# Strong Storm-effect behaviors of topside and bottom-side ionosphere under low solar activity: Case study in the geomagnetic storm during 25-27 August 2018

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#### Abstract

The 25-27 August geomagnetic storm was the third largest storm in 24th solar cycle, which was a surprising space event that generated in the background of very low solar activity. This study presents an overview of temporal-spatial behaviors of ionospheric plasma irregularities as functions of geographic longitude, latitude and altitude by ground-based (GNSS receivers and ionosonde) instruments and space-borne (Swarm-A and Swarm-B) satellites. The results not only reveal the enhanced equatorial ionization anomaly (EIA) and hemispheric asymmetry over the Asian-Australian and American sectors in a particular time, but also discover the development of hemispheric asymmetric features of global ROTI in the main and recovery phases. In addition, this storm also triggered positive plasma irregularities in altitudes of 100 to 150km near Auroral zone, and the changed ratio of bottom-side plasma irregularities exceeded 250%, which has been cross validated by multiple instrument and TIE-GCM's simulation. Furthermore, the thermospheric density ratio  $O/N_2$ , equatorial electrojet and vertical  $E \times B$  drifts suffered from the storm largely, the equatorial and mid-latitude plasma irregularities may be a combined action of thermospheric composition change, equatorial electrojet and vertical  $E \times B$  drifts. Finally, the storm also induced positive Joule heating irregularities in Auroral ionosphere in altitudes of 100 to 400km with a maximum changed ratio of >200%, as well as the cross Polar voltage enhanced to ~90kv. The Polar ionospheric irregularities may be associated with the additional energy input through the ways of particle precipitation, Joule heating and ionospheric currents intensification.

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18	Key points:
19	• First time to give an overview of the development of global Rate of total electron content
20	(TEC) Index change in the whole phase of storm
21	• Positive ionospheric plasma irregularities in the altitudes of 100 to 150km were only
22	detected near Auroral zone (>~50°N/S)
23	• Low-middle latitude plasma irregularities were a combined action of density ratio $O/N_2$ ,
24	equatorial electrojet and vertical $E \times B$ drift
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26	
27	

# 28 Abstract

The 25-27 August geomagnetic storm was the third largest storm in 24th solar cycle, which 29 was a surprising space event that generated in the background of very low solar activity. This 30 study presents an overview of temporal-spatial behaviors of ionospheric plasma irregularities 31 as functions of geographic longitude, latitude and altitude by ground-based (GNSS receivers 32 and ionosonde) instruments and space-borne (Swarm-A and Swarm-B) satellites. The results 33 not only reveal the enhanced equatorial ionization anomaly (EIA) and hemispheric asymmetry 34 over the Asian-Australian and American sectors in a particular time, but also discover the 35 development of hemispheric asymmetric features of global ROTI in the main and recovery 36 phases. In addition, this storm also triggered positive plasma irregularities in altitudes of 100 37 to 150km near Auroral zone, and the changed ratio of bottom-side plasma irregularities 38 exceeded 250%, which has been cross validated by multiple instrument and TIE-GCM's 39 simulation. Furthermore, the thermospheric density ratio O/N<sub>2</sub>, equatorial electrojet and 40 vertical  $E \times B$  drifts suffered from the storm largely, the equatorial and mid-latitude plasma 41 42 irregularities may be a combined action of thermospheric composition change, equatorial electrojet and vertical E×B drifts. Finally, the storm also induced positive Joule heating 43 irregularities in Auroral ionosphere in altitudes of 100 to 400km with a maximum changed ratio 44 of >200%, as well as the cross Polar voltage enhanced to ~90kv. The Polar ionospheric 45 irregularities may be associated with the additional energy input through the ways of particle 46 precipitation, Joule heating and ionospheric currents intensification. 47

Key words: ionospheric disturbances, geomagnetic storm, hemispheric asymmetry, TIE-GCM,
O/N<sub>2</sub>

Plain language: Large amounts of charged particles deposited in the thermosphere-ionosphere system during a strong geomagnetic storm, this process could change global ionospheric convection and weaken the activities of positioning, navigation, radio communication, etc. This study tries to discover the spatial-temporal changes of global ionosphere under the strong geomagnetic storm during 25-27, August 2018. From the observations of Global Navigation Satellite System (GNSS) receivers and radars, we first time to discover the spatial-temporal evolutions of global plasma irregularities, and reveal the storm-enhanced equatorial ionization

anomaly and hemispheric asymmetry in the Asian-Australian and American sectors. Besides, 57 positive plasma irregularities in the altitudes of 100 to 150km were only near Auroral zone (> 58 ~50°N/S), rather than in low-middle latitudes. Furthermore, the potential drivers are 59 investigated for explaining the plasma irregularities. The equatorial and mid-latitude 60 irregularities may be a combined action of thermospheric composition change, equatorial 61 electrojet and vertical E×B drifts. The Polar ionospheric irregularities may be associated with 62 the additional energy input through the ways of particle precipitation, Joule heating and 63 ionospheric currents intensification. 64

# 65 **1. Introduction**

Ionosphere has a serious effect on absorbing, scattering and refracting radio signals, which 66 is a main error source in the navigation and positioning service. Now the global navigation 67 satellite system (GNSS) differential technique and empirical/theoretical models are usually 68 69 used to correct the ionospheric delay. However, ionospheric plasma during severe geomagnetic storms suddenly increases or decreases violently, which easily reduces the accuracy of 70 positioning and navigation. Now available techniques are not good at correcting severe storm-71 effect ionospheric perturbations; therefore, it is necessary and valuable to investigate the 72 73 ionospheric spatial-temporal behaviors response to strong magnetic storms and to discuss the probable drivers. 74

Geomagnetic storm usually has a severe effect on the ionospheric system, auroral particle 75 precipitation, Polar ionospheric currents and convection are reinforced largely during a 76 77 geomagnetic storm. The enhanced Polar ionospheric ionization and electric fields penetrate to low-middle latitudes, this process affects global electrodynamics and changes the structure of 78 the thermospheric-ionospheric system remarkably. The Joule heating and Auroral particle 79 precipitation heat and expand the thermosphere, which further changes the composition and 80 81 dynamics of the thermospheric-ionospheric system [Astafyeva et al., 2015]. With rapid 82 development of Global Navigation Satellite System (GNSS) and radio occultation, the stormeffect ionospheric behaviors have been paid more attention by multiple ground-based and 83 space-borne techniques. The "Halloween" storm erupted on 29-30, October 2003 was one of 84 strongest geomagnetic storms in this century. After a few hours when the interplanetary 85

magnetic field suddenly turned southward, the dayside ionospheric total electron content (TEC) 86 increased about 40%, and the Challenging Minisatellite Payload (CHAMP) profiles indicated 87 the dayside TECs over mid-latitudes increased ~900% on 30 October [Mannucci et al., 2005]. 88 The significant increments of TEC and peak density (NmF2) were also observed over the 89 European and North African sectors during a following stronger geomagnetic storm occurred 90 on November 20, 2003[Crowley et al., 2006]. Moreover, the sudden ionospheric irregular 91 behaviors response to geomagnetic storms were also been reported over Jicamarca [Zhang et 92 al., 2019], Brazilian equatorial-low latitudes [de Paula et al., 2019], China and adjacent areas 93 [Aa et al., 2018], Asian-Australian sector [Lei et al., 2018], Indian sector [Ramsingh et al., 94 2015], Turkey [Karatay, 2020], Arctic and Antarctic [Durgonics et al., 2017; Mitchell et al., 95 2005; Shreedevi et al., 2020] and in a global scale [Atici and Sağır, 2020; Li et al., 2022]. These 96 reports revealed that strong geomagnetic storms easily triggered large-scale positive or 97 negative traveling ionospheric disturbances (TIDs), and sometimes the TIDs had a significant 98 latitudinal asymmetric structure in northern-southern hemispheres that is caused by the 99 displaced magnetic poles and seasonal asymmetries in the thermosphere-ionosphere system. 100 101 Multi-instruments observations and theoretical model simulations were conducted for explaining the disturbed ionospheric dynamic convections, and the results concluded that the 102 negative TIDs were primarily attributed to a decrement of the thermospheric density ratio O/N<sub>2</sub> 103 [Dmitriev et al., 2017; Fuller - Rowell et al., 1994]. However, the drivers for positive TIDs 104 were various, thermospheric neutral winds, disturbance dynamo electric fields (DDEF), prompt 105 penetration electric field (PPEF) as well as charged particle precipitation, had been reported to 106 107 be the potential factors in enhancing plasma densities [Atici and Sağir, 2020; Crowlev et al., 2006; Nava et al., 2016; Qian et al., 2019; Richmond and Lu, 2000]. 108

Ionospheric storms primarily occur as a consequence of strong coronal mass ejection (CME), such as the magnetic storms in 22-23, June 2015 [*Ngwira et al.*, 2019] and 7-8, September 2017 [*Li et al.*, 2018a]. However, the 25-27, August 2018 space event that is the third largest magnetic storm in 24<sup>th</sup> solar cycle happened after a slowly moving CME on 20 August, it is a huge surprise that the weak CME that even didn't show a sudden impulse could trigger a strong ionosphere-thermosphere response. The positive and negative ionospheric perturbations over North America [*Cherniak and Zakharenkova*, 2022], Brazil [*Spogli et al.*,

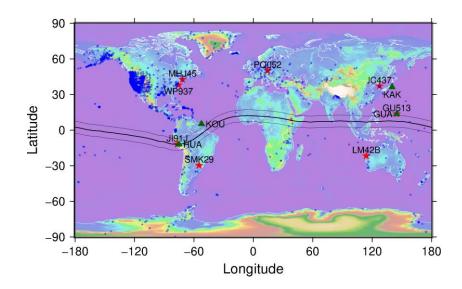
2021], Asia [Lissa et al., 2020], Middle latitudes [Chang et al., 2022], as well as the global 116 [Astafyeva et al., 2020] response to the surprising space event have been reported. This 117 geomagnetic storm induced penetration electric fields created favorable conditions for strong 118 fountain effect enhanced equatorial ionization anomaly (EIA) and generated equatorial plasma 119 bubbles (EPBs). The EPBs appeared over a larger latitudinal extent of EIA crests, while the 120 plasma density in the western coast of North America depleted in the northwestward direction 121 [Cherniak and Zakharenkova, 2022]. The Defense Meteorological Satellite Program (DMSP) 122 also detected midlatitude plasma depletion in the Asian sector, the local TIDs were responsible 123 for the midlatitude plasma depletion in Asia and United States, rather than the absence of EPBs 124 [Chang et al., 2022]. Different from the mid-latitude plasma depletion in the northern 125 hemisphere, Spogli et al. [2021] used the situ measurements provided by China Seismo-126 Electromagnetic Satellite and by Swarm-A satellite with ground-based observations to reveal 127 the ionospheric response at low-middle latitudes over Brazil, and found that significant foF2 128 increments appeared over the ionosondes located at both of dip-equator and southern crest of 129 the EIA. The decrease of the eastward electric field was the main driver for the equator station, 130 131 while it was resulted from the storm-induced equatorward thermospheric winds for the crest station. In addition, unprecedented hemispheric asymmetries of the thermospheric-ionospheric 132 responses were also observed during the main and recovery phases of the storm, which 133 expressed that strong positive plasma storms occurred in the northern hemisphere at the 134 beginning of the space event, while an extreme expansion of the thermospheric composition 135 ratio O/N<sub>2</sub> appeared in the opposite hemisphere during the recovery phase. The seasonal 136 137 asymmetry in the high-latitude plasma and neutral mass density distributions along with the asymmetries in the geomagnetic field played a decisive role for the hemispheric asymmetric 138 structure of disturbed plasma [Astafyeva et al., 2020]. 139

Most of previous studies focused on the planar ionospheric response (TEC) to the storm, but the altitudinal plasma behavior was still not clear. Besides, the storm-effect ionospheric response is controlled by multiple drivers at a particular moment of time and in particular location. The present study has the objectives to examine: (1) the development of hemispheric asymmetry of plasma irregularities in the main and recovery phase; (2) the altitudinal behaviors of plasma irregularities in low-middle latitudes and auroral zone; (3) the potential drivers for the latitudinal plasma irregularities over the Asian-Australian and American sectors. To address
these objectives, the latitudinal ionospheric irregularities are detected by a set of ground-based
(GNSS receiver and digital-ionosonde) and space-borne (SWARM) instruments, and the
potential drivers for the equatorial-auroral irregularities are discussed by a set of magnetometer,
SuperDARN, Global Ultraviolet Imager (GUVI) and TIE-GCM's simulation.

# 151 **2. Datasets**

In order to analyze planar-vertical behaviors of global plasma irregularities during the 152 magnetic storm on 25-27, August 2018 completely, the ground-based (GNSS receiver and 153 ionosonde) observations and the plasma profiles derived from the space-borne Swarm 154 constellation are used. In addition, the horizontal components of the magnetic field, 155 thermospheric density ratio O/N2 and the TIE-GCM's simulations are also adopted for 156 explaining the drivers of plasma irregularities. The geographic locations of ground-based 157 instruments are shown in Figure 1, and more detail information of several kinds of datasets are 158 introduced as follows. 159

The GNSS observations are obtained from the University NAVSTAR Consortium 160 (UNAVCO) that provides access to geodetic GPS/GNSS data used for geoscience research and 161 education. The UNAVCO provides about 2500 Receiver Independent Exchange Format 162 (RINEX) files daily through the link https://www.unavco.org/data/gps-gnss/gps-gnss.html. It 163 should be noted that the vertical total electron content (VTEC) estimated by the carrier-phase 164 smoothed pseudo-range method is used to investigate the ionospheric perturbations during the 165 geomagnetic storm [Li et al., 2018b]. For similarity, the TEC signifies VTEC in the whole 166 study. In addition, two chain digital-ionosondes located at the Asian-Australian and American 167 sectors are also utilized to investigate the vertical behaviors of storm-induced plasma 168 irregularities. Due to the influence of geomagnetic storm, the digital-ionosonde in the Asian 169 sector fails to observe the auroral plasma irregularities. The ionosonde PQ052 (14.6°E, 50°N) 170 located at Pruhonice in Europe is selected. The sounder profiles of two chains can be obtained 171 from the Lowell DIDBase (jdbc:firebirdsql://didbase.giro.uml.edu/didb) via the SAO explorer. 172



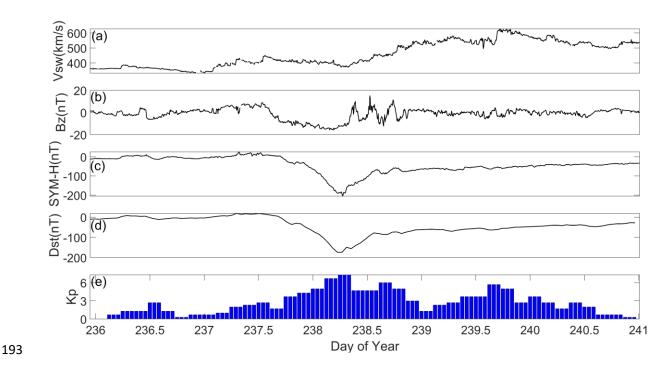
#### 173

Figure 1. Spatial distribution of GNSS stations, ionosondes and magnetometers. The blue pentagrams
 signify global GNSS stations, the red pentagrams signify the ionosondes located at the Asian-Australian
 and American sectors, the green triangles signify magnetometers, and the solid and dashed curves depict
 the location of the magnetic equator and the region of equatorial electrojet.

Swarm consists of three microsatellites (Alpha, Bravo and Charlie) that are placed in two 178 different orbital planes, among them the Swarm-A and Swarm-C fly at a mean altitude of 179 450km, and the satellite Swarm-B places in a mean altitude of 530km. In this study, the electron 180 density profiles of Swarm-A and Swarm-B are selected to analyze the plasma irregularities, 181 182 and the electron density is derived from the high gain ion current that is determined by the Langmuir probe. The Swarm profiles be obtained from the website 183 can https://swarmdiss.eo.esa.int/#swarm%2FLevel2daily%2FEntire mission data%2FTEC%2FT 184 MS. In addition, the daily F10.7 index, 81-day mean F10.7 and Kp index are imported to the 185 TIE-GCM model as input parameters, and the output parameters include electron density, 186 neutral winds, thermospheric composition and electric field. More details about the TIE-GCM, 187 the readers please refer to https://www.hao.ucar.edu/modeling/tgcm/tie.php. For improving the 188 simulated accuracy of the TIE-GCM products response to geomagnetic storms, the period of 189 the TIE-GCM products is from 00:00 (Universal Time, UT), 23 August to 29 August 2018. 190

191 **3. Results** 

### **3.1 Solar-terrestrial environment**



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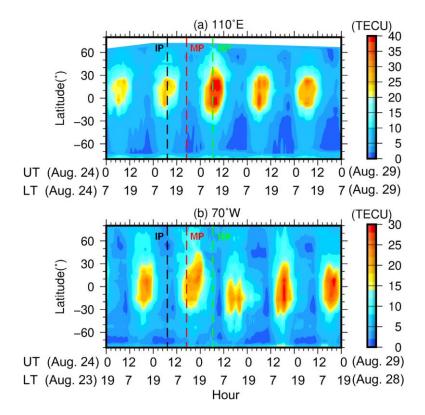
Figure 2. Variations of solar-terrestrial indices during day of year (DOY) 236 to 240 in 2018

Solar-geomagnetic change is an important factor in effecting the process of ionospheric 195 plasma irregularities during storms. Therefore, the data of solar wind speed (Vsw), 196 interplanetary magnetic field, and geomagnetic indices obtained from the Goddard Space 197 Flight Center (https://omniweb.gsfc.nasa.gov/form/dx1.html) is analyzed. The record third 198 largest geomagnetic storm in the 24<sup>th</sup> cycle was initiated from a slow CME on 20 August 2018, 199 which arrived at the earth thermosphere-ionosphere system on 25 August. As shown in Figure 200 201 2, the solar-terrestrial indices were in a quiet level before UT14, DOY 237 (25 August), and a strong geomagnetic storm happened from UT14, DOY237 to DOY 239 (27 August). Figure 202 2(a) shows the Vsw was low with a mean velocity of 400 to 450km/s in the main phase (UT14, 203 DOY237 - UT7, DOY238) of the geomagnetic storm, while it enhanced abruptly in the 204 205 recovery phase with a maximum speed of 620km/s. Figure 2(b) shows the Bz component of interplanetary magnetic field (IMF) had an abrupt southward excursion. From UT14, DOY237, 206 the Bz turned southward with a minimum value of -18nT in the forenoon of DOY 238. The 207 horizontal component of longitudinally symmetric disturbances (SYM-H) is essentially the 208 same as the Dst index to describe the mid-latitude geomagnetic disturbances. Figure 2(c) and 209 2(d) express that both SYM-H and Dst had a significant negative excursion since the afternoon 210 of DOY 237, and these indices reached to minimum values of ~-200nT and ~-180nT at UT06-211

08, DOY 238, respectively. As well as the global geomagnetic field indexed by Kp enhanced
to a maximum level of 7. According to the classification of geomagnetic storm released by

214 National Oceanic and Atmospheric Administration (NOAA), this storm is classed as "strong".

**3.2** Ionospheric irregularities over the Asian-Australian and American sectors



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Figure 3. Latitudinal TEC changes along (a) 110°E, (b) 70°W longitudes during August 24-28, the dashed
black line signify the initial phase (IP) of the geomagnetic storm, the dashed red line signify the main
phase (MP), and the dashed green line signify the recovery phase (RP). White color depicts empty cells
due to lack of actual observations.

The dual-frequency observations of about 2500 GNSS receivers provided by UNAVCO 221 are used to estimate the global TEC map, then the TEC grid maps are constructed by Kriging 222 interpolation method. The north-south cross-sections (keograms) of the GNSS TEC maps along 223 the 110°E and 70°W longitudes during 24-28 August are plotted to illustrate the temporal 224 evolution of the storm-effect TEC changes over two sectors. These keograms, plotted as a 225 function of UT time and geographic latitude, as shown in Figure 3. Figure 3(a) demonstrates 226 that in the Asian-Australian sector, the ionospheric TEC kept in a low level of 20 to 25TECU. 227 In the main phase of the storm, the equatorial ionospheric TEC enhanced significantly with a 228

double peak structure that coincided to the Dst index decreased to a maximum value of -174nT.
The enhanced TEC primary occurred in the ending of main phase and the beginning of recovery
phase (UT04-10, 26 August), the maximum TEC reached to 40TECU.

Different from the TEC change over the Asian-Australian sector, Figure 3(b) found a 232 significant hemispheric asymmetric structure of TEC irregularity in the American sector. For 233 example, significant TEC depletion happened over the American sector in the recovery phase. 234 In the afternoon of 26 August, the TEC in the northern hemisphere depleted from 15-20TECU 235 to 10TECU, while the equatorial and mid-latitude TEC in the opposite hemisphere enhanced 236 about 5TECU. The positive TEC storm was stronger in the afternoon of following two days, 237 and the maximum enhanced TEC exceeded 30TECU. Figure 3 concludes that the geomagnetic 238 storm triggered an asymmetric ionospheric storm, and the occurred phases of ionospheric storm 239 over different longitudes had a particular diversity. 240

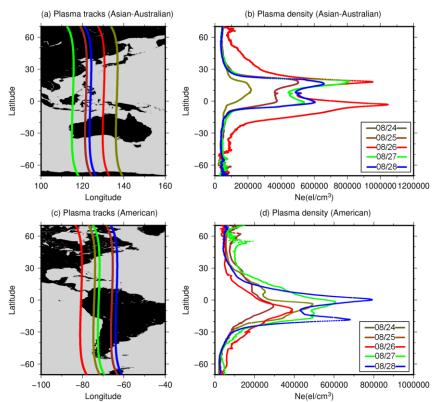




Figure 4. Daily changes of Ionospheric electron density detected by Langmuir probe instrument onboard
 Swarm-A satellite during (a-b) the main phase over the Asian-Australian sector (c-d) and the recovery
 phase over the American phase from 24 to 28 August.

Figure 4 gives an overview on the daily development of storm-effect ionospheric electron density (Ne) as measured by the space-borne Swarm-A satellite that flies at an orbital altitude

of ~450km. Here, the electron density signifies the amounts of plasma at the Orbital altitude of 247 Swarm-A, which is useful to reveal the change of topside ionosphere during the storm. The 248 result in Figure 3 demonstrates that the storm-induced plasma irregularities over the Asian-249 Australian sector was larger in the main phase, but in the recovery phase, the American sector 250 was larger. Therefore, five Ne profiles derived from Swarm-A during 24-28 August are plotted 251 to describe the Asian-Australian ionospheric response during the main phase. These profiles 252 not only have an adjacent longitude in ~120°E to 140°E span, but also have a nearly observed 253 254 time within UT05:10 to UT06:30, as shown the Ne tracks in Figure 4(a). Figure 4(b) shows the variation of daily Ne profiles as a function of geographic latitude, we can find that the plasma 255 density kept in a low level with the peak density of  $\sim 5 \times 10^5$  el/cm<sup>3</sup> during 24-25 August. In the 256 main phase, the topside Ne over the Asian-Australian sector significantly enhanced. The red 257 curve expresses a significant structure of EIA on 26 August, the double plasma crests located 258 at the 10°N to 20°N and 0° to 10°S latitudinal spans, respectively. Compared to the background 259 values, the storm-enhanced Ne increased ~2 times with a maximum value of ~ $1.05 \times 10^{6}$  el/cm<sup>3</sup>. 260 The enhanced EIA is believed caused by a daytime "superfountain" effect that driven by the 261 262 PPEFs. During strong geomagnetic storms, the PPEFs of eastward polarity could largely uplift the equatorial ionosphere over the sunlit and post-sunset sectors that drive the equatorial plasma 263 along the geomagnetic field line to higher altitudes and expanded poleward latitudes with a 264 significant enhancement of the EIA [Cherniak and Zakharenkova, 2022]. In the following two 265 days, the intensity of EIA gradually decreased to a normal level. 266

Different from the sudden enhanced TECs over the Asian-Australian sector, the TEC's 267 change over the American sector in the main phase (25-26 August) was not significant. 268 However, remarkable TEC enhancements were observed in the recovery phase, especially on 269 28 August, the peak density occurred in the North America with a value of  $\sim 8 \times 10^{5}$  el/cm<sup>3</sup>. The 270 hemispheric asymmetric structure of ionospheric TEC agrees well with the observations of 271 ground-based radars and space-borne Swarm-A. Finally, the TEC's changes over both the 272 Asian-Australian and American sectors reveal that the magnetic storm not only enhanced the 273 equatorial plasma density, but also triggered drastic polar ionospheric disturbances. The 274 development of storm-induced polar ionospheric disturbances will be investigated in the 275 following section. 276

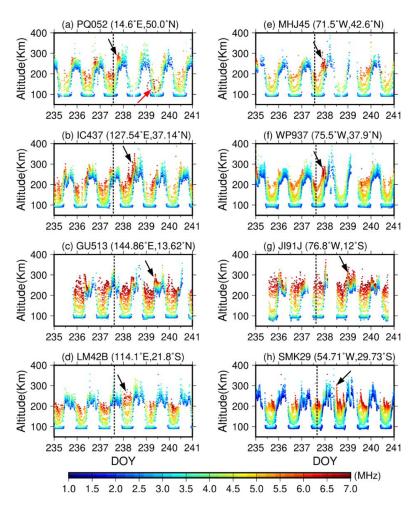


Figure 5. Vertical sounder profiles of two chain ionosondes located at the Asian-Australian (left) and
 American sectors in a whole phase of geomagnetic storm, the dashed line signifies the onset of the main
 phase of the storm

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An ionosonde sounder is a radar that sweeps the high-frequency (HF) band signals and 281 receives the echoes for examing the ionosphere and monitoring HF propagation conditions. 282 283 Ionosonde primarily operates between 1.6MHz to 12 MHz. With this advantage, the two ionosonde chains located at the Asian-Australian and American sectors could detect the vertical 284 dynamic propagation of storm-induced plasma irregularities, the results as shown in Figure 5. 285 From the left panels, the radars located at the Asian-Australian sector detected HF signals with 286 frequencies between 5.5 to 7MHz in the midnight-dawn (for local time, it was around the noon) 287 during DOY235-240, and the HF signals accumulated under 200km. The equatorial radar 288 detected HF plasma signal accumulated at a higher altitude. For example, Figure 5(c) shows 289 the ionosonde GU513 located at the Guam measured a maximum frequency of 7MHz around 290 291 the altitude of 250km in a minor solar-geomagnetic activity. After the onset of geomagnetic

storm, the plasma density during the midnight-dawn in the next day (26 August) suddenly 292 enhanced and uplifted. As shown in the black arrows, enhanced HF plasma density were all 293 detected over the equatorial and midlatitude radars (IC437, GU513, LM42B) in the main phase. 294 Especially over the radar IC437, the peak height of ionosphere was uplifted above 350km with 295 a peak frequency of 7MHz, and the intensity of Ne profiles in the northern hemisphere was 296 significant stronger than that over dip-equator and southern hemisphere (GU513 and LM42B), 297 which agreed well with the hemispheric asymmetric structure of plasma irregularities reported 298 299 by [Astafyeva et al., 2020].

The enhanced plasma was also detected by the radars located at the American sector. It 300 should be noted that the blank areas over PQ052, GU513, MHJ45, WP937, JI91J signify the 301 ionosondes failed to receive the HF echoes. It is found that the storm-enhanced plasma 302 irregularities were observed in both of the main and recovery phase, Figure 5(h) had enough 303 vertical profiles to describe the pattern. Also, this storm uplifted the peak height of ionosphere, 304 but in the dawn-forenoon (it's around the midnight for local time). The largest intensity of 305 plasma irregularities was detected over the dip-equatorial radar JI91J with a peak frequency of 306 307 7MHz, and the ionospheric peak height was uplifted about 50-80km. Compared to the Asian-Australian sector, the hemispheric asymmetric structure was not significant. Finally, the 308 detecting results of radars in low-middle latitudes manifested that the positive plasma 309 irregularities primary accumulated between the altitudes of 200 to 300km. Finally, the radar 310 PQ052 near the Arctic detected an interesting result, see the red arrow in Figure 5(a). During 311 the recovery phase, significant positive plasma irregularities at the altitudes of 110 to 150km 312 were observed near noon, DOY 239 (August 27). The different altitudinal behaviors of plasma 313 irregularities indicated that significant positive plasma irregularities may be triggered only in 314 315 the bottom-side of auroral ionosphere, rather than in the equatorial and mid-latitude ionosphere. Due to the TEC maps and ionosondes fail to reveal the altitudinal structures of plasma 316 irregularities, hence the plasma densities during DOY 235-240, 2018 are simulated by the TIE-317 GCM for solving this problem. The averaged plasma during DOY 235-236 are selected as 318 background value, and the altitudinal changing percent of storm-induced plasma irregularities 319 compared to background value is shown in Figure 6. Figure 6(a)-6(f) express the temporal 320 variations of storm-induced plasma as a function of geographic latitude at the layers span from 321

150 to 500km along the meridian 110°, and the vertical scale is proportional to the changing
percent, which is represented by a color bar for better understand the storm-enhanced plasma
behaviors.

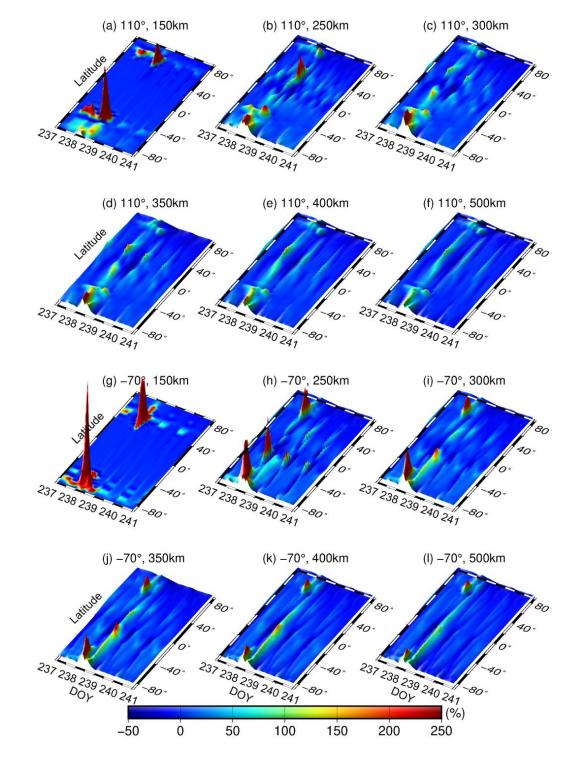


Figure 6. Ionospheric plasma disturbances at the layers span from 150 to 500km along the meridians 110°
(a-f) and -70° (g-l) during day of year (DOY) 237-240, 2018

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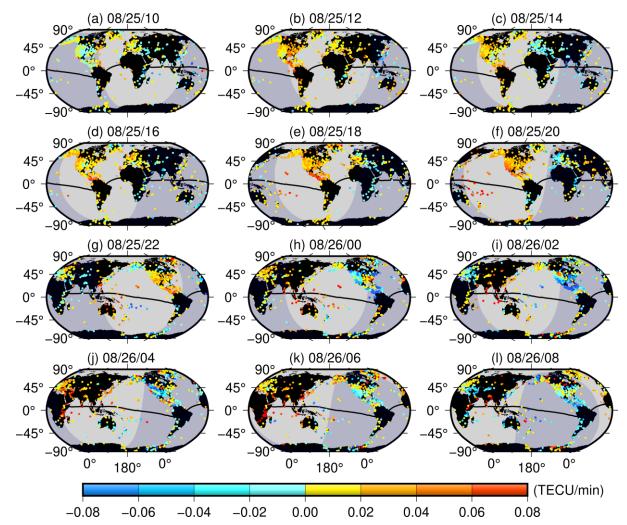
Figure 6 successes to simulate the development of double crests of the EIA at the altitudes 328 of 150 to 500km. At the layer of 150km, since the main phase on DOY 237, two plasma 329 increments appeared in 60°-65°N and 40°-45°S geographic latitudes along the meridian 110°, 330 respectively. On the next day, the amplitude of plasma enhancements enhanced to 250%. At 331 the same time, some tiny increments also occurred in the Antarctic. In the 250 km layer, the 332 two plasma crests that located in 60°-65°N and 40°-45°S latitudes were weakened, while the 333 plasma densities in the Antarctic were enhanced. With the increasing altitude, the plasma 334 increments had an equatorward movement. For example, at the layer of 250 km, a plasma 335 enhancement with a percent of ~200% appeared in 30°-40°N latitude, and a weaker increment 336 located in 40°S latitude. The two low latitudinal enhancements moved equatorward within  $\pm 20^{\circ}$ 337 latitude at 350km layer. Above 350km, the two plasma crests merged into one unit and the EIA 338 phenomena disappeared. 339

The change of ionospheric plasma along the meridian -70° agreed well with that over the Asian-Australian sector. Figure 6(g) shows two plasma increments appeared in 50°-60°N and 60°-70°S latitude at 150km layer, and the maximum percent exceeded 250%. The EIA phenomenon was also observed within the layers of 250 to 350km, and the crests of the EIA enhanced about 200% in DOY 238. Compared to the Asian-Australian sector, the storminduced plasma irregularities over the American sector were larger.

# **346 3.3 Global ROTI in the main and recovery phase**

The results of Figure 3-5 reveal hemispheric asymmetric structures of plasma irregularities 347 over two sectors in different phases. In order to further investigate the development of global 348 storm-induced ionospheric irregularities, the Rate of total electron content Index change (ROTI) 349 that expresses sharpness of the GNSS phase fluctuations caused by ionospheric irregularities 350 and by strong spatial gradients of TEC is estimated by the ground-based receivers. Figure 7 351 presents an overview of global GNSS ROTI maps during the main phase, and the time 352 resolution of GNSS ROTI maps is one minute. The large positive and negative ROTI 353 magnitudes are marked by red and blue, respectively; correspondingly, the small ROTI 354 magnitudes are marked by yellow and cyan, respectively. At UT10, August 25, the ROTI map 355 shows a low intensity of global ionospheric irregularities with an averaged value between -0.02 356

to 0.02TECU/min. From UT12, 25 August, the global ionospheric regularities abruptly
intensified. The positive plasma irregularities primary occurred in the sunlit sector, and the
magnitude of plasma irregularities over the Eastern Coast region of US and Mexico was largest
with a value of 0.06TECU/min. The plasma over the nighttime Greenland was also enhanced
about 0.02TECU/min.



**Figure 7.** Global Navigation Satellite System Rate of total electron content (TEC) Index change (ROTI) maps in the main phase (mm/dd/hh) of the geomagnetic storm. The black line signifies the magnetic

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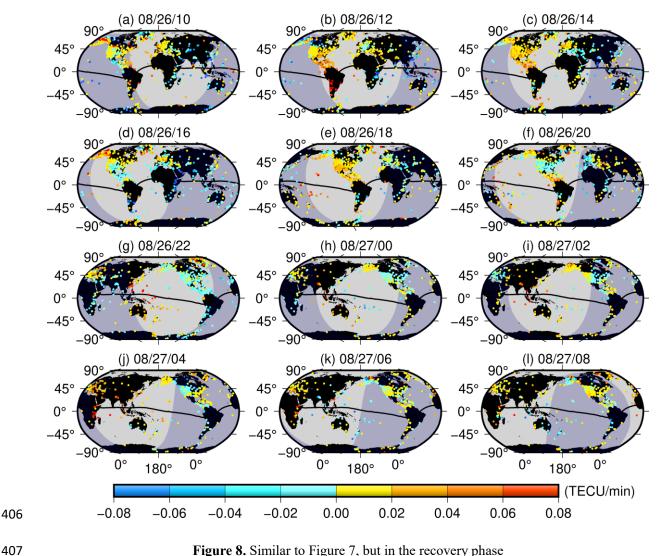
After ~UT18, 25 August, the AE index rapidly increased above 500nT, even reached to a peak of 1500 to 200nT in the main phase. Correspondingly, the daytime plasma irregularities suddenly enhanced from UT18, 25 August, and significant hemispheric asymmetry of plasma irregularities was observed in the American sector. Figure 7(e) shows the equatorial and auroral plasma irregularities over the North America enhanced with a maximum magnitude of

equator, and the shaded area shows nighttime.

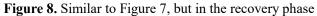
0.08TECU/min, as well as the intensified mid-latitude plasma with a lower magnitude of 371 0.06TECU/min. A narrow channel of positive ionospheric irregularities was registered along 372 the western coast of North America in the northwestward direction, the latitudinal extent of 373 EPBs reached to 20-25°N with a maximum value of 0.1TECU/min. In addition, the GNSS 374 ROTI observations over some ground-based receivers located at several islands in the Pacific 375 Ocean revealed an occurrence of positive storm-induced EPBs over 160°-140°W longitudinal 376 span, the positive ionospheric irregularities were observed at both sides of the magnetic equator 377 and their latitudinal extent was up to 25°- 27°N/S. The feature of equatorial ionospheric ROTI 378 agreed well with the results reported by [Cherniak and Zakharenkova, 2022]. In the nighttime 379 hemisphere, the ionospheric irregularities were negative with a low intensity. From 26 August, 380 the GNSS ROTI over the Western Coast of The North America and Greenland turned negative, 381 though some negative irregularities were under sunlit sector. However, the ROTI over the 382 383 European-African sector gradually turned positive and enhanced to a largest magnitude during UT04-06, 26 August. The positive irregularities with a maximum value of 0.1TECU/min were 384 observed in Europe, Africa and Asia, rather than Australia, though it was also under the sunlit 385 386 sector. The results conclude that the hemispheric asymmetry of plasma irregularities was also significant during the ending of the main phase, which expresses the ROTI over Africa was 387 larger than that over Europe, while it was opposite in the Asian-Australian sector. 388

Figure 8 shows an overview of global ionospheric irregularities during the recovery 389 phase of the geomagnetic storm, it is found that the storm also induced strong GNSS ROTI, 390 especially over the American sector. At UT10, 26 August, strong plasma irregularities with a 391 level of 0.06TECU/min appeared over the South America, as well as the equator and mid-high 392 latitudes of the North America. In addition, it is interesting that a narrow channel of positive 393 394 ionospheric irregularities was observed in the northwestward direction over the nighttime Alaska. Two hours later, a significant hemispheric asymmetry of plasma irregularities 395 developed over the American sector, which expresses the storm-induced plasma over 20°S to 396 45°S latitudinal span enhanced larger than 0.08TECU/min, while the GNSS ROTI over the 397 North America kept in a low level of 0.02 to 0.04TECU/min. Except the American sector, the 398 plasma irregularities over other daytime or nighttime Continents maintained in a low level. 399 From UT14, 26 August, the hemispheric asymmetry was reversed, which means high-400

magnitude GNSS ROTI concentrated in the North America. Here, the signatures of the 401 ionospheric irregularities persisted for many hours till the midnight. Furthermore, significant 402 equatorial and mid-latitude positive plasma irregularities also appeared in the daytime Asian-403 Australian and European-African sectors, as see Figure 8(g) and 8(j). In the following hours, 404 the global GNSS ROTI gradually recovered to a low level. 405









#### 3.4.1 Potential drivers of equatorial and mid-latitude ionospheric irregularities 409

Thermospheric composition change is an important driver in inducing positive or negative 410 ionospheric irregularities. Therefore, the thermospheric density ratio O/N<sub>2</sub> measured by GUVI 411 on board the space-borne TIMED satellite ( $\sim$ 625km) is analyzed. It is noted that the density 412 ratio O/N<sub>2</sub> is a height integrated quantity within the orbit altitudes of the GNSS constellation 413

and the GUVI satellite. At the same time, the global topside TECs derived from Swarm-A and 414 Swarm-B are also investigated, and the TEC signifies the integrated electrons within the 415 altitudes from the orbit of Swarm microsatellite to the orbit of GNSS constellation. Figure 9 416 give an overview of daily topside TEC and O/N<sub>2</sub> during the storm. One can also notice that the 417 TEC distribution is slightly different in the data of two satellites, which is most likely due to 418 the ~80km of difference in altitude. Both the profiles of Swarm-A and Swarm-B conclude that 419 the topside TECs over the Asian-Australian and American sectors were quiet before UT12, 25 420 August, the averaged TEC was under 6TECU. Figure 9(b) shows the TEC profile over the 421 American sector suddenly enhanced, and the expanded profiles covered the eastern Pacific, 422 this phenomenon was also validated by the observation of Swarm-B. During UT00-12, 26 423 August, the TEC over the Asia-Australia sector strengthened remarkably with a maximum 424 value of exceeded 12TECU. After that, the enhanced TEC profiles gradually decreased and 425 recovered to a normal level. The profiles derived from Swarm-B agreed well with that of 426 Swarm-A that large-scale positive TEC irregularities appeared over the Asian-Australian and 427 American sectors. 428

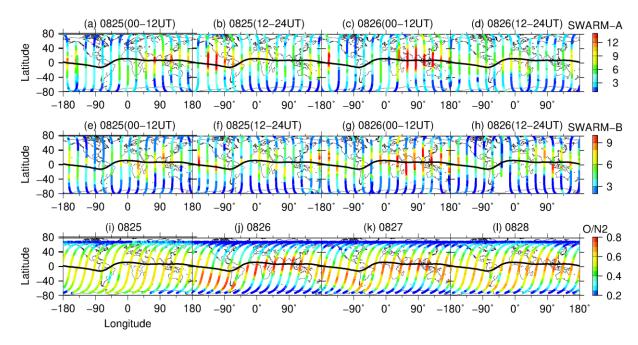


Figure 9. Topside TEC variations measured by GPS receivers on board the (a) – (d) Swarm-A and (e) – (h)
Swarm-B on 25-26 August, (i) – (l) thermospheric density ratio O/N<sub>2</sub> as measured by the GUVI satellite
during August 25-28. The black thin curve signifies the magnetic dip equator.

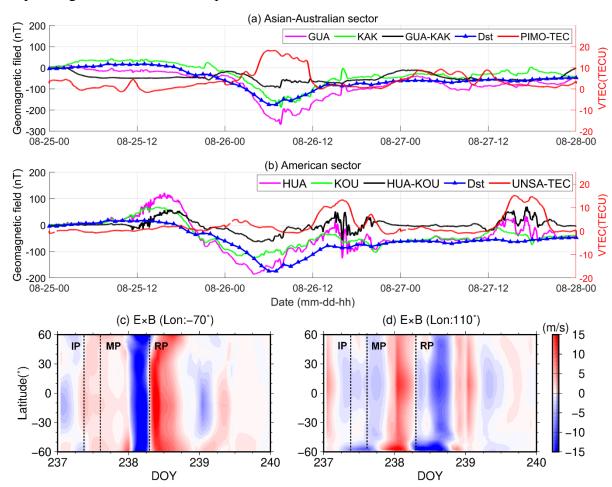
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Figure 9(i) shows the density ratio  $O/N_2$  was inversely proportional to geographic latitude, 433 the O/N<sub>2</sub> ranged from 0.3 to 0.6 in quiet days. However, the O/N<sub>2</sub> ratio had a suddenly change 434 with the eruption of a storm. Specifically, the O/N2 in low-middle latitudes increased 435 remarkably with a maximum value of 0.8, while the  $O/N_2$  in Polar regions decreased to 0.2. 436 Besides, the enhanced O/N<sub>2</sub> had a southward excursion in the America sector. For example, see 437 panel 9(j), the O/N<sub>2</sub> in North America was about 0.4, while this ratio increased to 0.8 in South 438 America. In the following two days, the storm-effect O/N<sub>2</sub> gradually decreased with the 439 geomagnetic field recovered to a normal level. The change of density ratio O/N<sub>2</sub> agreed well 440 with the TEC irregularities, which implied the change of density ratio O/N2 may be an 441 important driver in generating plasma disturbances. O/N<sub>2</sub> has a good positive correlation with 442 plasma density, and it has proven to be a successful indicator of a neutral composition 443 disturbance for analyzing ionospheric storms [Strickland et al., 2001]. The ionospheric ion 444 density loss rate is proportional to the molecular concentration, an increment of the mean 445 molecular mass causes a decrement in electron density, while a decrement of molecular 446 concentration provokes a positive disturbance. 447

As we know, at the altitudes of 90 to 130 km, many electrons move westward driven by 448 dayside electric field. According to the equatorial dynamo effect, the westward electron flow 449 generates a dayside eastward electric current, the electric current is defined as equatorial 450 electrojet (EEJ). The EEJ could be changed severely suffered from the disturbed electric field 451 penetrated from magnetosphere under a strong geomagnetic storm. The EEJ signatures can be 452 453 estimated by taking the difference between the horizontal components performed by a pair of off-the-equator and at-the-equator magnetometers. The horizontal components of the 454 magnetometers PHU, DLT, SJG and HUA located at the Asian-Australian and American sectors 455 are used to investigate the storm-effect EEJ changes, and the observations of magnetometers 456 obtained from the International Real-time Magnetic Observatory Network 457 are (https://intermagnet.org/index-eng.php). 458

Figure 10(a) and 10(b) show the changes of EEJ signatures along with the Dst index over the Asian-Australian and American sectors. It is found that the Dst abruptly decreased with a geomagnetic storm erupted on UT14, 25 August 25 2018, as well as the horizontal components of the magnetometers PHU, DLT, SJG and HUA. In Figure 10(a), the equatorial magnetometer GUA decreased to a minimum value of ~-270nT during UT 06-08, 26 August, and the

differential component between GUA and KAK had a negative perturbation with a minimum 464 value of -100nT. Correspondingly, the TEC over the station PIMO enhanced 18TECU in the 465 severest moment. The EEJ changes estimated by the difference between HUA and KOU in 466 Figure 10(b) agreed well with that over the Asian-Australian sector. The EEJ signature had two 467 distinct perturbations in the afternoons of 26-27 August with an amplitude of ~50nT. The 468 differential TEC over station UNSA was consistent with the EEJ signature. Two TEC 469 enhancements appeared in the recovery phase of geomagnetic storms, the maximum delta TEC 470 reached to 14TECU. The results demonstrate that the change of EEJ may be an important driver 471 in triggering storm-effect TEC disturbances. However, the slight EEJ fluctuations cannot fully 472 responsible for the strong TEC enhancements. Therefore, more drivers should be analyzed for 473 explaining the remarkable ionospheric disturbances. 474



475

476 Figure 10. Variations of the horizontal intensity of the geomagnetic field, equatorial electrojet (EEJ),
477 differential TECs and Dst value in the Asian-Australian (a) and American (b) sectors, as well as the
478 differential latitudinal vertical E×B drifts along the -70° (c) and 110° (d) longitudes.

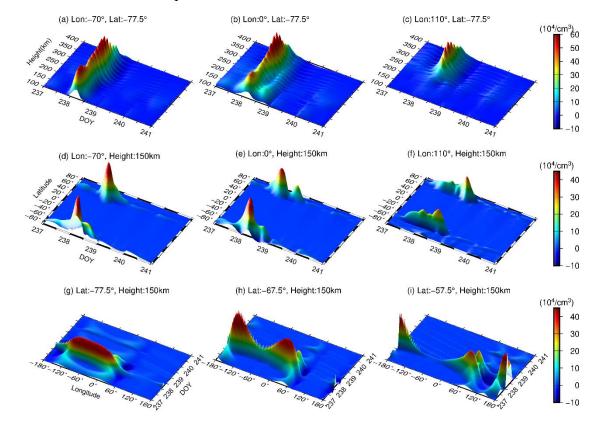
The equatorial plasma fountain effect plays a dominating role in generating the EIA. The 479 vertical E×B drift drives the equatorial plasmas upward to higher altitudes, and the accumulated 480 plasmas diffuse down to higher latitudes along the geomagnetic field lines, which results two 481 high concentrated plasma crests distributed on both sides of the magnetic equator [Balan and 482 Bailey, 1995; Li et al., 2021]. In the fountain process, a stronger E×B drift could lift more 483 plasmas to higher altitudes, and the crests of EIA generated by the plasma diffused process are 484 stronger and more poleward. Therefore, the changes of vertical E×B drift may be another driver 485 for the equatorial and mid-latitude plasma perturbations. The latitudinal changes of vertical E 486  $\times$ B drifts along the -70° and 110° longitudes are simulated by the TIE-GCM in Figure 10(c) 487 and 10(d). Figure 10(c) expresses the  $E \times B$  drifts along the meridian -70° began to increase 488 489 from UT14, 25 August, the slight E×B enhancement was 3 to 5m/s. Then the nighttime E×B drifts suddenly weakened with a maximum decrement of -15m/s in the dawn, 26 August (LT, 490 it was at night). Subsequently, the daytime differential E×B drifts turned positive from UT8, 491 26 August with a maximum increment of 15m/s. In the forenoon, 27 August (DOY 239), the 492 differential  $E \times B$  had a hemispheric asymmetric structure, which expressed that the differential 493  $E \times B$  on DOY 239 enhanced ~5m/s in the southern hemisphere. In the Asian-Australian sector, 494 the E×B drift enhanced from UT20, 25 August with a magnitude of 5 to 10m/s (LT, it was in 495 daytime). In the following day, a slight positive  $E \times B$  irregularity was also observed. 496

The results in Figure 9-10 reveal that the thermospheric density ratio  $O/N_2$ , equatorial 497 electrojet and vertical E×B drift were suffered from the strong geomagnetic storm seriously. 498 Among them, the equatorial electrojet was activated by the disturbed electric field penetrated 499 from magnetosphere, and the changes of vertical E×B drifts may be associated with PPEFs and 500 DDEFs. Therefore, it is believed that the equatorial and mid-latitude ionospheric irregularities 501 are a combined action of multiple physical-chemical processes. The enhanced density ratio 502 O/N<sub>2</sub>, vertical E×B drift and equatorial electrojet played a decisive role in inducing the positive 503 irregularities. In the recovery phase, the hemispheric asymmetric  $O/N_2$  and  $E \times B$  drift on 504 August 27 may be responsible for the asymmetric TEC over the American sector in Figure 3. 505

# 506 3.4.2 Potential drivers of Auroral ionospheric irregularities

507 The GNSS ROTI, sounder density profiles and TIE-GCM's simulations demonstrate 508 significant Auroral ionospheric irregularities induced by the storm. For further to reveal the 509 vertical structures of Auroral ionospheric irregularities, the plasma irregularities within the

altitude of 96 to 400km along the -70°, 0° and 110° longitudes are simulated by the TIE-GCM. 510 Figure 11(a) - 11(c) show the temporal variations of differential plasma density as a function 511 of altitude. In Figure 6, the TIE-GCM's simulation find that at the layer of 150km, largest 512 plasma irregularities with a changing percent of >250% located in the 70°S – 80°S latitude 513 span. Therefore, the geographic latitude is selected as  $77.5^{\circ}$ S in Figure 11(a) – 11(c). It should 514 be noted that the vertical scale of each panel is similar to Figure 7, but for the amplitude of 515 plasma irregularities. We can find that significant plasma enhancements occurred in topside 516 and bottom-side of the Antarctic ionosphere along three longitudes. In the main phase, the 517 increment along the meridian  $-70^{\circ}$  was maximum with a value up to  $6 \times 10^{5}$  el/cm<sup>3</sup>, followed by 518 the meridian  $0^{\circ}$ , the last was the weakest increment of  $4 \times 10^{5}$  el/cm<sup>3</sup> along the meridian  $110^{\circ}$ . In 519 addition, the plasma fluctuations were also observed in bottom-side ionosphere along the 520 meridians  $-70^{\circ}$  and  $0^{\circ}$ , except the meridian  $110^{\circ}$ . 521



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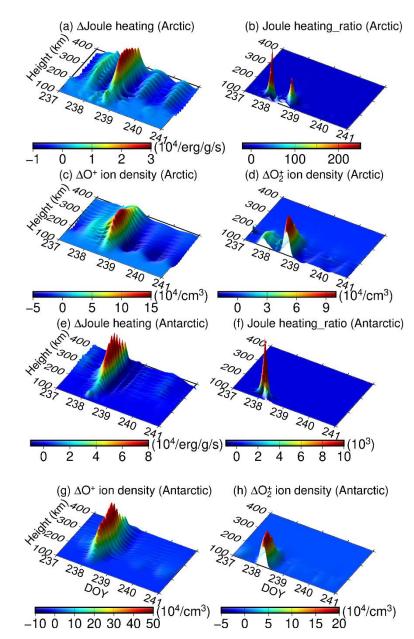
Figure 11. Structures of differential plasma irregularities along the -70°, 0° and 110° longitudes as a
function of height (a-c), geographic latitude (d-f), and geographic longitude (g-i) during DOY 237 – 240.
Did the geomagnetic storm only disturb the bottom-side ionosphere in the western
Antarctic? To address the question, the temporal variations of differential bottom-side plasma

(150km) as a function of geographic latitude along the meridians -70°, 0° and 110° are shown 527 in panels 11(d) - 11(f). It is found that significant storm-effect plasma increments occurred in 528 bottom-side ionosphere over all three longitudinal sectors, but the geographic latitudes of 529 plasma increments were not stationary. Panel 11(d) shows two plasma crests were located in 530 60°N and 80°S latitudes, respectively. The two crests had a northward movement in Eastern 531 hemisphere. For example, in the Asian-Australian sector, the plasma crests moved to 80°N and 532  $60^{\circ}$ S latitudes with a weaker value of  $3 \times 10^{5}$  el/cm<sup>3</sup>. The law of latitudinal motion of bottom-533 side plasma enhancements in different longitudinal sectors was associated with the asymmetric 534 structure of geomagnetic field. 535

Figure 11(g) - 11(i) express the longitudinal structures of differential plasma (150km) as a 536 function of day of year (DOY), the geographic latitudes in three sectors are 77.5°S, 67.5°S, 537 57.5°S, respectively. Panel 11(g) reveals positive plasma irregularities occurred in -120°  $\sim 0^\circ$ 538 longitudinal span on 26 August with a value of 4×10<sup>5</sup>el/cm<sup>3</sup>. Along the 67.5°S latitude, the 539 plasma irregularities had a double-peak structure that occurred in  $-180^{\circ} \sim 60^{\circ}$  longitudinal span. 540 Along the 57.5°S latitude, the double-peak plasma increments were observed in  $0^{\circ} \sim 90^{\circ}$  and 541  $150^{\circ} \sim -120^{\circ}$  longitudinal span. Compared to Auroral plasma irregularities, the intensity of 542 bottom-side plasma irregularities decreased in middle geographic latitude. The results conclude 543 that the strong storm not only induced topside plasma fluctuations, but also triggered positive 544 bottom-side plasma irregularities near the Auroral zone ( $\sim > 50^{\circ}$ N/S), which is consistent with 545 the sounder profiles of the radar PQ052 in Figure 5(a). 546

In order to explain the development of Auroral ionospheric irregularities, the Joule heating, 547  $O^+$  and  $O_2^+$  ion densities within the altitudes of 100 to 400 km are simulated by the TIE-GCM. 548 In addition, the neutral mass density decreases exponentially with height, and the Joule heating 549 per unit mass at higher altitude is much larger than that at the lower altitude. The bottom-side 550 change will be neglected if only focus on the differential Joule heating per unit mass, thus the 551 ratio of the changed Joule heating during storms compared to the quiet background values is 552 also investigated. In the Arctic, the study area is selected at 110°E, 67.5°N. The vertical changes 553 of Joule heating, ratio of enhanced Joule heating,  $O^+$  and  $O_2^+$  ion density over the Arctic are 554 shown in Figure 12(a) - 12(d). It is found that the Joule heating enhanced from the main phase 555 with a magnitude of  $1 \times 10^4$  erg/g/s, then the enhanced Joule heating reached a maximum in the 556

- recovery phase with a value of  $3 \times 10^4$  erg/g/s. After that, the Joule heating gradually recovered
- 558 to backgrounds.



559

Figure 12. Vertical changes of Joule heating, ratio of enhanced Joule heating, O<sup>+</sup> and O<sub>2</sub><sup>+</sup> ion density over
the locations (110°E, 67.5°N, (a)-(d)) and (70°W, 77.5°S, (e)-(h)) during DOY237-240

Similar to the variation of topside Joule heating, the positive Joule heating disturbance was also observed under 200km on DOY237-238 with a slight value of  $1 \times 10^4$  erg/g/s. Different from the absolute change of the differential Joule heating in Figure 12(a), Figure 12(b) shows the Joule heating in the main and recovery phases enhanced more 200 times than the background values, and the maximum Joule heating enhancements were located in the altitudes of 100-150km. The changes of  $O^+$  ion density in panel 12(c) agreed well with the Joule heating, the  $O^+$  ion density enhanced from DOY237 and grew stronger on DOY238 above the 200 km layer, the maximum value reached to  $1.5 \times 10^5$ /cm<sup>3</sup>. An  $O_2^+$  increment generated from UT14, DOY237, and grew to  $1 \times 10^5$ el/cm<sup>3</sup> on DOY238. Different from  $O^+$  ion, the  $O_2^+$  increments were mainly occurred under the 200 km layer, which is consistent with the behaviors of bottomside enhanced Joule heating in Figure 12(b).

In the Antarctic, the study area is selected at 70°W, 77.5°S. The changes of Joule heating 573 over the Antarctic agreed well that over the Arctic, but the positive Joule heating disturbances 574 only appeared in the main phase, which is consistent with the variation of storm-effect Antarctic 575 ionospheric plasma in Figure 11(a). The magnitude of enhanced Joule heating over the 576 Antarctic was several times larger than that over the Arctic. Figure 12(e) expresses the topside 577 and bottom-side Joule heating enhanced about  $8 \times 10^4 \text{erg/g/s}$  and  $2 \times 10^4 \text{erg/g/s}$ , respectively. 578 However, Figure 12(f) indicates that the changed ratio of bottom-side Joule heating in the main 579 phase was larger than that in the topside significantly, and the maximum enhanced ratio exceed 580  $1 \times 10^4$ . Similar to Figure 12(c) - 12(d), in the main phase, the O<sup>+</sup> ion density over the Antarctic 581 enhanced about  $5 \times 10^{5}$ /cm<sup>3</sup> above the 200km layer, and the bottom-side O<sub>2</sub><sup>+</sup> ion density 582 enhanced about  $2 \times 10^5$  el/cm<sup>3</sup>. 583

The changed amplitudes of Joule heating,  $O^+$  and  $O_2^+$  ion density over the Antarctic were much stronger than that over the Arctic, which agrees well with the magnitude of Polar ionospheric plasma disturbance in Figure 11. During a space weather event, the sudden enhanced energy could ionize the main neutral gases  $O_2$  and  $N_2$  that leads to an increment in ion density [*Gordon et al.*, 2020]. Our simulations confirmed the theory that the enhanced Joule heating could accelerate the Polar ionospheric ionization process, and the enhanced  $O^+$  and  $O_2^+$ ion densities are responsible for the topside and bottom-side plasma increments, respectively.

Geomagnetic storms not only form storm-enhanced densities (SEDs) in low-middle latitudes and tongues-of-ionization at the polar cap, but also change the global magnetic field and strength ionospheric-magnetospheric current systems [*Walach et al.*, 2021]. Ionosphere is a conductor, and the Polar ionosphere contains significant electric fields. The electric fields could drive the ionospheric current that close field-aligned currents flowing in the ionosphericmagnetospheric system, generate Joule heating in the upper atmosphere, and even control the circulation of ionospheric plasma that change the Polar ionospheric electron density structure.
The ionospheric electric potential contour maps calculated with the Super Dual Auroral Radar
Network (SuperDRAN) Assimilative Mapping procedure (<u>http://vt.superdarn.org/tiki-</u>
<u>index.php?page=ASCIIData</u>) are used for investigating the spatial-temporal variations of the
Polar convection patterns in the main phase of the geomagnetic storm.

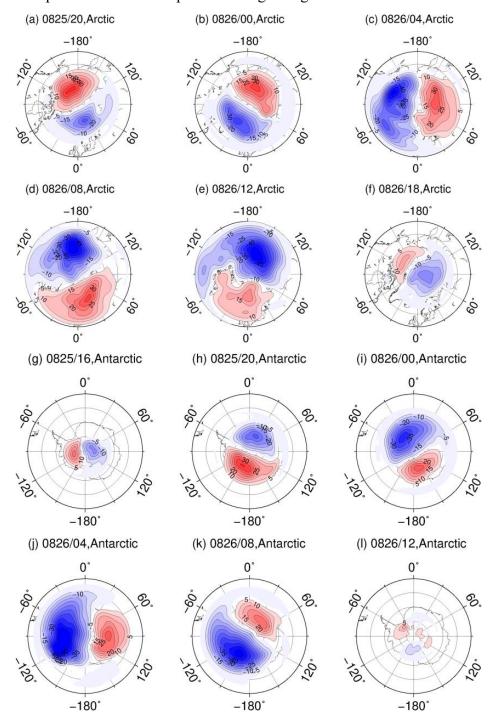




Figure 13. Maps of the electric potentials from SuperDARN over the Arctic (a-f) and Antarctic (g-l)
during August 25-26, 2018, the red and blue contours signify positive and negative potentials, respectively.

As shown in Figure 13, the positive and negative ionospheric electric potentials are 605 indicated by red and blue contours. Usually the potential pattern has a maximum near dawn 606 and a minimum near dusk. The difference between the maximum and minimum of potential is 607 called the cross polar voltage. Figure 13(a) - f(f) gives an overview of the spatial-temporal 608 evolution of electric potentials over the Arctic from UT20, 25 August to UT18, 26 August. At 609 UT20, 25 August, a positive electric potential with a maximum value of 41kv distributed in -610  $90^{\circ} \sim -180^{\circ}$  longitudinal span, and the negative electric potential with a minimum magnitude 611 of -32kv occurred in  $-60^{\circ} \sim 120^{\circ}$  longitudinal span. From 26 August, the Polar electric potential 612 intensified remarkably with the decreasing Dst index. The strongest electric potential occurred 613 in UT04-08, 26 August, concurrently with the AE index reached to a peak of 1500 to 2000nT. 614 For example, the negative potential in Figure 13(d) dropped to -53kv, on the contrary, the 615 positive potential enhanced to 36kv, and the cross polar voltage reached to 89kv. One can see 616 that during the ending of the main phase, the convection zone in the Arctic extended  $\sim 50^{\circ}$ N, 617 the scope of enhanced convection zone was consistent with the geographic latitude of radar 618 PQ052 (Figure 2(a)) that was the station with a minimum latitude could detect the bottom-side 619 620 ionospheric regularities. Similar to the Arctic electric potential, the scale and scope of electric potential over the Antarctica also enhanced and expanded remarkably. The negative potential 621 dropped from -32kv at UT20, 25 August to -61kv at UT04, 26 August, while the corresponding 622 positive potential reduced from 40kv to 33kv, the maximum cross polar voltage at the UT04, 623 26 August reached to 94kv. The scale of Antarctic storm-effect electric potential was stronger 624 than that over the Arctic, but the scope was smaller. 625

During active space weather events, the sudden energy and momentum deposited in the 626 high-latitude ionosphere and thermosphere, mostly in the forms of particle precipitation and 627 Joule heating. The incident precipitating particles gradually transfer energy to the different 628 layers of the ionosphere, and ionize more charged particles as the stronger deposited energy. 629 The accelerated ionization process enhances the ionospheric current flowing in the medium. 630 The particle precipitation and Joule heating control the variations of the short-scale structures 631 of the ionosphere-thermosphere, which results in an increment in the electric conductivity and 632 heating of the ionosphere-thermosphere system. During the April 5, 2010 geomagnetic storm, 633 the TIE-GCM simulations concluded that additional particle precipitation not only largely 634

ionospheric conductivity, but also causes remarkable 635 increases Joule heating enhancements[Sheng et al., 2017]. The enhanced conductivity, electric field, and a combination 636 of both could intensify the ionospheric electric currents. The current density is proportional to 637 the ionospheric conductivity directly, and the ionospheric conductivity is proportional to the 638 plasma density directly [Cherniak and Zakharenkova, 2018]. Therefore, there is a close 639 connection between magnetosphere energy deposition, particle precipitation, ionospheric 640 currents intensification, Joule heating, and SEDs generation. Figure 12 - 13 reveal the plasma 641 density, Joule heating and ionospheric electric potential affected by the storm all enhanced 642 significantly, which further confirms the charged particles diffusion process reported by 643 previous literatures. Thus, it is believed that the storm-induced Polar plasma irregularities are 644 associated with the additional energy input through the ways of particle precipitation, Joule 645 heating and ionospheric currents intensification. 646

# 647 **4.** Conclusion

The 25-27 August geomagnetic storm was a surprising space event that generated in the 648 649 background of very low solar activity. The prominent features of global ionospheric response to the strong geomagnetic storm that occurred at low solar activities are analyzed by ground-650 based instruments (GNSS receivers and ionosondes) and space-borne constellation (Swarm). 651 This geomagnetic triggered several unusual ionospheric plasma irregularities depend on 652 geographic longitude, latitude and altitude, and the potential drivers for explaining these 653 irregularities are also discussed using the observations of magnetometers, GUVI profiles and 654 TIE-GCM's simulations. Some important conclusions are drawn as follows: 655

656 (1) In the Asian-Australian sector, the observations of global GNSS receivers find that the storm enhanced the equatorial and mid-latitude TEC to a maximum value of 40TECU in the 657 ending of main phase and the beginning of recovery phase. While in the American sector, this 658 storm triggered a remarkable TEC hemispheric asymmetry in the recovery phase, which 659 expresses TEC depletion occurred in North America, and low-level TEC enhancements 660 occurred in mid-latitudes of South America. In the following two days, the equatorial and mid-661 latitude TEC over the American sector significant enhanced ~10TECU. The phenomenon was 662 also validated by the ionospheric topside profiles derived from the Swarm-A, the space-borne 663

observations not only confirmed the plasma density enhancements over the Asian-Australian and American sectors that happened in the main and recovery phases, respectively, but also detected an enhanced double-peak crests of EIA that caused by a daytime "superfountain" effect that driven by the PPEFs.

(2) The sounder profiles of ionosondes found that the storm induced positive plasma 668 irregularities in equatorial and mid-latitude ionosphere, and the enhanced plasma irregularities 669 primary accumulated in altitudes of 200 to 300km with a maximum frequency of 7MHz. 670 Different from TEC's change, a hemispheric asymmetric structure of ionospheric vertical 671 frequency was observed in the Asian-Australian sector, which expresses the topside (> 300km) 672 plasma over the ionosonde IC437 located in northern hemisphere increased a maximum 673 frequency of 7MHz. In addition, the profiles of the ionosonde PQ052 near the Arctic zone 674 revealed an interesting finding, that is the storm could trigger positive plasma irregularities in 675 the bottom-side (<150km) ionosphere near Auroral zone. Furthermore, the TIE-GCM 676 succeeded to simulate the temporal variation of differential plasma density as a function of 677 geographic latitude in the altitudes of 150 to 500km. The simulation not only discovered 678 679 positive plasma irregularities with a ratio of > 250% at the layer of 150km, but also captured the development of the double crests of EIA in the altitudes of 250 to 400km. 680

(3) This study first time to give an overview of the development of global ROTI in the 681 whole phase. The global ROTI maps found remarkable hemispheric asymmetry of plasma 682 irregularities in a particular time. In the beginning of the main phase (UT12 - 22, 25 August), 683 the ROTI in the American sector had a hemispheric asymmetric structure, which expressed the 684 plasma irregularities in North America were larger than that in South America, the maximum 685 irregularities appeared in Auroral zone and a narrow channel along the western coast of North 686 America with a value of 0.1TECU/min. In the ending of the main phase, the plasma 687 irregularities over Africa were larger than that in Europe, while it was opposite in the Asian-688 Australian sector. In the recovery phase, the GNSS receivers not only detected large plasma 689 irregularities in nighttime Alaska, but also found a new hemispheric asymmetry in the 690 American sector, which expressed the mid-latitude plasma with a positive ROTI of 691 0.08TECU/min in South America was significantly larger than that in North America. The 692 latitudinal plasma irregularities agreed well with TEC enhancements. 693

(4) The GUVI profiles indicated that the storm also induced significant thermospheric 694 composition change during 26-27 August, which expressed positive density ratio O/N<sub>2</sub> change 695 occurred in equatorial ionosphere, and negative change appeared in Auroral zone. A 696 hemispheric asymmetry of enhanced density ratio O/N2 was observed in South America. In 697 addition, the EEJ were suffered from the enhanced equatorial electric field caused by 698 geomagnetic storm slightly. The observations of magnetosphere demonstrated slight positive 699 EEJ fluctuations occurred in the Asian-Australian and American sectors. Furthermore, the 700 701 simulations of TIE-GCM concluded that the daytime E×B drifts enhanced exceeded 15m/s in two sectors on 25-26 August. The enhanced E×B drifts reinforced the equatorial fountain effect 702 and strengthen the ionospheric double-peak structure at the layers of 250 to 350 km. The 703 equatorial and mid-latitude plasma irregularities are believed to be a combined action of 704 thermospheric composition change, equatorial electrojet, vertical E×B drifts. 705

(5) The simulations of TIE-GCM demonstrated that the storm not only enhanced topside 706 ionospheric plasma density, but also triggered positive plasma irregularities in bottom-side 707 ionosphere near Auroral zone, which agreed well the observation of ionosonde PQ052. The 708 709 bottom-side plasma irregularities had a poleward excursion along the magnetic equator, which was associated with the longitudinal offset of geomagnetic field. In addition, the positive Joule 710 heating irregularities in the altitudes of 100 to 400km were observed in both Arctic and 711 Antarctic, and the changed ratio of bottom-side Joule heating enhanced > 250%. The temporal-712 spatial changes of Joule heating were consistent with the behaviors of Polar plasma 713 irregularities. The enhanced O<sup>+</sup> ion density was responsible for the topside plasma irregularities, 714 and the increment of  $O_2^+$  ion density may be a dominating driver for the positive bottom-side 715 plasma irregularities. Furthermore, the Polar ionospheric electric potential suffered from the 716 storm severely, the cross polar voltage abruptly enhanced to 89kv and 94kv in the Arctic and 717 Antarctic, respectively, and the enhanced electric potential expanded remarkably that the 718 boundary reached to ~50°N geographic latitude. The Polar ionospheric irregularities may be 719 associated with the additional energy input through the ways of particle precipitation, Joule 720 heating and ionospheric currents intensification. 721

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