First Observations of Large Scale Traveling Ionospheric Disturbances Using Automated Amateur Radio Receiving Networks

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Abstract

We demonstrate a novel method for observing Large Scale Traveling Ionospheric Disturbances (LSTIDs) using high frequency (HF) amateur radio reporting networks, including the Reverse Beacon Network (RBN), Weak Signal Propagation Reporter Network (WSPRNet), and PSKReporter. LSTIDs are quasi-periodic variations in ionospheric densities with horizontal wavelengths > 1000 km and periods between 30 to 180 min. On 3 Nov 2017, LSTID signatures were observed simultaneously over the continental United States in amateur radio, SuperDARN HF radar, and GNSS Total Electron Content with a period of $^{2.5}$ hr, propagation azimuth of $^{163^{\circ}}$, horizontal wavelength of 1680 km, and phase speed of 1200 km/hr. SuperMAG SME index enhancements and Poker Flat Incoherent Scatter Radar measurements suggest the LSTIDs were driven by auroral electrojet intensifications and Joule heating. This novel measurement technique has applications in future scientific studies and for assessing the impact of LSTIDs on HF communications.

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Key Points:

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15	٠	Amateur radio data provides a new method for studying Large Scale Traveling Iono-
16		spheric Disturbances and HF communications impacts
17	•	Large Scale Traveling Ionospheric Disturbances are seen for the first time simul-
18		taneously in amateur radio, SuperDARN, and GNSS TEC data
19	•	Observed midlatitude Large Scale Traveling Ionospheric Disturbances are likely
20		driven by auroral zone electrojet surges and Joule heating

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21 Abstract

We demonstrate a novel method for observing Large Scale Traveling Ionospheric Distur-22 bances (LSTIDs) using high frequency (HF) amateur radio reporting networks, includ-23 ing the Reverse Beacon Network (RBN), Weak Signal Propagation Reporter Network 24 (WSPRNet), and PSKReporter. LSTIDs are quasi-periodic variations in ionospheric den-25 sities with horizontal wavelengths > 1000 km and periods between 30 to 180 min. On 26 3 Nov 2017, LSTID signatures were observed simultaneously over the continental United 27 States in amateur radio, SuperDARN HF radar, and GNSS Total Electron Content with 28 a period of ~ 2.5 hr, propagation azimuth of $\sim 163^{\circ}$, horizontal wavelength of ~ 1680 km, 29 and phase speed of ~ 1200 km hr⁻¹. SuperMAG SME index enhancements and Poker 30 Flat Incoherent Scatter Radar measurements suggest the LSTIDs were driven by auro-31 ral electrojet intensifications and Joule heating. This novel measurement technique has 32 applications in future scientific studies and for assessing the impact of LSTIDs on HF 33 communications. 34

35 Plain Language Summary

Large Scale Traveling Ionospheric Disturbances (LSTIDs) are variations in the iono-36 sphere with wavelengths greater than 1000 kilometers, periodicities between 30 minutes 37 to 3 hours, and speeds greater than about 1400 kilometers per hour. Auroral zone dis-38 turbances are generally cited as the energy source for LSTIDs. In this paper, we show 39 for the first time that LSTIDs can cause variations in the distances amateur (ham) ra-40 dio operators can communicate using data from the Reverse Beacon Network (RBN), 41 Weak Signal Propagation Reporter Network (WSPRNet), and PSKReporter amateur 42 radio networks. The LSTID signatures in the amateur radio data are in excellent agree-43 ment with LSTID observations from two well-established instruments: the Blackstone, 44 Virginia SuperDARN radar and a large scale network of GNSS based ionospheric To-45 tal Electron Content receivers. The observed LSTIDs appear 2 to 3 hours after auroral 46 zone disturbances are detected by ground magnetometers in the SuperMAG network and 47 the Poker Flat Incoherent Scatter Radar (PFISR) in Alaska. Results suggest that au-48 roral zone disturbances were the ultimate cause of the observed LSTIDs. This paper pro-49 vides a foundation for using large-scale, crowd-sourced amateur radio observations of LSTIDs 50 as a new method for the study of LSTIDs. 51

52 1 Introduction

Traveling ionospheric disturbances (TIDs) are quasi-periodic variations of ionospheric 53 densities in Earth's upper atmosphere, believed to be the ionospheric signatures of at-54 mospheric gravity waves (AGWs) (Hines, 1960). TIDs are generally categorized as ei-55 ther Large Scale TIDs (LSTIDs, horizontal speeds between 400 to 1000 m s⁻¹, periods 56 between 30 min to 3 hr, horizontal wavelengths greater than 1000 km) or Medium Scale 57 TIDs (MSTIDs, horizontal speeds between 100 to 250 m s⁻¹, periods between 15 min 58 to 1 hr, and horizontal wavelengths of several hundred km) (e.g., Francis, 1975; Georges, 59 1968; Ogawa et al., 1987). LSTIDs are typically associated with AGWs generated by Joule 60 heating and particle precipitation from auroral zone disturbances (Hunsucker, 1982; Lyons 61 et al., 2019). These AGWs may propagate equatorward for long distances, transport-62 ing energy from the auroral zone to middle and low latitudes (Richmond, 1979) and can 63 even reach the opposite hemisphere (Zakharenkova et al., 2016). 64

Since first reported by Munro (1948), TIDs have been studied using many different techniques. These include ionosondes (e.g., Galushko et al., 1998, 2003; Altadill et
al., 2020), incoherent scatter radars (e.g., Thome, 1964; Kirchengast et al., 1996; Nicolls
& Heinselman, 2007; S.-R. Zhang et al., 2021), HF Doppler radars (e.g., Samson et al.,
1989, 1990; Bristow et al., 1994; Frissell, Baker, et al., 2014; Frissell et al., 2016), broadcast AM Doppler receivers (Chilcote et al., 2015), global navigation satellite system (GNSS)

total electron content (TEC) receivers (e.g., Tsugawa et al., 2007; Zakharenkova et al.,
 2016; Dinsmore et al., 2021), and airglow imagers (e.g. Mendillo et al., 1997; Otsuka et

al., 2004; Ogawa et al., 2009). Each of these different techniques provides a unique and
 complementary view into understanding the nature of TIDs.

In addition to their scientific value, TIDs are of interest technologically due to their 75 impact on high frequency (HF, 3-30 MHz) terrestrial communications systems. The time-76 dependent variations in ionospheric electron density associated with TIDs cause a focus-77 ing and de-focusing of ionospherically refracted HF radio signals (Samson et al., 1990; 78 79 Bristow et al., 1994; Frissell, Baker, et al., 2014). These effects can manifest as quasiperiodic fading and enhancements of HF communications signals. Work by the Ham Ra-80 dio Science Citizen Investigation (HamSCI; hamsci.org) collective have demonstrated that 81 data collected by global-scale, automated HF receiving systems built and operated vol-82 untarily by amateur (ham) radio operators can be used for both scientific study of iono-83 spheric phenomena and as a way to assess ionospheric impacts on real communications 84 systems. This work includes the impacts of solar flares and geomagnetic storms (Frissell, 85 Miller, et al., 2014; Frissell et al., 2019), total solar eclipses (Frissell et al., 2018), and 86 plasma cutoff and single-mode fading (Perry et al., 2018). 87

In this paper, we present the first observations of LSTIDs in the ionosphere through 88 data collected from the Reverse Beacon Network (RBN), Weak Signal Propagation Re-89 porter Network (WSPRNet), and PSKReporter amateur radio networks. These obser-90 vations are compared to Blackstone Super Dual Auroral Radar Network (SuperDARN) 91 radar and Global Navigation Satellite System (GNSS) differential Total Electron Con-92 tent (TEC) observations. Enhancements of the SuperMAG Electrojet (SME) Index and 93 electron densities observed by the Poker Flat Incoherent Scatter Radar (ISR) prior to 94 TID observation suggest auroral activity as the main driver for the observed LSTIDs. 95

⁹⁶ 2 Datasets and Methodology

2.1 Amateur Radio

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Amateur radio operators are communications hobbyists licensed to transmit on am-98 ateur radio frequencies. Radio signals that occur in the high frequency (HF, 3–30 MHz) 99 bands can be refracted back to Earth by the ionosphere, thereby enabling long-distance, 100 over-the-horizon communications. Variability in received signals may be related back to 101 the variations in the ionospheric state. Amateur radio observations have been previously 102 used to show the impacts of solar flares and geomagnetic storms (Frissell, Miller, et al., 103 2014; Frissell et al., 2019), and also to study the impact of a total solar eclipse (Frissell 104 et al., 2018). 105

In this paper, we use observations from the RBN (Sinanis et al., 2022), PSKRe-106 porter (Gladstone, 2022), and WSPRNet (Walker, 2022) amateur radio networks to study 107 LSTIDs. Each of these networks consists of geographically distributed automated receiv-108 ing stations that are able to identify and log Morse code and/or digital amateur radio 109 transmissions. Each observed radio transmission is referred to as a "spot" that includes 110 the observation time, frequency, call signs of the transmitter and receiver, and sometimes 111 the transmitter and receiver locations as reported by the radio operator. When station 112 location is not provided, it is determined by looking up the station's licensed callsign in 113 the HamCall Database (2022). 114

115 2.2 SuperDARN

The Super Dual Auroral Radar Network (SuperDARN) is a global network of coherentscatter HF Doppler radars that operates between 8 and 20 MHz (Greenwald et al., 1995; Chisham et al., 2007; Nishitani et al., 2019). Although SuperDARN is primarily designed to study ionospheric convection by measuring the Doppler velocity of field aligned ionospheric irregularities, it also routinely observes ground scatter. Ground scatter occurs
when radar signals undergo ionospheric refraction back to Earth, reflect off the ground,
and then return back to the radar via an ionospheric path. Although the radar returns
are from ground reflections, ground scatter can still be used for ionospheric study because the ionosphere will modulate the signals as they propagate through the medium.

Samson et al. (1990); Bristow et al. (1994); Frissell, Baker, et al. (2014) and Frissell 125 et al. (2016) have shown that TIDs moving through the field of view of a SuperDARN 126 127 radar focus and de-focus the radar rays such that the ground scatter range and signalto-noise ratio (SNR) vary with the period of the TID. In many ways, SuperDARN ground 128 scatter observations are analogous to amateur radio HF communication links. In both 129 cases, HF radio signals are modulated by the ionosphere before being returned to Earth. 130 Therefore, TIDs have the potential to affect amateur radio HF communications range 131 and SNR in a manner similar to SuperDARN ground scatter observations. 132

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2.3 GNSS Total Electron Content

Dual-frequency Global Navigation Satellite System (GNSS) receiver measurements 134 are now routinely used to measure the Total Electron Content (TEC) in a column be-135 tween a ground receiver and a satellite in space by measuring the phase difference be-136 tween the two signals (Coster et al., 1990, 1992). In this paper, we use GNSS TEC data 137 from the CONUS region processed according to the algorithms by Rideout and Coster 138 (2006) and Vierinen et al. (2016). Using a similar approach as Coster et al. (2017), S.-139 R. Zhang et al. (2017), and S. Zhang et al. (2019), we use a differential TEC (dTEC) 140 analysis rather than absolute TEC to observe the TIDs. In this approach, dTEC val-141 ues were calculated by subtracting a background TEC variation computed with a low-142 pass Savitzky-Golay filter (Savitzky & Golay, 1964) using successive windows of 60 min 143 length. Only GNSS satellite-to-ground ray paths with elevations $\geq 30^{\circ}$ were used. 144

2.4 Geomagnetic and Auroral Measurements

We use the SuperMAG electrojet (SME) index and Poker Flat Incoherent Scatter 146 Radar (PFISR) observations to quantify possible driving of LSTIDs from auroral sources. 147 SuperMAG is an international collaboration of institutions that combines the observa-148 tions from over 200 ground-based magnetometers (Gjerloev, 2012). To observe auroral 149 electrojet intensifications, we use the SuperMAG-derived SME index. This value is cal-150 culated using data from all available magnetometers between 40°N and 80°N magnetic 151 latitude. Over this range, ~ 110 stations are available, providing sufficient sampling den-152 sity to allow for the geographic localization of SME intensifications (Newell & Gjerloev, 153 2011a, 2011b). The SME index is comparable to the traditional auroral electrojet (AE) 154 index derived by Davis and Sugiura (1966). We employ the SME index because the AE 155 index is derived from only 12 magnetometer stations, making geographic localization of 156 AE enhancements difficult. 157

PFISR is located near Fairbanks, Alaska (Geographic: 65.13°N, 147.47°W; Mag-158 netic: 65.3° N, 92.1° W). Magnetic midnight occurs at ~UT-9.8 hours. For the interval 159 of interest, two radar modes were used: GPSAC5 and IPY27. The GPSAC5 mode is com-160 prised of alternating code observations providing sufficient range resolution to measure 161 E-region electron density (Lehtinen & Häggström, 1987). The IPY27 mode is a low duty 162 cycle background mode composed of both alternating code and uncoded (long) pulse ob-163 servations. The long pulse observations in the F-region are further processed to estimate 164 the electric field vector using the methodology described by Heinselman and Nicolls (2008). 165 The GPS mode and the IPY27 mode were integrated to 1 minute and 5 minute resolu-166 tion, respectively. 167

Electron density observations spanning the interval of interest are provided by the 168 high range resolution alternating code data. For this study, we use the electric field ob-169 servations, which are limited to the IPY mode observations only, to quantify the pas-170 sive energy deposition rate, $Q(z) = \sigma_P(z)E^2$, where $\sigma_P(z)$ corresponds to the altitude 171 resolved Pedersen conductivity. More details regarding how the passive energy deposi-172 tion rate was calculated using PFISR observations can be found in Zhan et al. (2021). 173 The passive energy deposition rate is a proxy for the Joule heating rate, although it ex-174 cludes the effects from the neutral winds. 175

176 **3 Observations**

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3.1 LSTID Observations

Figures 1a and 1b show combined 14 MHz observations from the RBN, PSKRe-178 porter, and WSPRNet networks on 3 Nov 2017 for the period of 1200–2359 UT for transmitter-179 receiver (TX-RX) pairs with great circle distances < 3000 km (to avoid multi-hop sit-180 uations). This event was selected by making daily summary plots of amateur radio data 181 in a format similar to Figures 1a and 1b and identifying a period with clear LSTID sig-182 natures. Figure 1a shows a map of the distribution of TX-RX midpoints of communi-183 cations observed over the continental United States (CONUS). TX-RX midpoints are 184 calculated based on the assumption that ionospheric refraction occurs at the half-way 185 point between the two stations. Figure 1b presents a time series showing the TX-RX dis-186 tance for the number of 14 MHz amateur radio spots in 2 min by 25 km bins. The bot-187 tom edge of the green-yellow region shows the communications skip focusing distance 188 varying with time, especially between 13 and 18 UT. This skip-distance variation is high-189 lighted by red dots overlaid on the data, which shows a manually-fit fiducial sinusoid with 190 a 2.5 hr period centered around 1050 km range with a 150 km amplitude. A version of 191 this figure without the overlaid sinusoid is presented in Figure 4e. 192

Figures 1c and 1d show observations from the Blackstone, Virginia (BKS) Super-193 DARN radar in a format comparable to the amateur radio observations of Figures 1a 194 and 1b. Figure 1c shows the location of the BKS radar and its field-of-view (FOV). Com-195 parison of Figure 1c with Figure 1a reveals that BKS Beam 13 (highlighted in red) looks 196 northwest over a region of dense amateur radio spot coverage. Figure 1d shows power 197 parameter observations from BKS Beam 13. The radar transmit frequency ranged be-198 tween 10.802 – 11.736 MHz during this time. The scatter is predominantly ground scat-199 ter, which is analogous to the amateur radio TX-RX communications distances shown 200 in Figure 1b. Large-scale features can be observed that are common to both the ama-201 teur radio and SuperDARN observations. Most importantly, skip distance oscillations 202 are observed in the SuperDARN data that match the those observed in the amateur ra-203 dio data. The large-scale component of these oscillations is highlighted in Figure 1d with 204 a red dotted sinusoid with identical parameters as the sinusoid in Figure 1b. In both Fig-205 ures 1b and 1d, it is noted that the sinusoid best matches the data at skip distance max-206 imum, and less so at skip distance minimum. This can be attributed to smaller-scale vari-207 ations consistent with MSTIDs mixing with the LSTID activity. 208

We next compare the amateur radio and SuperDARN observations with GNSS dTEC 209 measurements. Figure 1e shows a map of CONUS dTEC at 1343 UT, corresponding to 210 the time of the first skip-distance maximum of the sinusoid in Figure 1b. LSTID wave-211 fronts occurred with a southwest to northeast orientation, especially in the central and 212 Eastern portions of the CONUS. 1343 UT corresponds to a negative phase of the LSTIDs 213 over the CONUS, as indicated by dTEC values of \sim -0.2 for a large portion of the map. 214 The black inset box in Figure 1e indicates the region from 30° to 50° N latitude and 70° 215 to 120°W longitude. A time series of the median dTEC values within this region is pre-216 sented in Figure 1f (blue line). The dotted orange line shows the data filtered with a 2 217 -4 hr bandpass filter. Significant wave activity occurred in this time series data, and 218

comparison to amateur radio and SuperDARN observations show a general trend: depressions in median dTEC correspond to increases in skip distance, and vice-versa. The
FFT spectral analysis of the median dTEC time series shown in Figure 1g shows that
the dominant spectral component of dTEC had a period of 2.4 hr, in excellent agreement
with 2.5 hr oscillations measured with the sinusoid fit to the amateur radio and SuperDARN data.

Figure 2 further shows the spatial and temporal relationship between amateur ra-225 dio and GNSS dTEC LSTID observations. Figure 2a presents the amateur radio data 226 227 first shown in Figure 1b. The 2.5 hr sinusoid from Figure 1b is overlaid with red dots and the median dTEC values from Figure 1f are overlaid as a solid white line. The dot-228 ted sinusoid and the median dTEC values exhibit an anti-correlated relationship. Four 229 times, corresponding with the maxima and minima of the 2.5 hr sinusoid, are identified 230 with vertical dashed lines. Maps of GNSS dTEC observations corresponding to these times 231 are shown in Figures 2b - 2e. Results show an inverse relationship between the amateur 232 radio skip distances observed in Figure 2a and the dTEC measurements in Figures 2b 233 - 2e. Specifically, when maxima in amateur radio skip distances occur at 1343 and 1613 234 UT, a decrease of ~ 0.20 TECu is observed in the central regions of the maps. Conversely, 235 when minima in amateur radio skip distances occur at 1458 and 1728 UT, an increase 236 of ~ 0.20 TECu is observed. Wavefronts oriented from southwest to northeast can be 237 observed in the dTEC maps. This is most clearly seen in Figure 2d, where a black ar-238 row indicates the estimated horizontal wavelength ($\lambda_h \approx 1680$ km) and propagation 239 azimuth ($\alpha \approx 163^{\circ}$) of the largest wave feature in the map. Movie versions of Figure 240 2 showing the full progression of dTEC with time are provided in Supporting Informa-241 tion S1 and S2. Using this movie and the open-source Tracker Video Analysis and Mod-242 eling tool (Brown & Cox, 2009), the phase speed of the southeastward propagating LSTID 243 trough between 1300 and 1400 UT was estimated to be ~ 1220 km hr⁻¹. 244

In order to estimate the phase speed of the LSTIDs in the amateur radio data, Fig-245 ure 3 shows time series of latitudinal and longitudinal data slices plotted using a satu-246 rated filled contour from 1400 to 1800 UT. This time range is centered around the 1618 247 UT skip distance maximum identified in Figure 2a. The top four rows show 1° latitu-248 dinal slices that range from 42° to 38° N and extend from 88° to 74° W longitude. The 249 bottom four rows show 2° longitudinal slices that range from 85° to 79°W and extend 250 from 37° to 44° N latitude. Red arrows indicate the time of the skip distance maxima 251 manually identified in each time series plot. The arrows in the latitudinal slices indicate 252 a steady forward progression in time from the slice centered at 41.5° N to the one cen-253 tered at 39.5°N, consistent with a north-to-south propagating LSTID. Using a linear re-254 gression of distance traveled versus time, the north-to-south phase velocity was estimated 255 to be $\sim 1206 \text{ km hr}^{-1}$. Note that the skip distance maximum in the 38.5°N slice appeared 256 to move backwards in time. We ascribe this non-coherent behavior (compared to higher 257 latitude bins) to the multi-dimensional complexity of the wave field the radio signals prop-258 agated through at that time. The red arrows in the longitudinal slices (Figure 2b) show 259 almost no progression with time, which is consistent with a predominantly north-south 260 propagating LSTID that has east-west oriented wavefronts spanning the entire longitu-261 dinal observational range. 262

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3.2 Geomagnetic Conditions and Auroral Zone Drivers

Figure 4 shows evidence of auroral zone activity preceding the midlatitude observation of LSTIDs by amateur radio. Signatures of two auroral electrojet enhancements can be seen in Figure 4a, where the SME index first peaked to ~500 nT between 10 and 12 UT, and then subsequently increased to ~700 nT between 12 and 13 UT. Figure 4b presents the regional SME index, which indicates that these enhancements occurred within the 22 - 04 magnetic local time (MLT) sector. This occurs during the recovery phase of a minor geomagnetic storm with Kp \leq 3+ and minimum Sym-H \approx -30 nT at 00

UT. Figure 4c shows that these electrojet surges were associated with electron density 271 enhancements from 90 to 150 km altitude as measured by PFISR, whose observation point 272 migrated from 2125 MLT at 8 UT to 0230 MLT at 13 UT. These electron density en-273 hancements were also associated with significant Joule heating measured by PFISR at 274 these same altitudes, as shown in Figure 4d. Figure 4e again presents the amateur ra-275 dio data in a similar format to Figure 1a, now starting at 08 UT. Figure 4e shows that 276 the TIDs are observed at midlatitudes by the amateur radio networks ~ 2 to 3 hours af-277 ter the onset of auroral zone activity. Note that no radio spots were observed between 278 8 and 12 UT because of a lack of 14 MHz radio propagation, due to lower nighttime mid-279 latitude ionospheric electron densities. 280

The large-scale nature of the observed mid-latitude LSTIDs and their predominantly 281 equatorward propagation suggest that an auroral zone source is likely. We can relate the 282 TID observations to the auroral zone disturbances by estimating the location of the source 283 region using the measurements of LSTID phase speed, propagation azimuth, and tim-284 ing. To estimate the location of the LSTID source region, we start at the point corre-285 sponding to the arrow tail in Figure 2d at the top of the LSTID observation region (44°N, 286 93° W). We then project backwards from the 163° propagation azimuth at the phase speeds 287 determined using the amateur radio and GNSS dTEC data. A low estimate using a 2 288 hr propagation time and 1100 km hr⁻¹ speed places the source region at geographic (62° N, 289 105°W) and magnetic (70°N, 43°W, 0239 MLT). A high estimate using a 3 hr propa-290 gation time and 1300 km hr^{-1} speed places the source region at geographic (76°N, 132°W) 291 and magnetic (77°N, 87°W, 2341 MLT). Both of these estimates place the source region 292 in areas consistent with the auroral electrojet enhancement observed using SME index 293 and the Joule heating enhancement observed using the PFISR radar. This supports the 294 hypothesis that LSTIDs are generated by AGWs generated by auroral zone Joule heat-295 ing and particle precipitation (e.g., Hunsucker, 1982; Lyons et al., 2019). 296

²⁹⁷ 4 Discussion

We used data from large-scale, automated, crowd sourced amateur radio networks 298 to observe the effects of LSTIDs on 14 MHz HF communications paths over the conti-299 nental United States. These observations are in excellent agreement with skip-distance 300 measurements made by the Blackstone, VA SuperDARN radar and dTEC measurements 301 made by ground-based GNSS receivers. Observations of LSTIDs by these amateur ra-302 dio networks are significant for two reasons. First, these observations demonstrate a novel 303 technique for the scientific study and characterization of LSTIDs. The RBN, WSPRNet, 304 and PSKReporter amateur radio networks have global-scale data that extends over an 305 entire solar cycle back to 2008 and simultaneously observes multiple frequency bands from 306 1.8 to 30 MHz. These datasets have the potential to complement and extend existing 307 professional instrumentation networks both in geographic and spectral extent. The re-308 sults here indicate that these datasets are appropriate for statistical searches of LSTIDs 309 similar to Frissell et al. (2016), and such analyses can provide further understanding of 310 the nature of LSTIDs and their connection to space and the neutral atmosphere. Sec-311 ondly, this new technique now allows LSTID impacts on actual HF communications sys-312 tems to be assessed and directly related to measurements made by professional scien-313 tific instrumentation. This has the potential to enable the future development of method-314 ologies to better understand and potentially predict the impacts of space weather and 315 the atmosphere on HF communications systems. 316

While we have highlighted the agreement of the amateur radio, SuperDARN, and GNSS dTEC TID observations, it is also important to note some of the differences and recognize that each technique does in fact provide a unique view of the ionosphere. Amateur radio and SuperDARN both sense TIDs through bottomside oblique HF ionospheric sounding and therefore have similar measurements. Still, the amateur radio observations are able to show a continental-scale ionospheric behavior that may not be appreciated with SuperDARN radars. Conversely, SuperDARN is able to better resolve fine-scale TID structures than the amateur radio technique.

It is also not reasonable to assume a strict one-to-one mapping of TIDs observed 325 with the bottomside HF sounding techniques and the GNSS dTEC technique. GNSS TEC 326 is a height integrated measurement that is not guaranteed to be sensitive to ionospheric 327 structures at the same altitudes as the HF systems. This is evidenced in the backward 328 phase progression seen at 38.5° N in Figure 3. This dichotomy has been reported in other 329 studies. Chilcote et al. (2015) showed, for example, that TIDs detected using Doppler 330 331 shift observations of AM broadcast signals propagated in the opposite direction of TIDs detected with GNSS dTEC. In general, we emphasize that different techniques may be 332 useful for extracting greater information through collective multi-technique study of a 333 single event, and also emphasize that each technique may provide unique information 334 in its own right. 335

This paper demonstrates only the first example of using amateur radio networks 336 to study LSTIDs. Future work includes automating amateur radio LSTID detection, im-337 proving the ability to localize LSTID measurements and estimate propagation direction, 338 conducting statistical studies, and working towards the development of methods to bet-339 ter resolve smaller-scale features such as MSTIDs. There are also important implications 340 for ionospheric citizen science, as the HamSCI Personal Space Weather Station project 341 (Collins et al., 2021) will be capable of contributing to both the WSPRNet and PSKRe-342 porter datasets. We recognize that the amateur transmissions employed here are not guar-343 anteed to be regular or continuous. Therefore, certain challenges exist in extending this 344 technique to a larger statistical study. However, these data gaps are similar in nature 345 to gaps due to propagation conditions in the SuperDARN data set and could be addressed 346 in a manner similar to Frissell, Baker, et al. (2014) and Frissell et al. (2016). Addition-347 ally, a manual LSTID climatology for 2017 by Sanchez et al. (2021) shows that sufficiently 348 regular amateur radio observations exist to support statistical studies, especially over 349 North America and Europe. 350

351 5 Summary

We demonstrated a novel method for observing Large Scale Traveling Ionospheric 352 Disturbances (LSTIDs) using high frequency (HF) amateur radio reporting networks, 353 including the Reverse Beacon Network (RBN), Weak Signal Propagation Reporter Net-354 work (WSPRNet), and PSKReporter. On 3 Nov 2017, LSTID signatures were observed 355 simultaneously over the continental United States in amateur radio, SuperDARN HF radar, 356 and GNSS Total Electron Content with a period of ~ 2.5 hr, propagation azimuth of $\sim 163^{\circ}$, 357 horizontal wavelength of ~ 1680 km, and phase speed of ~ 1200 km hr⁻¹. SuperMAG SME 358 index enhancements and Poker Flat Incoherent Scatter Radar measurements suggest the 359 LSTIDs were driven by auroral electrojet intensifications and Joule heating. This novel 360 measurement technique has applications in future scientific studies and for assessing the 361 impact of LSTIDs on HF communications. 362

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Open Research

Amateur radio data from the Reverse Beacon Network (RBN), Weak Signal 384 Propagation Reporter Network (WSPRNet), and PSKReporter used in this paper 385 has been aggregated and deposited into a Zenodo repository (Frissell & Engelke, 386 2021). SuperDARN data used in this paper is available from Ruohoniemi et al. 387 (2022) and can be visualized with the open-source pyDARN toolkit (Schmidt et 388 al., 2021). SuperMAG data are available from SuperMAG database (SuperMAG 389 Database, 2022). The Kp and SymH indices were accessed through the OMNI 390 database at the NASA Space Physics Data Facility (NASA CDAWeb, 2022). We 391 acknowledge the use of the Free Open Source Software projects used in this analy-392 sis: Ubuntu Linux, python (van Rossum, 1995), matplotlib (Hunter, 2007), NumPy 393 (Oliphant, 2007), SciPy (Jones et al., 2001), pandas (McKinney, 2010), xarray 394 (Hoyer & Hamman, 2017), iPython (Pérez & Granger, 2007), and others (e.g., Mill-395 man & Aivazis, 2011). 396

The GNSS TEC and Poker Flat Incoherent Scatter Radar data are available 397 for download from the Madrigal database system maintained by Massachusetts In-398 stitute of Technology's Haystack Observatory (CEDAR Madrigal Database, 2022). 399 Data for the TEC processing is provided from the following organizations: UN-400 AVCO, Scripps Orbit and Permanent Array Center, Institut Geographique Na-401 tional, France, International GNSS Service, The Crustal Dynamics Data Information 402 System (CDDIS), National Geodetic Survey, Instituto Brasileiro de Geografia e 403 Estatística, RAMSAC CORS of Instituto Geográfico Nacional de la República Argentina, Arecibo Observatory, Low-Latitude Ionospheric Sensor Network (LISN), 405 Topcon Positioning Systems, Inc., Canadian High Arctic Ionospheric Network, 406 Institute of Geology and Geophysics, Chinese Academy of Sciences, China Meteo-407 rology Administration, Centro di Ricerche Sismologiche, Système d'Observation du 408 Niveau des Eaux Littorales (SONEL), RENAG : REseau NAtional GPS permanent, 409 GeoNet - the official source of geological hazard information for New Zealand, GNSS 410 Reference Networks, Finnish Meteorological Institute, SWEPOS - Sweden, Harte-411 beesthoek Radio Astronomy Observatory, TrigNet Web Application, South Africa, 412 Australian Space Weather Services, RETE INTEGRATA NAZIONALE GPS, Esto-413 nian Land Board, Virginia Tech Center for Space Science and Engineering Research, 414 and Korea Astronomy and Space Science Institute. 415

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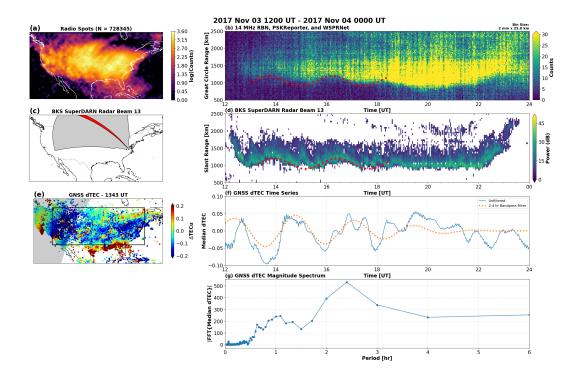


Figure 1. LSTIDs observed using amateur radio networks, the BKS SuperDARN radar, and GNSS dTEC. (a) Geographic distribution of TX-RX midpoints of amateur radio communications observed over the continental United States on 3 Nov 2017 from 1200-2359 UT. (b) Time series showing the TX-RX distance for 14 MHz amateur radio spots in 2 min by 25 km bins. (c) Location and FOV of the BKS SuperDARN radar; Beam 13 is highlighted in red. (d) Ground scatter power observations of BKS Beam 13 with ~11 MHz transmit frequency. (e) GNSS dTEC measurements at 1343 UT. (f) Time series (blue line) of GNSS dTEC median values calculated from measurements in the black box region in (e). Dotted orange line shows data filtered with a 2 - 4 hr bandpass filter. (g) FFT Magnitude spectrum of the unfiltered data in (f). Red dots overlaid on (b) and (d) show a sinusoidal 2.5 hr oscillation in skip distance common to both the amateur radio and SuperDARN measurements.

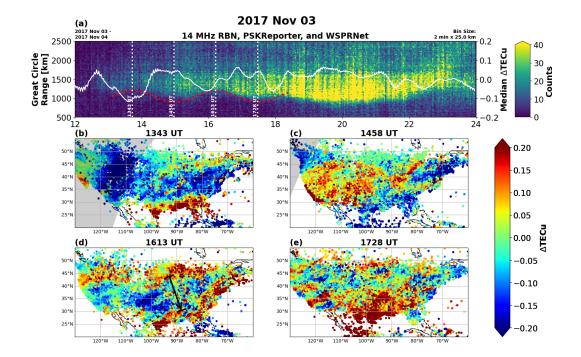


Figure 2. Amateur radio and GNSS dTEC observations of the 3 Nov 2017 LSTIDs. (a) CONUS amateur radio observations in the same format as Figure 1b. The red dashed sinusoid highlights the 2.5 hour skip distance oscillation; the white solid line shows the median dTEC values first presented in Figure 1f. Vertical white dashed lines indicate sinusoid maxima and minima times. (b – e) GNSS dTEC maps corresponding to the skip distance maxima and minima times indicated in (a). A black arrow is drawn on (d) indicating the estimated horizontal wavelength ($\lambda_h \approx 1680$ km) and direction of travel ($\alpha \approx 163^\circ$) of the GNSS LSTIDs corresponding with the amateur radio LSTIDs. A decrease of ~0.2 TECu is observed in the central region of the maps during skip distance maxima, while an increase of ~0.2 TECu is observed during skip distance minima. Movie versions of this figure are provided in Supporting Information S1 and S2.

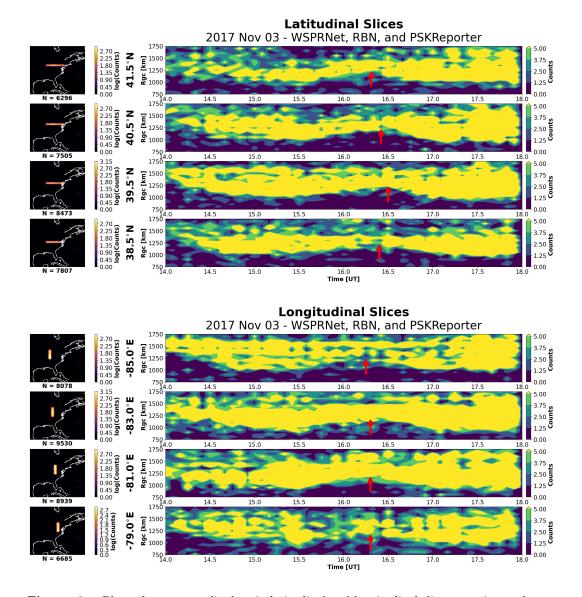


Figure 3. Plots of amateur radio data in latitudinal and longitudinal slices to estimate the phase speed of the LSTIDs. The top four rows show 1° latitudinal slices that range from 42° to 38° N and extend from 88° to 74° W longitude. The bottom four rows show 2° longitudinal slices that range from 85° to 79° W and extend from 37° to 44° N latitude. Red arrows indicate the time of the skip distance maxima manually identified in each time series plot.

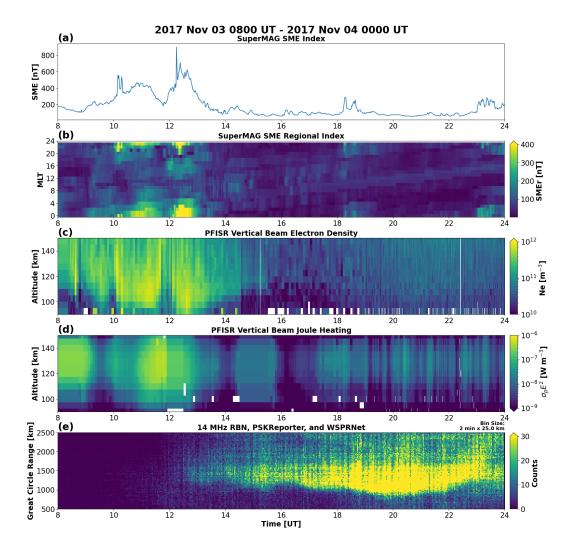


Figure 4. Measurements of auroral zone activity followed by midlatitude amateur radio LSTID observations for 0800 UT 3 Nov 2017 – 0000 UT 4 Nov 2017. (a) SuperMAG Electrojet (SME) Index. (b) Regional SuperMAG Electrojet Index. (c) Poker Flat Incoherent Scatter Radar (PFISR) vertical beam electron density measurements. (d) PFISR vertical beam Joule heating measurements. (e) Time series showing the TX-RX distance for continental U.S. 14 MHz RBN/WSPRNet/PSKReporter amateur radio spots in 2 min by 25 km bins.