A Spatiotemporal-Aware Climate Model Ensembling Method for Improving Precipitation Predictability

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Abstract

Multimodel ensembling has been widely used to improve climate model predictions, and the improvement strongly depends on the ensembling scheme. In this work, we propose a Bayesian neural network (BNN) ensembling method, which combines climate models within a Bayesian model averaging framework, to improve the predictive capability of model ensembles. Our proposed BNN approach calculates spatiotemporally varying model weights and biases by leveraging individual models' simulation skill, calibrates the ensemble prediction against observations by considering observation data uncertainty, and quantifies epistemic uncertainty when extrapolating to new conditions. More importantly, the BNN method provides interpretability about which climate model contributes more to the ensemble prediction at which locations and times. Thus, beyond its predictive capability, the method also brings insights and understanding of the models to guide further model and data development. In this study, we apply the BNN weighting scheme to an ensemble of CMIP6 climate models for monthly precipitation prediction over the conterminous United States. In both synthetic and real case studies, we demonstrate that BNN produces predictions of monthly precipitation with higher accuracy than three baseline ensembling methods. BNN can correctly assign a larger weight to the regions and seasons where the individual model fits the observation better. Moreover, its offered interpretability is consistent with our understanding of localized climate model performance. Additionally, BNN shows an increasing uncertainty when the prediction is farther away from the period with constrained data, which appropriately reflects our predictive confidence and trustworthiness of the models in the changing climate.

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6 Key Points:

- We develop a spatiotemporal-aware weighting scheme using Bayesian neural networks for improving model ensemble predictions
 The method calculates model skill-consistent weights, provides interpretability, and
 - quantifies uncertainty
 - We demonstrate the method's superior performance over three baseline ensembling methods in predicting precipitation in CONUS

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14 Abstract

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³⁶ Plain Language Summary

Precipitation is one of the key climatic factors affecting fluxes of water, energy, and 37 biogeochemical cycles. Global climate models (GCMs) are usually used for improving 38 precipitation prediction and advancing understanding of precipitation's responses to cli-39 mate change. A large set of GCMs are available and they show large uncertainties in phys-40 ical process representations, varying prediction skills at different locations and times, and 41 are usually not constrained by observations. Here, we propose a Bayesian neural network 42 ensembling method to address these challenges and thus improve precipitation predictabil-43 ity by providing accurate and uncertainty-aware predictions. 44

45 **1** Introduction

Precipitation is one of the key climatic factors affecting fluxes of water, energy, and 46 biogeochemical cycles. It has been observed that climate change non-uniformly shifts re-47 gional and seasonal distributions of the precipitation, where dry regions/seasons get drier 48 and wet regions/seasons get wetter (Stegall & Kunkel, 2019). This shift of precipitation 49 patterns significantly affects natural ecosystem health and human society development 50 (E. Martin, 2018; Greve et al., 2014). For instance, in humid regions, the heavy precip-51 itation can increase flood and landslide risks, degrade water quality for human consump-52 tion, and disrupt regional ecosystem balance. In arid regions, the decreased precipita-53 tion can exacerbate droughts, which leads to water shortages, agricultural production 54 loss, and energy supply risks. Therefore, improving our ability to accurately predict cur-55 rent and future patterns in precipitation is vital for assessing the vulnerability of ecosys-56 tems, preparing for extreme precipitation events, and concurrently enhancing water re-57 sources management (Konapala et al., 2020). 58

Global climate models (GCMs) have been used for improving precipitation prediction and advancing understanding of precipitation's responses to climate change (Weigel et al., 2021; Demory et al., 2020). One of the most inclusive sets of GCMs is from the Coupled Model Intercomparison Project (CMIP), initialized by the Working Group on Coupled Modeling under the organization of the World Climate Research Program (Eyring

et al., 2016; Taylor et al., 2012). CMIP is now in its sixth phase. CMIP6 consists of about 64 100 GCMs produced by 49 different modeling groups/institutes (Zelazowski et al., 2018). 65 These GCMs have large uncertainties in physical process representations, show varying 66 prediction skills at different locations and times, and are usually not constrained by ob-67 servations (Eyring et al., 2019). Each of these aforementioned factors affects the accu-68 rate prediction of precipitation at regional scales. One strategy that can improve the pre-69 cipitation prediction is a comprehensive multi-model ensembling approach that lever-70 ages each individual model's spatiotemporally varying predictive skill, integrates obser-71 vations to reduce prediction bias, and quantifies predictive uncertainty using a formal 72 calibration and uncertainty quantification (UQ) framework (Que et al., 2020; Fothering-73 ham et al., 2015). 74

Several multi-model ensembling methods have been developed. Some approaches 75 assume model independence and model democracy, in which each model is weighted equally. 76 Although studies have demonstrated that under certain conditions equal-weight model 77 averaging could produce better prediction performance than the individual models (Gleckler 78 et al., 2008; Knutti et al., 2010; Pincus et al., 2008), the assumption on model indepen-79 dence and democracy is not true. Many GCMs in CMIP share components or are vari-80 ants of other models in the ensemble, and these models have large inconsistency in their 81 skills at a given location and time (Alexander & Easterbrook, 2015; Abramowitz & Bishop, 82 2015; Sanderson et al., 2015; Bishop & Abramowitz, 2013). Even an individual model 83 shows considerably inconsistent skills in different locations and at different times. By rec-84 ognizing the distinct capabilities among the models, some studies assigned unequal weights 85 to individual ensemble members (Amos et al., 2020; Brunner et al., 2019; Wenzel et al., 86 2016; Karpechko et al., 2013; Räisänen et al., 2010). One of the most frequently adopted 87 ensemble weighting schemes was proposed by Sanderson et al. (2015). It calculates model 88 weights by balancing the model skill and model uniqueness; the coefficient controlling 89 the balance is determined subjectively, and its value could significantly impact the en-90 semble results (Knutti et al., 2017; Sanderson et al., 2017). 91

Although some weighted average methods have been proposed, the unequal weights 92 assigned to the individual models are mostly uncalibrated against the observations, an 93 uniform weight is assigned to a model across the space and time, and the same weight 94 is applied for future projections without UQ. Since the model skill varies at regional and 95 seasonal scales, the spatiotemporally uniform weight does not fully leverage each indi-96 vidual model's capability, resulting in the loss of information and possibly large biases 97 in predicting the distribution of precipitation (G. M. Martin et al., 2017; Kumar et al., 98 2014). Stegall and Kunkel (2019) discovered that assigning unequal but spatiotemporally uniform weights to individual models can improve the mean prediction of the pre-100 cipitation, but the estimated regional precipitation distribution still had a large incon-101 sistency with the observations. Additionally, the model weights need to be calibrated against 102 observations in each grid cell at each time step to reasonably reflect the individual model's 103 spatiotemporally varying skill in fitting the observed data and produce observationally 104 constrained ensemble predictions. Studies have shown that many models contributing 105 to CMIP yielded large discrepancies compared with observations, and these model bi-106 ases should be reduced by calibration before being used for prediction (Ukkola et al., 2020; 107 Lorenz et al., 2018; Mueller & Seneviratne, 2014). Finally, UQ is required for the ensem-108 ble prediction to avoid overconfidence—especially when we project the precipitation in 109 the future changing climate. 110

In this work, we propose a Bayesian neural network (BNN) ensembling method to improve precipitation predictability by providing accurate and uncertainty-aware predictions. The BNN ensembling approach combines GCMs within a Bayesian model averaging framework. It calculates spatiotemporally varying model weights and biases, calibrates the weights and biases against observations, and accounts for the varying quality of the observed data. Additionally, the BNN method quantifies epistemic uncertainty when extrapolating the prediction to new conditions. More importantly, BNN also provides interpretability about each individual model's contribution to the ensemble prediction in different regions and at different times.

The proposed BNN ensembling scheme overcomes the limitations of existing meth-120 ods by leveraging the power of machine learning (ML) in data analytics and predictive 121 analytics. ML techniques have been applied for predicting precipitation (Jose et al., 2022; 122 Heinze-Deml et al., 2021; Li et al., 2021; Ahmed et al., 2020). Most of these applications 123 used ML methods either as a surrogate model of an individual GCM to reduce compu-124 tational costs in simulation or as a data-driven, black-box regression model to simulate 125 the precipitation directly. The former application considers only a single GCM, and the 126 latter regression model simulation lacks mechanical interpretation and process under-127 standing. Here, we use ML techniques in the context of multiple model analysis to cal-128 culate the model weights of an ensemble of GCMs. The proposed BNN weighting strat-129 egy sufficiently leverages each individual GCM's diverse performance in heterogeneous 130 geography and different seasons by calculating spatiotemporally varying model weights 131 and biases. By fusing diverse GCMs, the BNN ensembling embeds our best physical knowl-132 edge; and by constraining the ensemble prediction with the observations, BNN enables 133 accurate predictions that match the historical data. Additionally, the BNN method quan-134 tifies both aleatoric uncertainty from the data noise and epistemic uncertainty when pro-135 jecting to the unknown future. Furthermore, besides providing high-quality ensemble pre-136 dictions with UQ, our method also brings insights and understanding of the climate model 137 performance to guide further model development and prioritize data collection. 138

We apply the BNN ensembling method for monthly precipitation prediction over the conterminous United States (CONUS). We consider an ensemble of GCMs from CMIP6 and use the European Centre Reanalysis Data (ERA5) as "observations" for model calibration and performance evaluation. We perform both synthetic and real case studies to verify, evaluate, and demonstrate the method's capability with respect to prediction accuracy, interpretability, and UQ. The main contributions of this effort are as follows.

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• We propose a BNN ensembling approach for precipitation prediction by leveraging individual GCM's spatiotemporally varying skill and calibrating the model weights and biases against the observations.

 We demonstrate the superior prediction performance of the proposed method in comparison with three widely used ensembling approaches on GCMs from CMIP6 and additionally show that BNN can reasonably calculate the epistemic uncertainty in extrapolation to avoid overconfident projections.

• We investigate the interpretability of the BNN method in terms of which GCMs contribute more to the ensemble prediction at which locations and times and demonstrate that the calculated spatiotemporally varying weights are consistent with the GCMs' simulation skill.

¹⁵⁷ 2 Methods and Data

In this section, we introduce the BNN ensembling method and describe the climate models and the precipitation data. Next, we briefly introduce three state-of-the-art ensembling schemes with which we compare the BNN for performance evaluation. Lastly, we discuss some evaluation metrics.

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2.1 Bayesian Neural Networks for Ensemble Model Predictions

We assume that observations $y(\mathbf{x}, t)$ at a given location \mathbf{x} and time t can be represented as a sum over an ensemble of m GCM predictions $M_i(\mathbf{x}, t)$ weighted by their

respective weights $\alpha_i(\mathbf{x}, t)$, a bias term $\beta(\mathbf{x}, t)$, and a data noise term $\epsilon(\mathbf{x}, t)$:

$$y(\mathbf{x},t) = \sum_{i=1}^{m} \alpha_i(\mathbf{x},t) M_i(\mathbf{x},t) + \beta(\mathbf{x},t) + \epsilon(\mathbf{x},t).$$
(1)

The model weights are positive and their sum over the ensemble models is one: $\alpha_i(\mathbf{x},t) >$ 163 0, and $\sum_{i=1}^{m} \alpha_i(\mathbf{x}, t) = 1$. The model bias $\beta(\mathbf{x}, t)$ represents the discrepancy of the weighted 164 ensemble model simulations from the observation. The data noise $\epsilon(\mathbf{x}, t)$ considers the 165 observation quality varying across the location and time, which is assumed following a 166 Gaussian distribution with a zero mean and a heteroscedastic standard deviation $\sigma(\mathbf{x}, t)$. 167 The combination of the first two terms at the right-hand side of Eq. (1) forms the BNN ensemble model prediction: $\hat{y}(\mathbf{x},t) = \sum_{i=1}^{m} \alpha_i(\mathbf{x},t) M_i(\mathbf{x},t) + \beta(\mathbf{x},t)$. This ensembling scheme expresses the model weights as a function of location and time to leverage in-168 169 170 dividual models' spatiotemporally varying simulation skills. The ensemble prediction ad-171 ditionally considers a bias term that is also a function of space and time. Incorporat-172 ing the bias term in ensembling is crucial, especially when all the individual models have 173 an over- or under-prediction. In this situation, the weighted ensemble model simulations 174 $\sum_{i=1}^{m} \alpha_i(\mathbf{x}, t) M_i(\mathbf{x}, t)$ would not perform better than the best-performing individual model, 175 no matter what their weights are. Incorporating the spatiotemporally varying model bias 176 into the ensemble prediction reflects the ensemble model deficiency. 177

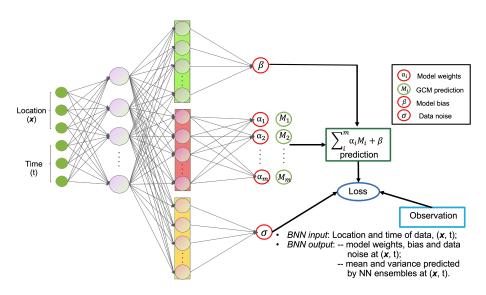


Figure 1. Architecture of the proposed Bayesian neural networks (BNNs).

In implementation, BNN reads the data location (\mathbf{x}) and time (t) as inputs and estimates the model weights, biases, and data noises at the given (\mathbf{x}, t) by calibrating the ensemble prediction against the observations. As illustrated in Figure 1, BNN first uses a set of dense layers to extract common information of the model weights, biases, and data noises. Then, three sets of dense layers are designed to learn the information specific to each component. Next, BNN incorporates the multiple GCM predictions $M_i(\mathbf{x}, t)$ and combines them with the estimated model weights, biases, and data noises in the loss function for optimization. The weights, biases, and noises are calibrated as probabilistic functions by specifying distributions over the parameters of the neural networks (NNs) (i.e., we perform the optimization in the Bayesian context). For computational efficiency, we train the BNN using the randomized maximum a posteriori (MAP) sampling (Pearce et al., 2018) instead of the computationally intractable full Bayesian inference, which may require Markov chain Monte Carlo simulation. The MAP sampling approach uses multiple NNs to quantify the ML model parameter uncertainty. Specifically, for the *j*-th network, we draw a sample from the prior distribution over the network parameters (assumed Gaussian) $\theta_{anc,j} \sim N(\mu_{prior}, \Sigma_{prior})$, and compute the MAP estimate corresponding to a prior re-centered at $\theta_{anc,j}$. When we consider a dataset of N observations y_k where k = 1, ..., N and specify the data likelihood by assuming a Gaussian noise with the heteroscedastic standard deviation of $\sigma(\mathbf{x}_k, t_k)$, the calculation of the MAP estimate is equivalent to minimize the following loss function for the *j*-th network:

$$Loss_{j} = \sum_{k=1}^{N} \frac{(y_{k} - \hat{y}_{j}(\mathbf{x}_{k}, t_{k}))^{2}}{\sigma_{j}^{2}(\mathbf{x}_{k}, t_{k})} + \sum_{k=1}^{N} log(\sigma_{j}^{2}(\mathbf{x}_{k}, t_{k})) + ||\Sigma_{prior}^{-1/2}(\theta_{j} - \theta_{anc,j})||_{2}^{2}.$$
 (2)

After training, the output of an ensemble of n_e such networks is thus a mixture of n_e Gaussians, $N(\hat{y}_j(\mathbf{x}_k, t_k), \sigma_j^2(\mathbf{x}_k, t_k))$. Then, the mean prediction of these networks $\frac{1}{n_e} \sum_j \hat{y}_j$ is the BNN prediction result. The variance $\frac{1}{n_e} \sum_j \sigma_j^2 + \frac{1}{n_e} \sum_j \hat{y}_j^2 - (\frac{1}{n_e} \sum_j \hat{y}_j)^2$ quantifies predictive uncertainty where the first term quantifies aleatoric data uncertainty, and the combination of the second and third terms quantifies the epistemic uncertainty describing the model's ignorance about the conditions outside the observational records.

Attributed to this special NN design and Bayesian training, the ensembling strat-184 egy of BNN not only calculates spatiotemporally varying model weights and biases, but 185 it also calibrates the weights and biases against observations to fully leverage each in-186 dividual model's simulation capability, allowing for more accurate and observationally 187 constrained predictions. Furthermore, we trained the BNN using the computationally 188 efficient randomized MAP sampling, which enables rapid quantification of the aleatoric 189 and epistemic uncertainty. Last but not the least, a key strength of this BNN approach 190 is the models interpretability, which can explain which models perform well in which lo-191 cations at which times. This interpretability extends the usage of ML techniques beyond 192 its predictive capabilities to bring insight and understanding to the climate models. 193

To enable the BNN to produce physically consistent results, we encoded our do-194 main knowledge into the network design and network training. First, in terms of net-195 work design, we chose tanh activations for the hidden layers in Figure 1 because their 196 mean output is zero-centered, which stabilizes the training. Furthermore, the tanh ac-197 tivations result in a predictably flat extrapolation outside the training set, which ensures 198 a realistic estimation of the model bias and data noise. For the set of dense layers in sim-199 ulating the model weights, we use a softmax layer at the end to ensure that the model 200 weights sum to unity. Additionally, in terms of network training, we first transform lat-201 itude, longitude, and time of each data point to a 6 dimensional space-time input. In a 202 climate model, we usually use latitude and longitude to represent a location and use a 203 scalar of t to represent the time (no matter what the unit is). However, directly inputting 204 the three numbers—latitude (lat), longitude (lon), and time (t)—to the BNN would be 205 problematic because the model weights, biases, and data noises generated by such a net-206 work would be discontinuous and would not respect seasonality. To address this prob-207 lem, we first represent the location input \mathbf{x} by its Euclidean coordinate [cos(lat)sin(lon),208 $\cos(lat)\cos(lon), \sin(lat)$ and warp the time input t onto a 3D helix $\left[\cos(2\pi t/T), \sin(2\pi t/T)\right]$ 209 t], where T is the time scale of the climate model simulation (here T = 1 month). This 210 transformation of the time variable makes the network generate model weights and bi-211 ases with both a strong monthly periodicity and a slow variation over the year, which 212 is more consistent with reality. Next, we rescale each column of space-time inputs to the 213 range [-a, a] to appropriately represent the varying frequency of the model weights and 214 biases across the space and time. A larger value of a results in a higher changing frequency. 215 In this study, the spatial coordinates are scaled into the [-2, 2] range, and the tempo-216 ral coordinates are scaled into the [-1, 1] range. The network complexity (e.g., the num-217 ber of layers and the number of nodes in each layer) and the number of networks for Bayesian 218 training are problem specific, depending on the GCM resolution, the model ensemble size, 219 and affordable computing resources. Generally speaking, a large number of complex NNs 220 is needed for an ensemble analysis of many high-resolution GCMs to calculate the spa-221 tiotemporally varying weights and quantify the uncertainty, which meanwhile requires 222

a high computational cost. In this study, we use an NN structure in which each set of
 dense layers in Figure 1 has a single hidden layer with 100 nodes, and we use 50 such
 NNs for UQ.

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2.2 Precipitation Data and Models

We apply the BNN ensembling method for precipitation prediction based on the 227 GCMs from CMIP6. The simulated precipitation data from the CMIP6-GCMs are down-228 loaded from the Earth System Grid Federation (ESGF) archives (https://esgf-node.llnl.gov/search/cmip6). 229 We consider monthly data from 53 GCMs during the period of 1980–2014, and our anal-230 yses focus on the CONUS area. The details of these models are listed in Table 1. We 231 use the European Centre for Medium-Range Weather Forecasts (ERA5) reanalysis data 232 from the same periods and regions as the reference or "observations" for model calibra-233 tion and performance evaluation (Muñoz-Sabater et al., 2021). The original ERA5 data 234 are at 33 km horizontal grid spacing and the hourly scale. We aggregate the data to the 235 monthly scale to be consistent with the GCMs simulation data. Both the simulation and 236 reference data are remapped to a common 1° latitude–longitude grid using the bilinear 237 interpolation method. 238

2.3 Three Widely Used Ensembling Schemes

In this section, we introduce three state-of-the-art ensembling schemes, which serve as baselines to evaluate the BNN's prediction performance. The simple average method is straightforward and normally used for multiple model analysis. The weighted average (Knutti et al., 2017) and spatially weighted average methods (Amos et al., 2020) have an increasing application because of their consideration of model skills and model independence and their good prediction performance. In the following, we briefly describe these three methods where the symbols are consistent with those in Section 2.1.

Simple Average The simple average method performs weighted averaging by assigning individual models with equal weights. The ensemble prediction is calculated as

$$\hat{y}(\mathbf{x},t) = \frac{1}{m} \sum_{i=1}^{m} M_i(\mathbf{x},t).$$
(3)

Weighted Average The weighted average method was introduced by Knutti et al. (2017), who used model ensembles to project the future sea ice change in the Arctic. This weighted average considered model skill and model independence in calculating the weights. For an ensemble of m models, the weight w_i for model i is calculated as

$$w_i = \exp\left(-\frac{D_i^2}{\sigma_D^2}\right) / \left(1 + \sum_{i \neq i}^m \exp\left(-\frac{S_{ij}^2}{\sigma_S^2}\right)\right),\tag{4}$$

where D_i^2 represents the discrepancy between the model *i* and the observation, and $S_{i_i}^2$ 247 describes the difference of the model i from the model j. Here, the model outputs and 248 observations in calculation of D_i^2 and S_{ij}^2 are an averaged value over space and time. This method uses D_i^2 and S_{ij}^2 to consider model skill and model uniqueness. It also introduces 249 250 two constants, σ_D and σ_S , to control the influence of the model skill and uniqueness on 251 the weights calculation and, consequently, on the ensemble predictions. For example, when 252 σ_D is assigned a small value, only a small number of models obtain weights, whereas when 253 σ_D is assigned a large value, this weighted average converges to the simple average with 254 equal weights. Although the values of σ_D and σ_S significantly affect the ensemble pre-255 dictions, it is unknown how to assign an appropriate value for a specific problem; cur-256 257 rently, the values are determined in a heuristic way. Additionally, although this weighting method considers model skill and independence, it does not consider the model's spa-258 tiotemporally varying skill and assigns a uniform weight across space and time. 259

Country	Research institute	Model name
Australia	Commonwealth Scientific and Industrial	ACCESS-ESM1-5 ACCESS-
	Research Organization	CM2
Canada	Canadian Centre for Climate Modelling and Analysis	CanESM5 CanESM5-CanOE
China	Beijing Climate Center	BCC-ESM1 BCC-CSM2-MR
	Chinese Academy of Meteorological Sciences	CAMS-CSM1-0
	Chinese Academy of Sciences	CAS-ESM2-0
	The State Key Laboratory of Numerical Modeling for LASG	FGOALS-g3 FGOALS-f3-L
	Nanjing University	NESM3
	Research Center for Environmental Changes	TaiESM1
	The First Institute of Oceanography, SOA	FIO-ESM-2-0
France	Institut Pierre Simon Laplace	IPSL-CM6A-LR
France	Centre National de Recherches Meteo-	CNRM-CM6-1 CNRM-CM6
		1-HR CNRM-ESM2-1
Germany	rologiques The Alfred Wegener Institute Helmholtz	AWI-ESM-1-1-LR AWI-CM-
Germany	Centre for Polar and Marine Research	1-1-MR
	Max Planck Institute for Meteorology	MPI-ESM1-2-LR MPI-ESM
	Max Flanck Institute for Meteorology	1-2-HAM MPI-ESM1-2-HR
Inner	The University of Tokyo, National Insti-	MIROC-ES2L MIROC6
Japan	tute for Environmental Studies, and Japan	MIROC-ES2L MIROCO
	Agency for Marine-Earth Science and Tech-	
	nology	
	Meteorological Research Institute	MRI-ESM2-0
Italy	Fondazione Centro Euro-Mediterraneo sui	CMCC-CM2-HR4 CMCC-
Italy	Cambiamenti Climatici	CM2-SR5
Korea	Korea Meteorological Administration	KACE-1-0-G
Rolea	Seoul National University	SAM0-UNICON
Nothorland	s EC-EARTH consortium published at Irish	EC-Earth3-Veg-LR EC-
/Ireland	Centre for High-End Computing	Ec-Earth3-Veg EC-Earth3
Norway	Bjerknes Centre for Climate Research, Nor-	NorESM2-MM NorESM2-LN
Norway	wegian Meteorological Institute	NorCPM1
Russia	Institute of Numerical Mathematics	INM-CM4-8 INM-CM5-0
UK	Met Office Hadley Center	HadGEM3-GC31-LL
UK	Net Onlee Hadley Center	HadGEM3-GC31-MM
	Natural Environment Research Council	UKESM1-0-LL
USA	National Center for Atmospheric Research	CESM2-WACCM-FV2
	National Center for Atmospheric Research	CESM2 CESM2-FV2 CESM2-
		WACCM
	Geophysical Fluid Dynamics Laboratory	GFDL-CM4 GFDL-ESM4
	University of Arizona	MCM-UA-1-0
	NASA/GISS (Goddard Institute for Space	GISS-E2-1-G GISS-E2-1-G-
	Studies)	CC GISS-E2-1-H
	Department of Energy	E3SM-1-0 E3SM-1-1 E3SM-1-

Table 1. The 53 GCMs from 28 institutes in CMIP6 are considered in this study. The 28 models in bold from each institute are used for ensembling in the real case application in Section 3.3.

Spatially Weighted Average Recognizing that the calculation of D_i^2 and S_{ij}^2 in Eq. (4) did not consider the difference in space and time, Amos et al. (2020) proposed a spatially weighted average method that calculates D_i^2 and S_{ij}^2 as a function of location **x** and time t. Specifically, for an ensemble of m models, the spatially weighted average is defined by

$$w_i = \exp\left(-\frac{D_i^2(\mathbf{x},t)}{n\sigma_D^2}\right) / \left(1 + \sum_{i\neq i}^m \exp\left(-\frac{S_{ij}^2(\mathbf{x},t)}{n\sigma_S^2}\right)\right),\tag{5}$$

where *n* is the number of data in calculating $D_i^2(\mathbf{x}, t)$ and $S_{ij}^2(\mathbf{x}, t)$. Although this method considers model–observation discrepancy and model–model difference across space and time in computing the model weights, it still assigns a uniform weight w_i to an individual model *i*.

2.4 Evaluation Metrics of Prediction Performance

We used several statistics and visualization tools to evaluate the prediction per-265 formance. For assessing the overall performance, we used root mean square error (RMSE), 266 density plots, and box plots. A better performing ensembling method would have a smaller 267 RMSE value, a closer density/box plot to that of the reference. To evaluate the perfor-268 mance in each grid cell, we present the prediction error across the simulation domain. 269 Additionally, we evaluate the BNN's spatiotemporal-aware weighting scheme by plot-270 ting the weights over the spatial domain, in specific regions, and along the simulation 271 time. 272

²⁷³ **3 Results and Discussions**

To validate and evaluate our proposed BNN ensembling scheme, we applied it to 274 three case studies and compare its prediction performance and weight calculation with 275 the three state-of-the-art methods introduced in Section 2.3. First, we designed a sim-276 ple numerical experiment in which we know the ground truth to evaluate whether the 277 BNN can accurately calculate the model weights reflecting the individual model's spa-278 tiotemporally varying skill. Secondly, we designed a synthetic study where the "obser-279 vations" come from one of the CMIP6 GCMs to further validate the BNN's capability. 280 In the last real case study, we applied the BNN for ensemble precipitation prediction us-281 ing 28 CMIP6 GCMs from different institutes and use the ERA5 reanalysis data for cal-282 ibration and evaluation. We analyze the results from three aspects: prediction perfor-283 mance, interpretability, and UQ. 284

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3.1 A Simple Numerical Experiment

In the simple numerical experiment, we used 35 years of monthly ERA5 reanaly-286 sis data over CONUS in 1980–2014 as the ground truth, based on which we designed four 287 individual models for ensemble analysis. Figure 2(a) shows the averaged ERA5 precip-288 itation data over the 35 years. We divided the simulation domain into four equal regions: 289 region I, II, III, and IV. Model i has the ground truth data in region i (where i repre-290 sents I, II, III, and IV) and has random noises in the other three regions. We generated 291 the random noise from the uniform distribution in the [10, 15] range, which is beyond 292 the ground truth, having the maximum average value of 8 mm/d. We trained the BNN 293 using the first 20 years of data and evaluated its performance on the remaining 15 years. 294

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After training, the BNN calculates the model weights for each grid cell at each month. Figure 2(b) summarizes its averaged model weights over the 15 years of the unseen test period for the four individual models in the entire domain. We observed that the BNN successfully recovered the expected model weights; it assigns weights of 1.0 to the regions

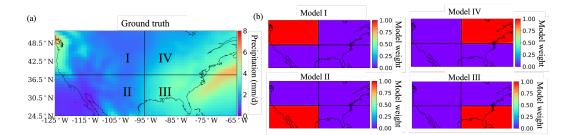


Figure 2. (a) The ERA5 precipitation data used as a ground truth in the numerical experiment, in which we divide the domain into four regions to design four individual models for the ensemble analysis; (b) The BNN ensembling scheme accurately assigns the weight of 1.0 to the regions where the model is accurate and assign the weight of 0.0 to the regions where the model produces random noise.

where the individual model is accurate and weights of 0.0 to those regions where the model 300 produces random noises. Because our BNN reasonably leverages each individual model's 301 prediction skill by accurately calculating the spatially varying weights, its ensemble pre-302 dictions have a great agreement with the ground truth. As shown in Figure 3, the prob-303 ability density function (PDF) of the BNN prediction for the out-of-sample test period 304 closely overlaps with the PDF of the ground truth. In contrast, the prediction from the 305 simple average differs dramatically from the truth by assigning equal weights to the mod-306 els and uniform weights to the entire domain. This numerical example validates this BNN's 307 capability in successfully capturing individual model's spatiotemporally varying skill and 308 demonstrates its competence in accurate ensemble predictions. 309

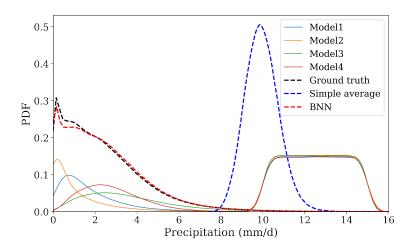


Figure 3. Probability density functions (PDFs) of the precipitation over the entire domain in the out-of-sample test period estimated by the simple average and BNN ensembling approaches, along with the data from the four individual models and the ground truth in the same period.

310 3.2 A Synthetic Study

In this second synthetic case study, we purposely selected seven CMIP6 GCMs from Table 1 to investigate the BNN's capability. Those seven GCMs are the Alfred Wegener Institute Climate Model (AWI-CM-1-1-MR), Manabe Climate Model v1.0 - University

of Arizona (MCM-UA-1-0), Community Earth System Model Version 2 (CESM2, CESM2-314 WACCM), and Energy Exascale Earth System Model (E3SM-1-0, E3SM-1-1, E3SM-1-315 1-ECA). We chose the simulation data of model CESM2-WACCM as the synthetic truth 316 to calibrate the BNN in the training period and evaluate the BNN's ensemble predic-317 tion in the test period. Figure 4(a) shows the precipitation data of model CESM2-WACCM 318 in CONUS averaged over the 35 years, and Figure 4(b) summarizes the PDFs of the pre-319 cipitation from the synthetic truth and the six models for ensemble analysis. We can see 320 that models AWI-CM-1-1-MR, MCM-UA-1-0, and CESM2 produce close predictions to 321 the synthetic truth, and the three E3SM models show similar performance, all perform-322 ing differently from the other four models. In this selection of the individual models and 323 the synthetic truth, we expect that a good-performing ensembling scheme should assign 324 a large weight to those three models, AWI-CM-1-1-MR, MCM-UA-1-0, and CESM2, which 325 produce similar precipitation simulations with the synthetic truth, and assign a small 326 weight to the three E3SM models which have a relatively large discrepancy from the "truth". 327 To further investigate the BNN's capability in generating reasonable spatiotemporally 328 varying weights, we divided the simulation domain into four regions—North, East, South, 329 and West (Figure 4(a))—to evaluate whether its regional weights reflect the individual 330 model's simulation skill locally. We used the first 20 years of data for training and the 331 last 15 years for out-of-sample testing. 332

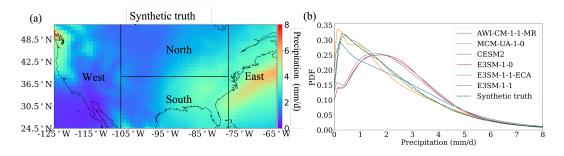


Figure 4. (a) Precipitation data of the synthetic truth averaged over 35 years; (b) The PDFs of the precipitation data from the synthetic truth and the six GCMs for ensemble analysis in the synthetic case study.

In the following, we analyze the ensemble prediction results. We first discuss the ensemble prediction accuracy and compare the BNN performance with the three stateof-the-art baselines. Next, we analyze the BNN's weighting scheme in detail by looking at its weights spatially and temporally and investigate the influence of the calculated model biases on prediction performance. In the analysis, we additionally demonstrate the BNN's interpretability. Lastly, we explore the BNN's capability in UQ.

Figure 5 shows the absolute prediction errors of the four ensembling approaches 339 averaged over the test period. The figure indicates that BNN produces more accurate 340 ensemble predictions than the other three methods by showing smaller prediction errors 341 in the simulation domain. Figure 6(a) summarizes the predictions of the six individual 342 models and the four ensembling methods in box plots. The box plots again demonstrate 343 that the ensemble predictions of BNN are closer to the synthetic truth, with similar me-344 dian and quantiles. On the other hand, the three baseline ensembling methods produce 345 quite similar results, all showing a relatively large difference from the synthetic truth. 346 In this case study, we fine-tuned the hyperparameters of σ_D and σ_S in the weighted av-347 erage and spatially weighted average methods and show here the best prediction results 348 we obtained after fine-tuning. However, the resulting ensemble predictions from these 349 two weighting schemes do not seem to bring much improvement from the simple aver-350 age. Their RMSEs are close to each other, with values of 1.44, 1.44 and 1.45 for simple 351

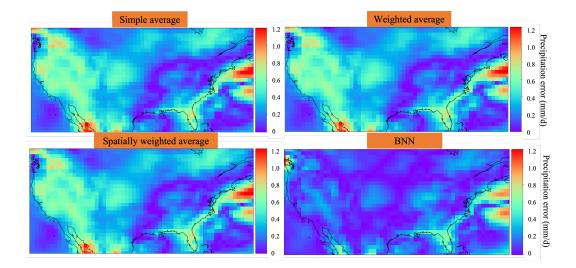


Figure 5. Absolute precipitation errors of the four ensembling methods averaged over the test period in the synthetic study.

average, weighed average and spatially weighted average method, respectively. Due to
the computational costs, we do not perform hyperparameter tuning for the BNN in this
work. However, the current BNN architecture and the set of hyperparameters already
show a great improvement in prediction accuracy compared to the three ensembling baselines and the individual models. A higher improvement of BNN is expected after its hyperparameter tuning and architecture optimization.

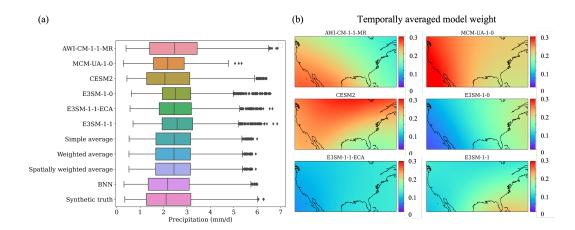


Figure 6. (a) Boxplot of the precipitation data in the test period for the six GCMs and the four ensemble predictions in the synthetic study; (b) Temporally averaged model weights over the test period for the six GCMs.

The superior prediction performance of our BNN is partially attributed to its spatiotemporally varying weights. Figure 6(b) presents the temporally averaged weights over the test period for the six individual GCMs in CONUS. We can see that the three topperforming models—AWI-CM-1-1-MR, MCM-UA-1-0, and CESM2—receive higher weights than the others overall, and the weights in each individual GCM vary spatially. We then divided the simulation domain into four regions (see Figure 4(a)) to closely examine the

BNN's spatial weighting and investigate whether its weighting aligns with the GCM's 364 skill. Figure 7(a) summarizes the temporally averaged weights in the four regions and 365 the entire CONUS domain for the six individual models, and it also presents the equal 366 weights as a baseline. The figure indicates that although models MCM-UA-1-0 and CESM2 367 have the highest weights overall in CONUS, MCM-UA-1-0 contributes highly in the West, 368 and CESM2 is the dominant GCM in the North and East. This spatially varying weight 369 aligns well with each individual model's spatially varying skill. Take the West region, 370 for example: Figure 7(b) indicates that model MCM-UA-1-0 performs better than E3SM-371 1-0 with smaller prediction errors in the West, and the BNN also assigns a higher weight 372 to MCM-UA-1-0 in this region. This suggests that the BNN's spatially varying weight-373 ing reasonably reflects GCMs' geographically heterogeneous prediction skill. Addition-374 ally, we investigate the BNN's temporally varying weights. Figure 8(a) plots the spatially 375 averaged weights of the 20 years for the six individual models. The figure indicates that 376 all the models present a seasonally changing weight, and no individual model performs 377 the best all the time. This suggests the importance of calculating temporally varying weights 378 in the ensembling. We picked a timestamp, August 1991, for a detailed analysis and present 379 the absolute prediction errors of model CESM2 and MCM-UA-1-0 at this specific time 380 in Figure 8(b). The figure indicates that model CESM2 predicts more accurate precip-381 itation in August 1991 than MCM-UA-1-0 by producing smaller prediction errors. More-382 over, the BNN accurately estimates the temporal-aware weights by assigning a larger value 383 to model CESM2 at this time step, which reasonably leverages the model's seasonally 384 distinct skill. 385

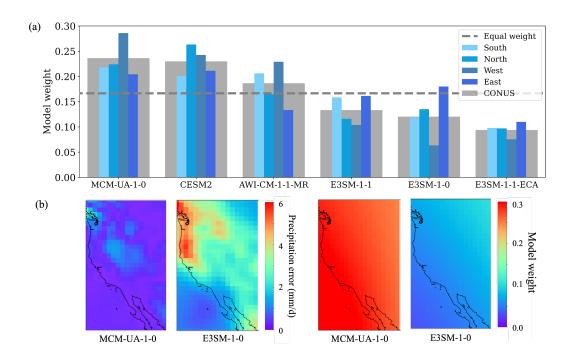


Figure 7. (a) Temporally averaged model weights over the test period in CONUS and the four sub-regions (see Figure 4(a)) for the six GCMs in the synthetic study; (b) Prediction errors and model weights of model MCM-UA-1-0 and E3SM-1-0 in the West region.

By calculating the spatiotemporally varying weights that accurately reflect the individual model's diverse skill across space and time, our BNN provides interpretability about which model contributes more to the ensemble prediction in which region and at which time. This scientific insight improves our understanding of each GCM's predic-

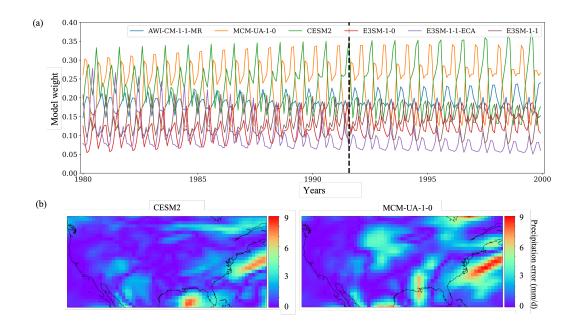


Figure 8. (a) Spatially averaged model weights over the simulation domain in the 20 year period for the six GCMs in the synthetic study; (b) Prediction errors of model CESM2 and MCM-UA-1-0 in the timestamp of August 1991 as highlighted in the black line of (a)

tive performance and help the model development by leveraging each model's merits. For 390 example, BNN identified that the model MCM-UA-1-0 is more accurate in predicting 391 the precipitation in the West region of CONUS, and CESM2 performs better in the Sum-392 mer season. Then we can go back to explore the mechanisms of the two models to in-393 vestigate why they yield better performance in the specific region at the specific time. 394 On the other hand, we can also examine why a certain model performs poorly in a cer-395 tain region at a certain time. Combining this comprehensive analysis, we can take ad-396 vantage of each individual model's strength to build a more powerful GCM for precip-397 itation prediction. And we can also explain that the BNN results in the superior ensem-398 ble predictions because it assigns higher weights to the regions and times where the model 399 performs better. In essence, we are confident in that we are getting right answers for the 400 right reasons. 401

Besides the smart weighting scheme, the spatially varying bias term in the BNN 402 ensembling also plays an important role for accurate precipitation prediction. As shown 403 in Figure 9(a), which presents the weighted prediction errors of the six GCMs, the north-404 west region has a relatively large positive prediction error. To compensate for the error 405 and make the ensemble prediction fit the calibration data well, the BNN estimates the 406 bias with a relatively large negative value in the region, as depicted in Figure 9(b). This 407 bias compensation scheme is particularly important when all the individual GCMs generate overestimation or underestimation, in which case the ensemble prediction will hardly 409 perform better than the best-performing individual GCMs despite the ensembling schemes. 410 In this situation, by introducing the bias term and calibrating its value against the data, 411 we can improve the ensemble predictions. Additionally, this bias term is a function of 412 space and time, so its calculation reflects the spatiotemporally varying model skill. 413

The BNN performs ensemble prediction in the Bayesian context, so it can quantify the data uncertainty to consider the data noise and quantify the epistemic uncertainty to consider the extrapolation error. Because the "observations" come from model

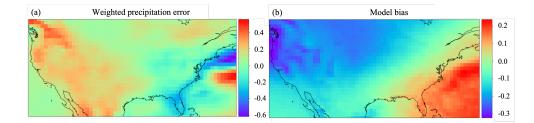
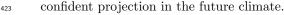


Figure 9. (a) Weighted precipitation errors (mm/d) of the six GCMs in the synthetic study; (b) The estimated bias (mm/d) (β in Figure 1) of BNN to compensate the weighted precipitation errors to enable a better ensemble prediction.

simulation data in this synthetic study, we do not have data noise. We focus more on the epistemic uncertainty discussion. Figure 10 shows the cumulative density function (CDF) of the epistemic uncertainty for the training and out-of-sample test data. The figure indicates that the BNN can reasonably quantify the uncertainty, where the epistemic uncertainty of the test data in the extrapolation regime is greater than that of the training data. This is highly desirable behavior and crucial in practice to prevent overaonfident projection in the future elimete.



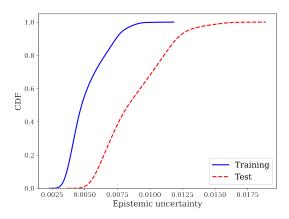


Figure 10. Epistemic uncertainty of the training and out-of-sample test data calculated by BNN in the synthetic study.

In this synthetic study, we demonstrate that our BNN produces superior predic-424 tion performance compared to that of the three state-of-the-art ensembling methods. BNN 425 can accurately calculate the spatiotemporally varying model weights and biases, which 426 can be justified by the model's prediction skill. This spatiotemporal-aware weighting scheme 427 meanwhile provides the interpretability of the BNN to help us understand which mod-428 els contribute more to the ensemble prediction at which locations and times. Addition-429 ally, we demonstrate that the BNN can reasonably quantify the epistemic uncertainty 430 by producing a larger uncertainty bound in the extrapolation regime to avoid overcon-431 fident predictions. 432

433 **3.3 A Real Case Application**

After verification and validation of the BNN method, we applied it to a real case study for precipitation prediction where the "observations" come from the ERA5 reanalysis data. We considered the 53 GCMs from CMIP6 as the model set; these are described

in Section 2.2. The 53 GCMs are from 28 institutes. Given that the models from the same 437 institute have strong dependence/similarities, we first performed data screening by se-438 lecting one model from one institute to roughly consider the model independence before 439 the ensemble analysis (Leduc et al., 2016; Ashfaq et al., 2022). For the models in the same 440 institute, we chose the one with the smallest RMSE compared to the ERA5 reference 441 data. The final selected 28 models are highlighted in bold in Table 1. Figure 11 shows 442 the PDFs of the precipitation data for the 28 GCMs and the ERA5 reference. As shown, 443 the 28 GCMs produce different precipitation simulations, and the major difference hap-444 pens at the small precipitation values ($\leq 2mm/d$). Some GCMs have close PDFs to the 445 "observations," and some others deviate significantly from the reference. 446

We applied the BNN to the 28 GCMs for ensemble predictions and investigated 447 whether our model can leverage each individual model's spatiotemporally varying skill 448 to produce an accurate prediction—and meanwhile reasonably quantify the predictive 449 uncertainty. We used the first 20 years of data for training and the remaining 15 years 450 for out-of-sample testing. In the training, we used the ERA5 data for model calibration; 451 in the testing, the ERA5 data were used as reference to evaluate the prediction perfor-452 mance. In the following results discussion, we first evaluate the BNN's prediction accu-453 racy in comparison with the three baseline ensembling methods. Next, we analyze the 454 BNN's model weights across the space and time and examine its interpretability. Lastly, 455 we present the UQ results and discuss the computational costs. 456

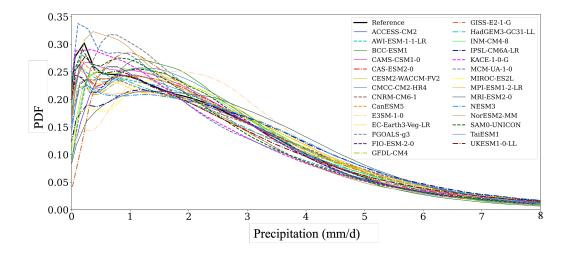


Figure 11. The PDFs of the precipitation data from the 28 GCMs for ensemble analysis and the reference data from the ERA5 reanalysis product.

Table 2 summarizes the RMSEs of the 15 years' test data in the entire simulation 457 domain and the four regions for the four ensembling methods, where the four regions are 458 divided in the same manner as shown in Figure 4(a). The BNN produces the smallest 459 RMSEs in CONUS and in the West, North, and South regions, demonstrating the best 460 prediction performance. Additionally, the BNN produced consistently smaller predic-461 tion errors than the simple average method, whereas in some cases, the advanced weighted 462 average and spatially weighted average even produced larger RMSEs than the simple av-463 464 erage. Please note that the ensembling results of the weighted average and spatially weighted average methods come from a fine-tuning of their hyperparameters, whereas for the BNN 465 approach, we did not perform a hyperparameter and architecture optimization. A fur-466 ther improvement in the BNN prediction performance is expected after a better choice 467 of its hyperparameters and network architectures. 468

	Simple average	Weighted average	Spatially weighted average	BNN
CONUS	1.48	1.48	1.51	1.45
West	0.71	0.68	0.67	0.57
North	0.25	0.27	0.28	0.23
South	0.50	0.48	0.49	0.43
East	0.46	0.44	0.43	0.44

Table 2. The RMSEs of the 15 years' precipitation data (mm/d) in the test period at CONUS and the four sub-regions (Figure 4(a)) for the four ensembling methods.

The superior prediction performance of the BNN benefits from its spatiotemporal-469 aware weighting scheme. Figure 12 shows the temporally averaged model weights in CONUS 470 for the 28 GCMs. We organized the models from the largest weights to the smallest weights 471 in row-wise order. Each GCM presents geographically heterogeneous weights. Overall, 472 model KACE-1-0-G, HadGEM3-GC31-LL, and NorESM2-MM in the top row gain the 473 highest weights, and model MIROC-ES2L, CESM2-WACCM-FV2, and BCC-ESM1 on 474 the bottom row obtain the lowest weights. However, the model's overall higher weight 475 does not necessarily show a uniform higher weight across the domain at each grid cell. 476 For example, in the second column of Figure 12, although model SAM0-UNICON has 477 a smaller weight than MRI-ESM2-0 in most areas, it shows a higher weight in the West 478 region. 479

Figure 13(a) summarizes the temporally averaged model weights in CONUS and 480 the four regions for the 28 GCMs. We can clearly see that the models present region-481 ally varying weights. Ten models have a higher weight than the equal value (i.e., 1/28). 482 For some models whose overall weights are below the equal weight, they could still show 483 a higher weight in a certain region. For example, model SAM0-UNICON presents a larger 181 weight in the West region than MRI-ESM2-0, although its overall weight in CONUS is 485 smaller than the latter. This spatially varying weight assignment is consistent with the 486 individual model's prediction skill. As shown in Figure 13(b), model SAM0-UNICON 487 shows a smaller prediction error than MRI-ESM2-0 in the West, where the RMSEs of 488 SAM0-UNICON and MRI-ESM2-0 are 0.7 and 1.2, respectively; we also observe a higher 489 spatial weight of SAM0-UNICON in this region. Additionally, Figure 13(c) illustrates 490 that model KACE-1-0-G and NorESM2-MM have similar weights in the West region, 491 and these two models indeed show similar prediction performance: the RMSE of KACE-492 1-0-G is 0.65 close to that of the NorESM2-MM value of 0.63. 493

The BNN not only gives reasonable spatially aware weights that accurately reflect 494 the individual model's spatially varying skill, but it also produces skill-consistent weights 495 in the temporal dimension. To avoid a busy figure and for a better demonstration, Fig-496 ure 14(a) plots the spatially averaged model weights in the out-of-sample test period for 497 three GCMs, which show the top prediction performance and have the highest spatial 498 weights. All of the three models demonstrate a seasonally changing weight, and none of 499 them obtain the highest weights all the time. The weights of model KACE-1-0-G show 500 a decreasing annual trend regardless of the seasonality, the weights of model NorESM2-501 MM present an increasing annual trend, and there is no much annual change in the weights 502 of model HadGEM3-GC31-LL. We picked two timestamps—at the beginning and at the 503 end of the test period—to analyze the weights of model KACE-1-0-G and NorESM2-MM 504 in detail. Figure 14(b) shows that KACE-1-0-G has a smaller prediction error than NorESM2-505 MM in February 2000, which justifies its higher model weights at this time. Addition-506 ally, the lower weight of KACE-1-0-G in July 2014 once again aligns with its relatively 507 higher prediction errors at this time. 508

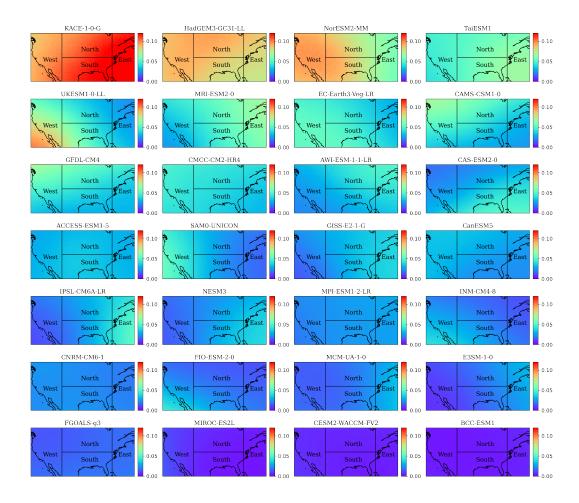


Figure 12. Temporally averaged model weights over the 15 years of the test period in CONUS for the 28 GCMs in the real case application.

The BNN accurately calculates the model weights for each individual model in each 509 grid cell at each time step. Its weighting scheme sufficiently leverages models' predic-510 tion skill and produces skill-consistent weights. This smart weighting not only improves 511 model prediction accuracy, but it also provides interpretability about each GCM's con-512 tribution to the ensemble prediction. Please note that all the results and weights anal-513 ysis presented in this real case application are based on the out-of-sample test data, so 514 when deploying the BNN method in practice for future projection where the ground truth 515 is unknown, its verified interpretable and skill-consistent weights increase our confidence 516 in the BNN's ensemble prediction. Certainly, when projecting to the future unknown con-517 ditions, besides the point estimate, we are also interested in the predictive uncertainty. 518 BNN can reasonably quantify the epistemic uncertainty caused by the model ignorance 519 and data shortage. Figure 15 shows the CDFs of the epistemic uncertainty for the train-520 ing and out-of-sample test data. The figure indicates that BNN produces a larger epis-521 temic uncertainty of the test data than that of the training data, accurately reflecting 522 our lesser confidence in the unknown conditions and thus preventing overconfident ex-523 trapolation. 524

In this real case application, we successfully apply the BNN ensembling method to 28 GCMs from CMIP6 for precipitation predictions in CONUS. We demonstrate BNN's superior prediction performance regionally and locally in comparison to the three baseline methods. We investigate BNN's spatiotemporal-aware weighting scheme, verify its

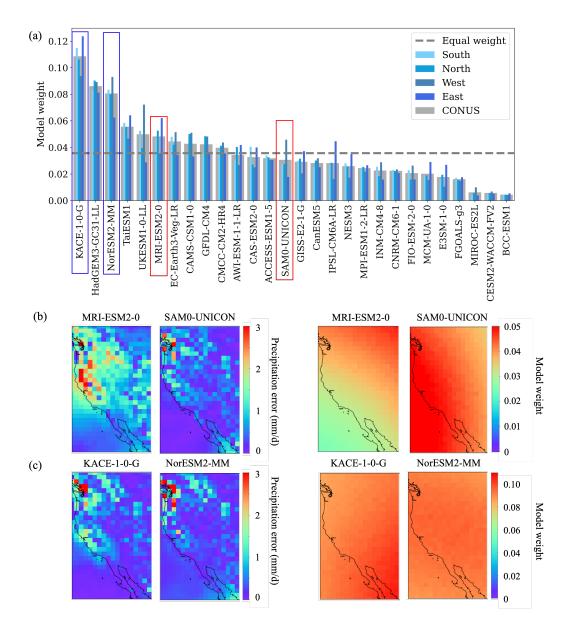


Figure 13. (a) Temporally averaged weights over the test period in CONUS and the four regions (see Figure 12) for the 28 GCMs considered in the real case application; (b) Prediction errors and model weights of model MRI-ESM2-0 and SAM0-UNICON in the West region; (c) Prediction errors and model weights of model KACE-1-0-G and NorESM2-MM in the West region.

weight's consistency with the model prediction skill, and interpret the individual models' contribution to the ensemble prediction spatially and temporally. Lastly, we analyze the reasonableness of BNN's UQ capability.

One possible limitation of the BNN ensembling scheme is the high computational cost. In this work, all the training ends at 1000 epochs when the loss function shows marginal decay. In the synthetic study in which six GCMs are considered, it takes about 35 minutes to train one NN and 29.17 hours to finish the training of 50 NNs in the BNN ensembling. For the real case application where 28 GCMs are analyzed, it takes about 40

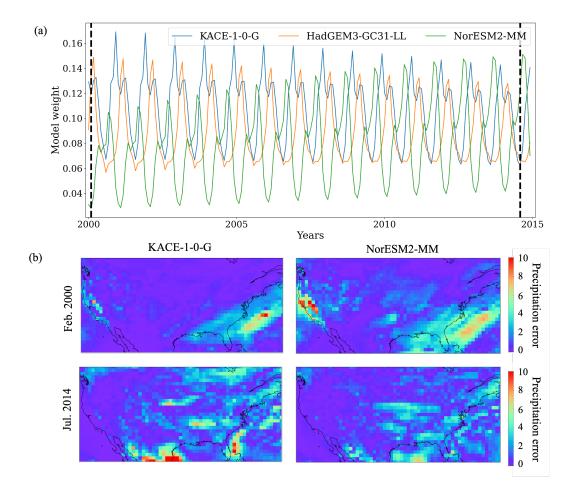


Figure 14. (a) Spatially averaged weights in the 15 years of test period for the three top performing GCMs in the real case application; (b) Prediction errors (mm/d) of model KACE-1-0-G and NorESM2-MM in February 2000 and July 2014.

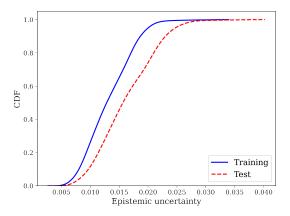


Figure 15. Epistemic uncertainty of the training and out-of-sample test data calculated by BNN in the real case application.

minutes to train one NN and 33.33 hours to train the 50 NNs. All the experiments were performed on a 2.3 GHz Quad-Core Intel Core i7 CPU. Roughly speaking, the computational cost increases with increasing numbers of networks in BNN training and ensemble GCMs, as well as with the resolution of the GCMs; this is because the BNN calculates weights at each time step in each grid cell. In spite of the relatively high computational cost of the BNN compared to other ensembling schemes, the cost is affordable
(e.g., within one or two days); more importantly, the BNN provides better prediction performance, interpretable ensembling results, and UQ.

⁵⁴⁵ 4 Conclusions and Future Work

In this work, we propose a BNN ensembling method for multiple model analysis 546 to enhance the predictive capability. The method improves prediction accuracy by learn-547 ing spatiotemporally varying model weights and biases based on the individual models 548 skill in simulating the observations across space and time. Additionally, the BNN method 549 accounts for the varying quality of the observations by incorporating their aleatoric un-550 certainty and avoids overconfident extrapolating predictions by quantifying the epistemic 551 uncertainty. More importantly, the method offers interpretability about which models 552 contribute more to the ensemble prediction at which locations and seasons. This insight 553 advances predictive understanding, guides process-based model development, and pri-554 oritizes data collection. 555

We apply the BNN ensembling method for precipitation prediction in CONUS based 556 on the GCMs from CMIP6. In both synthetic and real case studies, we demonstrate that 557 the BNN produces a better prediction performance than the three baseline ensembling 558 approaches; it can correctly assign a higher weight to the regions and the seasons where 559 the individual GCM fits the "observations" better; and it gives a reasonable bias value 560 to compensate for the error of the weighted average to enable a better ensemble predic-561 tion than the individual models. Additionally, we verify that the proposed BNN's inter-562 pretability is consistent with our prior knowledge in the synthetic design and with our 563 understanding of localized GCM performance in the real case application. Finally, the 564 BNN shows an increasing uncertainty when the prediction is farther away from the pe-565 riod with constrained data, which appropriately reflects our predictive confidence and 566 the trustworthiness of the models in the changing climate. Although the BNN ensem-567 bling method produces high-quality, interpretable, and uncertainty-aware predictions at 568 the expense of high computational costs in calculating the grid-specific and time-specific 569 model weights and biases, the cost is affordable: for example, about 33 hours are spent 570 in application of the 28 GCMs. More importantly, the provided high predictive accuracy 571 and the insights of the model performance are significant. In the future, we will apply 572 the BNN ensembling technique for other Earth system modeling problems, including pre-573 dictions of other response variables from the GCMs and problems in other disciplines 574 such as hydrology and ecology. 575

576 5 Data Availability Statement

The simulated precipitation data used in this work are available from the CMIP6 archive https://esgf-node.llnl.gov/search/cmip6. The Python scripts of our proposed BNN method can be found in https://github.com/patrickfan/BNN.

580 6 Author Contributions

MF implemented the numerical experiments, prepared the figures and analyzed the results. DL developed the algorithms, contributed to the research plan, and interpreted the results. DR processed the data and interpreted the results. EMP formulated the problem and interpreted the results. All the four authors contributed to the manuscript writing.

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- ⁵⁹³ ological and Environmental Research.

594 **References**

- Abramowitz, G., & Bishop, C. (2015). Climate model dependence and the ensemble dependence transformation of cmip projections. *Journal of Climate*, 28(6), 2332–2348.
- Ahmed, K., Sachindra, D. A., Shahid, S., Iqbal, Z., Nawaz, N., & Khan., N. (2020).
 Multi-model ensemble predictions of precipitation and temperature using ma chine learning algorithms. *Atmospheric Research*, 236(104806).
- Alexander, K., & Easterbrook, S. M. (2015). The software architecture of climate
 models: a graphical comparison of cmip5 and emicar5 configurations. *Geoscientific Model Development*, 8(4), 1221–1232.
- Amos, M., Young, P. J., Hosking, J. S., Lamarque, J.-F., Abraham, N. L., Akiyoshi,
 H., ... others (2020). Projecting ozone hole recovery using an ensemble of
 chemistry-climate models weighted by model performance and independence.
 Atmospheric Chemistry and Physics, 20(16), 9961–9977.
- Ashfaq, M., Rastogi, D., Abid, M. A., & Kao, S.-C. (2022). Evaluation of cmip6 gcms over the conus for downscaling studies.
- Bishop, C. H., & Abramowitz, G. (2013). Climate model dependence and the replicate earth paradigm. *Climate dynamics*, 41(3), 885–900.
- Brunner, L., Lorenz, R., Zumwald, M., & Knutti, R. (2019). Quantifying uncertainty
 in european climate projections using combined performance-independence
 weighting. *Environmental Research Letters*, 14 (12), 124010.
- Demory, M.-E., Berthou, S., Fernández, J., Sørland, S. L., Brogli, R., Roberts, M. J.,
 ... others (2020). European daily precipitation according to euro-cordex regional climate models (rcms) and high-resolution global climate models (gcms)
 from the high-resolution model intercomparison project (highresmip). Geoscientific Model Development, 13(11), 5485–5506.
- Eyring, V., Bony, S., Meehl, G. A., Senior, C. A., Stevens, B., Stouffer, R. J., &
 Taylor, K. E. (2016). Overview of the coupled model intercomparison project
 phase 6 (cmip6) experimental design and organization. Geoscientific Model
 Development, 9(5), 1937–1958.
- Eyring, V., Cox, P. M., Flato, G. M., Gleckler, P. J., Abramowitz, G., Caldwell, P.,
 ... others (2019). Taking climate model evaluation to the next level. Nature
 Climate Change, 9(2), 102–110.
- Fotheringham, A. S., Crespo, R., & Yao, J. (2015). Geographical and temporal weighted regression (gtwr). *Geographical Analysis*, 47(4), 431–452.
- Gleckler, P. J., Taylor, K. E., & Doutriaux, C. (2008). Performance metrics for cli mate models. *Journal of Geophysical Research: Atmospheres*, 113(D6).
- Greve, P., Orlowsky, B., Mueller, B., Sheffield, J., Reichstein, M., & Seneviratne,
 S. I. (2014). Global assessment of trends in wetting and drying over land.
 Nature geoscience, 7(10), 716–721.
- Heinze-Deml, C., Sippel, S., Pendergrass, A. G., Lehner, F., & Meinshausen, N.
 (2021). Latent linear adjustment autoencoder v1. 0: a novel method for estimating and emulating dynamic precipitation at high resolution. *Geoscientific Model Development*, 14(8), 4977-4999.
- Jose, D. M., Vincent, A. M., & Dwarakish, G. S. (2022). Improving multiple model

639	ensemble predictions of daily precipitation and temperature through machine
640	learning techniques. Scientific Reports, $12(1)$, 1-25.
641	Karpechko, A. Y., Maraun, D., & Eyring, V. (2013). Improving antarctic total
642	ozone projections by a process-oriented multiple diagnostic ensemble regres-
643	sion. Journal of the Atmospheric Sciences, $\gamma_0(12)$, 3959–3976.
644	Knutti, R., Furrer, R., Tebaldi, C., Cermak, J., & Meehl, G. A. (2010). Challenges
645	in combining projections from multiple climate models. Journal of Climate,
646	23(10), 2739-2758.
647	Knutti, R., Sedláček, J., Sanderson, B. M., Lorenz, R., Fischer, E. M., & Eyring, V.
648	(2017). A climate model projection weighting scheme accounting for perfor- mence and interdemondance. Combusing Research Letters $1/(4)$ 1000–1018
649	mance and interdependence. Geophysical Research Letters, $44(4)$, 1909–1918. Konnende C. Michre A. K. Wede V. & Mann M. E. (2020, 6). Climate change
650	Konapala, G., Mishra, A. K., Wada, Y., & Mann, M. E. (2020, 6). Climate change will affect global water availability through compounding changes in sea-
651	sonal precipitation and evaporation. <i>Nature Communications</i> , 11(1). doi:
652 653	10.1038/s41467-020-16757-w
654	Kumar, D., Kodra, E., & Ganguly, A. R. (2014). Regional and seasonal intercom-
655	parison of cmip3 and cmip5 climate model ensembles for temperature and
656	precipitation. Climate dynamics, 43(9), 2491–2518.
657	Leduc, M., Laprise, R., De Elia, R., & Šeparović, L. (2016). Is institutional democ-
658	racy a good proxy for model independence? Journal of Climate, 29(23), 8301-
659	8316.
660	Li, D., Liu, Y., & Chen, C. (2021). Msdm v1. 0: A machine learning model for pre-
661	cipitation nowcasting over eastern china using multisource data. Geoscientific
662	$Model \ Development, \ 14(6), \ 4019-4034.$
663	Lorenz, R., Herger, N., Sedláček, J., Eyring, V., Fischer, E. M., & Knutti, R. (2018).
664	Prospects and caveats of weighting climate models for summer maximum tem-
665	perature projections over north america. Journal of Geophysical Research:
666	Atmospheres, 123(9), 4509-4526.
667	Martin, E. (2018). Future projections of global pluvial and drought event character-
668	istics. Geophysical Research Letters, 45(21), 11–913.
669	Martin, G. M., Klingaman, N. P., & Moise, A. F. (2017). Connecting spatial and
670	temporal scales of tropical precipitation in observations and the metum-ga6.
671	Geoscientific Model Development, $10(1)$, $105-126$. Mueller, B., & Seneviratne, S. I. (2014). Systematic land climate and evapotran-
672	spiration biases in cmip5 simulations. <i>Geophysical research letters</i> , 41(1), 128–
673	134.
674 675	Muñoz-Sabater, J., Dutra, E., Agustí-Panareda, A., Albergel, C., Arduini, G., Bal-
676	samo, G., others (2021). Era5-land: A state-of-the-art global reanalysis
677	dataset for land applications. <i>Earth System Science Data</i> , 13(9), 4349–4383.
678	Pearce, T., Zaki, M., Brintrup, A., Anastassacos, N., & Neely, A. (2018). Uncer-
679	tainty in neural networks: Bayesian ensembling. <i>stat</i> , 1050, 12.
680	Pincus, R., Batstone, C. P., Hofmann, R. J. P., Taylor, K. E., & Glecker, P. J.
681	(2008). Evaluating the present-day simulation of clouds, precipitation, and
682	radiation in climate models. Journal of Geophysical Research: Atmospheres,
683	113(D14).
684	Que, X., Ma, X., Ma, C., & Chen, Q. (2020). A spatiotemporal weighted regression
685	model (stwr v1. 0) for analyzing local nonstationarity in space and time. Geo-
686	scientific Model Development, $13(12)$, $6149-6164$.
687	Räisänen, J., Ruokolainen, L., & Ylhäisi, J. (2010). Weighting of model results
688	for improving best estimates of climate change. Climate dynamics, $35(2)$, 407–
689	422.
690	Sanderson, B. M., Knutti, R., & Caldwell, P. (2015). A representative democracy to
691	reduce interdependency in a multimodel ensemble. Journal of Climate, $28(13)$,
692	5171-5194.
693	Sanderson, B. M., Wehner, M., & Knutti, R. (2017). Skill and independence weight-

694	ing for multi-model assessments. $Geoscientific Model Development, 10(6),$
695	2379–2395.
696	Stegall, S. T., & Kunkel, K. E. (2019). Simulation of daily extreme precipitation
697	over the united states in the cmip5 30-yr decadal prediction experiment. Jour-
698	nal of Applied Meteorology and Climatology, 58(4), 875–886.
699	Taylor, K. E., Stouffer, R. J., & Meehl, G. A. (2012). An overview of cmip5 and
700	the experiment design. Bulletin of the American meteorological Society, $93(4)$,
701	485 - 498.
702	Ukkola, A. M., De Kauwe, M. G., Roderick, M. L., Abramowitz, G., & Pitman,
703	A. J. (2020). Robust future changes in meteorological drought in cmip6 pro-
704	jections despite uncertainty in precipitation. Geophysical Research Letters,
705	47(11), e2020GL087820.
706	Weigel, K., Bock, L., Gier, B. K., Lauer, A., Righi, M., Schlund, M., others
707	(2021). Earth system model evaluation tool (esmvaltool) v2. 0-diagnostics for
708	extreme events, regional and impact evaluation, and analysis of earth system
709	models in cmip. Geoscientific Model Development, $14(6)$, $3159-3184$.
710	Wenzel, S., Eyring, V., Gerber, E. P., & Karpechko, A. Y. (2016). Constraining
711	future summer austral jet stream positions in the cmip5 ensemble by process-
712	oriented multiple diagnostic regression. Journal of Climate, $29(2)$, 673–687.
713	Zelazowski, P., Huntingford, C., Mercado, L. M., & Schaller, N. (2018). Climate
714	pattern-scaling set for an ensemble of 22 gcms–adding uncertainty to the
715	imogen version 2.0 impact system. Geoscientific Model Development, $11(2)$,
716	541 - 560.