# Experimental evaluation of predicted undrained pore pressure generation as function of stress path and orientation in the Draupne shale

Magnus Soldal<sup>1</sup>, Jung Chan Choi<sup>1</sup>, and Elin Skurtveit<sup>1</sup>

<sup>1</sup>Norwegian Geotechnical Institute

January 3, 2023

#### Abstract

Injection or production of fluids from subsurface reservoirs lead to stress changes affecting both reservoir and surrounding rocks. For low-permeable caprocks overlying such reservoirs, the movement of pore fluids to or from the formation is restricted and the immediate and short-term response to changes in the stress field will be undrained. Consequently, stress changes transfer partly into pore pressure changes. The aim of the current study is to investigate theoretical means of forecasting the undrained pore pressure generation in the Draupne Formation shale and to compare predictions with experimental results. Predictions are based on measurements from a single undrained triaxial test on a sample with known orientation, and a combination of Skempton's classical formulation and anisotropic poroelastic theory. The predicted pore pressures are compared to measured pore pressures from a series of triaxial tests on samples with various orientations exposed to different total stress paths. First, it is confirmed that the normalized undrained pore pressure measured is linearly connected to the total stress path. Then it is demonstrated that a tensorial pore pressure parameter can be used to accurately predict the influence of stress orientation on generated pore pressure. Lastly, it is experimentally confirmed that the two predictions can be combined to predict the pore pressure arising from stress changes along any compressional stress path and orientation. The observations herein may contribute significantly to the understanding of induced pore pressure in low-permeable materials and provide valuable input to geomechanical modeling of various field operations.

#### Hosted file

951553\_0\_art\_file\_10512499\_rmmnsy.docx available at https://authorea.com/users/565285/ articles/612266-experimental-evaluation-of-predicted-undrained-pore-pressure-generationas-function-of-stress-path-and-orientation-in-the-draupne-shale

1	
2	Experimental evaluation of predicted undrained pore pressure generation
3 4	as function of stress path and orientation in the Drauphe shale
5	
6	
7	M. Soldal <sup>1,2</sup> , J. C. Choi <sup>2</sup> , and E. Skurtveit <sup>2</sup>
8	
9	<sup>1</sup> Department of Geosciences, University of Oslo, Oslo 0371, Norway
10	<sup>2</sup> NGI – Norwegian Geotechnical Institute, Oslo 0806, Norway
11	
12	Corresponding author: Magnus Soldal (magnus.soldal@ngi.no)
13	
14	
15	
16	Key Points:
17	• Effects of stress path and orientation on pore pressure generation are predicted from
18	theory and minimal experimental data
19	• The predictions are compared to measurements made in a series of undrained, triaxial
20	tests on the North Sea Draupne Formation shale
21	• Results show that pore pressure can be accurately predicted using a combination of
22	Skempton's formulation and anisotropic poroelasticity

#### 23 Abstract

24 Injection or production of fluids from subsurface reservoirs lead to stress changes affecting both 25 reservoir and surrounding rocks. For low-permeable caprocks overlying such reservoirs, the movement of pore fluids to or from the formation is restricted and the immediate and short-term 26 27 response to changes in the stress field will be undrained. Consequently, stress changes transfer 28 partly into pore pressure changes. The aim of the current study is to investigate theoretical means 29 of forecasting the undrained pore pressure generation in the Draupne Formation shale and to 30 compare predictions with experimental results. Predictions are based on measurements from a 31 single undrained triaxial test on a sample with known orientation, and a combination of 32 Skempton's classical formulation and anisotropic poroelastic theory. The predicted pore 33 pressures are compared to measured pore pressures from a series of triaxial tests on samples with various orientations exposed to different total stress paths. First, it is confirmed that the 34 35 normalized undrained pore pressure measured is linearly connected to the total stress path. Then 36 it is demonstrated that a tensorial pore pressure parameter can be used to accurately predict the 37 influence of stress orientation on generated pore pressure. Lastly, it is experimentally confirmed 38 that the two predictions can be combined to predict the pore pressure arising from stress changes 39 along any compressional stress path and orientation. The observations herein may contribute 40 significantly to the understanding of induced pore pressure in low-permeable materials and 41 provide valuable input to geomechanical modeling of various field operations.

#### 43 Plain Language Summary

44 Fluid production from or injection into geological reservoirs usually cause changes in stresses 45 experienced by both reservoir and surrounding rocks. In the low-permeable caprocks found above reservoirs, fluids occupying the pore space are, at least in the short-term, restricted from 46 47 moving in or out of the formation in response to these stress changes. The fluids instead carry 48 parts of the stress change and consequenctly experience changes in pore pressure. The aim of the 49 current study is to investigate means of predicting pore pressure generation following changes in 50 surrounding stresses and to compare them to experimental data from a North Sea shale. Using 51 only limited experiemental data and available poroelastic theory, predictions of pore pressure 52 generation under various total stress change scenarios are made. A series of triaxial tests are 53 then performed on the same shale core, and the measured pore pressures are compared to those predicted. The results show that the procedure followed enables accurate prediction of pore 54 55 pressure development as a function of both material orientation and stress path. These findings 56 highlight how pore pressure changes within anisotropic rocks can be predicted from minimal 57 experimental testing campaigns and provide valuable input to geomechanical modelling of 58 subsurface stress changes.

59

60

61

#### 63 **1** Introduction

64 Reservoir pore pressure alterations arising from fluid injection or production can cause 65 deformational processes and changes in total stresses within both reservoir and surrounding 66 rocks due to the coupling that exists between pore pressure and stress (e.g. Addis, 1997; Hettema 67 et al., 2000; Teufel et al., 1991; Zoback & Zinke, 2002). These pore pressure- and total stress 68 variations together determine the effective stresses that eventually cause rock deformation 69 (Terzaghi, 1936). Estimation of effective stress evolution for reservoir and surrounding rocks 70 during production or injection is needed to predict mechanical consequences such as reservoir 71 compaction and subsidence/uplift, caprock fracturing, fault reactivation, borehole collapse or 72 casing deformation (e.g. Altmann et al., 2014; Altmann et al., 2010; Angus et al., 2016; 73 Castelletto et al., 2013; Holt et al., 2004; Lynch et al., 2013; Santarelli et al., 1998; Segura et al., 74 2011; Aadnoy, 1991). Permeable reservoir rocks typically display drained behavior in response 75 to changes in total stresses. However, very little fluid movement occurs in the low-permeable 76 caprocks overlying such reservoirs. Consequently, the caprocks' response to alterations in the 77 stress field will be undrained and can lead to significant pore pressure changes even without fluid 78 movement to or from the reservoir. The amount of undrained pore pressure generation depends 79 on the magnitude and direction of the total stress change and the material's poroelastic 80 parameters. Poroelastic theory addresses the coupling between fluid flow and mechanical 81 deformation of fluid-saturated rocks (e.g. Biot, 1941; Cheng, 2016; Wang, 2017). Within the 82 poroelastic framework, Skempton's pore pressure parameters, As and Bs, predict the undrained 83 pore pressure generation in response to deviatoric and isotropic stress changes, respectively 84 (Skempton, 1954). Understanding these pore pressure parameters and what affects them is thus 85 crucial to explain effective stress changes in caprocks. In anisotropic poroelastic theory, the

directional dependency of properties found in most sedimentary rocks are included (e.g. Biot, 1955; Carroll, 1979; Cheng, 2016; Holt, Bakk, & Bauer, 2018; Thompson & Willis, 1991). As a result, Skempton's pore pressure parameter Bs is generalized into a second rank tensor, where the effect of pore pressure by deviatoric stresses is similar to that of As.

90

91 One consequence of Skempton's original formulation is that the amount of undrained 92 pore pressure generated depends on the ratio between changes in minor ( $\Delta\sigma$ 3) and major ( $\Delta\sigma$ 1) 93 principal stresses (i.e. the stress path). In fact, a linear relationship is predicted between the pore 94 pressure change normalized to the total vertical stress change and the stress path (Skempton, 95 1954). For anisotropic rocks, the influence of relative orientation between principal stresses and 96 material symmetry axis on pore pressure generation from deviatoric stress changes can be 97 predicted from the components of the tensorial Bij (Cheng, 2016; Holt, Bauer, et al., 2018). Both 98 the relationship between stress path and normalized pore pressure and the directional dependence 99 of Skempton's As have previously, to a limited extent, been experimentally evaluated for shales 100 (Holt, Bakk, Stenebråten, et al., 2018; Holt, Bauer, et al., 2018). However, to our knowledge, the 101 literature is missing studies combining these two dependencies and thereby enabling a more 102 comprehensive pore pressure forecasting on a well characterized shale. In the current study, the 103 aim is to demonstrate experimentally that the effects of both stress path and stress orientation on 104 pore pressure generation in a low-permeable, anisotropic rock can be predicted using poroelastic 105 theory and require only minimal data form laboratory testing. The experimental work is 106 performed in the triaxial apparatus with shale samples from the North Sea Draupne Formation. 107 Parameters needed to construct the pore pressure predictions are first derived from a single test 108 on a sample with known orientation. Then, a series of undrained tests following different total

stress paths are performed on samples subcored with different sample axis orientations relative to the originally horizontal rock layering. The target is to investigate how measured undrained pore pressure alterations match those predicted by poroelastic theory for various stress paths and stress orientations. If reliable predictions can be made, it demonstrates that relatively simple experimental procedures can be used to obtain data needed to model pore pressure changes under a variety of different field conditions.

#### 115 2 Poroelastic theory and undrained pore pressure generation

116 The mechanical properties of rocks are affected by the presence of fluids that can move 117 within their porous frames, and poroelastic theory can describe the constitutive behavior of fluid-118 saturated rocks. Whether or not fluids can leave or enter the pores of rocks subjected to stress 119 changes determine if the material behaves drained or undrained, respectively. If drainage is 120 permitted, the fluid will not take part in the load-bearing process, and any deformation occurring 121 will be the result of deformation of the solid constituents and the porous structure. If drainage is 122 prohibited, fluids occupying the pore space provide additional resistance to stress changes. In 123 such cases, stress variations will be partly transferred to the fluid pressure which will change in 124 response. Skempton (1954) introduced the pore pressure parameters  $A_s$  and  $B_s$  to quantify the 125 undrained pore pressure change ( $\Delta u$ ) of saturated soils during undrained loading. The 126 formulation in Equation 1 was made for stress conditions known as compressional triaxial 127 conditions in which  $\Delta \sigma_l$  and  $\Delta \sigma_3$  define changes in major and minor principal stresses, 128 respectively (and  $\Delta \sigma_2 = \Delta \sigma_3$ ):

$$\Delta u = B_{s}[\Delta \sigma_{3} + A_{s}(\Delta \sigma_{1} - \Delta \sigma_{3})]$$
 Equation 1

130 Skempton's  $B_s$ -parameter represents the ratio of undrained pore pressure to change in 131 mean stress  $\left(\left(\frac{2\Delta\sigma_3+\Delta\sigma_1}{3}\right)\right)$  and varies with the rock frame compressibility (*C*) relative to that of the 132 pore fluid (*C<sub>f</sub>*) and solid constituents (*C<sub>s</sub>*). In Equation 2, Skempton's *B<sub>s</sub>*-parameter is expressed 133 as a function of these compressibilities and the porosity (*n*) (Kümpel, 1991):

134

$$B_s = \frac{1}{1 + n\left(\frac{C_f - C_s}{C - C_s}\right)}$$
Equation 2

135

136 Skempton experimentally showed that the parameter  $B_s$  is close to unity for saturated and 137 unconsolidated sediments where the frame compressibility significantly exceeds that of the 138 fluids. Materials containing more compressible pore fluids (e.g. gas) or a relatively increased 139 frame stiffness compared to the solids (e.g. rocks) will have  $B_s$ -values less than 1.

140

141 Skempton's  $A_s$  describes the undrained pore pressure response to deviatoric stress 142 changes (i.e. the difference between  $\Delta \sigma_1$  and  $\Delta \sigma_3$ ). If a rock is assumed to be isotropic and to 143 follow linear poroelasticity, the pore pressure response is entirely controlled by mean stress 144  $(\Delta u = 1/3(\Delta \sigma_1 + 2\Delta \sigma_3))$  and Skempton's  $A_s$  equal to 1/3. However, from Skempton's original 145 experiments on various types of clays, a range of  $A_s$ -values from -0.5 to 1.5 was measured. The 146 variation in  $A_s$  was attributed to the non-elastic nature of clays.

147

148 In the case of compressional triaxial stress conditions, it is convenient to replace the 149 major and minor principal stresses with vertical and horizontal stresses, respectively. In the

150 current study, discussions are limited to cases where the minor and intermediate principal

151 stresses (horizontal stresses) are equal and smaller than the vertical stress (i.e.  $\Delta \sigma_v > \Delta \sigma_H = \Delta \sigma_h$ ).

Furthermore, the ratio of changes in minor and major principal stresses is described by the stress
path parameter κ:

154

$$\kappa = \frac{\Delta \sigma_h}{\Delta \sigma_n}$$
 Equation 3

155

By rearranging Skempton's formulation given in Equation 1 and including the stress path
parameter κ, it is shown that the undrained pore pressure normalized to the vertical stress
increment is linearly connected to κ:

159

$$\frac{\Delta u}{\Delta \sigma_{v}} = (B_{s}(1 - A_{s}))\kappa + A_{s}B_{s}$$
 Equation 4

Equation 4 signify that different undrained stress paths generate varying pore pressure. A consequence of Equation 4 is therefore that if the pore pressure parameters  $A_s$  and  $B_s$  are known, the undrained pore pressure arising from stress changes along any compressional stress path can be predicted.

164

165 So far, the discussion has been limited to isotropic rocks. However, most rocks have a 166 directional dependency in their properties. For anisotropic materials, it is known that undrained

pore pressure generation varies with stress orientation relative to any symmetry direction that exists in the rock (Cheng, 2016). Horizontally layered rocks, such as shales (Piane et al., 2011), are often treated as vertical transverse isotropic (VTI) materials with equal properties in the horizontal plane and different properties in the vertical direction (Fjær et al., 2008). By utilizing anisotropic poroelastic stress-strain relations (and the inverse strain-stress relations) in a similar manner as for anisotropic elasticity, Biot's fluid strain parameter ( $\zeta$ ) can be expressed by the stress tensor ( $\sigma_{ij}$ ), pore pressure (u) and Hooke's law constants (C and  $B_{ij}$ ) (Cheng, 1997):

$$\zeta = C(u + \frac{1}{3}B_{ij}\sigma_{ij})$$
 Equation 5

175 The fluid strain parameter gives the fluid volume leaving or entering the solid frame per 176 unit volume of solid frame. In the case of undrained stress change with no shear-induced volume change of the solid frame,  $\zeta$  is zero and Equation 5 reduces to the first part of Equation 1 in 177 178 which  $B_{ii}$  can be viewed as a generalization of Skempton's  $B_s$  into a second rank tensor (Cheng, 179 1997; Cheng, 2016; Holt, Bakk, Stenebråten, et al., 2018; Thompson & Willis, 1991). For 180 vertical transverse materials, the tensorial  $B_{ii}$  consists of two independent components,  $B_v$  and  $B_h$ . 181 Subscripts v and h refer to vertical and horizontal, or perpendicular and parallel to the symmetry 182 plane in VTI materials, respectively (Cheng, 1997):

$$\Delta u = \frac{1}{3} B_{ij} \Delta \sigma_{ij} = \left(\frac{2B_h + B_v}{3}\right) \left(\frac{2\Delta \sigma_h + \Delta \sigma_v}{3}\right)$$
Equation 6

184 According to Cheng (1997), the significance of Equation 6 is that pore pressure can be 185 generated by incremental normal as well as shear stress, the latter being an effect similar to the 186 concept of Skempton's  $A_s$  parameter. The undrained pore pressure response to an isotropic stress 187 change ( $\Delta \sigma_{iso} = \Delta \sigma_h = \Delta \sigma_v$ ) is independent of material orientation relative to the symmetry axis, since  $B_{ij} = B_s$  is expressed as a volume-weighted average of the two invariant components 188 189 (Holt et al., 2017). However, the undrained pore pressure generated from an anisotropic stress 190 change depends on stress orientation. If  $\theta$  denotes the angle between symmetry axis and major 191 principle stress direction, the undrained pore pressure response in a triaxial compression test can 192 be predicted by Equation 7 (Holt, Bakk, Stenebråten, et al., 2018):

193

$$\Delta u = B_s \left[ \Delta \sigma_3 + \frac{(B_h \sin^2 \theta + B_v \cos^2 \theta)}{2B_h + B_v} (\Delta \sigma_1 - \Delta \sigma_3) \right]$$
 Equation 7

194 Equation 7 resembles the formulation by Skempton in Equation 1, with  $A_s$  displaying a 195 directional dependence given by the  $B_s$  tensorial components:

196

$$A[\theta] = \frac{(B_h \sin^2\theta + B_v \cos^2\theta)}{2B_h + B_v} = \frac{(B_h \sin^2\theta + B_v \cos^2\theta)}{3B_s}$$
 Equation 8

197 Worth noting is that the subscript *s* is herein reserved for the original Skempton's  $A_s$ 198 parameter. Since the tensorial  $B_{ij}$  equals  $B_s$  for VTI materials under triaxial stress conditions, the 199 two are used interchangeably. However,  $A[\theta]$  is an elastic parameter, in contrast to Skempton's 200 original  $A_s$  (Raaen et al., 2019). Combining Equation 4 and Equation 8 thus allows for theoretical 201 predictions of normalized pore pressure to be made from the tensorial components of Skempton's 202  $B_s$  on VTI materials oriented at all directions relative to rock layering and exposed to elastic 203 loading along all different compressive stress paths:

204

$$\frac{\Delta u}{\Delta \sigma_v} = \left(\frac{2B_h + B_v}{3}\right) * (1 - A[\theta]) * \kappa + A[\theta] * \left(\frac{2B_h + B_v}{3}\right)$$
 Equation 9

205

206  $B_{\rm s}$ -value measurements on saturated shales found in the literature typically vary between 207 0.5 and 0.9 (e.g. Belmokhtar et al., 2016; Mohajerani et al., 2011; Favero et al., 2018; Giger et al., 208 2018; Lozovyi & Bauer, 2019; Wild et al., 2017; Holt, Bauer, et al., 2018; Ma & Gutierrez, 209 2020). Soldal et al. (2021b) previously reported  $B_s$ -values between 0.52 and 0.71 for the Draupne 210 shale. There are fewer reports of Skempton's  $A_s$  from laboratory testing on low-permeable rocks 211 in the literature. One example is from Lozovyi and Bauer (2019) who measured  $A_s$ -values 212 between 0.13 and 0.6 during triaxial testing on sandy and shalv facies of Opalinus Clay. Their 213 testing on samples subcored with different angles between sample axis and rock layering also 214 clearly showed that the undrained pore pressure during deviatoric loading varied with the sample 215 orientation. Even less experimental data has been reported from studies seeking to demonstrate 216 the directional variations in  $A[\theta]$  predicted by anisotropic poroelastic theory. Cheng (1997) used 217 the laboratory data collected by Aoki et al. (1993) to compute tensorial  $B_s$ -value components of 218 the Trafalgar shale from triaxial testing ( $B_h = 0.51$  and  $B_v = 0.63$ ). The components were not, 219 however, used to predict undrained pore pressure variation with orientation. Holt, Bakk, 220 Stenebråten, et al. (2018) reported anisotropic poroelastic coefficients from several different, 221 unspecified shales. For a field core of 'soft shale', the horizontal and vertical  $B_s$ -value

components were experimentally determined to 0.57 and 1.33, respectively.  $A[\theta]$ -values measured in undrained triaxial experiments on samples with different orientations were in line with those predicted from the  $B_s$ -value components. A similar experimental campaign on a more porous overburden field shale showed that Skempton's  $A[\theta]$  decreased from 0.6 to 0.2 between samples oriented parallel and perpendicular to the symmetry axis. The novelty of the present study is that the directional variation in  $A[\theta]$  is used to predict the pore pressure variation in stress path dependence as a function of orientation.

229 **3 Material** 

230 The shale material tested in the current study is from the Draupne Formation, which is 231 considered both one of the main petroleum source rocks and caprocks in the North Sea (Faleide 232 et al., 2010). It belongs to the Viking group and was deposited under anoxic conditions in several 233 over-deepened basins during the Upper Jurassic (Faleide et al., 2010; Færseth et al., 1995; 234 Underhill, 1998; Whipp et al., 2014). Samples tested were extracted from core material collected 235 from well 16/8-3S in the Ling depression. The Ling depression is located south of the Horda 236 platform and separates the basement highs of Utsira and Sele (Fossen & Hurich, 2005; Færseth et 237 al., 1995). From the well, 9 meters of Draupne core material was retrieved from a depth of 238 2574.5 – 2583.5 m MD. The mineralogy and mechanical properties of the Draupne core material 239 from this well have been characterized by several authors in recent studies (e.g. Bohloli et al., 240 2020; Koochak Zadeh et al., 2017; Skurtveit et al., 2015; Smith, 2019; Soldal et al., 2021a). The 241 clay content of the Draupne shale is approximately 50 % and the total organic content is between 6 and 8 wt.%. Due to the low hydraulic permeability  $(10^{-15} \text{ m/s})$  and high CO<sub>2</sub> capillary 242 243 breakthrough pressure ( $\approx 4$  MPa) (Skurtveit et al., 2012), intact Draupne shale is expected to 244 make an excellent caprock above potential CO<sub>2</sub> storage formations in the North Sea. The

245	porosity ranges between $13 - 17$ % and the average bulk density is 2.25 g/cm <sup>3</sup> . Assumed vertical
246	and horizontal effective in-situ stresses for the core material used in the current study is 26.1 and
247	17.2 MPa, respectively (Koochak Zadeh et al., 2017), and the material is considered normally to
248	slightly overconsolidated (Soldal et al., 2021a).

249 **4 Method** 

4.1. Sample preparation

251 Since shales are highly sensitive to changes in saturation (e.g. Ewy, 2015, 2018; Valès et 252 al., 2004), the Draupne core material has been kept submerged in mineral oil since retrieval to 253 prevent drying. Recent fluid content measurements after several years of storage showed no 254 change in saturation (Soldal et al., 2021a). Core sections from which triaxial rock samples were 255 prepared were initially separated from the remaining core using a circular saw. The end surfaces 256 of the core sections were then made parallel and planar using a grinding machine in a controlled 257 humidity environment. Next, cylindrical samples were sub-cored using an oil-cooled, custom-258 made drill bit with an internal, air-pressure supported piston maintaining a small, constant axial 259 load on the sample during drilling. Immediately after completion of the coring procedure, the 260 samples were inspected for visible fractures, before they were placed inside a small aluminum 261 rack providing some vertical support and submerged in oil until testing. The tested samples had 262 diameters of 25 mm and height to diameter ratio from 2 - 2.5:1. Prior to testing, the samples' 263 surfaces were wiped free of mineral oil using oil-only absorbent pads.

#### 265 4.2. Experimental device

266 Testing was done inside a traditional type of triaxial pressure cell where changes in 267 confining pressure normally cause equal changes in vertical and horizontal stresses. Deviatoric 268 stresses is supplied by a stepping motor located beneath the cell base, and confining pressure and 269 top and bottom pore pressures are independently controlled by hydraulic pressure controllers. 270 Vertical and horizontal sample deformations are measured internally inside the cell over the 271 middle third of the sample height using two vertical and two horizontal Linear Variable 272 Differential Transducers (LVDTs). Vertical load is measured internally by a vented load sensor 273 located just beneath the bottom end piece. The membrane surrounding the sample is made of a 274 material that prevents water diffusion from the rock sample without being too stiff. Four vertical 275 side drains connecting the porous filters at the top and bottom were used for radial drainage to 276 speed up pore pressure equilibration. A more detailed description of the equipment used can be 277 found in Berre (2011) and Soldal et al. (2021a). During undrained testing, the measured pore 278 pressure will be affected by the presence of a non-zero dead-volume with a certain 279 compressibility between the sample and the closed valve(s) (e.g. Bishop, 1976; Ghabezloo & 280 Sulem, 2008; Ghabezloo & Sulem, 2010; Wissa, 1969). Efforts have been made to reduce the 281 dead-volume by using only 1/16" steel tubing and by keeping the length of the tubing at a 282 minimum. The 70 MPa pore pressure sensor, having an accuracy of 0.05 % of full scale, between 283 the sample and the drainage valve has also been modified to contain a minimal internal fluid 284 volume. The compressibility of the drainage system was evaluated by measuring the volume 285 needed to pressurize the system with only a steel dummy inserted into the triaxial cell, both with 286 open and closed drainage valves. The difference between the two can be used together with the 287 absolute volume of the drainage system to calculate compressibility. The procedure described by

Ghabezloo and Sulem (2008) for system compliance evaluation indicate that compliance of the current testing system would result in a corrected Skempton's  $B_s$ -value less than 10 % higher than the measured. Due to the relatively small error introduced by the none-zero dead volume and the differences in testing procedures followed in the current study, however, the corrections are not applied to the results presented herein.

4.3. Test procedures

294 Results from a total of nine triaxial tests are presented in the current study (see Table 1). 295 The purpose of Test 1 was to derive the parameters needed to predict pore pressure variation as a 296 function of stress path and orientation, whereas the purpose of all the other tests was generate 297 experimental data to compare to the predictions. Test 1 was done on a sample with its axis 298 parallel with the symmetry axis, and the experimental procedure followed will be described in 299 detail later. To investigate potential effects of consolidation conditions on the generated pore 300 pressure, the anisotropic consolidation in Test 1 was replaced by isotropic consolidation in Test 301 2. In Tests 3 and 4, the same procedure as in Test 1 was followed for samples with orientation  $\theta$ 302  $= 90^{\circ}$ . The reason for doing two identical tests here was to evaluate potential influence of 303 heterogeneity between samples. Results from Tests 5-9 have previously been reported by 304 Skurtveit et al. (2015) and Soldal et al. (2021b) and will only be briefly described here. It is 305 emphasized that only Test 1 was needed to formulate the pore pressure predictions; the purpose 306 of the remaining tests was to evaluate experimentally the applicability of those predictions.

307

308

			Consolidation		Elastic stress cycles			
Test #	θ (°)*	φ (%)**	Iso.	Aniso.	1 <sup>st</sup> cycle к	2 <sup>nd</sup> cycle к	3 <sup>rd</sup> cycle к	4 <sup>th</sup> cycle κ
Test 1	0	14.2		$\checkmark$	1	0.7	0	-0.5
Test 2	0	13.4	~		1	0.5	0	-0.5
Test 3	90	15.9		✓	1	0.7	0	-0.5
Test 4	90	16.1		✓	1	0.7	0	-0.5
Test 5	0	13.9	✓		×			
Test 6	30	16.4	~		×			
Test 7	45	15.8	~		×			
Test 8	60	16.2	<b>√</b>		×			
Test 9	90	16.3	<b>√</b>		×			

Table 1: Overview of the 9 triaxial tests included in the current study. \* Orientation given as the

angle  $\theta$  between the major principal stress direction and the symmetry axis. \*\* Initial porosity

313 calculated from the initial and final fluid contents and a grain density of  $2.55 \text{ g/cm}^3$ .

314

315 The experimental protocol followed in Test 1 is schematically illustrated in Figure 1. The 316 confining pressure was immediately increased to 1 MPa after inserting the sample surrounded by 317 four side drains and the sleeve into the triaxial cell. The vertical stress was then increased to give 318 the correct in-situ ratio between principal stresses. Next, the stresses were simultaneously 319 increased to the absolute values of the in-situ effective stresses ( $\sigma_v$ ' = 26.1 MPa and  $\sigma_h$ ' = 17.2 320 MPa). Degassed analog pore fluid (NaCl = 37 g/l) could now enter the porous end filters and 321 flushed through the side drains. Then the backpressure was increased to the in-situ pore pressure 322 of 27.9 MPa (rate 0.6 MPa/hr), whilst maintaining constant horizontal and vertical effective 323 stresses. After this, a minimum of 40 hours was given for the sample pore pressure to equilibrate 324 before the tests proceeded.

326 The valve between the backpressure system and the sample was then closed in Test 1, 327 leaving the samples under undrained conditions. The sample was subsequently subjected to four 328 stress cycles, each following a different stress path; isotropic, uniaxial strain, uniaxial stress and 329 constant mean stress path. In all stress cycles, the total vertical stress increment was 4 MPa and 330 was applied over a period of 8 hours. After completion of the total vertical stress increase, a 331 minimum of 10 hours was granted for equilibration of excess pore pressure before the stress was 332 decreased again at the same rate ( $\Delta \sigma_v = 0.5$  MPa/hour). In the isotropic stress cycle, the total 333 horizontal stress increment was equal to the total vertical stress increment, in the uniaxial strain 334 cycle the total horizontal stress was adjusted to prevent horizontal sample deformation, in the 335 uniaxial stress cycle there was no change in total horizontal stress, whereas in the constant mean 336 stress cycle the total horizontal stress increment was -0.5 times the total vertical stress increment (maintaining  $\frac{2\Delta\sigma_h + \Delta\sigma_v}{3} = 0$ ). Between all stress cycles, drainage was re-opened, and 337 338 the pore pressure controlled at the initial value of 27.9 MPa. The pore pressure was controlled for 339 a minimum of 8 hours before the drainage valve was again closed and the next stress cycle 340 initiated.



Figure 1: Schematic illustration of the triaxial testing procedure followed in Test 1 ( $\theta$ =0°). The effective in-situ stresses were  $\sigma_v' = 26.1$  MPa and  $\sigma_h' = 17.2$  MPa and the backpressure was 27.9 MPa. The total vertical stress increments in all four elastic load cycles were 4 MPa. The same procedure was also followed in Test 2, with the exception that consolidation was isotropic ( $\sigma' = 23.5$  MPa) and the backpressure was 20 MPa. In Tests 3 and 4, the procedure in Figure 1 was

- 347 reproduced on samples with ( $\theta = 90^{\circ}$ ).
- 348

349 In Test 2, the effective isotropic consolidation stress and backpressure was 23.5 and 20 350 MPa, respectively. Instead of a uniaxial strain load cycle, the second load cycle in Tests 2-4351 was replaced with a constant stress path parameter cycle ( $\kappa = 0.5$  and 0.7 for Tests 2 and 3 & 4, 352 respectively). In Tests 5-9 samples of different orientations were isotropically consolidated and 353 subjected to undrained shearing. The effective consolidation stresses and backpressures in Tests 354 5-9 were 20 and 30 MPa, respectively. Undrained shearing was done by increasing the vertical stress to give a specified axial strain rate. A very low axial strain rate  $(10^{-9} \text{ s}^{-1})$  was used to 355 356 ensure that the generated undrained pore pressure had ample time to equilibrate within the

sample pore volume (Soldal et al., 2021b). The only measurements from Tests 5 – 9 used in the
current study, are the pore pressure changes relative to the increase in total vertical stresses
between the start of shearing and sample failure.

360

361 5 Results

362 The undrained pore pressures ( $\Delta u$ ) generated in each of the four stress cycles in Tests 1 – 363 4 are plotted against the change in total vertical stress ( $\Delta \sigma_v$ ) in Figure 2. Figure 2a and b show the 364 results from Tests 1 and 2 with sample axis perpendicular to layering, whereas Figure 2c and d are 365 from tests 3 and 4 with samples oriented parallel with layering. In the same plots, the ratios of 366 maximum pore pressure change to total vertical stress change are given as single values (i.e. 367 'normalized pore pressure' =  $\Delta u / \Delta \sigma_v$ ). It is observed that the changes in undrained pore pressure 368 during the isotropic stress cycles in all tests have similar magnitude, demonstrating that 369 Skempton's  $B_s$  is independent of material orientation. Furthermore, results show that gradually 370 less pore pressure is generated going from the isotropic to the constant mean stress cycle (i.e. 371 reducing  $\kappa$ ). In Tests 1 and 2, the pore pressure generation relative to the total vertical stress 372 increment when the mean stress is kept constant is less than half of that during the isotropic 373 stress cycle. Results show very little difference in terms of pore pressure generation between the 374 anisotropically (Test 1) and isotropically (Test 2) consolidated tests on samples with identical 375 orientation. During the uniaxial strain cycle in Test 1, the change in total horizontal stress needed to prevent changes in horizontal deformation resulted in a secant stress path parameter  $\kappa$  of 0.7. 376 377 This stress path was used when controlling the second stress cycles in Tests 3 and 4. The 378 reduction in generated pore pressure from isotropic to constant mean stress cycle is larger in 379 Tests 3 and 4 compared to Tests 1 and 2. During loading (i.e. increasing total vertical stress)

380 under constant mean stress conditions, the pore pressure in fact decreased. Also, worth noting is

that pore pressure after unloading generally returned back to values close to the pore pressure



382 before loading was initiated.



393

One of the aims of the current study is to examine whether reliable predictions of

394 undrained pore pressure variation with orientation can be made using components of the

tensorial  $B_s$  parameter. To this end, the horizontal and vertical components need first to be

396 quantified. Figure 3 plots the undrained pore pressure normalized to the change in total vertical 397 stress against stress path parameter  $\kappa$  for Tests 1 and 2. Results from both tests display the same 398 linearity predicted by Skempton's formulation, and any possible influence of isotropic as 399 opposed to anisotropic consolidation on pore pressure generation is therefore considered 400 negligible. The resulting slope and intercept from linear regression can be used together with 401 Equation 4 to calculate Skempton's  $A_s$  and  $B_s$  values of 0.61 and 0.59, respectively. Since the 402 sample orientation is known ( $\theta = 0^{\circ}$ ), Skempton's  $B_s$  can then be divided into a horizontal 403 component of 0.34 and a vertical component of 1.09 using Equation 6 and Equation 8.



404

Figure 3: Normalized pore pressure plotted against the stress path parameter  $\kappa$  for Tests 1 and 2 ( $\theta = 0^{\circ}$ ). Open circles are drawn from the anisotropically consolidated Test 1 and open diamonds are drawn from the isotropically consolidated Test 2. The result of linear regression given above the dashed line.

410	The generated pore pressures during undrained shearing of samples with different
411	orientation in Tests $5-9$ are plotted against changes in total vertical stress in Figure 4a. As the
412	angle between the sample axis and the symmetry axis increases, the generated pore pressure
413	decreases. The total undrained pore pressure in Test 5 ( $\theta = 0^{\circ}$ ) is almost 2.5 times that in Test 9
414	( $\theta = 90^{\circ}$ ). To calculate $A_s$ from Tests 5 – 9, the secant total pore pressure change recorded
415	between the initiation of shearing and sample failure (defined as the highest measured shear
416	stress) was first divided by the corresponding change in total vertical stress. The result is equal to
417	the product of Skempton's two pore pressure parameters, and $A_s$ was then found by dividing this
418	by $B_s = 0.59$ . The calculated $A_s$ values are plotted as open circles in Figure 4b. The predicted
419	variation in $A[\theta]$ based on $B_s$ -value components and Equation 8 is represented by the dashed line
420	in the same figure. Figure 4b shows that the poroelastic expression of $A[\theta]$ made up of the $B_s$ -
421	value components can be used to accurately predict undrained pore pressure generation as a
422	function of material orientation, even in cases where significant plastic deformation most
423	probably has occurred (i.e. $A_s \approx A[\theta]$ ).



Figure 4: a) Measured pore pressure change plotted against change in total vertical stress during undrained shearing of the isotropically consolidated samples in Tests 5 – 9. The orientation of sample axis relative to symmetry axis for each test is given in the figure legend. b) Values of A<sub>s</sub> from Tests 5 – 9 plotted as open circles together with the predicted variation in A [ $\theta$ ] based on B<sub>s</sub>-value components.

433 Once it is demonstrated that the variation in Skempton's  $A_s$  with material orientation can 434 be predicted from the tensorial  $B_s$  components, the next step is to examine how the relationship between normalized pore pressure and stress path will vary with orientation. The dashed lines in 435 436 Figure 5 are predictions of how the normalized pore pressures will vary as functions of stress 437 path based on a constant  $B_s$ -value and  $A_s$ -values changing with orientation. The orange line is drawn for samples oriented with  $\theta = 0^{\circ}$  and the blue line for samples oriented with  $\theta = 90^{\circ}$ . The 438 439 diamond markers representing measured values in Tests 1 - 4 follow the predicted trends extremely well. Furthermore, the consistency between Tests 1 & 2 and 3 & 4 indicates both the 440 441 applicability of the current method and that there is no significant heterogeneity between the 442 tested samples.



Figure 5: Predicted normalized pore pressure as a function of stress path in dashed lines. Orange line predicts the linear relationship for samples with sample axis parallel to the symmetry axis, and blue line for samples with axis perpendicular to symmetry axis. Orange and blue diamonds show the measured normalized undrained pore pressures in Tests 1 - 4.

### 448 6 Discussion

449 The objective of the current study was to examine experimentally predictions of 450 undrained pore pressure in a typical North Sea shale exposed to varying stress paths and stress 451 orientations. Using components of the tensorial  $B_s$  parameter, the effect of stress orientation on 452 pore pressure generation from deviatoric stress changes were predicted. Subsequent 453 measurements of pore pressure in undrained triaxial tests on samples subcored at different angles 454 relative to the layering in the rock were in accordance with the predictions. Then a linear 455 relationship between undrained normalized pore pressure and stress path was demonstrated on a 456 sample with its axis parallel to the material symmetry axis. Knowing how Skempton's  $A_s$  varied 457 with orientation, the effect of stress direction on the relation between normalized undrained pore

458 pressure and stress path was predicted. Again, the predicted pore pressure generation was 459 reproduced in undrained tests on samples with their axis oriented perpendicular to the symmetry 460 axis. The importance of what is presented in Figure 5 is that from relatively simple undrained 461 stress cycling of one sample with a known orientation, the pore pressure generation for all 462 orientations and compressional stress paths of that same material can be accurately predicted 463 using Equation 9.

464

465 The results shown herein demonstrate that significant variation in induced pore pressure 466 generation as a function of stress orientation and stress path can be expected and should be 467 accounted for. The importance of including anisotropy in poroelastic parameters when predicting 468 undrained pore pressure generation has been examined in recent studies by Raaen et al. (2019). 469 They investigated numerically differences in predicted undrained pore pressure around a 470 horizontal borehole in an anisotropic material using isotropic versus anisotropic Skempton's  $B_s$ . 471 In the isotropic case, the components of  $B_s$  are equal, whereas for the anisotropic case the  $B_s$ -472 value components were estimated from a given set of elastic material properties. Simulations 473 showed that whether choosing isotropic scalar  $B_s$  or anisotropic tensorial  $B_s$  would significantly 474 affect pore pressure predictions. Since the formation considered was assigned anisotropic elastic 475 properties, it must be assumed that the anisotropic solution was closer to reality than the isotropic 476 simplification.

477

Asaka and Holt (2021) compared an anisotropic approach with a conventional approach
in terms of predicting unwanted borehole instability during drilling operations. The conventional
method assumes isotropic rock properties and ignores pore pressure changes resulting from total

481 stress changes. Both anisotropy in elastic properties and pore pressure parameters were respected 482 in the anisotropic approach. Considerable differences in predicted failure regions and modes 483 resulted from the comparison of the two approaches. Field observations of wellbore failure in 484 shale section at the sides of highly inclined boreholes were only forecasted if anisotropic 485 poroelasticity was included. Furthermore, sensitivity analysis showed that the risk of failure was 486 significantly influenced by the amount of undrained pore pressure, and, accordingly, the 487 anisotropic pore pressure parameters. Skempton parameters used by Asaka and Holt (2021) were 488 estimated from the inverse anisotropic Gassmann's equation. Consequently, there is an 489 uncertainty related to the pore pressure parameters which can only be reduced by more 490 experimental data addressing anisotropic poroelasticity of shales. 491 492 Finally, it is mentioned again that since the influence of plastic deformation on pore 493 pressure parameters is unknown, the predictions herein are strictly speaking limited to loading 494 within the elastic domain. Duda et al. (2021) recently performed experiments to examine how 495 Skempton's pore pressure parameters are affected by plastic deformation. Cyclic triaxial testing 496 of the relatively porous Pierre II shale ( $\varphi = 40$  %) indicated that although B<sub>s</sub>-values remained 497 unaffected throughout shearing, Skempton's  $A_s$  reduced quite significantly as the degree of plastic deformation increased. Future research should also consider the potentially variable 498 499 effects of plastic deformation on pore pressure parameters following different stress paths since 500 this may aid the description of stress evolution in faulted or fractured rock formations.

#### 502 7 Conclusion

503 This study has shown experimentally that the variation in undrained pore pressure as a 504 function of stress path and stress orientation can be predicted using Skempton's  $B_s$ -value 505 components and the stress path parameter  $\kappa$ . A series of undrained triaxial tests on the Draupne 506 Formation shale was designed and executed to enable predictions of pore pressure generation to 507 be made and to later examine experimentally the applicability of those predictions. The tensorial 508 components of  $B_s$  were first derived from the measured linear relationship between stress path 509 and normalized pore pressure in one triaxial test on a sample with its axis parallel with the 510 material symmetry axis. Subsequent measurements in five undrained triaxial tests on samples 511 with various relative orientations reproduced the predicted decrease in Skempton's  $A_s$  with 512 increasing  $\theta$ . The variation in  $A_s$  with orientation was then incorporated into the relation between 513 normalized undrained pore pressure and stress path. Results from another two tests on samples 514 with their axis perpendicular to the symmetry axis exposed to various stress paths demonstrated 515 that also the variation in stress path dependence as a function of orientation could be forecasted.

516

517 The significant variation in pore pressure measured as a function of orientation 518 emphasizes that anisotropic poroelasticity need to be considered to accurately predict undrained 519 pore pressure generation in materials with different properties in different directions. If not, risks 520 related to e.g. borehole stability or caprock fracturing cannot be properly evaluated prior to 521 operations involving fluid injection or production. The experiments herein have shown that 522 tensorial  $B_s$ -value components can be very useful in predicting the undrained pore pressure 523 response over a wide range of stress orientation- and stress path scenarios. Even though the 524 predictions do not incorporate any effect of plastic deformation on the pore pressure parameters,

525	the undrained pore pressure in the Draupne shale could be satisfactorily forecasted even in cases
526	involving loading with associated plastic deformations. Characterizing the influence of non-
527	elastic deformation on pore pressure generation is most likely relevant for already fractured or
528	faulted sections of the subsurface and will be subject for future research.
529	
530	
531	
532	
533	
534	
535	
536	
537	
538	
539	
540	
541	
542	
543	
544	

### 545 References

- 546 Addis, M. A. (1997). The Stress-Depletion Response Of Reservoirs. SPE Annual Technical
- 547 Conference and Exhibition,
- 548 Altmann, J. B., Müller, B. I. R., Müller, T. M., Heidbach, O., Tingay, M. R. P., & Weißhardt, A.
- 549 (2014). Pore pressure stress coupling in 3D and consequences for reservoir stress states and fault
- 550 reactivation. *Geothermics*, 52, 195-205. <u>https://doi.org/10.1016/j.geothermics.2014.01.004</u>
- 551 Altmann, J. B., Müller, T. M., Müller, B. I. R., Tingay, M. R. P., & Heidbach, O. (2010).
- 552 Poroelastic contribution to the reservoir stress path. International Journal of Rock Mechanics
- *and Mining Sciences*, *47*(7), 1104-1113.
- 554 <u>https://doi.org/https://doi.org/10.1016/j.ijrmms.2010.08.001</u>
- Angus, D. A., Fisher, Q. J., Segura, J. M., Verdon, J. P., Kendall, J. M., Dutko, M., & Crook, A.
- 556 J. L. (2016). Reservoir stress path and induced seismic anisotropy : results from linking coupled
- 557 fluid-flow/geomechanical simulation with seismic modelling. Petroleum Science, 13(4), 669-
- 558 684. <u>https://doi.org/10.1007/s12182-016-0126-1</u>
- Aoki, T., Tan, C. P., & Bamford, W. E. (1993). Effects of deformation and strength anisotropy
- 560 on borehole failures in saturated shales. *International Journal of Rock Mechanics and Mining*
- 561 Sciences & Geomechanics Abstracts, 30(7), 1031-1034. <u>https://doi.org/10.1016/0148-</u>
- 562 <u>9062(93)90067-N</u>
- 563 Asaka, M., & Holt, R. (2021). Anisotropic Wellbore Stability Analysis: Impact on Failure
- 564 Prediction. Rock Mechanics and Rock Engineering, 54, 1-23. https://doi.org/10.1007/s00603-
- 565 <u>020-02283-0</u>

- 566 Belmokhtar, M., Delage, P., Ghabezloo, S., Tang, A.-M., Menaceur, H., & Conil, N. (2016).
- 567 Poroelasticity of the Callovo–Oxfordian Claystone. Rock Mechanics and Rock Engineering,
- 568 50(4), 871-889. <u>https://doi.org/10.1007/s00603-016-1137-3</u>
- 569 Berre, T. (2011). Triaxial Testing of Soft Rocks. *Geotechnical Testing Journal*, 34(1), 61-75.
- 570 <u>https://doi.org/10.1520/GTJ102879</u>
- 571 Biot, M. A. (1941). General Theory of Three Dimensional Consolidation. Journal of applied
- 572 *physics*, *12*, 155-164.
- 573 Biot, M. A. (1955). Theory of Elasticity and Consolidation for a Porous Anisotropic Solid.
- 574 *Journal of applied physics*, 26(2), 182-185. <u>https://doi.org/10.1063/1.1721956</u>
- 575 Bishop, A. W. (1976). The influence of system compressibility on the observed pore-pressure
- 576 response to an undrained change in stress in saturated rock. 26(2), 371-375.
- 577 <u>https://doi.org/10.1680/geot.1976.26.2.371</u>
- 578 Bohloli, B., Soldal, M., Smith, H., Skurtveit, E., Choi, J. C., & Sauvin, G. (2020). Frictional
- 579 Properties and Seismogenic Potential of Caprock Shales. *Energies (Basel)*, 13(6275), 6275.
- 580 <u>https://doi.org/10.3390/en13236275</u>
- 581 Carroll, M. M. (1979). An effective stress law for anisotropic elastic deformation. J. Geophys.
- 582 *Res*, *84*(B13), 7510-7512. <u>https://doi.org/10.1029/JB084iB13p07510</u>
- 583 Castelletto, N., Gambolati, G., & Teatini, P. (2013). Geological CO2 sequestration in multi-
- 584 compartment reservoirs: Geomechanical challenges. Journal of Geophysical Research: Solid
- 585 *Earth*, *118*(5), 2417-2428. <u>https://doi.org/10.1002/jgrb.50180</u>
- 586 Cheng, A. H. D. (1997). Material coefficients of anisotropic poroelasticity. International journal
- 587 of rock mechanics and mining sciences (Oxford, England : 1997), 34(2), 199-205.
- 588 https://doi.org/10.1016/S0148-9062(96)00055-1

- 589 Cheng, A. H. D. (2016). Poroelasticity (1st ed. 2016. ed., Vol. 27). Springer International
- 590 Publishing : Imprint: Springer.
- 591 Duda, M. I., Holt, R. M., Stenebråten, J. F., & Stroisz, A. M. (2021). Effects of Plastic
- 592 Deformation on Poroelastic Pore Pressure Coefficients in Pierre II Shale. 55th U.S. Rock
- 593 Mechanics/Geomechanics Symposium,
- 594 Ewy, R. T. (2015). Shale/claystone response to air and liquid exposure, and implications for
- 595 handling, sampling and testing. International Journal of Rock Mechanics and Mining Sciences,
- 596 80, 388-401. <u>https://doi.org/https://doi.org/10.1016/j.ijrmms.2015.10.009</u>
- 597 Ewy, R. T. (2018). Practical approaches for addressing shale testing challenges associated with
- 598 permeability, capillarity and brine interactions. Geomechanics for energy and the environment,
- 599 14, 3-15. <u>https://doi.org/10.1016/j.gete.2018.01.001</u>
- 600 Faleide, J. I., Bjørlykke, K., & Gabrielsen, R. H. (2010). Geology of the Norwegian Continental
- 601 Shelf. In K. Bjorlykke (Ed.), Petroleum Geoscience: From Sedimentary Environments to Rock
- 602 *Physics* (pp. 467-499). Springer Berlin Heidelberg. <u>https://doi.org/10.1007/978-3-642-02332-</u>
- 603 <u>3\_22</u>
- 604 Favero, V., Ferrari, A., & Laloui, L. (2018). Anisotropic Behaviour of Opalinus Clay Through
- 605 Consolidated and Drained Triaxial Testing in Saturated Conditions. Rock Mechanics and Rock
- 606 Engineering, 51(5), 1305-1319. <u>https://doi.org/10.1007/s00603-017-1398-5</u>
- 607 Fjær, E., Holt, R. M., Horsrud, P., Raaen, A. M., & Risnes, R. (2008). Petroleum related rock
- 608 mechanics (2nd ed. ed., Vol. 53). Elsevier.
- 609 Fossen, H., & Hurich, C. A. (2005). The Hardangerfjord shear zone in SW Norway and the
- 610 North Sea; a large-scale low-angle shear zone in the Caledonian crust. Journal of the Geological
- 611 Society, 162(4), 675-687. <u>https://doi.org/10.1144/0016-764904-136</u>

- 612 Færseth, R. B., Gabrielsen, R. H., & Hurich, C. A. (1995). Influence of basement in structuring
- of the North Sea Basin, offshore southwest Norway. Norsk geologisk tidsskrift, 75, 105-119.
- 614 Ghabezloo, S., & Sulem, J. (2008). Stress dependent thermal pressurization of a fluid-saturated
- 615 rock. Rock Mechanics and Rock Engineering, 42(1), 1-24. <u>https://doi.org/10.1007/s00603-008-</u>
- 616 <u>0165-z</u>
- 617 Ghabezloo, S., & Sulem, J. (2010). Effect of the volume of the drainage system on the
- 618 measurement of undrained thermo-poro-elastic parameters. International journal of rock
- 619 mechanics and mining sciences (Oxford, England : 1997), 47(1), 60-68.
- 620 https://doi.org/10.1016/j.ijrmms.2009.03.001
- 621 Giger, S. B., Ewy, R. T., Favero, V., Stankovic, R., & Keller, L. M. (2018). Consolidated-
- 622 undrained triaxial testing of Opalinus Clay: Results and method validation. Geomechanics for
- 623 energy and the environment, 14, 16-28. <u>https://doi.org/10.1016/j.gete.2018.01.003</u>
- 624 Hettema, M. H. H., Schutjens, P. M. T. M., Verboom, B. J. M., & Gussinklo, H. J. (2000).
- 625 Production-Induced Compaction of a Sandstone Reservoir: The Strong Influence of Stress Path.
- 626 SPE Reservoir Evaluation & Engineering, 3(04), 342-347. <u>https://doi.org/10.2118/65410-PA</u> %J
- 627 SPE Reservoir Evaluation & Engineering
- 628 Holt, R., Bauer, A., & Bakk, A. (2017). Overburden pore-pressure changes and their influence
- on 4D seismic. 2017 SEG International Exposition and Annual Meeting,
- 630 Holt, R. M., Bakk, A., & Bauer, A. (2018). Anisotropic poroelasticity Does it apply to shale?
- 631 <u>https://doi.org/https://doi.org/10.1190/segam2018-2994785.1</u>
- 632 Holt, R. M., Bakk, A., Stenebråten, J. F., Bauer, A., & Fjær, E. (2018). Skempton's A A Key
- 633 to Man-Induced Subsurface Pore Pressure Changes. 52nd U.S. Rock Mechanics/Geomechanics
- 634 Symposium,

- 635 Holt, R. M., Bauer, A., & Bakk, A. (2018). Stress Path Dependent Velocities in Shales: Impact
- on 4D Seismic Interpretation. Stress Path Dependent Velocities in Shales: Impact on 4D Seismic
- 637 Interpretation. https://doi.org/https://doi.org/10.1190/geo2017-0652.1
- 638 Holt, R. M., Flornes, O., Li, L., & Fjaer, E. (2004). Consequences Of Depletion-Induced Stress
- 639 Changes On Reservoir Compection And Recovery. Gulf Rocks 2004, the 6th North America
- 640 Rock Mechanics Symposium (NARMS),
- 641 Koochak Zadeh, M., Mondol, N. H., & Jahren, J. (2017). Velocity anisotropy of upper jurassic
- 642 organic-rich shales, Norwegian continental shelf.
- 643 Kümpel, H. J. (1991). Poroelasticity: parameters reviewed. Geophys. J. Int, 105(3), 783-799.
- 644 <u>https://doi.org/10.1111/j.1365-246X.1991.tb00813.x</u>
- 645 Lozovyi, S., & Bauer, A. (2019). Static and dynamic stiffness measurements with Opalinus Clay.
- 646 Geophysical Prospecting, 67(4), 997-1019. <u>https://doi.org/10.1111/1365-2478.12720</u>
- 647 Lynch, T., Fisher, Q., Angus, D., & Lorinczi, P. (2013). Investigating Stress Path Hysteresis in a
- 648 CO2 Injection Scenario Using Coupled Geomechanical-fluid Flow Modelling. Energy Procedia,
- 649 37, 3833-3841. <u>https://doi.org/https://doi.org/10.1016/j.egypro.2013.06.280</u>
- 650 Ma, S., & Gutierrez, M. (2020). Determination of the poroelasticity of shale. Acta geotechnica,
- 651 16(2), 581-594. https://doi.org/10.1007/s11440-020-01062-z
- 652 Mohajerani, M., Delage, P., Monfared, M., Tang, A. M., Sulem, J., & Gatmiri, B. (2011).
- 653 Oedometric compression and swelling behaviour of the Callovo-Oxfordian argillite.
- 654 International journal of rock mechanics and mining sciences (Oxford, England : 1997), 48(4),
- 655 606-615. <u>https://doi.org/10.1016/j.ijrmms.2011.02.016</u>

- 656 Piane, C. D., Dewhurst, D. N., Siggins, A. F., & Raven, M. D. (2011). Stress-induced anisotropy
- 657 in brine saturated shale. *Geophysical journal international*, 184(2), 897-906.
- 658 <u>https://doi.org/10.1111/j.1365-246X.2010.04885.x</u>
- 659 Raaen, A. M., Larsen, I., Fjær, E., & Holt, R. M. (2019). Pore Pressure Response in Rock:
- 660 Implications of Tensorial Skempton B in an Anisotropic Formation. 53rd U.S. Rock
- 661 Mechanics/Geomechanics Symposium,
- 662 Santarelli, F. J., Tronvoll, J. T., Svennekjaier, M., Skeie, H., Henriksen, R., & Bratli, R. K.
- 663 (1998). Reservoir Stress Path: The Depletion and the Rebound. SPE/ISRM Rock Mechanics in
- 664 Petroleum Engineering,
- 665 Segura, J. M., Fisher, Q. J., Crook, A. J. L., Dutko, M., Yu, J. G., Skachkov, S., . . . Kendall, J.
- 666 M. (2011). Reservoir stress path characterization and its implications for fluid flow production
- 667 simulations. *Petroleum geoscience*, 17(4), 335-344. <u>https://doi.org/10.1144/1354-079310-034</u>
- 668 Skempton, A. W. (1954). The Pore-Pressure Coefficients A and B. 4(4), 143-147.
- 669 <u>https://doi.org/10.1680/geot.1954.4.4.143</u>
- 670 Skurtveit, E., Aker, E., Soldal, M., Angeli, M., & Wang, Z. (2012). Experimental investigation
- of CO2 breakthrough and flow mechanisms in shale. *Petroleum geoscience*, 18(1), 3-15.
- 672 <u>https://doi.org/10.1144/1354-079311-016</u>
- 673 Skurtveit, E., Grande, L., Ogebule, O. Y., Gabrielsen, R. H., Faleide, J. I., Mondol, N. H., ...
- Horsrud, P. (2015, 2015/11/13/). *Mechanical testing and sealing capacity of the Upper Jurassic*
- 675 Draupne Formation, North Sea 49th U.S. Rock Mechanics/Geomechanics Symposium, San
- 676 Francisco, California. <u>https://doi.org/</u>
- 677 Smith, H. (2019). Engineering parameters of Draupne shale Characterization of fractured
- 678 samples and integration with mechanical tests [Master's thesis, University of Oslo,

- 679 Soldal, M., Skurtveit, E., & Choi, J. C. (2021a). Laboratory Evaluation of Mechanical Properties
- of Draupne Shale Relevant for CO2 Seal Integrity. *Geosciences*, 11(6), 244.
- 681 <u>https://www.mdpi.com/2076-3263/11/6/244</u>
- 682 Soldal, M., Skurtveit, E., & Choi, J. C. (2021b). Poroelastic and Mechanical Anisotropy of the
- 683 Draupne Caprock. 55th U.S. Rock Mechanics/Geomechanics Symposium,
- 684 Terzaghi, K. v. (1936). The shearing resistance of saturated soils and the angle between the
- 685 planes of shear. First international conference on soil Mechanics, 1936,
- 686 Teufel, L. W., Rhett, D. W., & Farrell, H. E. (1991). Effect of Reservoir Depletion And Pore
- 687 Pressure Drawdown On In Situ Stress And Deformation In the Ekofisk Field, North Sea. The
- 688 32nd U.S. Symposium on Rock Mechanics (USRMS),
- Thompson, M., & Willis, J. R. (1991). A Reformation of the Equations of Anisotropic
- 690 Poroelasticity. Journal of applied mechanics, 58(3), 612-616. <u>https://doi.org/10.1115/1.2897239</u>
- 691 %J Journal of Applied Mechanics
- 692 Underhill, J. R. (1998). Jurassic. In Petroleum Geology of the North Sea (pp. 245-293).
- 693 https://doi.org/https://doi.org/10.1002/9781444313413.ch8
- 694 Valès, F., Nguyen Minh, D., Gharbi, H., & Rejeb, A. (2004). Experimental study of the influence
- 695 of the degree of saturation on physical and mechanical properties in Tournemire shale (France).
- 696 Applied clay science, 26(1), 197-207. <u>https://doi.org/10.1016/j.clay.2003.12.032</u>
- 697 Wang, H. F. (2017). Theory of Linear Poroelasticity with Applications to Geomechanics and
- 698 Hydrogeology. Princeton University Press.
- 699 Whipp, P. S., Jackson, C. A.-L., Gawthorpe, R. L., Dreyer, T., & Quinn, D. (2014). Normal fault
- array evolution above a reactivated rift fabric; a subsurface example from the northern Horda

- 701 Platform, Norwegian North Sea. 26(4), 523-549.
- 702 <u>https://doi.org/https://doi.org/10.1111/bre.12050</u>
- 703 Wild, K. M., Barla, M., Turinetti, G., & Amann, F. (2017). A multi-stage triaxial testing
- 704 procedure for low permeable geomaterials applied to Opalinus Clay. Journal of Rock Mechanics
- 705 and Geotechnical Engineering, 9(3), 519-530.
- 706 https://doi.org/https://doi.org/10.1016/j.jrmge.2017.04.003
- Wissa, A. E. Z. (1969). Pore Pressure Measurement in Saturated Stiff Soils. 95(4), 1063-1073.
- 708 https://doi.org/doi:10.1061/JSFEAQ.0001304
- 709 Zoback, M. D., & Zinke, J. C. (2002). Production-induced Normal Faulting in the Valhall and
- 710 Ekofisk Oil Fields. *pure and applied geophysics*, *159*(1), 403-420.
- 711 <u>https://doi.org/10.1007/PL00001258</u>
- 712 Aadnoy, B. S. (1991). Effects of reservoir depletion on borehole stability. Journal of petroleum
- 713 science & engineering, 6(1), 57-61. <u>https://doi.org/10.1016/0920-4105(91)90024-H</u>
- 714

#### 715 Acknowledgments

- This publication has been produced with support from the Research Council of Norway
- 717 (RCN) through the NCCS Centre (RCN# 257579), performed under the Norwegian research
- 718 programme Centres for Environment-friendly Energy Research (FME). The authors would like
- to thank the partners in PL360 with funding from CLIMIT (Project 223122) and Statoil
- 720 (Equinor) for the core material used in this study. Thanks also to Bjørnar Slensvik and Halvard
- 721 Smith for help with sample preparation and discussions on testing protocols.
- 722
- 723

### 724 **Open Research**

- 725 All experimental data used from triaxial testing herein is published at DataverseNO (Magnus
- 726 Soldal, 2022, "Replication Data for: Experimental evaluation of predicted undrained pore
- 727 pressure generation as function of stress path and orientation in the Draupne shale").

728