Selecting appropriate model complexity: An example of tracer inversion for thermal prediction in enhanced geothermal systems

Hui Wu¹, Zhijun Jin², Su Jiang³, Hewei Tang⁴, Joseph P. Morris⁵, Jinjiang Zhang¹, and Bo Zhang⁶

¹Peking University
²Institute of Energy, Peking University
³Stanford University
⁴Lawrence Livermore National Laboratory
⁵Lawrence Livermore National Laboratory (DOE)
⁶The Key Laboratory of Orogenic Belts and Crustal Evolution, School of Earth and Space Sciences, Peking University

December 7, 2022

Abstract

A major challenge in the inversion of subsurface parameters is the ill-posedness issue caused by the inherent subsurface complexities and the generally spatially sparse data. Appropriate simplifications of inversion models are thus necessary to make the inversion process tractable and meanwhile preserve the predictive ability of the inversion results. In the present study, we investigate the effect of model complexity on the inversion of fracture aperture distribution as well as the prediction of longterm thermal performance in a field-scale single-fracture EGS model. Principal component analysis (PCA) was used to map the original cell-based aperture field to a low-dimensional latent space. The complexity of the inversion model was quantitatively represented by the percentage of total variance in the original aperture fields preserved by the latent space. Tracer, pressure and flow rate data were used to invert for fracture aperture through an ensemble-based inversion method, and the inferred aperture field was then used to predict thermal performance. We found that an over-simplified aperture model could not reproduce the inversion data and the predicted thermal response was biased. A complex aperture model could reproduce the data but the thermal prediction showed significant uncertainty. A model with moderate complexity, although not resolving many fine features in the "true" aperture field, successfully matched the data and predicted the long-term thermal behavior. The results provide important insights into the selection of model complexity for effective subsurface reservoir inversion and prediction.

1	
2	Selecting appropriate model complexity: An example of tracer inversion for thermal
3	prediction in enhanced geothermal systems
4	

Hui Wu^{1,2}, Zhijun Jin^{1,2,3}, Su Jiang⁴, Hewei Tang⁵, Joseph P. Morris⁵, Jinjiang Zhang¹, Bo Zhang¹

- ⁷ ¹ School of Earth and Space Sciences, Peking University, Beijing, China.
- ⁸ ² Institute of Energy, Peking University, Beijing, China.
- ⁹ ³ Petroleum Exploration and Production Research Institute, SINOPEC, Beijing, 100083, China
- ⁴ Department of Energy Resources Engineering, Stanford University, Stanford, CA, USA.
- ⁵ Lawrence Livermore National Laboratory, Livermore, CA, USA.
- 12
- 13 Corresponding author: Hui Wu (hui.wu@pku.edu.cn)
- 14 Key Points:
- Combined dimensionality reduction and data assimilation to infer fracture aperture distribution from tracer recovery data.
- Quantitatively investigated the effect of model complexities on the aperture inversion and
 thermal prediction of a field-scale EGS.
- A moderate model complexity is sufficient to reproduce tracer recovery data and provide
 accurate thermal predictions.

21 Abstract

22 A major challenge in the inversion of subsurface parameters is the ill-posedness issue caused by

the inherent subsurface complexities and the generally spatially sparse data. Appropriate

simplifications of inversion models are thus necessary to make the inversion process tractable

and meanwhile preserve the predictive ability of the inversion results. In the present study, we

investigate the effect of model complexity on the inversion of fracture aperture distribution as

well as the prediction of long-term thermal performance in a field-scale single-fracture EGS
 model. Principal component analysis (PCA) was used to map the original cell-based aperture

field to a low-dimensional latent space. The complexity of the inversion model was

30 quantitatively represented by the percentage of total variance in the original aperture fields

30 quantitatively represented by the percentage of total variance in the original aperture fields
31 preserved by the latent space. Tracer, pressure and flow rate data were used to invert for fracture

32 aperture through an ensemble-based inversion method, and the inferred aperture field was then

used to predict thermal performance. We found that an over-simplified aperture model could not

34 reproduce the inversion data and the predicted thermal response was biased. A complex aperture

35 model could reproduce the data but the thermal prediction showed significant uncertainty. A

36 model with moderate complexity, although not resolving many fine features in the "true"

aperture field, successfully matched the data and predicted the long-term thermal behavior. The

results provide important insights into the selection of model complexity for effective subsurface

39 reservoir inversion and prediction.

40 **1 Introduction**

41 Flow and transport processes in geothermal reservoirs highly depend on spatially heterogeneous reservoir properties, such as permeability distribution in a hydrothermal system 42 (Cox et al., 2001; Dobson et al., 2003; Shi et al., 2018) and fracture aperture distribution in an 43 enhanced geothermal system (EGS) (Chen & Zhao, 2020; Guo, Fu, Hao, Peters, & Carrigan, 44 2016; Okoroafor et al., 2022; Wu, Fu, Morris, et al., 2021). Characterizing permeability/aperture 45 fields is important for the modeling, prediction, optimization and long-term risk management of 46 47 geothermal reservoirs. However, due to the high cost and technical difficulties in directly measuring subsurface fields, available permeability/aperture data are generally spatially sparse. 48 A comprehensive characterization is often performed through the inversion of indirect hydraulic 49 or geophysical data, such as hydraulic and tracer testing data (Berkowitz, 2002; Chen et al., 50 2013; Somogyvári et al., 2017; Vogt et al., 2012; Wu, Fu, Hawkins, et al., 2021), electrical 51 resistivity (Johnson et al., 2021; Wu et al., 2019), seismic (Emerick, 2018; Liu & Grana, 2020), 52 and so on. A key component of hydraulic/geophysical inversion is a reliable model that can 53 properly simulate the underlying physical processes and output model responses for given model 54 parameters. As analytical models are only applicable to idealized scenarios with over simplified 55 fields, numerical models are required for the inversion of representative fields in real-world 56 applications. The infinite-dimensional space of a heterogeneously distributed field is projected to 57 a finite-dimensional parameter space by discretizing the model on a finite element mesh. 58

A major challenge in permeability and aperture inversion is the ill-posedness issue caused by the high dimensionality of model parameter space and the scarcity of hydraulic/geophysical data. A reliable numerical approximation of model responses requires a relatively fine discretization, which inevitably leads to a high-dimensional parameter space. Practically available hydraulic/geophysical data are usually insufficient to constrain such a high-dimensional parameter space. To tackle this challenge, dimensionality reduction methods have been used to 65 map the discretization-dependent, cell-based high-dimensional parameter space to a low-

- dimensional latent space (Jiang et al., 2021; Laloy et al., 2013; Marzouk & Najm, 2009; Tang et
- al., 2021; Xiao et al., 2022; Zhu & Zabaras, 2018). Principal component analysis (PCA) is a
 conventional dimensionality reduction method, which learns spatial similarities in training
- 69 samples (prior permeability or aperture models) and compress the most salient features into a
- latent space defined by orthogonal principal components (Hawkins et al., 2020; Sarma et al.,
- 71 2008; Wu, Fu, Hawkins, et al., 2021; Zhang et al., 2020). As a linear transform method, PCA is
- applicable to Gaussian and log-normal fields that can be fully characterized by two-point
- 73 statistics. For fields that follow non-Gaussian distributions, nonlinear transform methods are
- required for effective dimensionality reduction, such as deep learning-based methods (e.g.,
- 75 generative adversarial network and variational autoencoder) that have been widely explored in
- the recent literature (Canchumuni et al., 2020; Jiang & Jafarpour, 2021; Laloy et al., 2018; Mo et al., 2020).

78 The reduction of model dimensionality essentially leads to the reduction of model complexity. In a cell-based parameter space, each cell value is tuned independently during 79 inversion, and the model has the maximum variance. Through dimensionality reduction, the prior 80 knowledge in training samples, such as spatial auto-correlation and statistical features, are 81 learned and used to reduce the model degree of freedom (model complexity). The learned prior 82 83 knowledge serves to constrain the heterogeneous distribution of permeability/aperture fields and regularize subsequent inversion. Inversion on a low-dimensional latent space not only mitigates 84 the ill-posedness issue and make the inversion computationally tractable, but also better honors 85 the spatial auto-correlation nature of permeability/aperture fields than inversion on a cell-based 86 parameter space does. The tuning of a latent parameter changes the overall spatial distribution of 87 a field rather than its value at a single cell, which is a highly desired feature for inversion in a 88 89 data-scarce environment.

Latent space dimensionality, as a quantitative measure of model complexity, is a 90 91 hyperparameter that needs to be carefully determined prior to dimensionality reduction. On the one hand, model complexity should be deliberately compromised to accommodate the limited 92 information in hydraulic/geophysical data. On the other hand, the model needs to capture 93 adequate variations in the unknown field to appropriately simulate the underlying physical 94 processes. An extremely complex model is prone to overfitting and may undermine the 95 predictive ability of the inferred permeability/aperture fields, while an over-simplified model 96 97 may fall into the opposite error of underfitting and be unable to reproduce hydraulic/geophysical data. Unfortunately, due to the many inherent complexities of subsurface reservoirs, it is 98 99 difficult, if not impossible, to predetermine an ideal model complexity for permeability/aperture 100 inversion from hydraulic/geophysical data. In most previous studies, the dimensionality of latent space is subjectively determined (Yang et al., 2021). Some studies used relatively large latent 101 spaces to preserve at least 90% of the total variance in original cell-based parameter spaces when 102 103 using PCA for dimensionality reduction (Hawkins et al., 2020; Laloy et al., 2013; Tang et al., 2021; Zhao & Luo, 2020). Some other studies, on the other hand, preserved $50\% \sim 60\%$ of the 104 total variance through relatively small latent spaces (Romary, 2009; Fernández-Martínez et al., 105 2012; Emerick, 2017). The dimensionality of the resultant latent spaces in these studies varies 106 from 30 to 1,000. 107

Although dimensionality reduction methods have been widely used in
 permeability/aperture inversion, how to select an appropriate model complexity/latent space to

circumvent the dilemma of overfitting and underfitting remains unclear. The effects of model 110 111 complexity on permeability/aperture inversion and subsequent reservoir performance prediction require further investigation. Several studies examined the effect of latent space dimensionality 112 on forward simulation accuracy by first generating permeability/facies fields from latent spaces 113 with different dimensionalities, and then performing forward simulations on these generated 114 fields (Romary, 2009; Fernández-Martínez et al., 2012). The results indicated that a small latent 115 space (preserving 50% - 60% of the total variance in the original cell-based parameter spaces) 116 was sufficient to accurately simulate the underlying physical processes. Li & Cirpka (2006) 117 investigated the effect of latent space dimensionality on the inversion of a 2D hydraulic 118 conductivity field. With the increase of latent space dimensionality, the consumed computational 119 resources increased, while the inversion error (defined as the root mean square error between 120 true and inferred fields) gradually decreased and converged to a stable value. These studies focus 121 on the effect of model complexity on forward and inversion modeling, but lack analysis of the 122 predictive ability of the inversion results. Results from these studies provide insights into the 123 lower limit of model complexity to prevent underfitting. However, the upper limit of model 124 complexity to avoid overfitting, which manifests as good data match but poor predictive ability, 125

126 remains unexplored.

The main goal of the present study is to investigate the effect of model complexity on the 127 128 inversion of fracture aperture distribution as well as the prediction of long-term thermal recovery in an EGS. PCA is used to map the original cell-based aperture distribution to a latent space. The 129 corresponding model complexity is quantitatively represented by the percentage of total variance 130 in the original aperture fields preserved by the latent space. An ensemble-based inversion 131 method, ensemble smoother with multiple data assimilation (ES-MDA), is used for aperture 132 inversion from practically available tracer, pressure and flow rate data. Through this 133 investigation, we aim to analyze not only the minimum model complexity required to 134 appropriately reproduce tracer/pressure/flow rate data, but also the impact of overfitting on 135 thermal performance prediction due to excessive model complexity. The paper is organized as 136 137 follows. In Section 2, we introduce PCA for the dimensionality reduction of spatially autocorrelated aperture fields. Aperture fields generated from latent spaces with different 138 dimensionalities are compared to demonstrate the effect of latent space dimensionality on model 139 complexity. Section 3 describes a field-scale synthetic EGS model with a predominant horizontal 140 fracture, followed by the introduction of forward simulation methods (flow, tracer and thermal) 141 as well as a data assimilation framework using ES-MDA. In Section 4, synthetic tracer, pressure 142 and flow rate data are provided to ES-MDA to invert for the latent space obtained from PCA. 143 The inverted latent space is then mapped back to a cell-based aperture field to predict the thermal 144 performance of the EGS model. A series of latent spaces with different dimensionalities are 145 analyzed to investigate the effect of model complexity. Section 5 provides discussions regarding 146 the implication of the results. 147

148 **2** Principal component analysis for dimensionality reduction

Principle component analysis (PCA), also known as Karhunen-Loève (KL) expansion, is a well-established dimensionality reduction method. It has been broadly used in many subsurface inversion problems to map Gaussian or log-normal fields (e.g., permeability and aperture) to low-dimensional latent spaces that follow the standard normal distribution (Hawkins et al., 2020; Wu, Fu, Hawkins, et al., 2021). To perform PCA on the field of interest, we first generate an ensemble of training fields based on our prior knowledge of the field obtained from 155 geological/geophysical investigations such as core logs, wellbore images and outcrop analysis.

The reduction of dimensionality is then achieved by first computing orthogonal principal

157 components from the training fields, and then retaining the most significant principal

components as the basis functions to generate new fields through linear combination. The principal components can be calculated through either computing the eigenvectors and

eigenvalues of the covariance matrix of the training fields, or directly performing singular value

decomposition (SVD) on the training fields. The significance of a principal component is

represented by the percentage of the total variance in the training fields preserved by the principal component. For a spatially auto-correlated field, most of the variance in the training

fields can be effectively preserved by a small number of principal components. New fields

165 generated from the linear combination of the retained principal components have the same

166 dimensionality as the training fields, and are fully controlled by the weights of the retained 167 principal components. These weights form the latent space to be inferred in subsequent

168 inversion. The detailed procedure of PCA has been widely reported in the literature (Liu &

169 Durlofsky, 2020; Wu, Fu, Hawkins, et al., 2021) and therefore not repeated here.

In the current study, the field of interest is the aperture distribution of a 2D fracture. To 170 demonstrate the relationship between latent space dimensionality (i.e., the number of retained 171 principal components) and model complexity, we perform PCA on training aperture fields and 172 then compare the aperture fields reconstructed/generated with different numbers of principal 173 components. We generate 5,000 aperture fields on an 800 m × 800 m domain discretized into a 174 160×160 regular grid. We use sequential gaussian simulations and assume a spherical 175 variogram with a mean of 0.6 mm, a standard deviation of 0.45 mm and a correlation length of 176 75 m. The generated aperture fields follow a log-normal distribution, and are provided to PCA as 177 training fields. After PCA, 5,000 principal components are obtained and ranked in a descending 178 179 order according to their significance, i.e., percentage of preserved variance. We then use the first *l* principal components to reconstruct training aperture fields as well as generate new aperture 180 fields. 181

We first analyze the effect of *l* on training aperture field reconstruction. A random aperture field is selected from the training ensemble and reconstructed with l = 2, 10, 50, 200 and 800 as shown in Fig. 1. The preserved percentage of total variance is 1%, 5%, 21%, 56% and 84% for l = 2, 10, 50, 200 and 800 respectively. With a small *l*, the reconstructed aperture field is almost uniform and misses most of the variance in the training aperture field. With the increase of *l*, the complexity of the reconstructed aperture field increases, manifesting as the capture of fine features in the training aperture field.

189 We then analyze the effect of l on new aperture field generation. A random l-dimensional latent parameter vector is first sampled from the standard normal distribution, and then used as 190 the weights of the retained *l* principal components to generate a new aperture field. For each *l*, 191 192 we generated 30 aperture fields and analyze their mean, standard deviation and correlation length (Fig. 2). The new aperture fields also follow log-normal distributions, but the mean, standard 193 deviation and correlation length are different from that of the training fields. Compared with the 194 195 training aperture fields, aperture fields generated from latent spaces exhibit smaller mean and standard deviation, and larger correlation length. For an extremely small latent space (l = 2), the 196 197 mean of the generated aperture fields is slightly smaller than that of the training aperture fields. 198 The standard deviation is significantly smaller than that of the training aperture fields, and the correlation length shows the opposite trend. A smaller standard deviation and a larger correlation 199

- 200 length lead to a smoother aperture distribution, and therefore a less complex aperture model.
- 201 With the increase of *l*, the mean, standard deviation and correlation length gradually approaches
- 202 to their corresponding values in the training aperture fields.

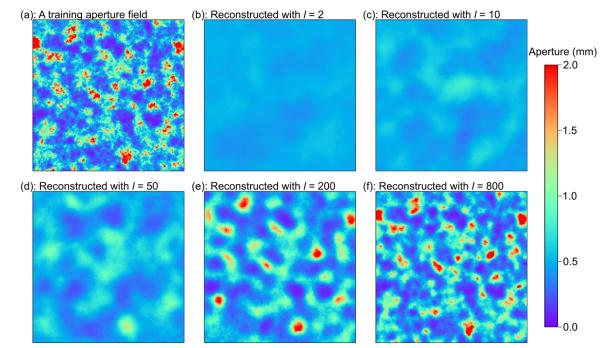
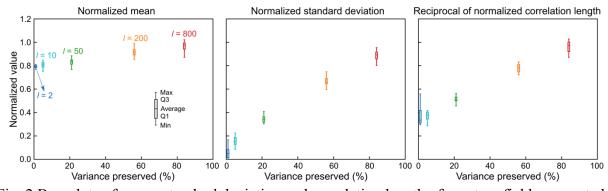


Fig. 1 Reconstruction of a training aperture field. (a) A randomly selected aperture field from the training ensemble. (b) \sim (f) Reconstructed aperture fields using different numbers of principal components.



207

Fig. 2 Box plots of mean, standard deviation and correlation length of aperture fields generated with different numbers of principal components. The box plots show the minimum, maximum, average, as well as the 25% (Q1) and 75% (Q3) percentiles. The mean, standard deviation and correlation length are normalized by their corresponding values used to generate the training aperture fields, i.e., 0.6 mm, 0.45 mm and 75 m respectively. For correlation length, we show the

213 reciprocal of the normalized correlation length.

214 **3 An EGS model and forward simulation/inversion methods**

In this section, we develop a field-scale single-fracture EGS model to demonstrate the effect of model complexity on aperture inversion and thermal prediction (Fig. 3). Data used for aperture inversion include practically available tracer, pressure and flow rate data. In what

- follows, we first describe the model details and then the numerical simulation of flow, tracer
- transport and thermal extraction in the EGS model. Finally, we briefly introduce a data
- assimilation framework developed in our previous work (Wu, Fu, Hawkins, et al., 2021), which
- has proven an effective method for aperture inversion and thermal prediction.
- 222 3.1 A field-scale single-fracture EGS model
- The developed EGS model is $3000 \times 3000 \times 3000$ m³ in dimension with a horizontal circular fracture 800 m in diameter, located at the center of the model. An injection well and two production wells are connected by the fracture (Fig. 3(a)). A vertical temperature gradient of 40 °C/km is assumed in the model with an initial temperature of 200 °C at the fracture depth.
 - (a) (b) (c) Aperture (mm) 2.0 1.5 1.0 -0.5 0.0 000 Flow rate (10⁻⁴ m²/s) 4.0 -3.0 -2.0 1.0 3000 m 0.0

Fig. 3 (a) A field-scale EGS model with a horizontal circular fracture located at 1,500 m depth.

(b) A Gaussian aperture field and the corresponding flow field under 20 L/s injection rate and
 constant pressure at the two production wells. (c) A non-Gaussian, two facies aperture field and

the corresponding flow field under 20 L/s injection rate and constant pressure at the two

232 production wells.

We investigate two "true" aperture fields, one is a spatially auto-correlated log-normal 233 field (Fig. 3(b)) and the other is a two facies field (Fig. 3(c)). The log-normal aperture field is 234 randomly generated from sequential gaussian simulation assuming a spherical variogram with a 235 mean of 0.6 mm, a standard deviation of 0.45 mm and a correlation length of 75 m. The two 236 facies aperture field is generated using the 'snesim' geostatistical algorithm (Strebelle, 2002) 237 from a geostatistical tool box SGeMS (Remy et al., 2009). The background aperture is 0.2 mm 238 239 and the aperture of flow channels is 1 mm. Note that the three wells are connected by flow channels. Although the two aperture fields follow different statistical distributions, we use the 240 same aperture model in subsequent inversion for them, i.e., log-normal aperture model. We use 241 the two facies aperture example to demonstrate the scenario where the statistical distribution of 242 the ground truth field fundamentally differs from that of the assumed aperture distribution in the 243 inversion process. This is commonly encountered in real-world problems as the ground truth 244 field is complex and we do not have sufficient data to correctly characterize its statistical 245 distribution. 246

247 3.2 Flow, tracer and thermal simulation

Flow and tracer simulation is performed to generate synthetic data for the "true" aperture fields, including tracer breakthrough curves (BTCs) and flow rates at the two production wells, as well as the pressure difference between the injection and production wells. The data are then provided to a data assimilation framework (Section 3.3) for aperture inversion, during which tracer simulation is used as forward model to simulate tracer, pressure and flow rate responses under various aperture scenarios. After inversion, thermal simulation is performed for both the "true" and inferred aperture fields to examine the predictive ability of the inferred aperture fields.

The discretization of the model is as follows. The fracture plane is represented by a thin 255 layer 4 mm in thickness, and the in-plane mesh resolution is 5×5 m² within the circular fracture 256 and gradually increases to 150×150 m² in the far field. For the surrounding rock formations, the 257 mesh resolution is $5 \times 5 \times 2.5$ m³ near the fracture plane and becomes progressively coarser in 258 the far field. The resulting computational domain consists of approximately 2,800,000 elements. 259 A massively parallel multi-physics simulation platform developed at the Lawrence Livermore 260 National Laboratory (Settgast et al., 2017), GEOS, is used for flow, tracer and thermal 261 simulation. GEOS provides a thermal-hydro-mechanical-chemical framework to simulate 262 various physical processes in subsurface reservoirs, such as fluid flow, mass and heat transport, 263 and hydraulic fracturing (Fu et al., 2013; Fu et al., 2016; Vogler et al., 2018; Wu, Fu, Morris, et 264 al., 2021). The implementation of flow, tracer and thermal modules relevant to the present study 265 266 has been described in the literature (Guo, Fu, Hao, & Carrigan, 2016; Guo, Fu, Hao, Peters, & Carrigan, 2016), and therefore not repeated here. 267

268 We first simulate the flow field and then solve the advection-dispersion-sorption equation based on the obtained flow field to simulate tracer transport processes. Note that we do not 269 consider mechanical simulation, indicating that the fracture aperture distribution does not evolve 270 271 during flow and tracer transport processes. Table 1 lists the parameters for flow and tracer modeling. As the fracture plane is represented by a thin layer, we calculate the equivalent 272 porosity and permeability of the fracture through $\phi = w/H$ and $k = w^3/12H$ respectively (Guo, Fu, 273 274 Hao, & Carrigan, 2016), where w is the aperture and H is the thickness of the fracture layer. Due to the relatively low rock formation permeability and the minor effect of matrix diffusion on 275 tracer transport (Wu, Fu, Hawkins, et al., 2021), we assume that tracer transport is confined 276 within the circular fracture and only consider the fracture for tracer modeling. Fracture 277 boundaries are assumed impermeable. A hydrostatic initial pressure is assumed in the model with 278 a pressure of 30 MPa at the fracture depth. The flow field in the fracture is simulated with an 279 injection rate of 20 L/s and a constant downhole pressure of 30 MPa at the two production wells 280 (Fig. 3(b) and (c)). According to the simulation results, the pressure difference between the 281 injection and production wells and flow rates at production well 1 and 2 are 824 kPa, 5.9 L/s, and 282 14.1 L/s respectively for the log-normal aperture field (Fig. 3(b)), and 178 kPa, 12.2 L/s, and 7.8 283 L/s respectively for the two facies aperture field (Fig. 3(c)). 284

We inject tracers into the fracture for one hour and simulate tracer transport for 20 hours to obtain tracer BTCs at the two production wells (Fig. 4). We consider both conservative and sorptive tracers. Note that for the simulation of a sorptive tracer, we assume an equilibrium sorption process with a typical partition coefficient of 1 mm. Compared with the conservative tracer BTCs, the sorptive tracer BTCs exhibit delayed peaks and reduced peak concentrations due to sorption effects. The tracer BTCs from the two facies aperture field show earlier arrival and larger peak concentration magnitude than that from the log-normal aperture field do,

especially for the BTCs at production well 1.

For thermal simulation, we circulate water among the injection and production wells for 50 years, with an injection rate of 20 L/s, an injection temperature of 50 °C, and a constant downhole pressure of 30 MPa at the two production wells. The upper, lower and lateral model boundaries are assumed impermeable to both fluid and heat. Parameters for thermal simulation are also listed in Table 1. For the log-normal aperture field, the production temperature at production well 1 decreases slower than that at production well 2 does (Fig. 4(a)), while for the two facies aperture field, the production temperature at production well 1 decreases faster.

Parameter	Value
Porosity of rock matrix	0.01
Permeability of rock matrix (m ²)	1×10^{-16}
Density of rock matrix (kg/m ³)	2500
Specific heat capacity of rock matrix (J/kg/K)	790
Thermal conductivity of rock matrix (W/m/K)	2.5
Density of water (kg/m ³)	887.2
Dynamic viscosity of water (Pa·s)	1.42×10^{-4}
Compressibility of water (Pa ⁻¹)	5×10^{-10}
Specific heat capacity of water (J/kg/K)	4460
Longitudinal dispersivity (m)	0.2
Transverse dispersivity (m)	0.02
Partition coefficient (mm)	1

Table 1 Parameters for flow, tracer and thermal simulations of the EGS model.

301

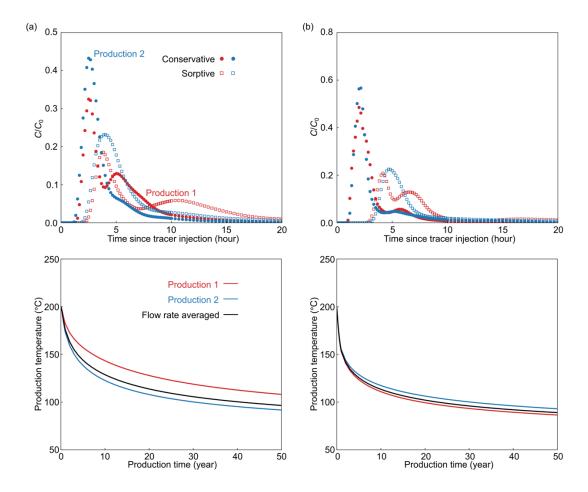


Fig. 4 Tracer (upper) and thermal (lower) breakthrough curves at the two production wells. (a) Results for the log-normal aperture field. (b) Results for the two facies aperture field. The simulated tracer concentration is normalized by injection concentration C_0 .

306 3.3 A data assimilation framework for aperture inversion and thermal prediction

The data assimilation framework developed in Wu, Fu, Hawkins, et al. (2021) is used in the current study for aperture inversion and thermal prediction. The framework includes three major components, i.e., parameterization, inversion and prediction. Here we briefly introduce the key procedures of applying the framework to the aperture inversion and thermal prediction in this study. We refer to Wu, Fu, Hawkins, et al. (2021) for more details of the framework.

312 3.3.1 Parameterization

The latent spaces generated in Section 2 are used as parameter spaces for aperture inversion. Five latent spaces with dimensionalities of 2, 10, 50, 200 and 800 (Fig. 1) are considered. Note that the aperture field generated from the latent spaces has a square shape (800 $m \times 800$ m), while the aperture field in the EGS model has a circular shape (800 m in diameter). Therefore, only the field within the inscribed circle of the generated 800 m \times 800 m aperture field is used for flow, tracer and thermal simulation.

319 3.3.2 Aperture inversion using ES-MDA

The synthetic tracer, pressure and flow rate data from the "true" aperture fields (Section 320 3.2) are used for latent space inversion through ES-MDA. An advantage of ES-MDA over 321 deterministic inversion methods is that a posterior ensemble of viable realizations (instead of a 322 single optimal realization) can be obtained to quantify the uncertainties associated with the 323 324 aperture field. To perform ES-MDA, we first generate 720 l-dimensional latent parameter sets as the prior ensemble by randomly sampling from the standard normal distribution. For each latent 325 parameter set in the ensemble, we use the retained principal components from PCA to map the 326 latent parameter set to an aperture field, and run flow and tracer simulation based upon the 327 aperture field. ES-MDA is then used to update the latent parameter sets according to the 328 simulated and "true" tracer, pressure and flow rate data. The two-step procedure (flow/tracer 329 simulation and latent parameter update) is repeated for 12 iterations to get the posterior 330 ensembles of latent parameter sets and aperture distributions. The major steps and update 331 equation of ES-MDA have been widely described in the literature (Emerick and Reynolds, 2013; 332 Wu, Fu, Hawkins, et al., 2021) and are not repeated here. The key parameters for ES-MDA, such 333 as data standard deviation and inflation coefficient, are the same as that in Wu, Fu, Hawkins, et 334 al. (2021). Note that 3% random Gaussian noise is added in the synthetic data before inversion 335 with ES-MDA. 336

337 3.3.3 Thermal prediction based on posterior aperture fields

After ES-MDA, the obtained posterior aperture fields are incorporated into the EGS model to perform thermal simulation and predict temperature responses at the two production wells. Predictions from different posterior aperture fields are used to analyze the associated uncertainties.

342 **4** Aperture inversion and thermal prediction with different model complexities

343

4.1 Comparison between "true" and simulated tracer, pressure and flow rate data

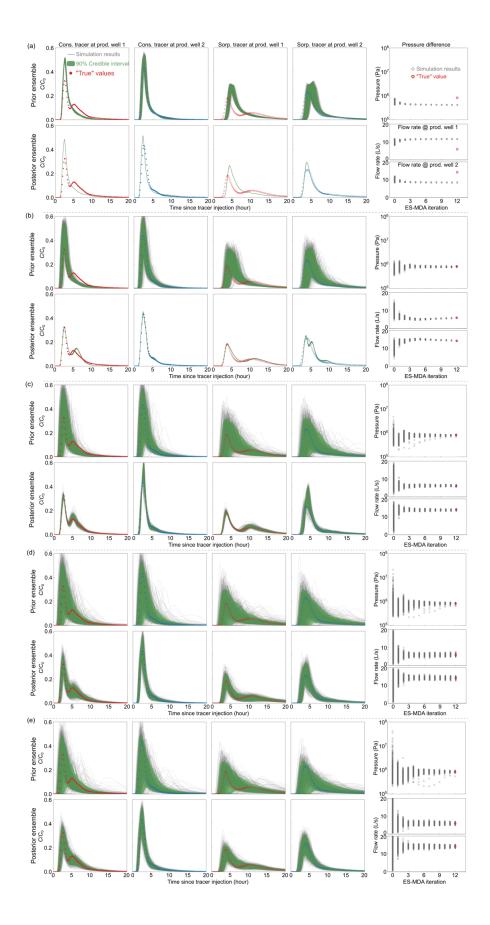
In each ES-MDA iteration, we record the simulation results (tracer, pressure and flow rate) to compare with the "true" data (Figs. 5 and 6). The latent space dimensionality (*l*) shows similar effects on the fit of tracer/pressure/flow rate data for the log-normal and two facies aperture fields, as summarized below.

For l = 2, the simulated tracer BTCs, pressure and flow rates show little variation 348 among the 720 prior realizations (before ES-MDA) (Figs. 5(a) and 6(a)). This is because 349 the aperture fields generated from such a low-dimensional latent space have little 350 variance and are relatively smooth (i.e., low model complexity), as shown in the first 351 column of Figs. 7 and 8. After ES-MDA, the 720 parameter sets collapse to the same 352 posterior parameter set (Fig. S1 in the Supporting Information), similar to the small 353 ensemble size-induced ensemble collapse phenomenon reported in many previous studies 354 (Nejadi et al., 2017; Xiao & Tian, 2020). However, the simulation results from this 355 posterior parameter set cannot correctly fit the "true" data (Figs. 5(a) and 6(a)), which is 356 an indicator of underfitting. The second peak of the tracer BTC at production well 1 is 357 not resolved as the underlying aperture model is unable to capture the complexities in the 358 "true" aperture fields. 359

• When *l* increases to ten, the complexity of the aperture model increases and the variation among the prior simulation results also increases (Figs. 5(b) and 6(b)). Nevertheless, the 720 parameter sets still collapse to the same posterior parameter set after ES-MDA (Fig. S1). The fit of the "true" data, especially the pressure and flow rate data, is better than that for l = 2. The second peak of the tracer BTC at production well 1 is successfully resolved by the posterior realizations, although there still exist some discrepancies between the "true" and simulated tracer BTCs.

When *l* further increases to 50, the collapse of parameter sets is greatly alleviated 367 and the posterior latent parameters show considerable uncertainties (Fig. S1). 368 Correspondingly, the uncertainty of the simulation results from the posterior ensemble 369 370 also increases, especially for the tracer BTCs (Figs. 5(c) and 6(c)). The 90% credible intervals of the simulated tracer BTCs properly match the "true" tracer BTCs, and both 371 the arrival time and magnitude of the second peak of the tracer BTC at production well 1 372 are correctly reproduced. An aperture model with moderate complexity is able to capture 373 the necessary variations in the "true" aperture field to reproduce the tracer, pressure and 374 flow rate data, even if the aperture model and the true" aperture field follow 375 fundamentally different statistical distributions. 376

• When *l* increases to 200 and 800, the latent parameter uncertainties in the obtained posterior ensemble further increase (Fig. S1), and the simulation results from the posterior ensemble show even larger uncertainties compared with that for l = 50 (Figs. S(d), 5(e), 6(d) and 6(e)). With a relatively large latent space dimensionality, the tracer BTCs, pressure and flow rate data can be matched but the associated uncertainties are significant. Further analysis of the predictive ability of the obtained posterior realizations is necessary to examine possible overfitting of the inversion results.



- Fig. 5 Comparison of tracer BTCs, pressure and flow rates between the "true" data and the
- 386 simulation results for the log-normal aperture scenario. (a) Inversion with a latent space
- dimensionality of l = 2. (b) l = 10. (c) l = 50. (d) l = 200. (e) l = 800. For tracer BTCs (the first to
- fourth columns), the upper row shows the results from prior realizations, and the lower row
- shows the results from posterior realizations. Note that we show the tracer BTCs from all the 720
- realizations (grey curves), as well as the corresponding 90% credible intervals (green shadings).
 For pressure difference and flow rate (the fifth column), we show the evolution of simulation
- results (grey circles) with respect to ES-MDA iterations. The "true" values are annotated by red
- 393 circles situated along the final iteration.

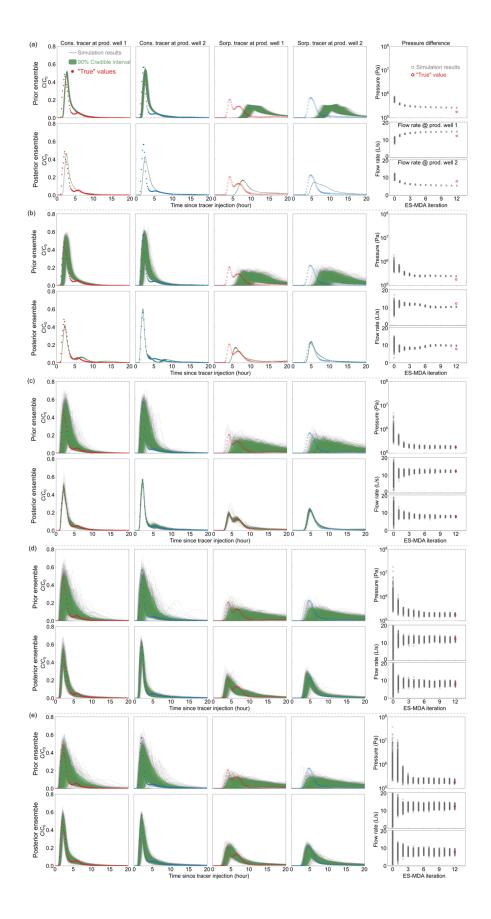


Fig. 6 Comparison of tracer BTCs, pressure and flow rates between the "true" data and the

simulation results for the two facies aperture scenario. (a) Inversion with a latent space dimensionality of l = 2. (b) l = 10. (c) l = 50. (d) l = 200. (e) l = 800.

398 4.2 Aperture distribution and flow field in the fracture

We now analyze the aperture distribution and flow field in the fracture (Figs. 7 and 8). 399 We generate aperture distributions from the prior and posterior ensembles, and then perform 400 flow simulations to obtain the corresponding flow fields. For both the log-normal and two facies 401 aperture scenarios, we observe some narrow flow channels connecting the injection and 402 production wells in the "true" flow fields (Fig. 3). When the latent space dimensionality is low (l 403 = 2), the aperture field generated from PCA is relatively smooth, and the posterior aperture 404 distributions cannot resolve these narrow flow channels (first column of Figs. 7 and 8). A larger 405 latent space dimensionality leads to a more heterogeneous aperture distribution and therefore a 406 407 more channelized flow field. When the latent space dimensionality is high, both the prior and posterior realizations exhibit some narrow channels between the injection and production wells 408 (Figs. 7 and 8). Of course, as the prior realization is not conditioned on the tracer, pressure and 409 410 flow rate data, the corresponding flow field is significantly different from the "true" flow field. For example, the flow field from a prior realization with l = 800 (second row, fifth column in 411 Fig. 7) shows two major channels connecting the injection well and production well 2, but 412 misses the channel between the injection well and production well 1. After ES-MDA, the 413 obtained posterior realization shows a flow field that resembles the "true" flow field better than 414 the prior realization does, especially for the high latent space dimensionality cases (fourth row in 415 Figs. 7 and 8). 416

However, not every flow channel in the "true" flow field is resolved by the posterior 417 realizations. For the log-normal aperture scenario, the "true" flow field shows four major and 418 several minor flow channels between the injection and production wells (Fig. 3(b)), while the 419 posterior realizations only resolve three major flow channels (sixth row in Fig. 7). For the two 420 facies aperture scenario, there are seven flow channels in the "true" flow field (Fig. 3(c)), but 421 only two major flow channels in the flow fields from the posterior realizations (sixth row in Fig. 422 8). Compared with the "true" flow fields, the flow fields from posterior realizations have fewer 423 flow channels but larger channel width. The relatively large channel width is necessary for the 424 posterior realizations to maintain comparable effective fracture areas as that in the "true" flow 425 fields, so that the tracer, pressure and flow rate data can be matched (especially the sorptive 426 tracer BTC which highly depends on the interaction area between fracture fluid and surrounding 427 rocks). The overall effect of the many narrow flow channels in the "true" flow field is 428 represented by the two or three relatively wide flow channels in the flow fields from posterior 429 realizations. 430

The latent space dimensionality shows significant effect on the posterior aperture 431 distributions. For a small latent space dimensionality (l = 2 or 10), the posterior aperture 432 distributions (flow fields) are almost identical, which is a direct consequence of latent parameter 433 collapse. Besides the randomly selected posterior realization in Figs. 7 and 8 (third and fourth 434 rows in Fig. 7 for the log-normal aperture scenario, and in Fig. 8 for the two facies aperture 435 scenario), we provide two additional, randomly selected posterior realizations for comparison in 436 the Supporting Information (Fig. S2). All the three aperture distributions (flow fields) are almost 437 the same as the average aperture distribution (flow field) displayed in the fifth and sixth rows in 438

439 Figs. 7 and 8. The standard deviation of the posterior realizations is negligible (Fig. S3 in the

440 Supporting Information). With the increase of latent space dimensionality, the variations among

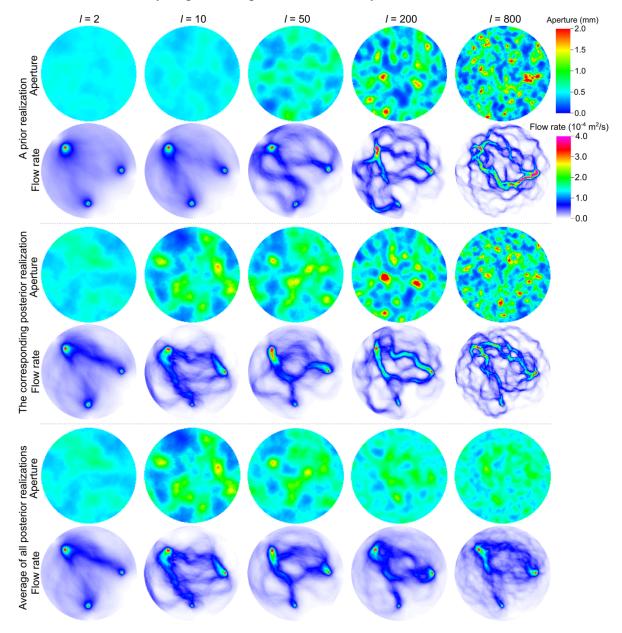
the posterior aperture distributions (flow fields) increase (Fig. S3). The uncertainty in the latent

442 parameters propagates to the aperture distribution and flow field. For a large latent space 443 dimensionality (l = 200 or 800), we observe considerable variations in aperture distribution

among the posterior realizations, but the major flow channels from these posterior realizations

445 are similar, as a result of conditioning on tracer, pressure and flow rate data.

In the next section, we will analyze how the uncertainty in the latent parameters further
propagates to thermal predictions and examine possible overfitting of the posterior realizations
obtained with a relatively large latent space dimensionality.



449

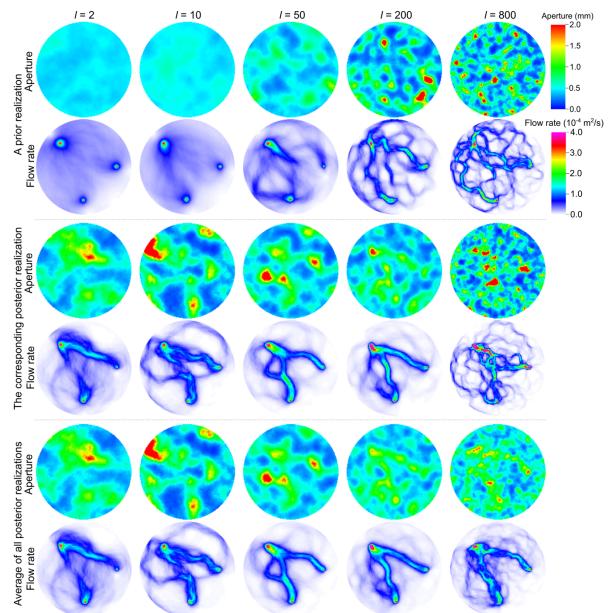
450 Fig. 7 Aperture distribution and flow field in the fracture for the log-normal aperture scenario.

451 The first and second rows are results from a randomly selected realization in the prior ensemble

452 (before ES-MDA), and the third and fourth rows are results from the corresponding posterior

453 realization (after 12 ES-MDA iterations). The fifth and sixth rows are average results of all the

454 realizations in the posterior ensemble.



455

Fig. 8 Aperture distribution and flow field in the fracture for the two facies aperture scenario.
The first and second rows are results from a randomly selected realization in the prior ensemble
(before ES-MDA), and the third and fourth rows are results from the corresponding realization in
the posterior ensemble (after 12 ES-MDA iterations). The fifth and sixth rows are average results

460 of all the realizations in the posterior ensemble.

461 4.3 Thermal performance prediction

In this sub-section, we perform thermal simulations with both prior and posterior realizations to analyze their abilities in predicting the long-term thermal performance of the EGS model. For each latent space dimensionality, we randomly select ten prior realizations and their corresponding posterior realizations to perform thermal simulation, and then compare the simulated temperature responses with the "true" temperature responses (Figs. 9 and 10).

We first analyze the prior predictions (first and third rows in Figs. 9 and 10). With the 467 increase of latent space dimensionality, the variations among the prior predictions of temperature 468 responses increase (first row in Figs. 9 and 10). For relatively large latent space dimensionalities, 469 the prior predictions vary in broad ranges and many predictions significantly underestimate the 470 temperature reductions. Interestingly, the variation among the prior predictions of the flow rate-471 averaged temperature response (third row in Figs. 9 and 10), although considerable for large 472 latent space dimensionality cases, is substantially smaller than that of the temperature responses 473 474 at individual production wells. To understand the reduced variation of the flow rate-averaged temperature response, we select a prior realization for further analysis (Fig. 11). For the prior 475 realization, the predicted flow rate at production well 1 is 2.7 L/s, much smaller than the "true" 476 flow rate (5.9 L/s). As a result, the predicted flow rate at production well 2 is much larger than 477 the corresponding "true" flow rate. The temperature decrease is highly related to the flow rate. 478 The underestimated flow rate at production well 1 results in a slow temperature decrease at 479 production well 1, and the overestimated flow at production well 2 leads to a fast temperature 480 decrease at production well 2. Since the flow rates at the two production wells are not 481 independent, the underestimation of flow rate at one production well means the overestimation of 482 flow rate at the other production well. Therefore, when the temperature decrease at one 483 production well is significantly underestimated (e.g.: production well 1 in Fig. 11(a)), the 484 temperature decrease at the other production well is likely to be overestimated (production well 2 485 in Fig. 11(a)). As a result, the variation of the temperature prediction at one production well 486 487 counteracts the variation of the temperature prediction at the other production well, causing a reduced variation in the flow-rated averaged temperature prediction. 488

Compared with prior realizations, the posterior realizations provide more accurate 489 predictions for both the individual and flow-rate averaged temperature responses (second and 490 fourth rows in Figs. 9 and 10). When latent space dimensionality is low (l = 2 or 10), temperature 491 predictions from the ten posterior realizations are almost identical due to the collapse of latent 492 parameters, and cannot match the "true" temperature responses. With the increase of latent space 493 dimensionality, the posterior predictions match the "true" temperature responses better but also 494 495 show larger uncertainties. When the latent space dimensionality increases to 800, the posterior predictions at individual production wells exhibit significant uncertainties, indicating the 496 overfitting of the obtained posterior realizations. However, the posterior prediction of the flow-497 498 rated averaged temperature response still successfully reproduces the "true" response with relatively small uncertainty (fourth row, fifth column in Figs. 9 and 10). Once again this is 499 mainly caused by the counteraction between the underestimation of temperature response at one 500 501 production well and the overestimation of temperature response at the other production well, as shown by the example in Fig. 11(b). 502

According to Figs. 9 and 10, to correctly predict the thermal responses at the two production wells, the latent space dimensionality should be in the range of $10 \sim 200$ for the lognormal aperture scenario, and $50 \sim 200$ for the two facies aperture scenario. While for the flow-

- ⁵⁰⁶ rated averaged temperature response, the latent space dimensionality should not be smaller than
- 507 ten for the log-normal aperture scenario, and a larger latent space dimensionality not smaller than
- 508 50 appears to be necessary for the two-facies aperture scenario.

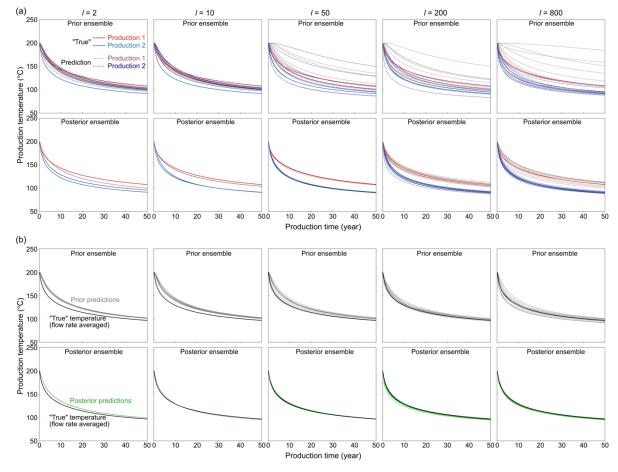


Fig. 9 Prediction of thermal responses from prior and posterior realizations for the log-normal

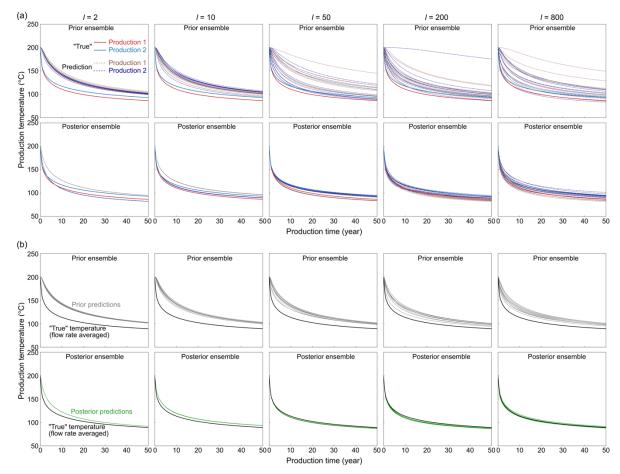
511 aperture scenario. (a) Temperature responses at the two production wells. The solid lines are 512 "true" temperature responses, and the dash lines are predictions. The upper row shows the

513 predictions from prior realizations, and the lower row shows the predictions from posterior

514 realizations. (b) Flow rate averaged temperature response. The black line is the "true"

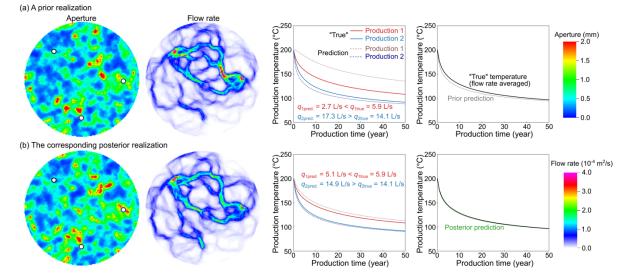
515 temperature response. The gray and green lines are predictions from prior (upper row) and

516 posterior (low row) realizations respectively.



518 Fig. 10 Prediction of thermal responses from prior and posterior realizations for the two facies

aperture scenario. (a) Temperature responses at the two production wells. (b) Flow rate averaged temperature response.



521

517

Fig. 11 Comparison of aperture distribution, flow field and temperature responses between a prior realization and the corresponding posterior realization. The prior and posterior realizations are from the ES-MDA case with a latent dimensionality of 800 for the log-normal aperture scenario.

Predicted $(q_{1\text{pred.}} \text{ and } q_{2\text{pred.}})$ and true $(q_{1\text{true}} \text{ and } q_{2\text{true}})$ flow rates at the two production wells are annotated.

527 **5 Discussion**

528 5.1 Selecting appropriate model complexity

Model reduction has been considered essential to tackle the challenges associated with 529 complex subsurface conditions and data scarcity in many subsurface inversion and 530 characterization problems (Jiang & Ou, 2017; Marzouk & Najm, 2009; Zhu & Zabaras, 2018). 531 The present study attempts to investigate the effect of model complexity on the inversion and 532 prediction of subsurface reservoirs, and more importantly, to provide some insights into the 533 selection of model complexity to avoid underfitting and overfitting. Through a field-scale EGS 534 model, we demonstrate both underfitting behavior under low model complexity (poor data 535 match) and overfitting behavior under high model complexity (good data match but poor 536 prediction). For the log-normal aperture scenario considered in the present study, an inversion 537 model that preserves 21% of the total variance in the "true" aperture field (corresponding to a 538 latent space dimensionality of l = 50) is sufficient to correctly reproduce tracer/pressure/flow rate 539 data and accurately predict long-term thermal performance. Increasing the model complexity to 540 preserve 56% of the total variance (l = 200) can also produce satisfactory data fit and thermal 541 prediction results but the associated uncertainties increase. Further increasing the model 542 complexity to preserve 84% of the total variance (l = 800) leads to significant uncertainties and 543 the thermal performance cannot be accurately predicted. Therefore, a model corresponds to a 544 latent space dimensionality between 50 and 200, i.e., with 21% to 56% of the total variance 545 preserved, is appropriate for tracer/pressure/flow rate data inversion and thermal prediction in the 546 present study. This is also true for the two facies aperture scenario where the "true" aperture field 547 and the aperture model used for inversion follow different statistical distributions. 548

549 The selection of model complexity actually depends on the purpose of inversion. For the presented EGS model, we note that although the posterior aperture fields obtained from the latent 550 space with l = 50 (Figs. 7, 8 and S2) can reproduce the "true" data and make accurate 551 predictions, they look different from the "true" aperture fields in Fig. 3. Many fine features in the 552 "true" aperture field could not be resolved due to the lack of necessary complexities in the 553 aperture model generated from such a low-dimensional latent space. Hence, if the primary goal is 554 to infer fracture aperture, a relatively high model complexity is required, but if the primary goal 555 is to predict thermal performance, then a moderate model complexity is sufficient. For many 556 subsurface characterization problems, people are mainly concerned with the predictive ability 557 rather than the realism of the inversion results, and therefore a moderate model complexity could 558 be employed. 559

560 The selection of model complexity also needs to consider the amount of information contained in the data for inversion. In general, the more information the data contain, the more 561 complex the model should be to avoid underfitting. In the present study, the inversion data 562 include tracer BTCs/flow rates at two production wells and pressure difference between injection 563 and production wells. The information in these data is spatially limited and far from sufficient to 564 characterize the aperture distribution in the 2D fracture plane. As a result, a relatively simple 565 model is able to retrieve the information and reproduce the data. Fortunately, since both tracer 566 transport and heat extraction processes are tightly related to the flow field among the injection 567

and production wells, the information retrieved from tracer BTCs, although limited, still provide

accurate thermal predictions. If more data, for example, tracer BTCs at other locations are

available, the model complexity needs to be increased to accommodate the increased amount ofinformation in the inversion data.

The result that a model preserving only 21% of the total variance in the "true" aperture 572 573 field is able to reproduce tracer/pressure/flow rate data and predict thermal responses is surprising as many studies preserved at least $50\% \sim 60\%$ total variance when using PCA for 574 model reduction (Fernández-Martínez et al., 2012; Hawkins et al., 2020; Laloy et al., 2013). An 575 important implication from the current study is that we should use a relatively simple model for 576 inversion/data assimilation in subsurface reservoirs, especially when the available measurements 577 are scarce and prediction is the primary goal. A low complexity model can not only mitigate the 578 579 overfitting pitfall but also alleviate the computational burden in many subsurface inversion problems. Of course, the model should not be too simple otherwise it may fail to reproduce 580 inversion data. 581

582 5.2 Geologic facies model

Geologic facies models have been used to describe highly channelized subsurface 583 reservoirs such as the two facies aperture model in Fig. 3(c). The characterization of such facies 584 models has been widely investigated in recent years (Chang et al., 2010; Jafarpour 585 & McLaughlin, 2009; Jiang & Jafarpour, 2021). To preserve the geologic realism of facies 586 models during inversion/data assimilation, a model reduction method that can directly generate 587 588 facies models from low-dimensional latent space is required. PCA is inappropriate as facies models do not follow a Gaussian or log-normal distribution. Many methods have been proposed 589 for the reduction of facies models, such as optimization-based PCA (Vo & Durlofsky, 2014), 590 discrete cosine transform (Jafarpour & McLaughlin, 2007), and deep learning algorithms such as 591 592 variational autoencoder (VAE) (Canchumuni et al., 2019; Laloy et al., 2017; Mo et al., 2020) and generative adversarial network (GAN) (Canchumuni et al., 2020; Laloy et al., 2018). 593

The current study provides an alternative strategy for the characterization of geologic 594 facies models. Instead of developing advanced model reduction methods for the two facies 595 aperture model in Fig. 3(c), we directly use a log-normal aperture model generated from PCA 596 latent space for data assimilation. Although the "true" and inversion aperture models follow 597 fundamentally different statistical distributions, the obtained posterior aperture models are able 598 to reproduce tracer/pressure/flow rate data and predict long-term thermal performance after data 599 assimilation. A log-normal aperture model with appropriate correlation length is capable of 600 inducing relevant channelized flow structures analogous to those of a facies-based model. The 601 posterior aperture distributions fail to preserve the geologic realism in the two facies aperture 602 model (Fig. 8). However, as Murray (2007) concluded, if prediction instead of explanation is the 603 primary goal, the realism of model parameters should not be considered an essential model-604 evaluation criterion. 605

6065.3 Prior realization

An interesting observation is that a prior aperture realization and the corresponding posterior aperture realization have many common features, especially when the aperture model is relatively complex (fourth and fifth columns in Figs. 7 and 8, Fig. 11). In another words, ES-MDA tends to perturb a prior realization as slightly as possible to match the data being assimilated. The obtained posterior realization largely depends on the prior realization provided

- to ES-MDA. If the prior realization is not well constrained or even physically unrealistic, then
- the corresponding posterior realization may also show unrealistic features. Therefore, it is of
- 614 great importance to constrain prior realizations with available prior knowledge. Fortunately, for 615 subsurface reservoirs, prior knowledge of the field of interest (aperture or permeability) can be
- subsurface reservoirs, prior knowledge of the field of interest (aperture or permeability) can be obtained from geological/geophysical measurements, such as core logs, wellbore images and
- 617 outcrop analysis. In the current study, the prior knowledge used to constrain prior aperture
- realizations includes the spatially autocorrelated nature as well as the mean, standard deviation
- and correlation length of the aperture field.

Interpretation results from other geophysical investigations, such as seismic and ground 620 penetrating radar (GPR), can also be used to constrain prior realizations. For example, Wu, Fu, 621 Hawkins, et al. (2021) used the results of GPR survey to constrain prior aperture realizations in a 622 horizontal fracture at a meso-scale field test site (the Altona Field Laboratory located in northern 623 New York State, USA). The GPR survey of the field test site indicated a narrow flow channel 624 between injection and production wells, which meant that the underlying aperture field was 625 anisotropic with larger correlation length in the direction from injection to production wells 626 (west to east) than that in the south to north direction. During subsequent tracer data assimilation, 627 such an anisotropic feature was used as prior knowledge to constrain prior aperture realizations. 628

629 6 Conclusion

We investigated the effect of model complexity on the inversion of fracture aperture 630 631 distribution as well as the prediction of long-term thermal recovery in a field-scale EGS model. Inversion models with different complexities were used to invert for fracture aperture 632 distribution through the assimilation of tracer/pressure/flow rate data using an ensemble-based 633 method (ES-MDA). Thermal simulations were then performed to examine the predictive ability 634 of the inferred aperture distributions. With a low model complexity, ensemble collapse occurred. 635 The inferred aperture distributions failed to reproduce tracer/pressure/flow rate data, and the 636 predicted long-term thermal response was biased. With a high complexity model, the data could 637 be properly matched, but the inferred aperture distribution and predicted thermal response 638 exhibit significant uncertainties. A moderate model complexity is sufficient to retrieve the 639 information contained in tracer/pressure/flow rate data and provide accurate thermal predictions. 640

An appropriate model complexity is essential to the inversion and prediction of 641 subsurface reservoirs, and deserves careful deliberation based on the primary purpose of the 642 inversion as well as the type and amount of the inversion data. In a real-world application, it is 643 difficult to predetermine model complexity and one might need to manually adjust model 644 complexity in a trial-and-error manner. According to the results in the current study, we 645 recommend starting with a relatively simple model rather than an extremely complex model, and 646 the quality of the fit to tracer BTCs appears to be a reasonable indicator of an appropriate model 647 complexity. 648

649

650 Data Availability Statement

The synthetic flow, pressure, tracer and thermal data used in this study is obtained from

numerical simulations with GEOS. The data assimilation framework is available in Wu, Fu,

653 Hawkins, et al. (2021).

Acknowledgments, Samples, and Data

We thank Dr. Pengcheng Fu for providing critical and constructive comments on various aspects of this study. This research was performed in support of the EGS Collab project and the EGS Collab team is gratefully acknowledged. This work was supported by U.S. Department of Energy, Geothermal Technologies Office, and performed under the auspices of the U.S.

660 Department of Energy by Lawrence Livermore National Laboratory under Contract DE-AC52-

661 07NA27344. This document is LLNL report LLNL-JRNL-838540. This work is also supported

by the National Key Research and Development Program of China (No. 2021YFA0716000), and

the China National Petroleum Corporation-Peking University Strategic Cooperation Project ofFundamental Research.

665 **References**

- Berkowitz, B. (2002). Characterizing flow and transport in fractured geological media: A review.
 Advances in Water Resources, 25, 861-884. https://doi.org/10.1016/S0309 1708(02)00042-8
- Canchumuni, S. W. A., Castro, J. D. B., Potratz, J., Emerick, A. A., & Pacheco, M. A. C. (2020).
 Recent developments combining ensemble smoother and deep generative networks for
 facies history matching. *Computational Geosciences*, 25, 433-466.
 https://doi.org/10.1007/s10596-020-10015-0
- Canchumuni, S. W. A., Emerick, A. A., & Pacheco, M. A. C. (2019). Towards a robust
 parameterization for conditioning facies models using deep variational autoencoders and
 ensemble smoother. *Computers & Geosciences*, 128, 87-102.
 https://doi.org/10.1016/j.cageo.2019.04.006
- Chang, H., Zhang, D., & Lu, Z. (2010). History matching of facies distribution with the EnKF
 and level set parameterization. *Journal of Computational Physics*, 229, 8011-8030.
 https://doi.org/10.1016/j.jcp.2010.07.005
- Chen, X., Hammond, G. E., Murray, C. J., Rockhold, M. L., Vermeul, V. R., & Zachara, J. M.
 (2013). Application of ensemble-based data assimilation techniques for aquifer
 characterization using tracer data at Hanford 300 area. *Water Resources Research*,
 49(10), 7064-7076. https:// doi.org/10.1002/2012WR013285
- Chen, Y., & Zhao, Z. (2020). Heat transfer in a 3D rough rock fracture with heterogeneous
 apertures. *International Journal of Rock Mechanics and Mining Sciences*, 134, 104445.
 https://doi.org/10.1016/j.ijrmms.2020.104445
- Cox, S. F., Knackstedt, M. A., & Braun, J. (2001). Principles of structural control on
 permeability and fluid flow in hydrothermal systems. *Reviews in Economic Geology*, 14,
 1-24. https://doi.org/10.5382/Rev.14.01
- Dobson, P. F., Kneafsey, T. J., Hulen, J., & Simmons, A. (2003). Porosity, permeability, and
 fluid flow in the Yellowstone geothermal system, Wyoming. *Journal of Volcanology and Geothermal Research*, 123, 313-324. https://doi.org/10.1016/S0377-0273(03)00039-8
- Emerick, A. A. (2017). Investigation on principal component analysis parameterizations for
 history matching channelized facies models with ensemble-based data
- assimilation. *Mathematical Geosciences*, 49(1), 85-120. https://doi.org/10.1007/s11004 016-9659-5

- Emerick, A. A. (2018). Deterministic ensemble smoother with multiple data assimilation as an
 alternative for history-matching seismic data. *Computational Geosciences*, 22(5), 1175 1186. https://doi.org/10.1007/s10596-018-9745-5
- Emerick, A. A., & Reynolds, A. C. (2013). Ensemble smoother with multiple data assimilation.
 Computers & Geosciences, 55, 3-15. https://doi.org/10.1016/j.cageo.2012.03.011
- Fernández-Martínez, J. L., Mukerji, T., García-Gonzalo, E., & Suman, A. (2012). Reservoir
 characterization and inversion uncertainty via a family of particle swarm optimizers.
 Geophysics, 77(1), M1-M16. https://doi.org/10.1190/geo2011-0041.1
- Fu, P., Hao, Y., Walsh, S. D. C., & Carrigan, C. R. (2016). Thermal drawdown-induced flow
 channeling in fractured geothermal reservoirs. *Rock Mechanics and Rock Engineering*,
 49(3), 1001–1024. https://doi.org/10.1007/s00603-015-0776-0
- Fu, P., Johnson, S. M., & Carrigan, C. R. (2013). An explicitly coupled hydro-geomechanical model for simulating hydraulic fracturing in arbitrary discrete fracture networks.
 International Journal for Numerical and Analytical Methods in Geomechanics, 37, 2278-2300. https://doi.org/10.1002/nag.2135
- Guo, B., Fu, P., Hao, Y., & Carrigan, C. R. (2016). Investigating the possibility of using tracer
 tests for early identification of EGS reservoirs prone to flow channeling. 41st workshop
 on geothermal reservoir engineering.
- Guo, B., Fu, P., Hao, Y., Peters, C. A., & Carrigan, C. R. (2016). Thermal drawdown-induced
 flow channeling in a single fracture in EGS. *Geothermics*, 61, 46-62.
 https://doi.org/10.1016/j.geothermics.2016.01.004
- Hawkins, A. J., Fox, D. B., Koch, D. L., Becker, M. W., & Tester, J. W. (2020). Predictive
 inverse model for advective heat transfer in a short-circuited fracture: Dimensional
 analysis, machine learning, and field demonstration. *Water Resources Research*, 56,
 e2020WR027065. https:// doi.org/10.1029/2020WR027065
- Jafarpour, B., & McLaughlin, D. B. (2007). Efficient permeability parameterization with the
 discrete cosine transform. SPE reservoir simulation symposium.
- Jafarpour, B., & McLaughlin, D. B. (2009), Estimating channelized-reservoir permeabilities with
 the ensemble Kalman filter: The importance of ensemble design. *SPE Journal*, 14(2),
 374-388. https://doi.org/10.2118/108941-PA
- Jiang, A., & Jafarpour, B. (2021). Deep convolutional autoencoders for robust flow model
 calibration under uncertainty in geologic continuity. *Water Resources Research*, 57(11),
 e2021WR029754. https://doi.org/10.1029/2021WR029754
- Jiang, L., & Ou, N. (2017). Multiscale model reduction method for Bayesian inverse problems of
 subsurface flow. *Journal of Computational and Applied Mathematics*, 319, 188-209.
 https://doi.org/10.1016/j.cam.2017.01.007
- Jiang, Z., Zhang, S., Turnadge, C., & Xu, T. (2021). Combining autoencoder neural network and
 Bayesian inversion to estimate heterogeneous permeability distributions in enhanced
 geothermal reservoir: Model development and verification. *Geothermics*, 97, 102262.
 https://doi.org/10.1016/j.geothermics.2021.102262
- Johnson, T. C., Burghardt, J., Strickland, C., Knox, H., Vermeul, V., White, M., et al. (2021). 4D
 proxy imaging of fracture dilation and stress shadowing using electrical resistivity
 tomography during high pressure injections into a dense rock formation. *Journal of*
- 740 *Geophysical Research: Solid Earth*, 126, e2021JB022298.
- 741 https://doi.org/10.1029/2021JB022298

Laloy, E., Hérault, R., Jacques, D., & Linde, N. (2018). Training-image based geostatistical 742 743 inversion using a spatial generative adversarial neural network. Water Resources Research, 54, 381-406. https://doi.org/10.1002/2017WR022148 744 745 Laloy, E., Hérault, R., Lee, J., Jacques, D., & Linde, N. (2017). Inversion using a new lowdimensional representation of complex binary geo-logical media based on a deep neural 746 network. Advances in Water Resources, 110, 387-405. 747 https://doi.org/10.1016/j.advwatres.2017.09.029 748 Laloy, E., Rogiers, B., Vrugt, J. A., Mallants, D., & Jacques, D. (2013). Efficient posterior 749 exploration of a high-dimensional groundwater model from two-stage Markov chain 750 Monte Carlo simulation and polynomial chaos expansion. Water Resources Research, 751 49(5), 2664-2682. https://doi.org/10.1002/wrcr.20226 752 Li, W., & Cirpka, O. A. (2006). Efficient geostatistical inverse methods for structured and 753 unstructured grids. Water Resources Research, 42, W06402. 754 https://doi.org/10.1029/2005WR004668 755 Liu, M., & Grana, D. (2020). Petrophysical characterization of deep saline aquifers for CO₂ 756 storage using ensemble smoother and deep convolutional autoencoder. Advances in 757 Water Resources, 142, 103634. https://doi.org/10.1016/j.advwatres.2020.103634 758 Liu, Y., & Durlofsky, L. J. (2020). 3D CNN-PCA: A deep-learning-based parameterization for 759 complex geomodels. Computers & Geosciences, 148, 104676. 760 https://doi.org/10.1016/j.cageo.2020.104676 761 Marzouk, Y. M., & Najm, H. N. (2009). Dimensionality reduction and polynomial chaos 762 acceleration of Bayesian inference in inverse problems. Journal of Computational 763 *Physics*, 228(6), 1862-1902. https://doi.org/10.1016/j.jcp.2008.11.024 764 Mo, S., Zabaras, N. J., Shi, X., & Wu, J. (2020). Integration of adversarial autoencoders with 765 residual dense convolutional networks for estimation of non-Gaussian hydraulic 766 767 conductivities. Water Resources Research, 56, e2019WR026082. https://doi.org/10. 1029/2019WR026082 768 Murray, A. B. (2007). Reducing model complexity for explanation and prediction. 769 Geomorphology, 90, 178-191. https://doi.org/10.1016/j.geomorph.2006.10.020 770 Nejadi, S., Trivedi, J. J., & Leung, J. (2017). History matching and uncertainty quantification of 771 discrete fracture network models in fractured reservoirs. Journal of Petroleum Science 772 773 and Engineering, 152, 21-32. https://doi.org/10.1016/j.petrol.2017.01.048 Okoroafor, E. R., Co, C., & Horne, R. N. (2022). Numerical investigation of the impact of 774 fracture aperture anisotropy on EGS thermal performance. Geothermics, 100, 775 102354. https://doi.org/10.1016/j.geothermics.2022.102354 776 Remy, N., Boucher, A., & Wu, J. (2009). Applied geostatistics with SGeMS: A user's guide. 777 Cambridge University Press, Cambridge, U. K. 778 779 Romary, T. (2009). Integrating production data under uncertainty by parallel interacting Markov 780 chains on a reduced dimensional space. Computational Geosciences, 13, 103-122. https://doi.org/10.1007/s10596-008-9108-8 781 Sarma, P., Durlofsky, L. J., & Aziz, K. (2008). Kernel principal component analysis for efficient 782 differentiable parameterization of multipoint geostatistics. Mathematical Geosciences, 783 40, 3-32. https://doi.org/10.1007/s11004-007-9131-7 784 785 Settgast, R. R., Fu, P., Walsh, S. D. C., White, J. A., Annavarapu, C., & Ryerson, F. J. (2017). A 786 fully coupled method for massively parallel simulation of hydraulically driven fractures

787 in 3-dimensions. International Journal for Numerical and Analytical Methods in 788 Geomechanics, 41, 627-653. https://doi.org/10.1002/nag.2557 Shi, Z., Zhang, S., Yan, R., & Wang, G. (2018). Fault zone permeability decrease following 789 790 large earthquakes in a hydrothermal system. Geophysical Research Letters, 45, 1387-1394. https://doi.org/10.1002/2017GL075821 791 792 Somogyvári, Y., Jalali, M., Jimenez Parras, S., Bayer, P. (2017). Synthetic fracture network 793 characterization with transdimensional inversion. Water Resources Research, 53(6), 794 5104-5123. https://doi.org/10.1002/2016WR020293 Strebelle, S. (2002). Conditional simulation of complex geological structures using multiple-795 796 point statistics. *Mathematical geology*, 34(1), 1-21. https://doi.org/10.1023/A:1014009426274 797 Tang, H., Fu, P., Sherman, C. S., Zhang, J., Ju, X., Hamon, F., et al. (2021). A deep learning-798 799 accelerated data assimilation and forecasting workflow for commercial-scale geologic carbon storage. International Journal of Greenhouse Gas Control, 112, 103488. 800 https://doi.org/10.1016/j.jggc.2021.103488 801 Vo, H. X., & Durlofsky, L. J. (2014). A new differentiable parameterization based on principal 802 component analysis for the low-dimensional rep- resentation of complex geological 803 models. Mathematical Geosciences, 46(7), 775-813. https://doi.org/10.1007/s11004-014-804 9541-2 805 Vogler, D., Settgast, R. R., Annavarapu, C., Madonna, C., Bayer, P., & Amann, F. (2018). 806 Experiments and simulations of fully hydro-mechanically coupled response of rough 807 fractures exposed to high-pressure fluid injection. Journal of Geophysical Research: 808 Solid Earth, 123, 1186-1200. https://doi.org/10.1002/2017JB015057 809 Vogt, C., Marquart, G., Kosack, C., Wolf, A., & Clauser, C. (2012). Estimating the permeability 810 distribution and its uncertainty at the EGS demonstration reservoir Soultz-sous-Forêts 811 812 using the ensemble Kalman filter. Water Resources Research, 48, W08517. https://doi. org/10.1029/2011WR011673 813 Wu, H., Fu, P., Morris, J. P., Mattson, E. D., Neupane, G., Smith, M. M., et al., & EGS. (2021). 814 Characterization of flow and transport in a fracture network at the EGS Collab field 815 experiment through stochastic modeling of tracer recovery. Journal of Hydrology, 593, 816 125888. https://doi.org/10.1016/j.jhydrol.2020.125888 817 Wu, H., Fu, P., Hawkins, A. J., Tang, H., & Morris, J. P. (2021). Predicting thermal performance 818 of an enhanced geothermal system from tracer tests in a data assimilation framework. 819 Water Resources Research, 57, e2021WR030987. 820 https://doi.org/10.1029/2021WR030987 821 Wu, H., Fu, P., Yang, X., Morris, J. P., Johnson, T. C., Settgast, R. R., & Ryerson, F. J. (2019). 822 Accurate imaging of hydraulic fractures using templated electrical resistivity 823 tomography. Geothermics, 81, 74-87. https://doi.org/10.1016/j.geothermics.2019.04.004 824 825 Xiao, C., & Tian, L. (2020). Surrogate-based joint estimation of subsurface geological and relative permeability parameters for high-dimensional inverse problem by use of smooth 826 local parameterization. Water Resources Research, 56, e2019WR025366. https://doi.org/ 827 10.1029/2019WR025366 828 829 Xiao, C., Zhang, S., Ma, X., Jin, J., & Zhou, T. (2022). Model-reduced adjoint-based inversion using deep- learning: Example of geological carbon sequestration modeling. Water 830 831 Resources Research, 58, e2021WR031041. https://doi.org/10.1029/2021WR031041

- Yang, H., Lin, Y., Wohlberg, B., & Tartakovsky, D. M. (2021). Consensus equilibrium for
 subsurface delineation. *Water Resources Research*, 57(10), e2021WR030151.
 https://doi.org/10.1029/2021WR030151
- Zhang, J., Zheng, Q., Chen, D., Wu, L., & Zeng, L. (2020). Surrogate-based bayesian inverse
 modeling of the hydrological system: An adaptive approach considering surrogate
 approximation error. *Water Resources Research*, 56, e2019WR025721. https://
 doi.org/10.1029/2019WR025721
- Zhao, Y., & Luo, J. (2020). Reformulation of Bayesian geostatistical approach on principal
 components. *Water Resources Research*, 55, e2019WR026732. https://doi.org/
 10.1029/2019WR026732
- Zhu, Y., & Zabaras, N. (2018). Bayesian deep convolutional encoder-decoder networks for
 surrogate modeling and uncertainty quantification. *Journal of Computational Physics*,
 366, 415-447. https://doi.org/10.1016/j.jcp.2018.04.018

AGU PUBLICATIONS

Water Resources Research

Supporting Information for

Selecting appropriate model complexity: An example of tracer inversion for thermal prediction in enhanced geothermal systems

Hui Wu^{1, 2}, Zhijun Jin^{1, 2, 3}, Su Jiang⁴, Hewei Tang⁵, Joseph P. Morris⁵, Jinjiang Zhang¹, Bo Zhang¹

¹ School of Earth and Space Sciences, Peking University, Beijing, China.

² Institute of Energy, Peking University, Beijing, China

³ Petroleum Exploration and Production Research Institute, SINOPEC, Beijing, 100083, China

⁴ Department of Energy Resources Engineering, Stanford University, Stanford, CA, USA.

⁵ Lawrence Livermore National Laboratory, Livermore, CA, USA.

Contents of this file

Figures S1 to S3

Introduction

This supporting information provides the box plots of prior and posterior latent parameters in Fig. S1. Fig. S2 shows two randomly selected posterior realizations in terms of aperture distribution and flow field. Fig. S3 shows the standard deviation of aperture distribution/flow field in posterior ensemble.

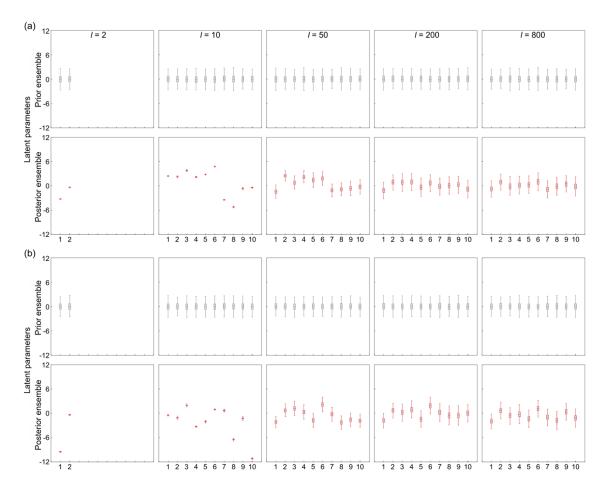


Figure S1. Box plots of ten latent parameters from prior (upper row) and posterior (lower row) ensembles. The ten latent parameters correspond to the ten most significant principal components after PCA. (a) Log-normal aperture field scenario. (b) Two facies aperture field scenario. For the latent space with a dimensionality (*l*) of two, only two latent parameters are shown.

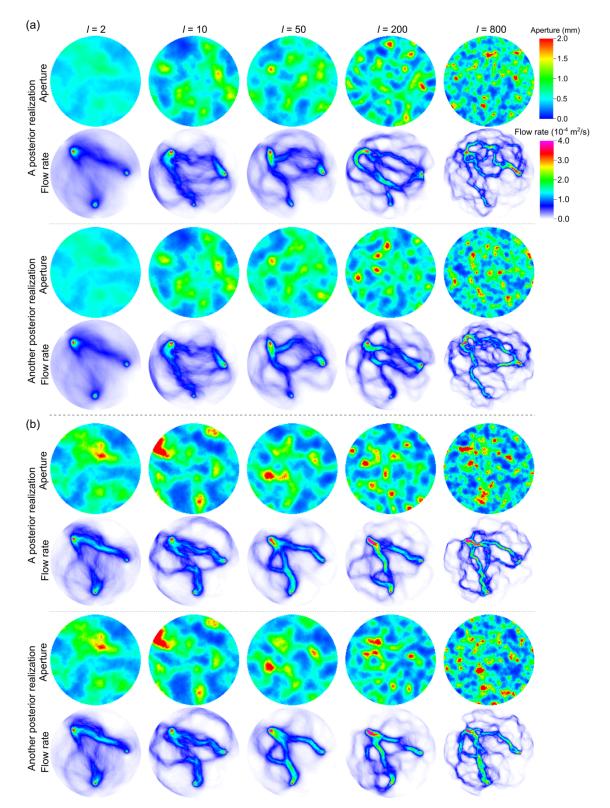


Figure S2. Aperture distribution and flow field from two randomly selected posterior realizations. (a) Log-normal aperture field scenario. (b) Two facies aperture field scenario.

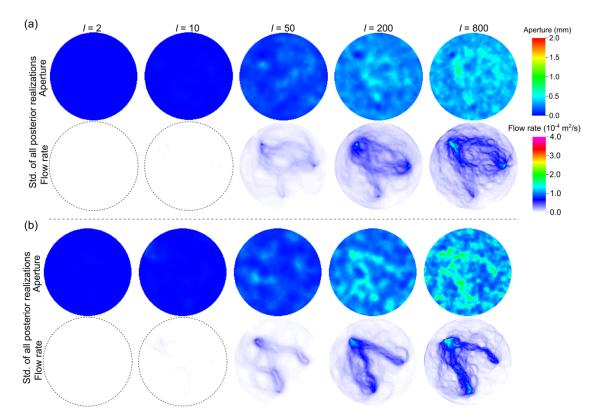


Figure S3. Standard deviation (Std.) of aperture distribution and flow field in the posterior ensemble. (a) Log-normal aperture field scenario. (b) Two facies aperture field scenario.