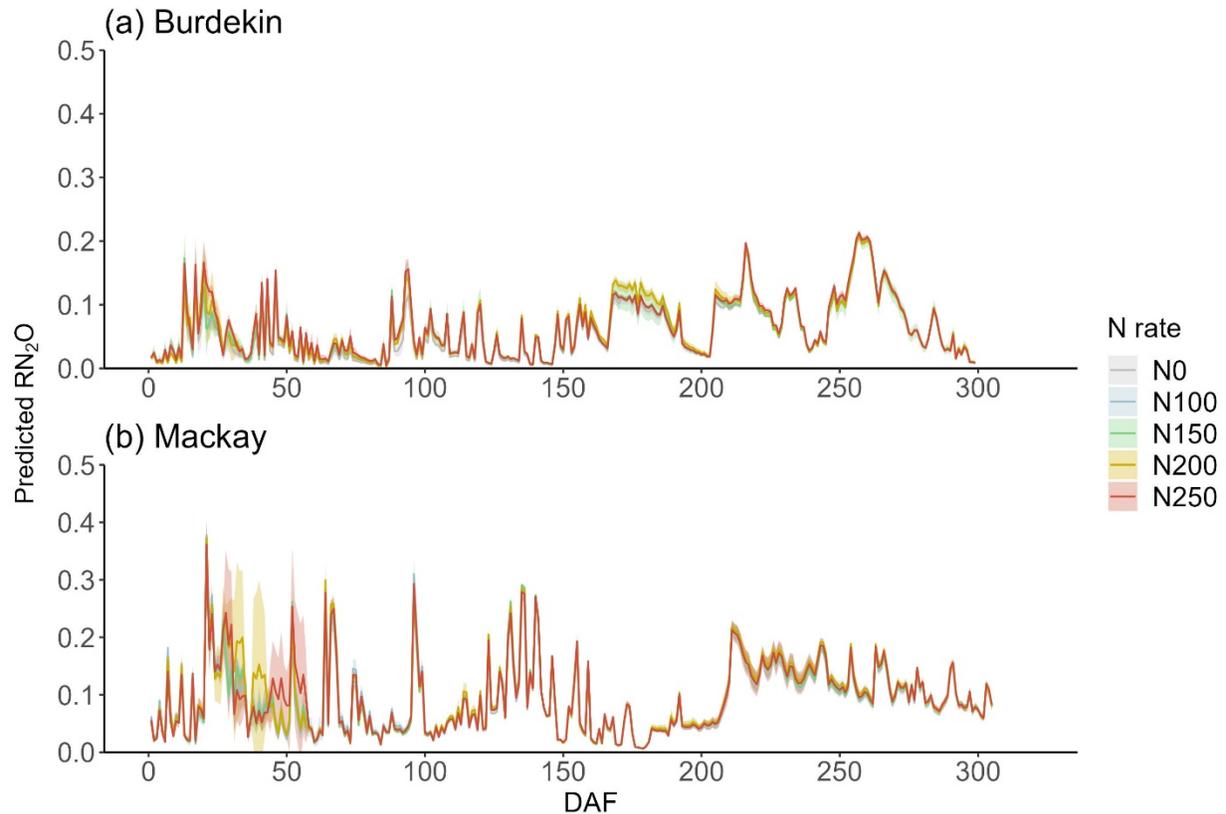
287 **Table 2** The RN₂O observed daily, RN₂O predicted daily for bed and furrow positions and RN₂O calculated
 288 with cumulative N₂O and N₂ emissions in response to N rates ranging from 0 to 250 kg N ha⁻¹, sites and
 289 positions

Site	N rate	Observed RN ₂ O	Predicted RN ₂ O		RN ₂ O at cumulative
			Bed	Furrow	
Burdekin	0		0.054 ± 0.001	0.054 ± 0.001	0.024 ± 0.002
	150	0.030 ± 0.01	0.060 ± 0.002	0.061 ± 0.002	0.032 ± 0.003
	200	0.092 ± 0.02	0.061 ± 0.001	0.063 ± 0.001	0.028 ± 0.002
	250	0.072 ± 0.02	0.061 ± 0.001	0.062 ± 0.001	0.035 ± 0.001
Mackay	0		0.091 ± 0.002	0.087 ± 0.002	0.050 ± 0.001
	100	0.082 ± 0.02	0.104 ± 0.003	0.087 ± 0.002	0.048 ± 0.007
	150	0.133 ± 0.04	0.097 ± 0.002	0.086 ± 0.002	0.051 ± 0.005
	200	0.093 ± 0.03	0.115 ± 0.003	0.087 ± 0.002	0.058 ± 0.003
	250	0.189 ± 0.06	0.109 ± 0.003	0.087 ± 0.002	0.047 ± 0.007
P value					
Site		< 0.001	< 0.001	< 0.001	
N rate		0.006	< 0.001	0.121	
Position			< 0.001		

290

291 Fitting the RN₂O observed near the fertiliser band in the micro plots using the GAMM with Site,
 292 soil temperature, WFPS, soil NH₄⁺ and NO₃⁻ contents and CO₂ emissions as predictors showed 51.7% of
 293 deviance explained and 0.151 of root mean square error (RMSE). The predicted RN₂O was larger at the
 294 Mackay site compared to the Burdekin site (P < 0.001) as well as on the bed compared to the furrow position
 295 (P < 0.001) (Table 2). The predicted RN₂O increased with increasing N rates (P < 0.001) (Table 2), which
 296 was apparent within 50 DAF (Fig. 2). The predicted RN₂O showed larger values during the late crop
 297 growing season compared to < 90 DAF (Fig. 2).

298



299

300

301 **Figure 2** Daily RN₂O predicted over the crop growing season across N rates 0, 100, 150, 200 and
 302 250 kg N ha⁻¹ at the Burdekin (a) and Mackay (b) sites. Lines and shaded areas indicate predicted mean
 303 values and 95% confidence intervals

303

304

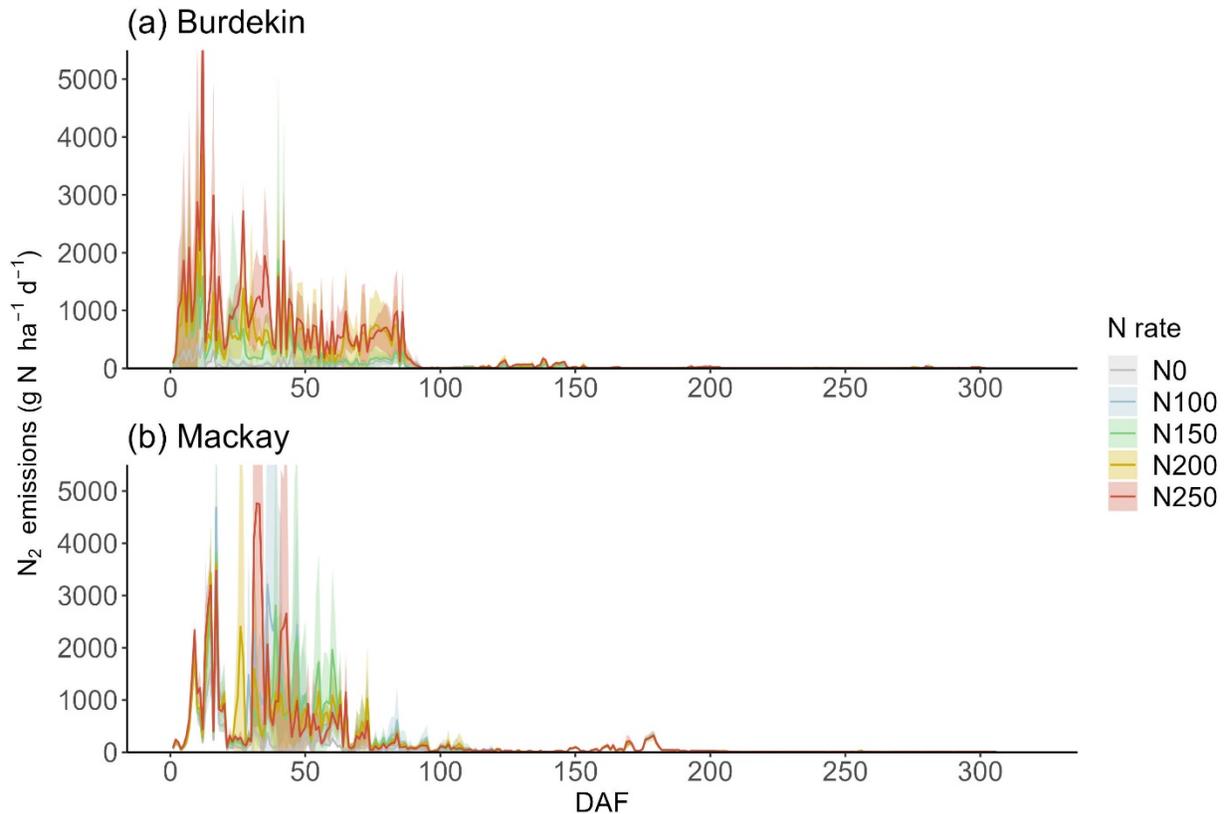
305

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308

Daily N₂ emissions reached up to 5 kg N ha⁻¹ d⁻¹ within 50 DAF and stayed elevated for
 approximately 100 DAF with minor emissions for the remainder of the season (Fig. 3). Daily N₂ emissions
 increased with increasing N rates ($P < 0.001$) and were on average larger at the Mackay site compared to
 the Burdekin site ($P < 0.001$).



309

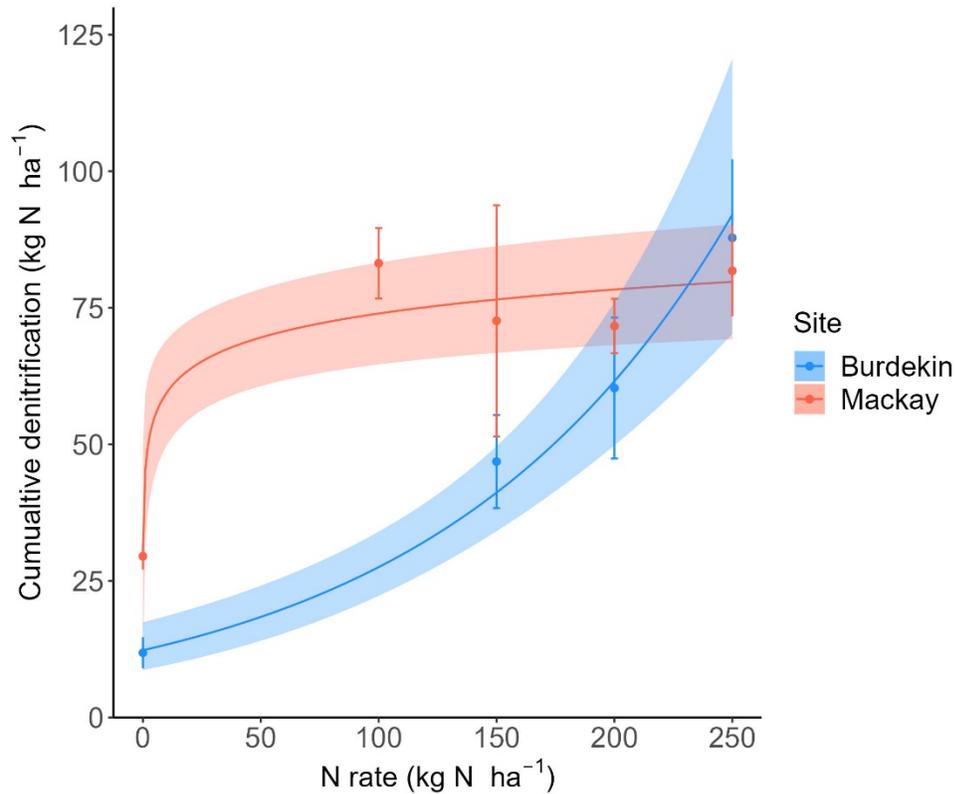
310 **Figure 3** Daily N_2 emissions estimated over the crop growing season at N rates of 0, 100, 150, 200 and
 311 250 $kg\ N\ ha^{-1}$ at the Burdekin (a) and Mackay (b) sites. Lines and shaded areas indicate predicted mean
 312 values and 95% confidence intervals, respectively

313

314 3.2 Cumulative denitrification losses (N_2O+N_2)

315 Cumulative denitrification losses (N_2O+N_2) for the whole growing season increased exponentially
 316 from 11.9 ± 2.9 to $87.8 \pm 14.4\ kg\ N\ ha^{-1}$ with increasing N fertiliser rates from 0 to 250 $kg\ N\ ha^{-1}$ at the
 317 Burdekin site (Fig. 4). At the Mackay site, cumulative N_2O+N_2 emissions increased from $29.5 \pm 2.5\ kg\ ha^{-1}$
 318 ¹ in the unfertilised treatment to a range from 71.7 ± 5.0 to $83.2 \pm 6.5\ kg\ N\ ha^{-1}$ observed across N rates
 319 from 100-250 $kg\ N\ ha^{-1}$, with no differences between N fertilised treatments (Fig. 4). Overall, cumulative
 320 N_2O+N_2 emissions were larger at the Mackay site compared to the Burdekin site ($P = 0.027$). Cumulative
 321 emissions of N_2O accounted for 2.4–3.5% of N_2O+N_2 emissions at the Burdekin site, which was lower than
 322 4.8–5.8% at the Mackay site ($P < 0.001$) (Table 2).

323

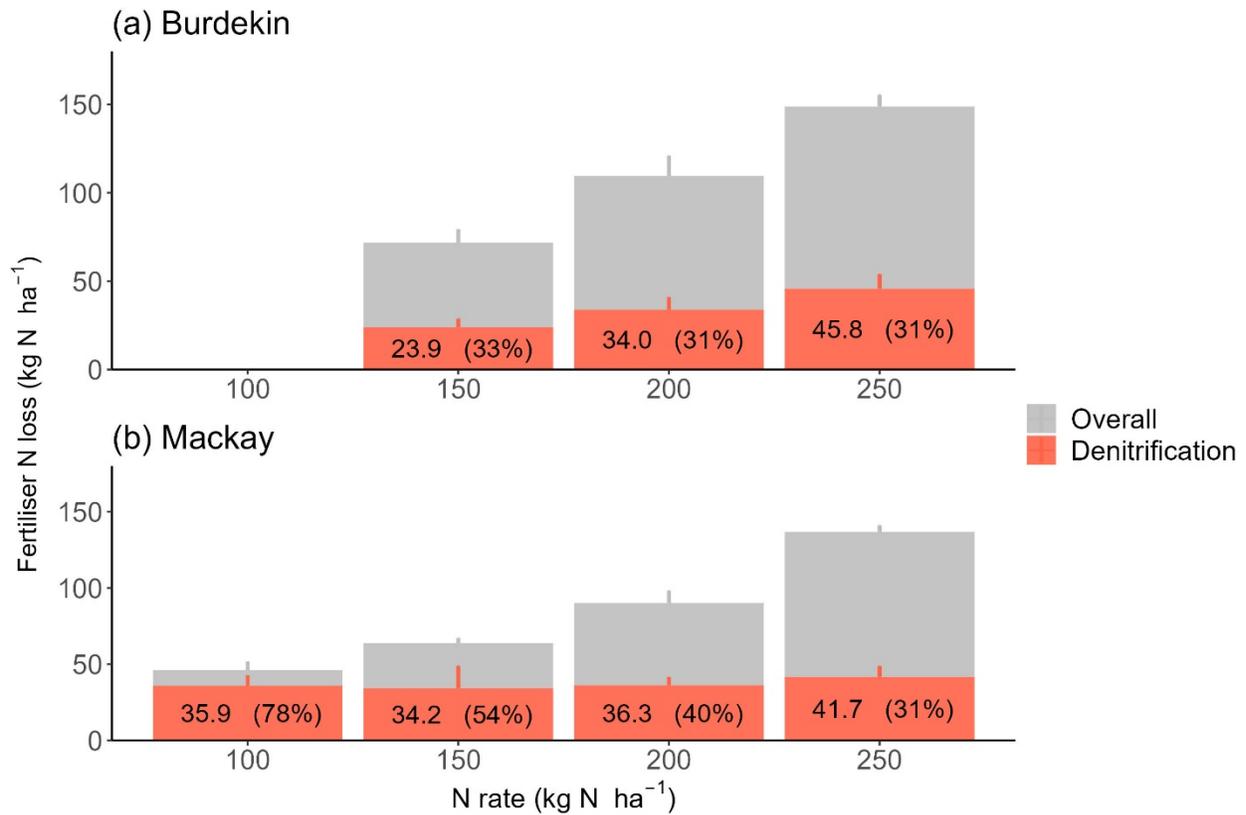


324
 325 **Figure 4** Cumulative denitrification losses over the crop growing season in response to N fertiliser rates at
 326 the Burdekin (blue) and Mackay (red) sites. Points and error bars indicate mean values and standard errors.
 327 Lines and shaded areas indicate fitted curves and 95% confidence intervals, respectively
 328
 329

330 3.3 Fertiliser N contribution to denitrification losses (N_2O+N_2)

331 Contribution of N fertiliser to N_2 emissions was high within 50 DAF, accounting for > 50% and
 332 70% of N_2 emissions at the Burdekin and at the Mackay site, respectively, with a diminishing contribution
 333 for the rest of the measurement period (Fig. S1). Of the cumulative N_2 emissions, 51.0–57.5% and 43.1–
 334 51.0% were derived from fertiliser N at the Burdekin and Mackay sites, respectively. Cumulative fertiliser-
 335 derived N_2O+N_2 emissions ranged from 23.9 to 45.8 and 34.2 to 41.7 kg N ha⁻¹ at the Burdekin and Mackay
 336 sites, respectively (Fig. 5). Cumulative fertiliser-derived N_2O+N_2 emissions accounted for 30.8–33.3% and
 337 30.5–77.5% of the overall fertiliser ¹⁵N loss, at the Burdekin and Mackay sites, respectively (Fig. 5). The
 338 percentage of fertiliser N lost as N_2O+N_2 was larger at the Mackay site ($P = 0.02$) and decreased with
 339 increasing N rates at both sites ($P = 0.009$). Contribution of fertiliser N to N_2O+N_2 emissions accounted for
 340 15.9–18.3% and 16.7–35.9% of the N applied at the Burdekin and Mackay sites, respectively.

341 Emissions of $\text{N}_2\text{O}+\text{N}_2$ derived from soil N in the fertilised treatments were 22.9–42.1 and 35.4–
 342 47.3 kg N ha^{-1} at the Burdekin and Mackay sites, respectively.
 343



344
 345 **Figure 5** Cumulative fertiliser-derived denitrification losses (red) in comparison to overall fertiliser ^{15}N
 346 loss (grey) in response to N fertiliser rates 100, 150, 200 and 250 kg N ha^{-1} at the Burdekin (a) and Mackay
 347 (b) sites. Error bars indicate standard errors
 348

349 4 Discussion

350 The unique combination of high-frequency N_2O and in-situ $\text{N}_2\text{O}/(\text{N}_2\text{O}+\text{N}_2)$ ratio (RN_2O)
 351 measurements using automated GHG monitoring systems and ^{15}N gas flux method together with GAMMs
 352 enabled us to quantify field-scale N_2O and N_2 emissions in response to N fertiliser rates in two sugarcane
 353 systems over the whole crop growing season. This method accounts for the dynamic nature of the RN_2O
 354 considering the overlapping effects of key drivers of N_2O and N_2 production, delivering robust estimates of
 355 N_2 emissions at the field scale. Furthermore, comparing fertiliser-derived $\text{N}_2\text{O}+\text{N}_2$ emissions to fertiliser
 356 ^{15}N loss allowed us to validate the estimated N_2 emissions at the cumulative scale. Applying this method
 357 across two intensively managed sugarcane systems showed a) $> 80 \text{ kg N ha}^{-1}$ lost as $\text{N}_2\text{O}+\text{N}_2$ over the

358 growing season, with b) emissions dominated by N_2 accounting for > 95% of N_2O+N_2 losses, and c) that
359 31-78% of ^{15}N fertiliser losses occurred in the form of N_2O+N_2 . The method proposed here can be used as
360 a blueprint approach to deliver seasonal denitrification estimates, targeting a key uncertainty in N budgets
361 of different agroecosystems.

362

363 4.1 Estimating N_2 emissions over the crop growing season using RN_2O

364 Daily The high temporal variability of observed RN_2O ranging from < 0.01 to 0.768 (Fig. 1)
365 emphasises the need to account for dynamic changes in RN_2O to estimate N_2 emissions. The use of GAMMs
366 in this study allowed us to express RN_2O as a function of soil water content, temperature, soil mineral N
367 content and CO_2 emissions, accounting for their effect on the RN_2O at both temporal and spatial scales (Fig.
368 2). Banding of N fertiliser on or beside the bed creates a distinct zone in and close to the band with high N
369 availability, decreasing towards the furrow. Direct measurements of RN_2O in the unfertilised furrow are
370 not possible with the ^{15}N gas flux method, as it requires the application of ^{15}N fertiliser, highlighting the
371 need for the GAMMs to estimate RN_2O accounting for changes in N availability in the furrow. Higher
372 values of RN_2O as a result of higher N-substrate availability are consistent with the increase in observed
373 RN_2O from the band with increasing N fertiliser rates (Table 2). This relationship is also shown by the
374 higher values of predicted RN_2O from the bed than the furrow at the Mackay site (Table 2), where the
375 application of a single N fertiliser band likely increased spatial differences in N availability as compared to
376 the Burdekin site with banding on both sides of the bed. Differences in RN_2O may be explained by
377 preferential NO_3^- reduction over N_2O in zones of high NO_3^- availability around the fertiliser band (Friedl
378 et al., 2020b; Senbayram et al., 2019). Since banding of N fertiliser is a common practice in intensively
379 managed cropping systems, accounting for its effects on RN_2O as demonstrated here is of therefore of great
380 importance to upscaling N_2 emissions.

381 It is noteworthy that in contrast to previous studies (Bizimana et al., 2022; Wang et al., 2020),
382 RN_2O data in the study presented here are based on field measurements, which removes the need for
383 measurements of the ratio between N_2O and N_2 using laboratory assays. In-situ measurements avoid a
384 potential bias due to the disturbance of the soil and the absence of plants in the laboratory incubation. An
385 incubation study using the soil samples from the Burdekin site without plants found much lower RN_2O <
386 0.03 across the whole measurement period compared to this study despite comparable ranges of soil factors
387 (Kirkby et al., personal communication). Both smaller (Bizimana et al., 2022) and larger (Wang et al., 2020)
388 N_2O emissions were reported under laboratory conditions compared to in-situ measurements, indicating an
389 inconsistent discrepancy in RN_2O between field and laboratory measurements. This discrepancy
390 emphasises the need for in-situ measurements as presented here. However, field measurements are likely

391 to show a higher degree of variability, which was reflected in 52% of deviance explained on average when
392 fitting GAMMs to the observed RN_2O with cross-validation. Fitting GAMMs to the entire dataset without
393 cross-validation resulted in 86% of deviance explained, comparable to the multivariate model of Wang et
394 al. (2020) which explained 92% of the variability of RN_2O . In this study, the cross-validated model by
395 replicate was used to extrapolate at both temporal and spatial scales. Setting the k-fold validation across
396 replicates considerably minimised the potential model overfitting observed when using the entire dataset
397 for model training (Dorich et al., 2020). Comparing the fertiliser-derived $\text{N}_2\text{O}+\text{N}_2$ with the overall fertiliser
398 ^{15}N loss allowed us to constrain the RN_2O modelling with GAMMs. This constraint at the cumulative scale
399 reduced the uncertainty in N_2 estimates, emphasising the advantage of in-situ N_2O and N_2 measurements
400 with the ^{15}N gas flux method combined with fertiliser ^{15}N recovery measurements.

401 Applying predicted values of RN_2O to high temporal-resolution N_2O measurements gave estimates of
402 daily N_2 emissions over the season (Fig 3). Similar to N_2O , the majority of N_2 emissions occurred within
403 100 days after fertilisation, which is consistent with peaks in soil NO_3^- availability (Takeda et al., 2021a).
404 High NO_3^- substrate availability for denitrification together with limited O_2 in the soil following intense
405 rainfall and/or irrigation promoted N loss in the form of N_2 , which accounted for > 95% of total $\text{N}_2\text{O}+\text{N}_2$
406 emissions over the crop growing season (Table 2). On the other hand, the average of observed RN_2O
407 without temporal and spatial upscaling demonstrated up to 9% and 19% of $\text{N}_2\text{O}+\text{N}_2$ losses as N_2O (Table
408 2). This discrepancy indicates an underestimation of N_2 emissions if the average of observed RN_2O was
409 directly applied to N_2O emissions. Using fixed RN_2O values from measurements with limited coverage of
410 environmental conditions may therefore lead to a bias in estimated N_2 emissions. In turn, this difference
411 emphasises the importance to include a range of soil conditions covering the spatio-temporal variability
412 observed within a cropping system and season when using the ratio between N_2O and N_2 to upscale N_2
413 emissions to the field scale.

414

415 4.2 Denitrification as a major N loss pathway in intensive sugarcane systems

416 Total $\text{N}_2\text{O}+\text{N}_2$ emissions over the season exceeded 80 kg N ha^{-1} at both sites (Fig. 4). Denitrification
417 losses have been regarded as a major portion of N budgets in intensively managed sugarcane systems (Bell
418 et al., 2014) but emissions were only measured from the fertiliser band in short-term trials (Warner et al.,
419 2019; Weier et al., 1996; Weier et al., 1998). The lack of seasonal estimates of denitrification losses in
420 sugarcane hinders the comparison to the range of $\text{N}_2+\text{N}_2\text{O}$ emissions observed in the study presented here.
421 In a simulation study, Thorburn et al. (2017) predicted denitrification losses up to 50 kg N ha^{-1} with N
422 fertiliser rates up to 200 kg N ha^{-1} from Australian sugarcane systems. This range is substantially lower
423 than the $\text{N}_2+\text{N}_2\text{O}$ emissions from both sites. Even though denitrification rates are subject to specific site

424 and environmental conditions, predictions of denitrification losses in biogeochemical models rely mostly
425 on N_2O data. The lack of N_2 data hinders the validation of overall rates, and changes in N_2O may be caused
426 by a change in denitrification rate and/or RN_2O (Del Grosso et al., 2020). Our estimates of seasonal $\text{N}_2\text{O}+\text{N}_2$
427 losses not only provide experimental evidence that denitrification is a major pathway of N loss from
428 intensively managed sugarcane systems, but also the opportunity to test and validate the representation of
429 denitrification in biogeochemical models.

430 Cumulative $\text{N}_2\text{O}+\text{N}_2$ losses responded exponentially to N fertiliser rates at the Burdekin site but
431 did not increase across the fertilised treatments at the Mackay site (Fig.4), indicating other factors but N
432 availability limited denitrification at the site. Mackay experienced less rainfall and received less irrigation
433 than the Burdekin site in the critical time window three months after fertilisation. Furthermore, irrigation
434 was applied via overhead sprinklers in Mackay, compared to furrow (flood) irrigation in Burdekin.
435 Considering the sandier soil texture (Table 1) at the Mackay site, the differences in management and rainfall
436 indicate an increased frequency of aerobic conditions in the soil at the Mackay site compared to the
437 Burdekin site (Takeda et al., 2022), limiting the response of denitrification to N rate. Regardless, relatively
438 large $\text{N}_2\text{O}+\text{N}_2$ losses $> 50 \text{ kg N ha}^{-1}$ were consistently observed at high N rates above the recommended N
439 rate ($\geq 200 \text{ kg N ha}^{-1}$) across the sites (Fig. 4), suggesting increased N substrate availability for N losses via
440 denitrification.

441 Denitrification was dominated by N_2 emissions (Table 2) and accounted for up to 33% and 78% of
442 the overall fertiliser ^{15}N loss (Fig. 5), showing that a large fraction of N fertiliser loss occurs in the form of
443 environmentally benign N_2 . The relative contribution of $\text{N}_2\text{O}+\text{N}_2$ losses to overall fertiliser ^{15}N loss however
444 decreased with increasing N rates (Fig. 5). This suggests increasing significance of other reactive N loss
445 pathways including ammonia volatilisation, leaching and runoff with increasing N rates, as denitrification
446 may become limited by factors other than N availability. Losses of $\text{N}_2\text{O}+\text{N}_2$ accounted for a smaller
447 proportion of fertiliser ^{15}N loss at the Burdekin site compared to the Mackay site, which is consistent with
448 furrow irrigation and severe flooding events likely causing greater losses of N fertiliser via leaching and
449 runoff at the Burdekin site. Loss of N via runoff and leaching from Australian sugarcane systems is currently
450 estimated to account for 46–65% of the total dissolved inorganic N load to the Great Barrier Reef (GBR)
451 (Bartley et al., 2017). Increasing N losses via runoff and leaching with increasing N rates have been mostly
452 demonstrated by simulation studies (Reading et al., 2019; Thorburn et al., 2017; Vilas et al., 2022). The
453 study presented here shows that even though a large proportion of N fertiliser loss from sugarcane systems
454 occurs as environmentally benign N_2 , more N is lost via environmentally harmful pathways of N loss
455 including ammonia volatilisation, leaching and runoff as N rates increase. These findings suggest that even
456 if $\text{N}_2\text{O}+\text{N}_2$ losses aren't responding to increasing N rates, environmental costs of sugarcane production are
457 likely to show a non-linear response to N fertiliser.

458 The large amounts of soil N contributing to $\text{N}_2\text{O}+\text{N}_2$ across N rates (23–47 kg N ha⁻¹) corroborate
459 the importance of mineralised N for N cycling in sugarcane soils (Takeda et al., 2022). These exports of
460 soil N, together with the plant N uptake derived from soil N (67–122 kg N ha⁻¹), largely exceeded the
461 fertiliser ¹⁵N remaining in the soil (40–60 kg N ha⁻¹) across N rates, even when accounting for N in the crop
462 residue which can be returned (~ 60 kg N ha⁻¹). This negative balance demonstrates the ineffectiveness of
463 increasing N fertiliser rates to compensate for soil N depletion. Higher rates of banded N fertiliser
464 application with the aim of carrying surplus N into subsequent seasons (“N-bank” concept) were reported
465 to be associated with high risks of N losses under wet conditions in sub-tropical sorghum systems (Rowlings
466 et al., 2022). The N balance in the study here suggests long-term soil N depletion despite high N inputs in
467 intensively managed sugarcane systems. Together with the non-linear responses of $\text{N}_2\text{O}+\text{N}_2$ losses and their
468 contribution to fertiliser ¹⁵N loss, these results indicate that increasing N fertiliser rates result in lower NUE
469 and higher environmental costs but also don’t prevent soil N mining. Maintaining crop productivity while
470 reducing environmental impacts therefore requires N fertiliser rate strategies integrated with additional
471 measures such as the use of enhanced efficiency fertilisers (Connellan & Thompson, 2022) and rotation
472 with legume crops (Otto et al., 2020).

473

474 4.3 Extrapolating RN_2O to a wider range of cropping systems towards the global N budget

475 Denitrification losses have been assumed to account for a significant portion of the global terrestrial
476 N budget despite uncertainties due to limited evaluation at the plot scale (Bouwman et al., 2013; Houlton
477 & Bai, 2009; Scheer et al., 2020). Given that measurements of N_2O emissions are relatively well established
478 and conducted globally, the values of RN_2O play a critical role in estimating the global N budget.
479 Nevertheless, agricultural systems or crop management practices have not been differentiated in most of
480 the reports to date. For example, Scheer et al. (2020) showed a mean RN_2O of 0.11 for agricultural soils
481 and 0.02 for wetlands by summarising the previously reported RN_2O values. The values of RN_2O 0.024–
482 0.058 (Table 2) based on the cumulative N_2 and N_2O emissions in the study presented herein are indicative
483 of intensively managed cropping systems with high N and water inputs. Compared to the range given by
484 Scheer et al. (2020), this would shift denitrification losses from agricultural soils towards the upper end of
485 the current uncertainty range. The method presented in this study provides a unique tool to estimate seasonal
486 denitrification losses accounting for spatial and temporal variability in intensive agroecosystems. This is
487 therefore well suited to generate data that can close the gap in current N budgets, helping to encourage
488 actions to mitigate N pollution.

489 Refinements of the global N budget require the effects of cropping systems and site conditions on
490 RN_2O to be incorporated. Within this study, the larger RN_2O at the Mackay site (Table 2) may reflect the

491 effect of the low pH (4.1) compared to the Burdekin site (pH 6.9) (Table 1) shifting the ratio towards N₂O
492 (Dannenmann et al., 2008; Russenes et al., 2016; Šimek & Cooper, 2002). The sandier soil texture may
493 have led to better drainage and larger gas diffusivity at the Mackay site, contributing to the larger RN₂O
494 (Friedl et al., 2017). On the other hand, GCTB management at the Mackay site possibly promoted
495 completion of denitrification and thus reduced RN₂O by preventing evaporation and thus promoting
496 anaerobic conditions (Weier et al., 1993). Accounting for these effects individually to generalise RN₂O
497 estimates requires further data collection across a wide range of environmental conditions such as cropping
498 systems, management practices, soil pH and texture. Controlling environmental factors in laboratory assays
499 can aid in disentangling such overlapping effects, highlighting the need to integrate both laboratory and in-
500 situ measurements of N₂O and N₂ in future research. Generalised estimation of RN₂O covering a wider
501 range of cropping systems and environmental conditions, together with increasing robust in-situ
502 measurements of N₂O emissions, will aid the accuracy of global N budget estimates as well as the
503 identification of hot spots of denitrification losses.

504

505 **5 Conclusions**

506 This is the first study establishing the response of cumulative denitrification losses (N₂O+N₂) to N
507 fertiliser rates over the whole crop growing season at the plot scale based on in-situ measurements. We
508 propose the integration of in-situ RN₂O with the ¹⁵N gas flux method, high-frequency N₂O with an
509 automated GHG monitoring system and fertiliser ¹⁵N recovery measurements as a novel and robust method
510 applicable to a wide range of cropping systems to quantify cumulative denitrification losses under field
511 conditions. In contrast to previous approaches, this method accounts for both temporal as well as spatial
512 variability of RN₂O and includes in-situ data for validation of denitrification losses at the cumulative scale.
513 The use of this method demonstrated that seasonal denitrification losses were dominated by N₂ emissions,
514 and accounted for 31–78% of total N fertiliser losses, providing critical evidence for its significance as an
515 N loss pathway from sugarcane systems. The non-linear response of cumulative denitrification losses to
516 increasing N rates, with > 80 kg N ha⁻¹ emitted as N₂ and N₂O emphasises the agronomic and environmental
517 inefficiency of excessive N fertiliser application. Even though a large proportion of N fertiliser loss
518 occurred as environmentally benign N₂, more N was lost via environmentally harmful pathways including
519 ammonia volatilisation, leaching and runoff with increasing N rates. These findings highlight that excessive
520 N rates not only increase agronomic inefficiencies, but also the environmental footprint of intensive
521 sugarcane production. This research delivers critical data targeting key uncertainties in biogeochemical
522 models and will aid parameterisation and improvement of denitrification algorithms, advancing our
523 understanding of N cycles across scales. These improvements are urgently needed to develop N fertiliser

524 rate strategies integrated with soil fertility management and simulate their long-term impacts, to maintain
525 crop productivity while reducing environmental impacts of intensive agroecosystems.

526

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533

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Global Biogeochemical Cycles

Supporting Information for

Denitrification losses in response to N fertiliser rates – a synthesis of high temporal resolution N₂O, in-situ ¹⁵N₂O and ¹⁵N₂ measurements and fertiliser ¹⁵N recoveries in intensive sugarcane systems

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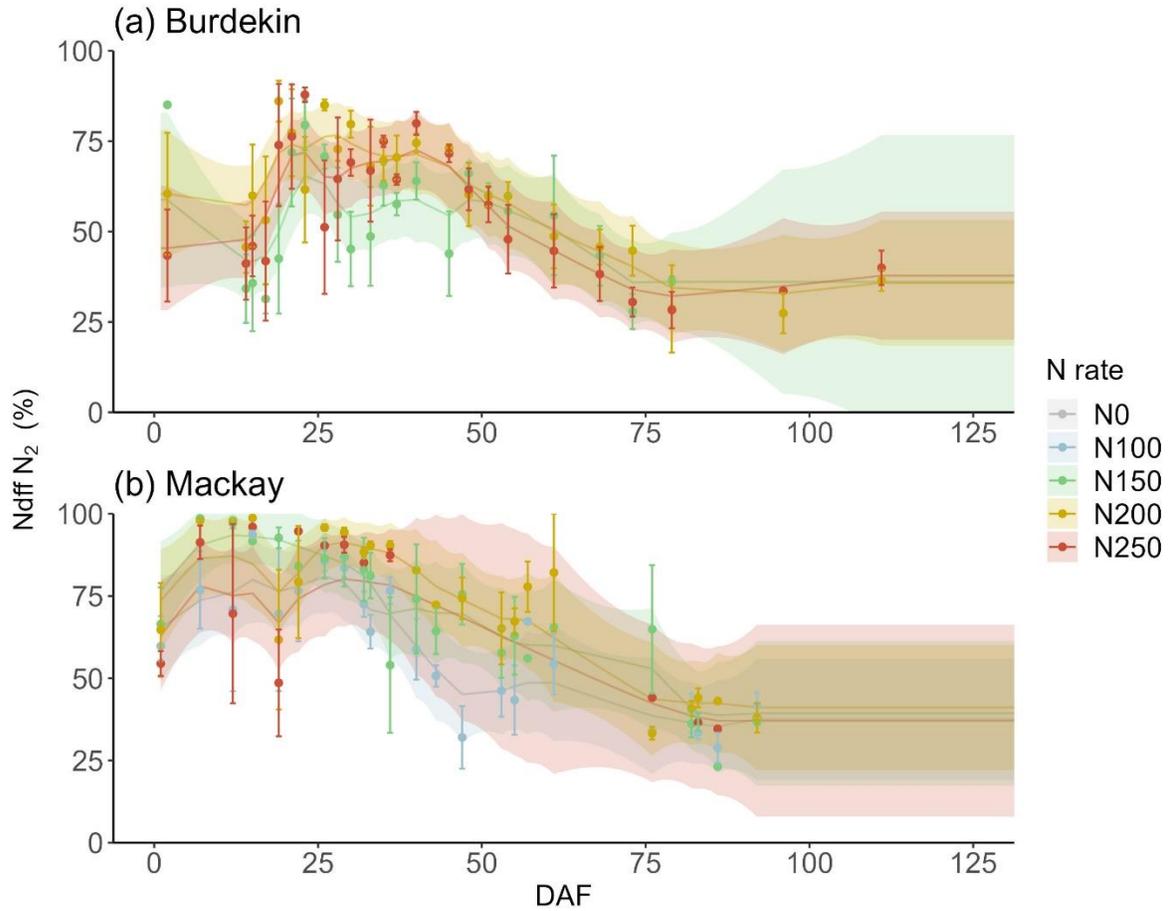


Figure S1 Relative contribution of fertiliser N to N₂ emissions (Ndff N₂) on the fertiliser band over the measurement period across N rates 100, 150, 200 and 250 kg N ha⁻¹ at the Burdekin (a) and Mackay (b) sites. Points and error bars indicate mean values and standard errors. Lines and shaded areas indicate fitted curves and 95% confidence intervals

Denitrification losses in response to N fertiliser rates – a synthesis of high temporal resolution N₂O, in-situ ¹⁵N₂O and ¹⁵N₂ measurements and fertiliser ¹⁵N recoveries in intensive sugarcane systems

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S1 Materials and methods

S1.1. Automated chamber system

Acrylic static chambers (0.5 m × 0.5 m × 0.15 m) were mounted on stainless steel frames inserted 10 cm into the soil. The lids of the chambers were opened and closed automatically with pneumatic pistons, and four chambers were closed at one time. Air samples were taken sequentially from each closed chamber, followed by a single-point known standard (Air Liquide, Dallas, TX, USA) of 0.5 ppm N₂O and 800 ppm CO₂ for calibration and drift correction (i.e., after every fourth sample). Changes in headspace N₂O and CO₂ concentration after chamber closure were measured with a gas chromatograph (SRI 8610C, SRI Instruments, Inc., Las Vegas, NV, USA) equipped with a ⁶³Ni electron capture detector (ECD) for N₂O analysis while an infrared gas analyser (LI-820, LI-COR Biosciences, Lincoln, NE, USA) was used for measurements of CO₂. In total, each chamber was sampled four times (every 15 min) over 60 min. This enabled up to eight single flux rates to be determined per chamber and day. The detection limit of the system was 2.0 μg N₂O-N m⁻² h⁻¹ for N₂O. Hourly N₂O and CO₂ fluxes were calculated from the slope of the linear change in gas concentration during the closure period (60 min) and corrected for air temperature, atmospheric

pressure and the ratio of chamber volume to surface area as described in detail by Grace et al. (2020). The coefficient of determination (R^2) for the linear regression was calculated and used as a quality check for the measurement. Flux rates were discarded if R^2 was < 0.80 .

S1.2. Manual chamber system

Manual gas sampling was conducted with sealed polyethylene chambers (0.5 m × 0.5 m × 0.15 m). Headspace gas samples were taken between 0900 and 1200 H analogous to the sampling regime of the GHG system, connecting a syringe to a 2-way luer-lock tap installed in the lid of the chamber. Gas samples were injected into pre-evacuated 12 mL glass vials with a double wadded Teflon/silicone septa cap (Labco Exetainer[®], UK) and analysed for N₂O and CO₂ using a Shimadzu GC-2014 Gas Chromatograph (Shimadzu, Kyoto, Japan) at the Central Analytical Research Facility of the Queensland University of Technology, Australia.

S1.3. Gas sampling procedure in the micro plots

Gas samples were taken 1–3 times a week between 0900-1200H with higher frequency after N fertiliser application and irrigation or rainfall events. At the Burdekin site, polyethylene chambers with a headspace height of 19.93 cm were placed on the steel frames, ensuring airtight conditions. Headspace gas samples (20 ml) were taken by connecting a syringe to a 2-way luer-lock tap installed in the lid of the chamber. Gas samples were then injected into a pre-evacuated 12 ml glass vial with a double wadded Teflon/silicone septa cap (Labco, UK). Headspace gas samples were collected at 0, 60 and 180 min after closure. At the Mackay site, airtight polyethylene chambers with a headspace height of 15 cm were placed on the base and connected to a battery-powered sampling unit. The chambers were automatically closed at sampling events according to the pre-programmed schedule and gas samples were collected at 0, 60 and 180 min after the chamber closure. Headspace gas samples (20 ml) were automatically taken and injected into the pre-evacuated 12 ml glass vials sequentially installed on a belt in the sampling unit.

S1.4. The ¹⁵N gas flux method

The ¹⁵N enrichment of the NO₃⁻ pool undergoing denitrification (a_p N₂ and a_p N₂O) and the fraction of N₂ and N₂O emitted from this pool (f_p) were calculated following the equations given by Spott et al. (2006)

$$f_p = \frac{a_m - a_{bgd}}{a_p - a_{bgd}} \quad [S1]$$

where a_{bgd} is the ^{15}N abundance of the atmospheric background and a_m is the measured ^{15}N abundance of N_2 from headspace gas samples taken 0 and 180 minutes after closure, respectively. Both a_{bgd} and a_m are calculated as

$$a_i = \frac{{}^{29}\text{R} + 2 * {}^{30}\text{R}}{2 * (1 + {}^{29}\text{R} + {}^{30}\text{R})} \quad [\text{S2}]$$

and ${}^{30}\text{x}_m$ is the measured fraction of m/z 30 in N_2 :

$${}^{30}\text{x}_m = \frac{{}^{30}\text{R}}{(1 + {}^{29}\text{R} + {}^{30}\text{R})} \quad [\text{S3}]$$

If only ${}^{29}\text{R}$ was > the detection limit (DL), f_p was calculated as

$$f_p = \frac{1}{1 - \frac{{}^{29}\text{R}(1 - a_p)^2 - 2a_p(1 - a_p)}{{}^{29}\text{R}(1 - a_{bgd})^2 - 2a_{bgd}(1 - a_{bgd})}} \quad [\text{S4}]$$

using a_p N_2O assuming that N_2 and N_2O were derived from the same NO_3^- pool undergoing denitrification.

S1.5. Plant and soil sampling

All the sugarcane plants in the 2.0 m section were sampled by cutting at ground level and the trash on the ground in the section was also collected, followed by fresh weight measurements. Roughly six stalks from the middle of the 2.0 m section were chosen and separated into tops (above 7th node), stalk and dead leaves. Two green leaves at the 3rd node (L+3) were sampled from the plants from the 2-m section (L3_CR) and the adjacent row (L3_AR) to estimate the fertiliser ^{15}N recovery in the plant in the adjacent rows. After sampling aboveground biomass, remaining stools and major roots were sampled from a 0.5 × 0.5 m square in the middle of the 2.0 m section by digging down to 0.15 m depth. Roots were washed to remove the attached soil and separated into stool and roots. The tops, stalks, dead leaves and trash samples were coarsely ground, subsampled (about 10% on a fresh weight basis) and weighed to calculate partitioning ratios. The plant subsamples, L+3 samples, root samples and stool samples were oven-dried at 60 °C and dry weights were recorded. The stalk samples were further dried with a vacuum oven at 40 °C for 48 hours before fine grinding to avoid aggregation due to sugar.

Soil samples were taken at three to four points between the bed and furrow centres using a soil corer and a post-hole driver down to 1.0 m. At the Burdekin site, the sampling points were 0, 0.25, 0.50, 0.75 m away from the bed centre. At the Mackay site, those were 0, 0.12, 0.40, 0.80 m away from the bed centre in 100N and 250N and the three points except for 0.12 m away from the bed centre in 150N and 200N. Each soil core sample of 1.0 m was separated into 0–0.2, 0.2–0.4, 0.4–0.7, 0.7–1.0 m soil depths and subsamples (about 100 g) were taken.

S1.6. ¹⁵N calculation in plant and soil

The percentage of N in the individual plant, soil, N₂O and N₂ samples ('sinks') derived from ¹⁵N-labelled fertiliser was calculated from

$$Ndff (\%) = \frac{\%^{15}N \text{ excess of sink}}{\%^{15}N \text{ excess of fertiliser}} \times 100 \quad [S5]$$

where the %¹⁵N excess used for all sources and sinks was the ¹⁵N abundance less an adjustment of %¹⁵N measured for the corresponding plant and soil samples in the 0N plots for background enrichment or the natural abundance (0.0036765) for N₂O and N₂ samples.

Fertiliser N recovered in each plant part (*PlantFN_i*) was calculated from

$$PlantFN_i = Biomass_i \times N \text{ content}_i \times \frac{Ndff_i}{100} \quad [S6]$$

where i indicates the plant part. Fertiliser N recovered in tops, stalk, dead leaves and trash were summed to calculate the fertiliser N recovered in the aboveground biomass (*PlantFN_{AG}*) in the centre row in each plot. Fertiliser N recovered in the belowground biomass (*PlantFN_{BG}*) in the centre row was calculated by summing the fertiliser N recovered in stool and roots. Fertiliser N recovered in the adjacent rows (*PlantFN_{AR}*) was calculated in each plot by multiplying the total N uptake in the centre row and the Ndff calculated from L3_AR sample. Fertiliser ¹⁵N uptake was then calculated as follows:

$$Fertiliser \ ^{15}N \text{ uptake} = (PlantFN_{AG} + PlantFN_{BG} + PlantFN_{AR} \times 2) \quad [S7]$$

To calculate fertiliser ¹⁵N recovery in the soil, the distribution of ¹⁵N fertiliser across the soil profile between bed and furrow was first analysed to account for the spatial variation caused by banding N fertiliser, following Takeda et al. (2021). (i) Fertiliser N recovered in the soil (*SoilFN_{ij}*) of each sample (four points from the bed centre and four depths down to 1.0 m per plot at the centre of the soil layer as the average) was calculated from

$$SoilFN_{ij} = BD_{ij} \times N \text{ content}_{ij} \times \frac{Ndff_{ij}}{100} \quad [S8]$$

where i and j indicate the sampling point and depth per plot. At the Mackay site, the missing *SoilFN_{ij}* values at the sampling point 0.12 m away from the bed centre in 150N and 200N treatments were substituted with the estimated values from regressions by N rates at each soil depth. (ii) The fertiliser ¹⁵N recovered in the soil (*SoilFN_{ij}*) was then interpolated across one side of the sugarcane row (0.75 m width, x-axis) and down to a soil depth of 1.0 m (y-axis) for each plot by fitting a thin-plate spline using a package "mgcv" (Wood, 2011). (iii) The interpolated fertiliser N recovered in the soil

(SoilFN_{xy}) was integrated and divided by the area of one side of the row (*Area*, 0.75 m²) to express fertiliser N recovered in kg N ha⁻¹. Fertiliser ¹⁵N recovery in the soil as a proportion to the N rate was then calculated as follows:

$$\text{Fertiliser } ^{15}\text{N recovery in the soil} = \frac{\sum \text{SoilFN}_{xy}}{\text{Area}} \quad [\text{S9}]$$

Overall fertiliser ¹⁵N loss was calculated by the difference between the N applied and fertiliser ¹⁵N recovered in the soil and plant.

Calculations for fertiliser ¹⁵N recovered in N₂O emissions are detailed in Takeda et al. (2022) and those for N₂ emissions followed the same procedure. Briefly, the proportion of N₂ emissions derived from fertiliser (*Ndff* N₂) at the fertiliser band was first calculated by Equation [S5]. Then, *Ndff* N₂ was gap-filled per N rate over the crop growing season on a daily basis at each site, which were then applied to daily N₂ emissions per plot in the main plots to calculate fertiliser-derived N₂ as follows:

$$\text{Fertiliser derived N}_2 \text{ emissions}_{i,j} = \text{N}_2 \text{ emissions}_{i,j} \times \frac{\text{Ndff N}_{2i,j}}{100} \quad [\text{S10}]$$

where *i* and *j* indicate days after fertilisation and chamber position (i.e. bed or furrow). At the Burdekin site, *Ndff* N₂ at the fertiliser band in micro plots was used for both bed and furrow chambers because both chambers covered the fertiliser band. At the Mackay site, *Ndff* N₂ at the furrow was assumed to be zero because the furrow chamber did not cover the fertiliser band. Contribution of fertiliser N to N₂ was calculated by area-weighted sum of fertiliser-derived N₂ emissions over the crop growing season.

S1.7. Use of generalised additive mixed models

The use of generalised additive mixed models (GAMMs) can quantify non-linear relationships without specifying the functional forms and GAMMs further allow repeated measurements. The distribution of RN₂O and *Ndff* N₂ was assumed Beta distributed, which is suitable to fit variables taking values between 0 and 1, and the logit function was specified as the link function, respectively. A dispersion parameter can further be specified in Beta regressions, which was useful to fit the RN₂O densely distributed near zero. The dispersion parameter was set at 23 based on a comparison of AIC.

To model RN₂O, the site factor was specified as a linear term and the CO₂ emissions measured in the micro plots, soil NH₄⁺ and NO₃⁻ contents measured near the band in the main plots and WFPS and soil temperature measured across the paddock were specified as smooth terms

and the micro plots as the random variable. For RN_2O modelling, cross-validation was performed by splitting datasets into training and testing by replicates ($k = 4$) and the predictive performance was evaluated by the averaged deviance explained and root mean square error (RMSE). The average of RN_2O predicted across k for each plot and bed/furrow position was used to calculate N_2 emissions.

In gap-filling of *Ndff* N_2 , days after fertilisation (DAF) and N rates were specified as the explanatory variables in a tensor product, allowing the changes in *Ndff* N_2 over time to differ between N rate treatments. To account for the repeated measurements of *Ndff* N_2 at the same chamber, 'chamber' was specified as the random variable nested in DAF. The estimated *Ndff* N_2 was then multiplied by 100 to show in percentage.