

# GPS Observations of Ionospheric TEC Variations during 2015 Mw 7.8 Nepal Earthquake

Monika Karki<sup>1</sup>, Ashok Silwal<sup>2</sup>, Narayan P. Chapagain<sup>3</sup>, Prakash Poudel<sup>2</sup>, Sujan Prasad Gautam<sup>4</sup>, Roshan Kumar Mishra<sup>5</sup>, Binod Adhikari<sup>6</sup>, and Yenca Olivia Migoya Orue<sup>7</sup>

<sup>1</sup>Amrit Campus, Tribhuvan University

<sup>2</sup>Patan Multiple Campus, Tribhuvan University

<sup>3</sup>University of Illinois at Urbana Champaign

<sup>4</sup>Central Department of Physics, Tribhuvan University

<sup>5</sup>St. Xavier's College

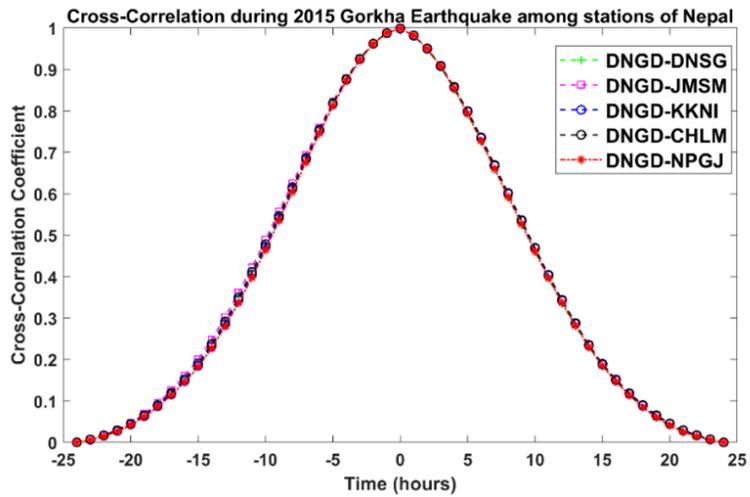
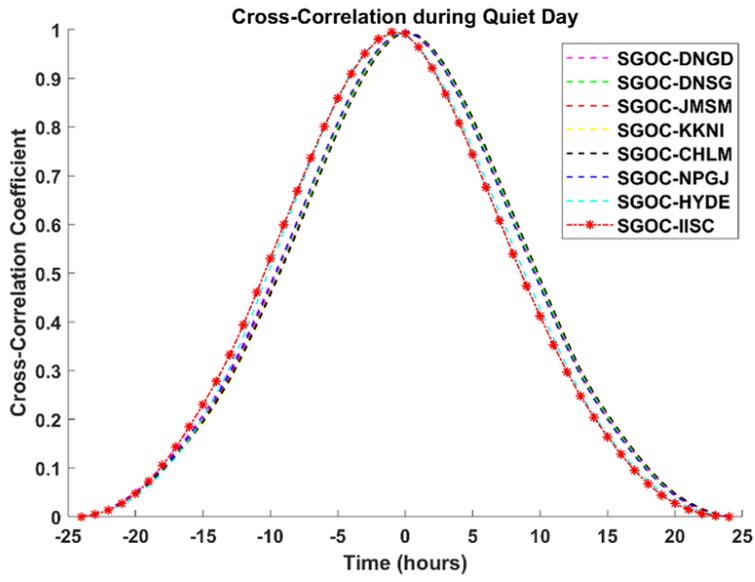
<sup>6</sup>St.Xavier's College, ,Maitighar

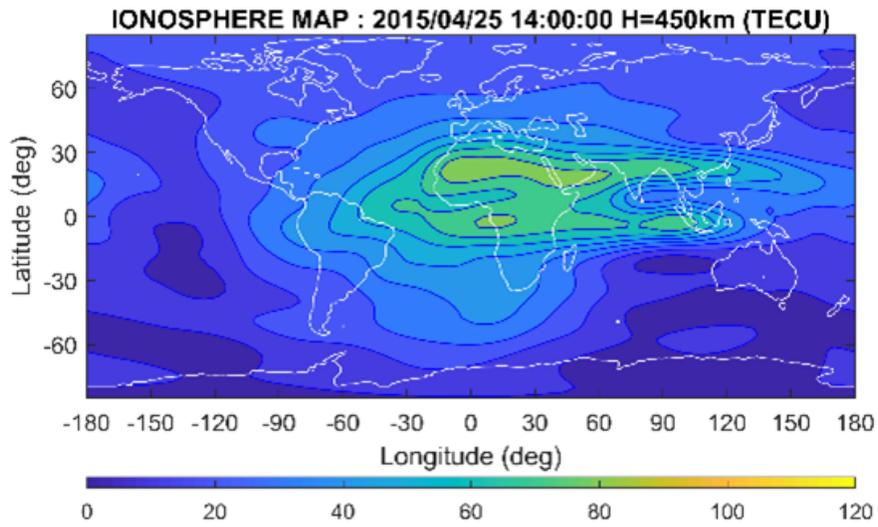
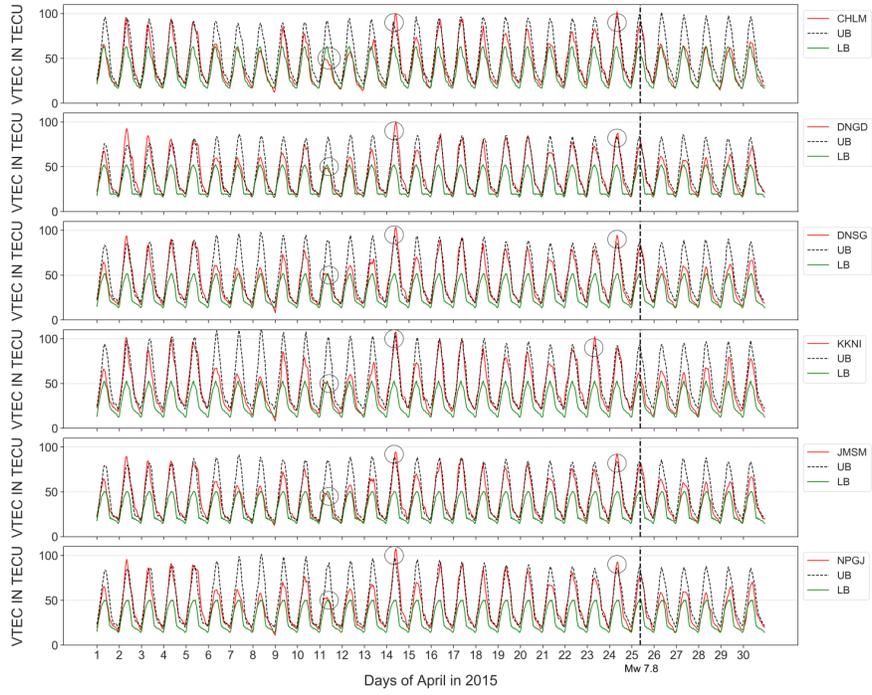
<sup>7</sup>International Centre for Theoretical Physics

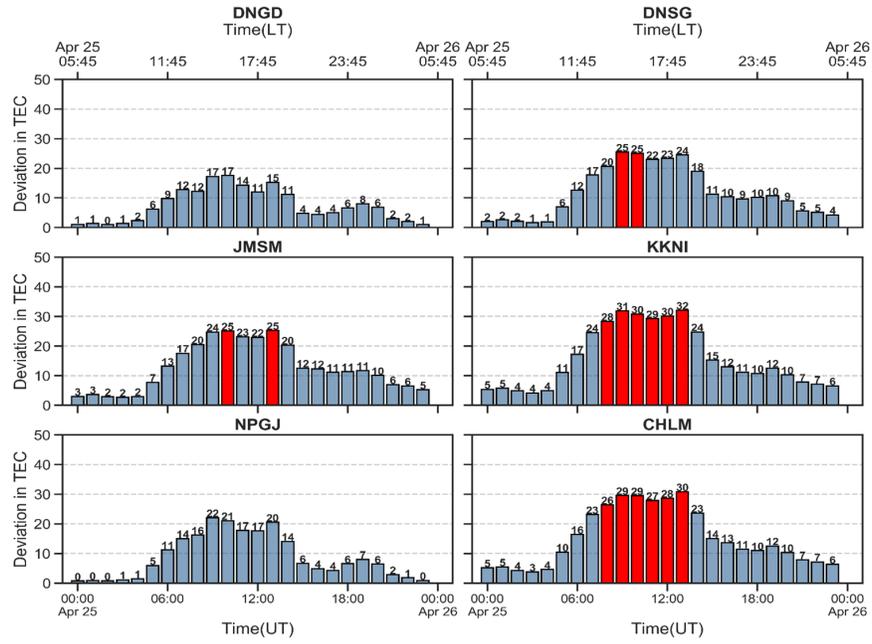
November 24, 2022

## Abstract

Ionospheric total electron content (TEC) variations prior and after to the great Gorkha Earthquake in Nepal (Mw = 7.8) on April 25, 2015, were analysed using measurements from widely distributed Global Positioning System (GPS) network. This study has been performed to understand the relationship between ionospheric TEC anomalies and earthquake occurrences. The analysis of vertical TEC (VTEC) time series from different GPS stations shows that the abnormal TEC variations appeared few days up to a few hours before the events. The results indicate that deviation in VTEC observed on the distant GPS station from the epicentre was found less relative to that of the stations near the epicentre, inferring that the variation in ionospheric VTEC nearly inversely relies upon the distance of GPS stations from the epicentre. Moreover, the pre-earthquake ionospheric anomalies were also observed in the geomagnetically conjugated region. In view of the solar-terrestrial environment, the pre-earthquake ionospheric anomalies could be associated with the Nepal earthquake. The VTEC anomaly was identified when it crosses the upper bound or lower bound. The outcomes additionally show that TEC variation was dominant in the vicinity of the earthquake epicentre. We also describe contrast in TEC throughout the globe using global ionospheric maps at regular 2-hour UT intervals, the day before, during and after the earthquake. In addition, we observed that areas heavily influenced by TEC were found to be transposed from eastern sectors to western sectors through the equatorial plane.







# GPS Observations of Ionospheric TEC Variations during 2015 Mw 7.8 Nepal Earthquake

M. Karki<sup>1</sup>, A. Silwal<sup>2</sup>, N. P. Chapagain<sup>1</sup>, P. Poudel<sup>2</sup>, S. P. Gautam<sup>3</sup>, R. K. Mishra<sup>4</sup>, B. Adhikari<sup>4,\*</sup>,  
Yenca Migoya-Orue<sup>5</sup>

<sup>1</sup> Amrit Campus, Tribhuvan University, Kathmandu, Nepal

<sup>2</sup> Patan Multiple Campus, Tribhuvan University, Lalitpur, Nepal

<sup>3</sup>Central Department of Physics, Tribhuvan University, Kirtipur, Nepal

<sup>4</sup>Department of Physics, St. Xavier College, Maitighar, Kathmandu

<sup>5</sup>T/ICT4D laboratory of the Abdus Salam International Center for Theoretical Physics, 34151  
Trieste, Italy

\*Corresponding author: binod.adhi@gmail.com

## Abstract

Ionospheric total electron content (TEC) variations prior and after to the great Gorkha Earthquake in Nepal (Mw = 7.8) on April 25, 2015, were analysed using measurements from widely distributed Global Positioning System (GPS) network. This study has been performed to understand the relationship between ionospheric TEC anomalies and earthquake occurrences. The analysis of vertical TEC (VTEC) time series from different GPS stations shows that the abnormal TEC variations appeared few days up to a few hours before the events. The results indicate that deviation in VTEC observed on the distant GPS station from the epicentre was found less relative to that of the stations near the epicentre, inferring that the variation in ionospheric VTEC nearly inversely relies upon the distance of GPS stations from the epicentre. Moreover, the pre-earthquake ionospheric anomalies were also observed in the geomagnetically conjugated region. In view of the solar-terrestrial environment, the pre-earthquake ionospheric anomalies could be associated with the Nepal earthquake. The VTEC anomaly was identified when it crosses the upper bound or lower bound. The outcomes additionally show that TEC variation was dominant in the vicinity of the earthquake epicentre. We also describe contrast in TEC throughout the globe using global ionospheric maps at regular 2-hour UT intervals, the day before, during and after the earthquake. In addition, we observed that areas heavily influenced by TEC were found to be transposed from eastern sectors to western sectors through the equatorial plane. TEC Maps indicate that most of the Indian areas, Northern China, Nepal, Bhutan, were heavily affected, addressing the earthquake's onset effect. During the day of the occurrence and one of the quietest days, we also tested the cross-correlation of TEC of SGOE from Srilanka station and various other stations; the correlation steadily increased by epicentral distance with the surrounding stations, which was another fascinating result. Thus, these findings showed that the influence of the earthquake was seen

36 globally, with the conclusion that the variation of TEC over various stations of the globe could turn  
37 out to be very helpful in predicting the earthquake.

38 **Keywords:** Earthquake, Ionosphere, Total Electron Content, Seismic Waves

### 39 **Introduction**

40 The ionosphere is a complex layer in the upper part of the atmosphere which undergoes both spatial  
41 and temporal change mainly because of the ionization effect of high-energetic solar radiation of  
42 extreme ultraviolet (EUV) and solar-related ionospheric disturbances (Rana et al., 2019). The Earth  
43 sends transient signals, sometimes stronger and more often subtle and more transitory, according to  
44 ionospheric studies, before major earthquakes which often showed the temporary disturbances  
45 within the ionosphere (Pulinets, 2007; Sharma et al., 2017). All energy input into the ionosphere  
46 was predominantly clarified by the sources from solar and geomagnetic influences, while other  
47 influences on the ionosphere were essentially ignored (Pulinets et al., 1998). A significant number  
48 of worldwide research groups are studying ionospheric perturbations prior to big earthquakes (Liu  
49 et al., 2000; Pulinets, 2002; Pulinets et al., 2005; Sharma et al., 2017; Shi et al., 2020). Earthquakes  
50 are one of nature's most destructive and unforeseen natural phenomenon. The unexpected fault  
51 rupture results into the sudden release of energy that stored in Earth's crust creating seismic waves  
52 (Ohnaka, 2013). Over the history of earthquakes, many people have been killed and many localities  
53 of the country have been devastated. Numerous attempts have also been made to use parameters to  
54 predict earthquakes using a variety of phenomena, such as Ionosonde, seismicity pattern (Wang et  
55 al. 2010), radon or hydrogen gas (Planini et al., 2004) and space-based observations of energetic  
56 electron precipitation (Kudela et al., 1992). Different techniques can be used to detect the  
57 ionospheric perturbations namely Global Positioning System (GPS)-Total electron content (TEC)  
58 measurements (Pulinets, 2002; Haase, 2011), COSMOS Satellite measurements, DEMETER  
59 Satellite measurements (Pisa et al., 2012) etc. TEC is the total number of electrons present along a  
60 path between two points, with units of electrons per square meter, where  $10^{16}$  electrons/m<sup>2</sup> =1 TEC  
61 unit (TECU) (Klobuchar, 1991, Adhikari et al., 2019). Based on these observations, ionospheric  
62 electron contents have been established as one of the key ionosphere parameters and can be applied  
63 to detect a pre-earthquake ionospheric anomaly. GPS / GNSS stations are being deployed globally  
64 and have become an important tool to research ionospheric TEC before major earthquakes. Jiang  
65 (2017) investigated the ionospheric VTEC before 2014, Mw 8.2 Chile earthquake and observed  
66 positive anomalies 4 days before the main earthquake.

67 The connection between ionospheric TEC anomalies and earthquake occurrences is ultimately  
68 governed by the lithosphere-atmosphere-ionosphere coupling mechanism (Pulinets, 2004). The  
69 ionosphere records the earthquake-induced by the change in the global electrical circuit triggered by

70 the accumulation of ions in the atmosphere, originating due to the development of stress in the  
71 crustal region prior to an earthquake (Pulinets, 2004; Friedemann and Kulahci, 2009; Sharma et al.,  
72 2017). Pullinets (2004) gave a detailed description of the process of ion cluster formation in the  
73 near-ground layer in the earthquake preparation zone. Ionospheric TEC has many difficulties for  
74 earthquake precursor studies to resolve, primarily due to lack of adequate observation stations and  
75 complications in the results of TEC data resulting from different factors. (Pullinets, 2007, 2009).  
76 Several studies reported the GPS derived TEC anomalies from 0 to 8 days before major seismic  
77 events (Pulinets, 2009; Liu et al., 2011; Sharma et al., 2017).

78 Nepal is a seismically active region of the world where there has been a devastating earthquake in  
79 the past. On April 25 2015, a large magnitude (Mw 7.8) earthquake occurred at 11:56 Nepal  
80 Standard Time (NST) (06:11:25 UTC) with an epicentre at 28.1473 °N and 84.708 °E, 8.2 km depth,  
81 34 km east-southeast of Lamjung, Nepal (Adhikari et al., 2020). The main-shock was followed by  
82 two large aftershocks of Mw 6.7 (on 26 April 2015 at 09:10 UTC) and Mw 7.3 (on 12 May at 07:05  
83 UTC).

84 In this present work, we report the disturbances of the ionosphere TEC observed by GPS TEC  
85 measurements from locations near and far from the epicentre during the April 2015 Nepal  
86 earthquake. Some new findings are drawn by studying the ionospheric TEC deviation and cross-  
87 correlation between the various GPS stations. In addition, we have used the Global Ionospheric  
88 Map (GIM) to detect TEC variation across the globe. Continuous monitoring of GIM provides an  
89 alternative way of measuring the global effects of natural disasters such as earthquakes on the  
90 ionosphere and could be used as a sophisticated tool to forecast the emergence of these natural  
91 events.

## 92 **2. Data and Method**

93 In this study, data from nine UNAVCO GPS stations have been used to recognize coseismic TEC  
94 perturbations caused by 2015 Nepal earthquake. Table 1 provides comprehensive information on the  
95 nine selected GPS stations with their geographic latitude and longitude. We analysed GPS data of  
96 30s sampling interval from two (HYDE, IISC) Indian sites, from one (SGOC) Srilanka site and 15s  
97 sampling interval from six sites (CHLM, NPGJ, KKNI, DNSG, JMSM, DNDG) of Nepal. The  
98 calibrated VTEC, captured with a time resolution of 15 sec and 30 sec was later converted to 1hr  
99 resolution by taking the average with the proper selection of time interval (1hr = 3600s). The  
100 vertical TEC data were taken from the dual-frequency carrier phases and differential code recorded  
101 by ground-based GPS receivers during the interval of 1 month, including both prior to the  
102 earthquake. The raw data from GPS station were pre-processed to convert into the receiver  
103 exchange format (RINEX). A RINEX file from the selected GPS stations was processed in order to

104 calculate vertical TEC by the help of the calibration algorithm (Ciraolo, 1993). It assumes that  
 105 ionosphere is concentrated within a thin shell located at around 357-457 km of altitude (Oikonomou  
 106 et al., 2016). The slant range TEC (STEC) can be obtained from a satellite to a receiver as the dual-  
 107 frequency carrier-phase observations L<sub>1</sub> and L<sub>2</sub> (Chen et al., 2016; Shi and Guo et al., 2019) as per  
 108 the following equation:

109 
$$\text{STEC} = \frac{(f_1 f_2)^2}{40.3(f_1^2 - f_2^2)} (L_1 - L_2 + \lambda_1 (N_1 + b_1) - \lambda_2 (N_2 + b_2) + \epsilon_L) \dots \dots \dots (1)$$
 where f<sub>1</sub> and f<sub>2</sub> are GPS

110 signal frequencies at 1575.42 MHz and 1227.60 MHz; λ is the signal wavelength; N is ambiguity; b  
 111 is the instrument biases for carrier phase; ε is residuals. The vertical TEC can be obtained by  
 112 converting STEC as:

113 
$$\text{VTEC} = \text{STEC} \times \cos \left( \arcsin \left( \frac{R \cdot \sin \alpha}{R + H} \right) \right) \dots \dots \dots (2)$$
 where R is the Earth radius; α is the elevation

114 angle of the satellite; H is the height of the ionosphere. We have chosen height at an altitude of 400  
 115 km for our analysis. For the identification of seismo-ionospheric signal, we have observed the  
 116 behaviour of VTEC for 30 running days by statistical analysis (Liu et al., 2004; Sharma et al.,  
 117 2017). To construct the upper and lower bounds, we have computed the median (μ) and standard  
 118 deviation (σ) of 15 running days written as:

119 Upper Bound (UB) = μ + 1.34 σ .....(3)

120 Lower Bound (LB) = μ - 1.34 σ .....(4)

121 If the VTEC value crosses the UB and LB, then the positive (high) and negative (low) anomalies  
 122 are considered respectively (Liu et al., 2009; Yao and Chen et al., 2012), an abnormal signal was  
 123 detected at a confidence level of about 82% (Kotz and Johnson, 1982; Oikonomou et al., 2016). In  
 124 addition to our analysis, for the same period of days, we also plotted the geomagnetic indices; Dst  
 125 (nT) and Ap (nT) along with the f<sub>10.7</sub> solar index in order to ensure that there were no inputs from  
 126 geomagnetic activities on the event period and the detected ionospheric perturbations can be  
 127 guaranteed as a consequence of the earthquake.

128

129

130 Table 1: Detailed information on GPS stations with their geographical latitude and longitude and  
 131 distance from the epicentre.

<b>GPS Station</b>	<b>Geographical Latitude (°N)</b>	<b>Geographical Longitude (°E)</b>	<b>Distance from the epicentre (km)</b>
<i>CHLM</i>	28.2072	85.3141	57.1

<i>NPGJ</i>	28.1172	81.5953	307.61
<i>KKNI</i>	27.8007	85.2788	71.9
<i>JMSM</i>	28.8053	83.7433	115.8
<i>DNGD</i>	28.7544	80.5818	409.7
<i>DNSG</i>	28.3451	83.7635	95.6
<i>HYDE</i>	17.4173	78.5579	1358.64
<i>SGOC</i>	6.8921	79.8742	2427.802
<i>IISC</i>	13.0212	77.5704	1148.146

132

### 133 **Cross-Correlation**

134 Cross-correlation is the standard statistical technique to study the relationship among different  
 135 variables along with timely analysis and draw the identical characters to probe the new information  
 136 (Usoro, 2015; Adhikari et al., 2017). In this present study, the cross-correlation technique has been  
 137 employed to establish the similarity pattern of different GPS stations in order to verify the abrupt  
 138 variation in VTEC value on 25<sup>th</sup> April 2015. The correlation coefficient is scaled from -1 to +1 and  
 139 the curve indicates strong correlation when it is close by  $\pm 1$ , whereas correlation coefficient around  
 140 zero displays moderate or less correlation between the variables (Katz, 1988; Adhikari et al., 2018).

### 141 **Results and Discussion**

142 In this section, we addressed the ionospheric response to April 25, 2015, Nepal earthquake ( $M_w$  7.8)  
 143 from GPS measurements recorded over a network of receivers which are distributed in the  
 144 epicentral region of the earthquake and sites quite far from the epicentre.

145

146

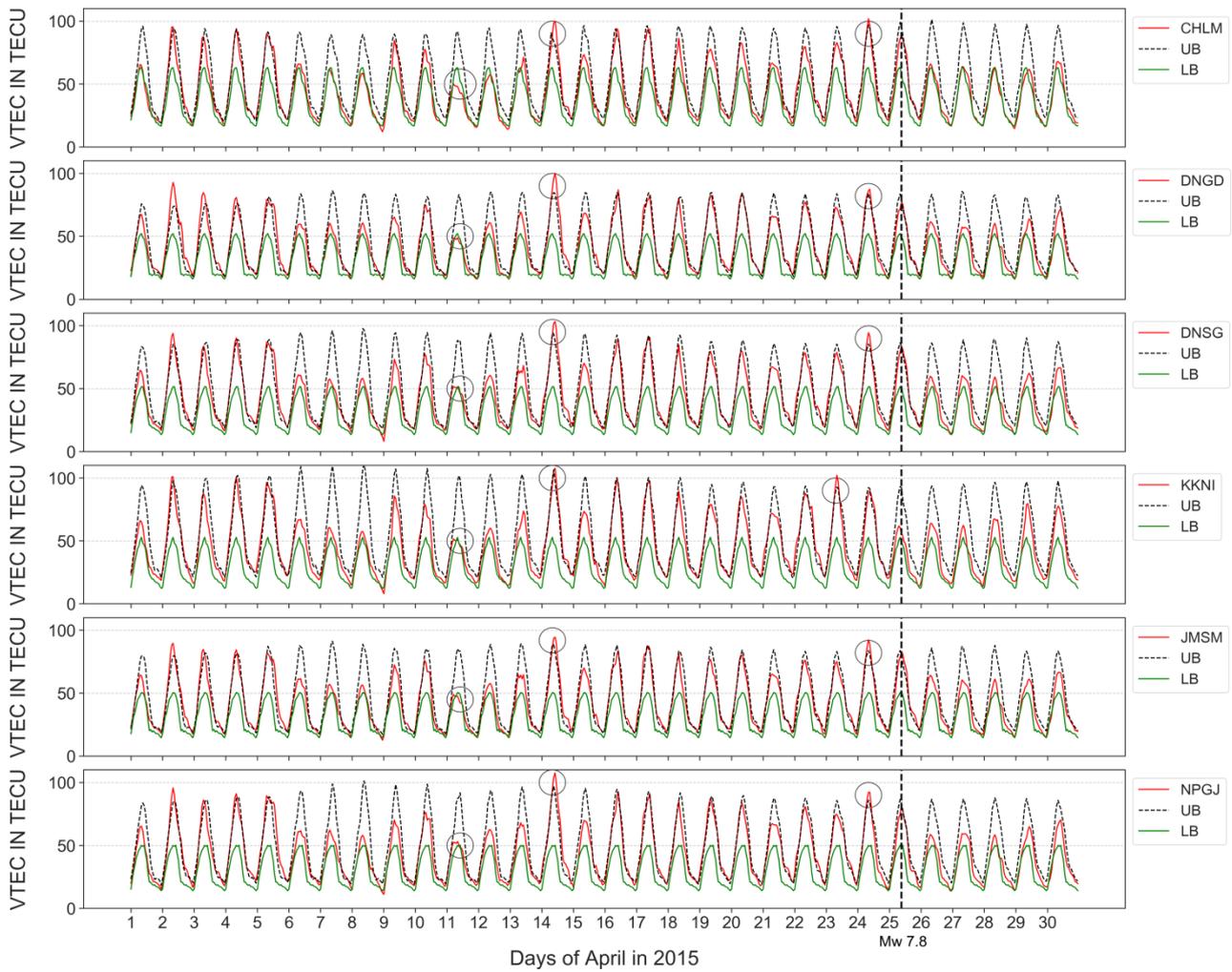
### 147 **Time Series Analysis for the month of April 2015**

148

149 In this study, we have attempted to investigate the time-series of TEC anomalies which are expected  
 150 for April 25 2015, Nepal earthquake. The TEC time-evolution analysis was carried out using six  
 151 GPS observations stations located at 57-500 km areal distance from the epicentre of the earthquake.  
 152 Time- series analysis of TEC for a period of 30 days was carried out to observe the diurnal TEC  
 153 variations, along with corresponding upper and lower bounds of TEC values. Fig 1 shows the 30  
 154 days TEC time-series observations with associated upper and lower bounds prior to and after to

155 earthquake. TEC variations preceding Nepal earthquake started almost a few days before the event.  
156 A prominent positive (high) anomaly crossing upper boundary layer was observed on 2<sup>nd</sup>, 3<sup>rd</sup>, 14<sup>th</sup>,  
157 16<sup>th</sup>, 17<sup>th</sup> and 24<sup>th</sup> prior to mainshock (25<sup>th</sup> April 2015) and low TEC value was observed on 11<sup>th</sup>  
158 April prior to and on 29<sup>th</sup> April 2015 after to mainshock at CHLM GPS station located at 57.1 km  
159 distance from the epicentre. The low TEC value observed was attributed to the reversal of electric  
160 field direction as postulated by [Pulinets, \(1998\)](#) and [Sharma et al., \(2017\)](#). In order to check the  
161 consistency of TEC anomalies, similar anomalies were expected to be observed from other stations.  
162 The analysis shows similar patterns in TEC anomalies from all six observations stations.  
163 Additionally, at other observatories GPS station, the analysis revealed high TEC anomalies crossing  
164 UBL on the same anomaly days that we observed at CHLM station except for 29<sup>th</sup> April, with  
165 including 4<sup>th</sup> April 2015 prior to the main event at DNGD located at a distance of 409.7 km from  
166 the epicentre. Observations at KKNi and NPGJ GPS stations located at a distance of 71.9 km and  
167 307.61 km respectively indicate high TEC anomaly on 5<sup>th</sup> and 23<sup>rd</sup> April before the earthquake.  
168 TEC derived from DNSG and JMSM located at a distance of 95.6 km and 115.8 km respectively  
169 showed similar anomalies like other stations. In addition to the mainshock on April 25, 2015, it was  
170 followed by big aftershocks on the same day as well as on 26<sup>th</sup> April with a magnitude of 6.7 at  
171 around 12:54:08 NST (07:08 UTC), with an epicentre located about 17 km south of Kodari, Nepal  
172 (USGS earthquake catalogue). These anomalies were also checked with other geomagnetic indices  
173 (Dst, Ap and f10.7) to detect the effect of a geomagnetic storm as shown in fig 2. It was observed  
174 that these anomalies were not affected by any geomagnetic phenomenon except for 16<sup>th</sup> and 17<sup>th</sup> of  
175 April. The change in TEC depends on solar activity, geomagnetic storm and receiving GPS station  
176 location and also varies with time and space. The energy and currents released by a geomagnetic  
177 storm increase the total height-integrated number of ionospheric TEC ([Sharma et. al., 2017](#)).  
178 Increase in ionospheric TEC results in the spatial variability of the ionosphere and cause  
179 ionospheric delays in the GPS signals ([Pulinets, 2009](#)). Overall, the TEC values corresponding to  
180 the anomaly time from all six stations was observed that the earthquakes had at least one high TEC  
181 value crossing UBL and low TEC value crossing LBL which was also observed in earlier studies  
182 ([Sharma et al., 2017, 2018, 2019, 2020](#)).

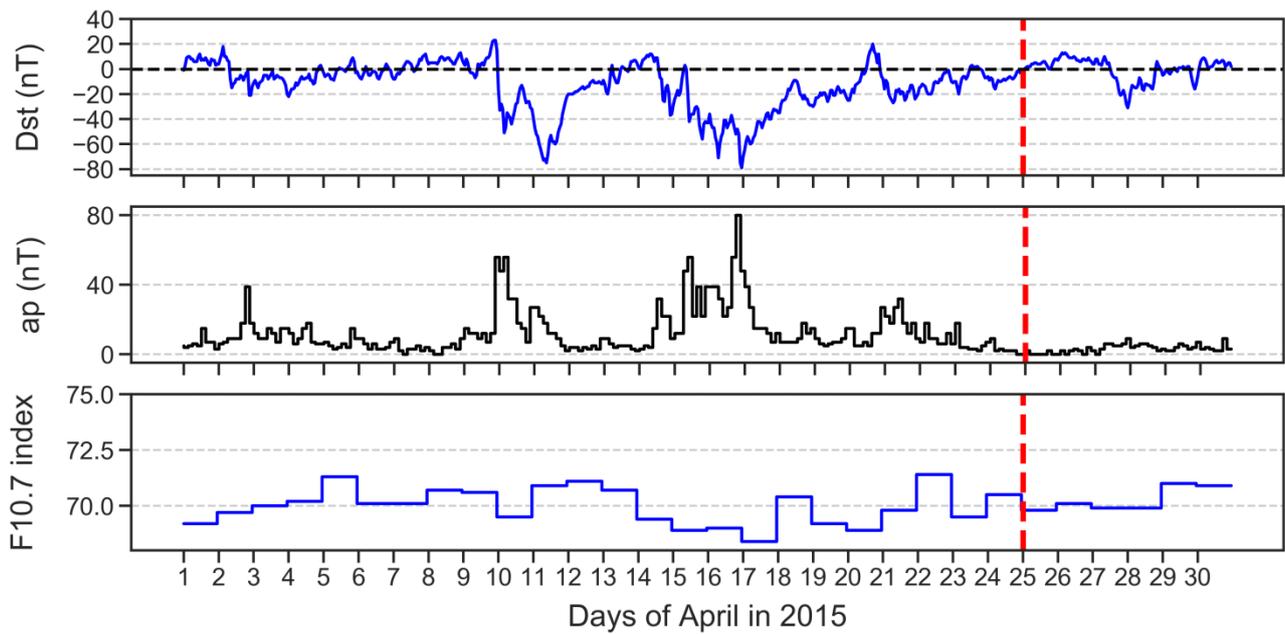
183



184

185 **Fig.1.** VTEC values observed in the TEC unit for the period of one month computed from CHLM,  
 186 DNGD, DNSG, KJNI, JMSM and NPGJ represent the TEC variation for the month of April. The  
 187 circles represent the anomalous TEC variation. Anomalies are marked when TEC crosses upper and  
 188 lower bounds.

189



190

191 **Fig.2.** Variation in Dst-index, ap-index and F10.7 index during the Nepal earthquake (25 April  
 192 2015)

193 **Comparison of quiet day TEC with that of earthquake day (25<sup>th</sup> April 2015)**

194 Fig 3 shows the variation of TEC level during the earthquake event of 25 April 2015 in comparison  
 195 to the mean TEC of topmost five quietest days of the same month at DNGD, DNSG, JMSM, KKNI,  
 196 NPGJ and CHLM stations. The trend of diurnal variation of average VTEC seems to be followed at  
 197 all the stations, minimum in a pre-dawn, a continuous inclination in the early morning followed by  
 198 an afternoon with maximum TEC value and then gradually decreases after the sunset. The value of  
 199 VTEC during “earthquake day” is exceeding the “quiet day” from some time before the earthquake  
 200 event. Significant increment on VTEC can be seen during the earthquake period. The VTEC  
 201 reached to the maximum point at ~ 8:00 UT (13:45 LT). It is found to be recovered slowly but was  
 202 greater than the quiet day value throughout the day. The increment in VTEC during earthquake  
 203 events ranges from ~ 1.8 to 4.6 TECU at the different stations at various distances from the  
 204 epicentre of the earthquake.

205

206

207

208

209

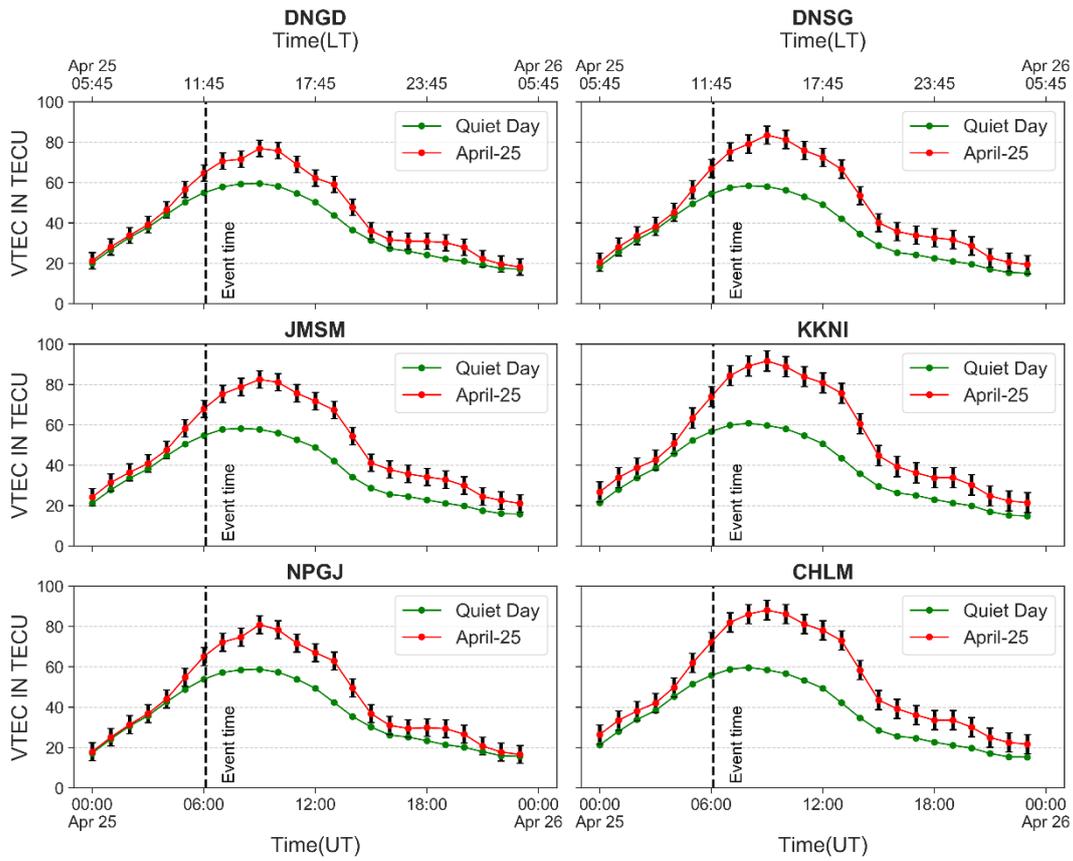
210

211

212

213

214  
 215  
 216  
 217  
 218  
 219  
 220  
 221  
 222  
 223  
 224  
 225  
 226  
 227  
 228  
 229  
 230  
 231  
 232  
 233  
 234  
 235  
 236  
 237  
 238  
 239  
 240  
 241  
 242



**Fig.3.** Comparison of TEC at six different stations during the earthquake of 25 April 2015 with the mean TEC of most five quietest days of the month. The vertical bar at earthquake day represents the standard error.

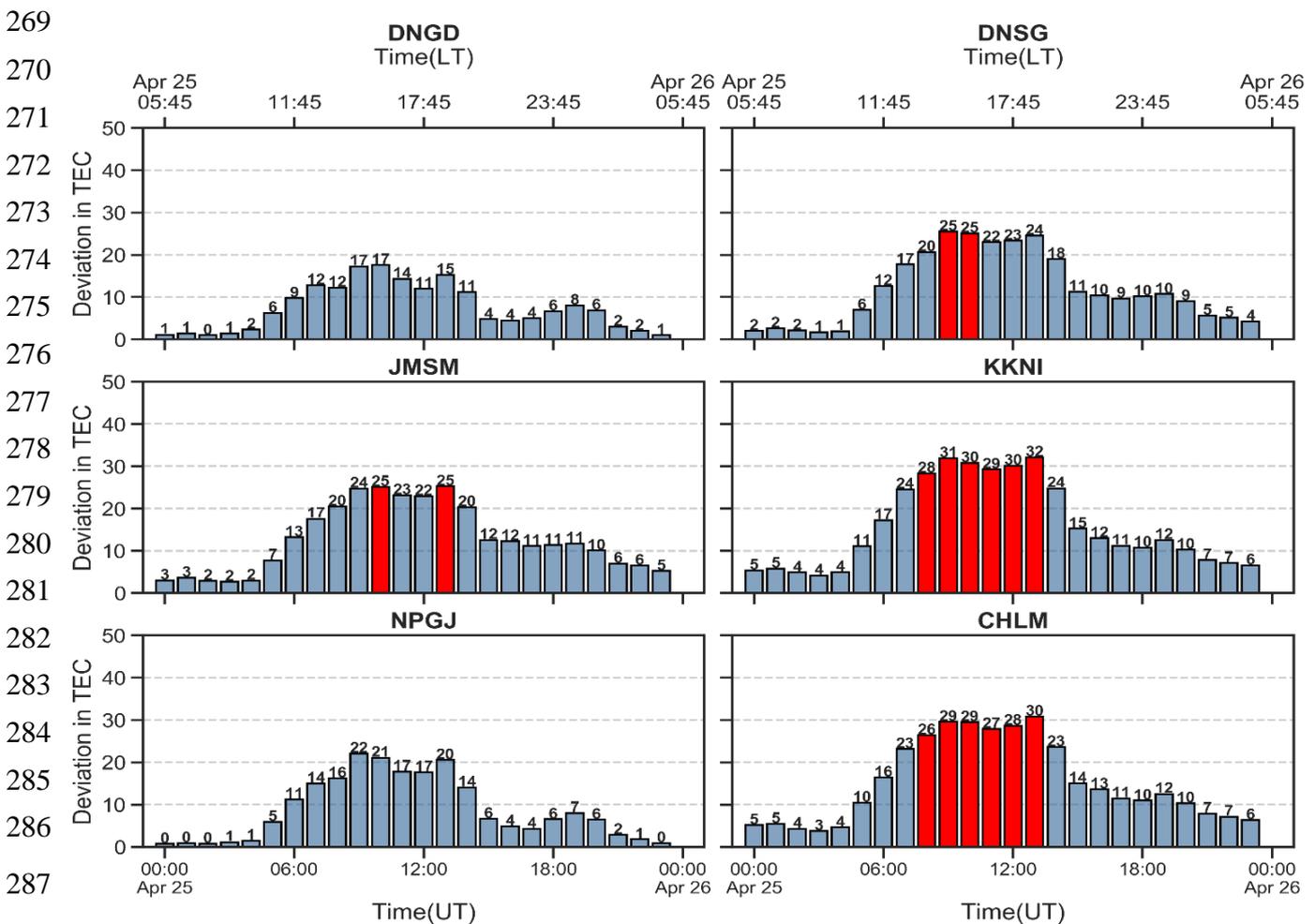
243  
 244  
 245  
 246  
 247  
 248  
 249  
 250  
 251  
 252  
 253  
 254  
 255  
 256  
 257

**Deviation in VTEC on earthquake day from a quiet day**

Fig 4 depicts that there was a significant increase in VTEC value during the earthquake event day compared to the quiet day. In this section, we study the deviation of GPS derived mean hourly VTEC in terms of percentage during the earthquake event day i.e. 25 April 2015 from mean of top 5 quietest days of the same month. The data of the top five quietest days were provided by the World Data Centre for Geomagnetism, Kyoto (<http://wdc.kugi.kyoto-u.ac.jp/qddays/>). Fig 4 describes that there was a high deviation in TEC right after the mainshock of earthquake event which was felt at 6:11 UT (~11:56 LT) and the almost constant increment remained for about 7 hours after the mainshock i.e. up to 14:00 UT. As observed from DNGD GPS station, the maximum VTEC difference between the day of the earthquake and the quiet day reached to 17 TECU whereas 25 TECU from DNSG and JMSM stations, around 32 TECU while observed from KKNI, 30 TECU from CHLM and 22 TECU from NPGJ GPS stations. The high deviation of VTEC on event day from the quiet day was observed during the time interval of 6:00 UT to 14:00 UT. However, there was a relatively small deviation of less than 10% before the mainshock and after 7 hours of mainshock i.e.

258 after 14:00 UT in all stations as witnessed from Fig 4. It is noteworthy that, throughout the earth-  
 259 quake event day the positive deviation of VTEC was recorded.

260 It was also noticed that the deviation decreased, with an increase in distance of the station from the  
 261 epicentre. The DNGD station which is ~409 km from the epicentre showed the maximum deviation  
 262 of only 17 TECU whereas deviation up to 32 TECU was spotted from KKNI GPS station which is  
 263 only 71.9 km away from the epicentre. TEC variation is found to increase as the epicentre distance  
 264 decreases which was also supported in earlier studies (Sharma et. al., 2017). The deviation of VTEC  
 265 from quiet days fairly relates with distance (see table 1) of the station from the epicentre of the  
 266 earthquake. Since earthquake event day was also the geomagnetically quiet day so, increment in  
 267 VTEC during and after the earthquake event compared to other quiet days disclosed that there was  
 268 momentous relation between seismic activity like earthquake and the ionospheric TEC.

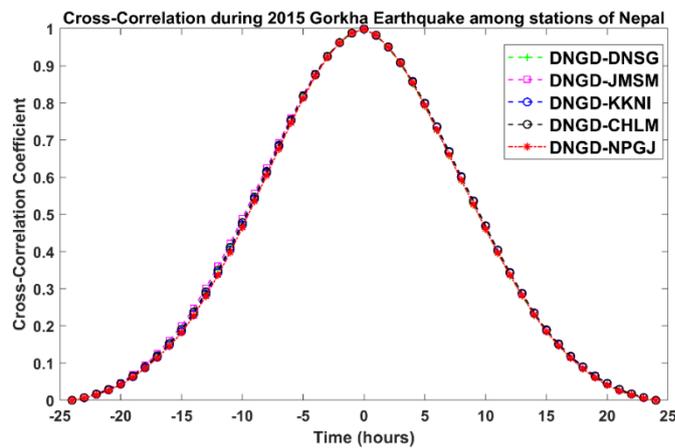


289 **Fig.4.** Deviation of TEC at six GPS stations on 25<sup>th</sup> April 2015 Nepal earthquake from the mean  
 290 TEC of five quietest days of the month. The red bars represent a high deviation in TEC values.

291  
 292  
 293  
 294

295 **Cross-Correlation Analysis**

296  
297 Fig 5(a) shows the plot of cross-correlation of mean hourly VTEC data of DNGD GPS station with  
298 mean hourly VTEC data of DNSG, JMSM, KKNi, CHLM and NPGJ GPS stations during the 2015  
299 Gorkha earthquake. The horizontal axis ranges from -25 to 25 represents the time scale in hours and  
300 the vertical axis represents the cross-correlation coefficient. Time scales help to understand the lead  
301 or lag between the indices after establishing their correlation (Adhikari et al., 2018). The Fig  
302 displays the overlapping of all curves throughout the time lag of -24 to +24 with cross-correlation  
303 coefficient +1 at a zero-time lag. This reveals that there was a very strong positive association of  
304 mean hourly VTEC data of DNGD GPS station with other mentioned stations of Nepal during the  
305 earthquake event day i.e. 25 April 2015. This verified the similar pattern of diurnal variation of  
306 VTEC over Nepal (earthquake-affected region) during the earthquake event day.

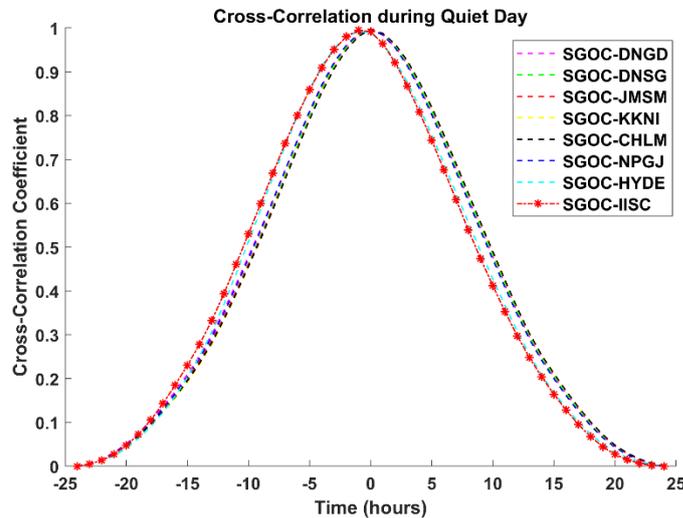


323 **Fig. 5(a):** Cross-correlation of VTEC derived from DNGD GPS station with DNSG, JMSM, KKNi,  
324 CHLM and NPGJ GPS station during the 2015 Gorkha event day.

325  
326 In order to correlate the hourly mean VTEC of different GPS stations during the quiet day, we had  
327 added SGOC GPS station of Srilanka, IISC and HYDE GPS stations of south India along with GPS  
328 stations of Nepal mentioned in the above section. All the stations are located on mid and lower  
329 latitude regions. SGOC, HYDE and IISC are the stations which were not affected by the Gorkha  
330 earthquake 2015. Fig 5(b) depicts the cross-correlation of hourly mean VTEC of top five quietest  
331 days of April 2015 derived from SGOC GPS station with hourly mean VTEC derived from other  
332 stations. It demonstrated the strong positive correlation between mean hourly VTEC derived from  
333 SGOC station and mean hourly VTEC derived from DNGD, DNSG, CHLM, NPGJ, KKNi, JMSM,  
334 HYDE and IISC with correlation coefficient value of +1 at zero time lag. This suggests that during

335 the quiet day the similar variation of ionospheric VTEC can be observed on mid and lower latitude  
336 regions irrespective of their geographical locations.

337  
338  
339  
340



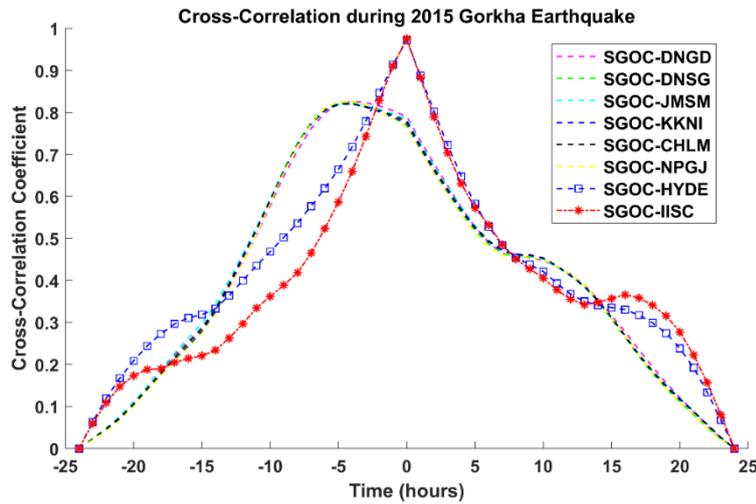
341  
342  
343  
344  
345  
346  
347  
348  
349  
350  
351  
352

353 **Fig. 5(b):** Cross-correlation of VTEC derived from SGOC GPS station with DNSG, JMSM, KKNI,  
354 CHLM, NPGJ, HYDE and IISC GPS station during the quiet day of April 2015.

355

356 Fig 5(c) shows the cross-correlation of SGOC GPS station with other stations during the earthquake  
357 event day of Gorkha earthquake 2015. In the Fig, the magenta, yellow, green, blue, black and cyan  
358 curves representing the cross-correlation of mean hourly VTEC derived from SGOC with DNGD,  
359 NPGJ, DNSG, KKNI, CHLM and JMSM GPS stations respectively are overlapped throughout the  
360 time interval of -24 to +24 with cross-correlation coefficient of 0.8 at a time lag of -5 hours. This  
361 suggests that VTEC derived from DNGD, NPGJ, DNSG, KKNI, CHLM and JMSM stations i.e.  
362 earthquake-affected stations lead SGOC by 5 hours before they get correlated. But the red starred  
363 (SGOC-IISC) and blue squared (SGOC-HYDE) line has reached the cross-correlation coefficient of  
364 +1 at zero time-lag. This means VTEC derived from SGOC and IISC as well as HYDE stations  
365 were strongly correlated during the earthquake event day. This result revealed that the VTEC from  
366 GPS stations which were not affected by earthquake shows less correlation with some lag in time  
367 with VTEC derived from earthquake-affected stations whereas it showed perfect correlation with  
368 VTEC from stations not affected by the earthquake. The results obtained from the cross-correlation  
369 analysis supports that the increase in TEC value reported in the above section was solely due to the  
370 Earthquake event.

371  
372



373  
374

375 **Fig. 5(c):** Cross-correlation of VTEC derived from SGOC GPS station with DNSG, JMSM, KKNI,  
376 CHLM, NPGJ, HYDE and IISC GPS station during the earthquake event day of Gorkha earthquake  
377 2015.

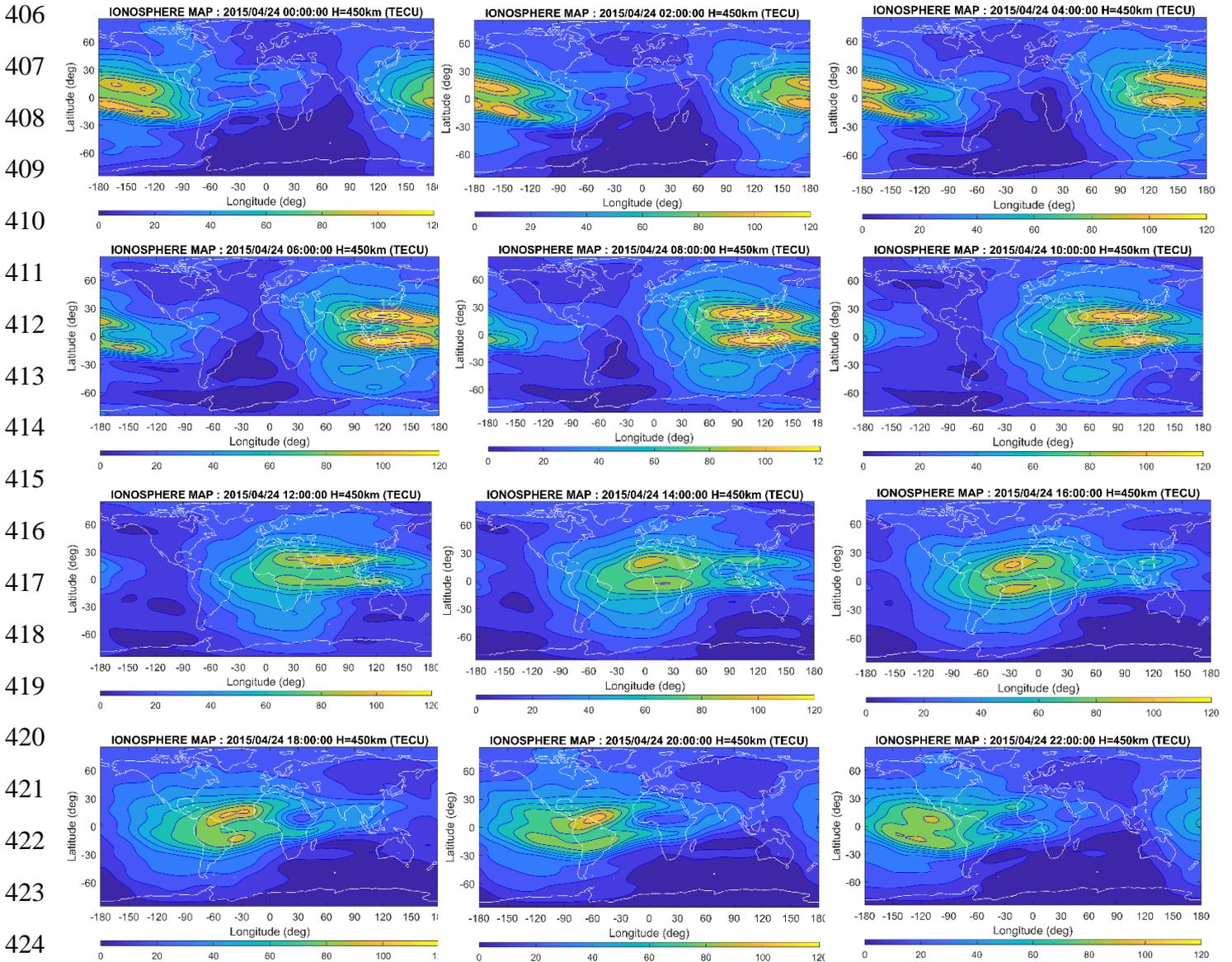
378

379 Since the cross-correlation is the standard method to estimate the degree to which two different  
380 series are correlated, this technique compares and evaluates the information between two time series  
381 of the included parameters as a function of a time lag (Finch & Lockwood, 2007; Mannucci et al.,  
382 2008). In addition, we found that the GPS TEC among the earthquake-affected stations was  
383 correlated with each other, as indicated by the cross-correlation technique. No noticeable lag has  
384 been found between the parameters discussed, which indicate a clear correlation between the  
385 specified stations. In these events, the GPS TEC response to the seismic activity is fairly rapid, as  
386 reflected by a zero-time lag value. All these stations that are affected by earthquake correspond to  
387 the similar change in global electric circuit produced by the gathering of ions in the atmosphere  
388 originating due to the development of stress in the crustal region prior to an earthquake (Pulinets,  
389 2004; Friedemann and Kulahci, 2009). On the other hand, the decreased cross-correlation  
390 coefficient with a low time lag value between the stations which are not affected by earthquake  
391 accounts for the fewer perturbations on the ionospheric TEC during the event day. This indicates  
392 that the findings obtained are encouraging because, in future days, the space weather associated  
393 with the impact of the earthquake can easily be predicted by precise measurements of lead or lag in  
394 the response of the GPS TEC linked with seismic events.

### 395 **Global Ionosphere Map**

396 Figs 6, 7 and 8 define the contrast across the globe in TEC at a daily 2-hour UT interval on the day  
397 before, during and after the 2015 April 25 Earthquake. The contour along with the colour gradient  
398 represents the global TEC intensity. The high to low TEC activity ratio can be compared at any  
399 spatiotemporal location for the assessment of changes in the ionosphere that are evolved during the  
400 day. Figs reveal ionospheric maps based on geographical longitude and latitude over a global scale,

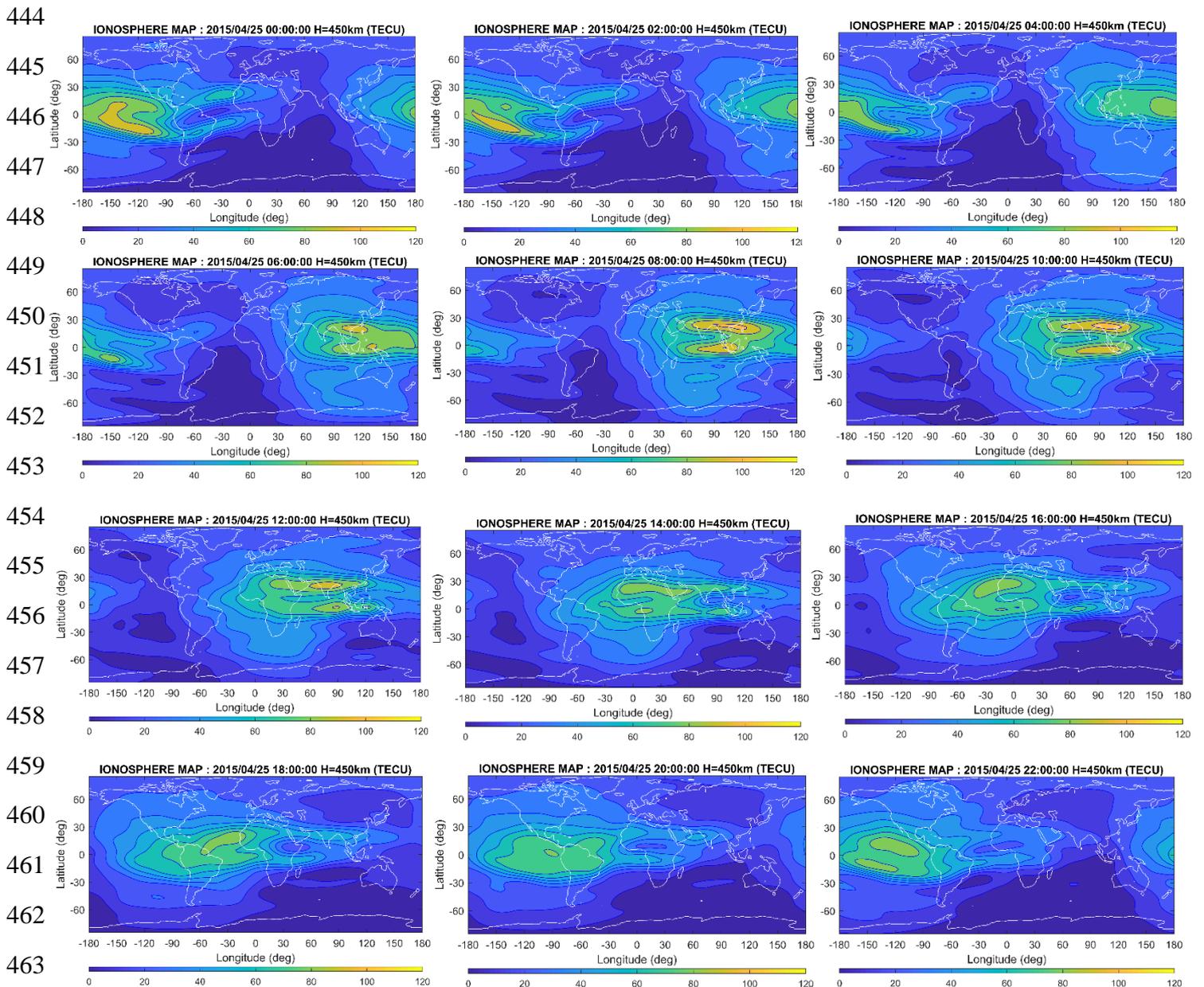
401 produced using various GPS stations in Nepal. Global contour plot indicates a transition in TEC  
 402 with a gap of 2 hours beginning at 00:00 UT. The equatorial area comprises optimum spatial and  
 403 temporal variations of the TEC and includes various characteristics of the ionosphere including the  
 404 anomaly of the equatorial system and scintillation (Paulo de Oliveira, 2009). Thus, the emphasis is  
 405 on the band of latitudes  $26^{\circ}$ - $31^{\circ}$ N, and longitudes  $80^{\circ}$ - $89^{\circ}$ E.



425 **Fig.6.** Global GPS-TEC maps on of 24<sup>th</sup> April 2015 at regular interval of 2 hours in UT.

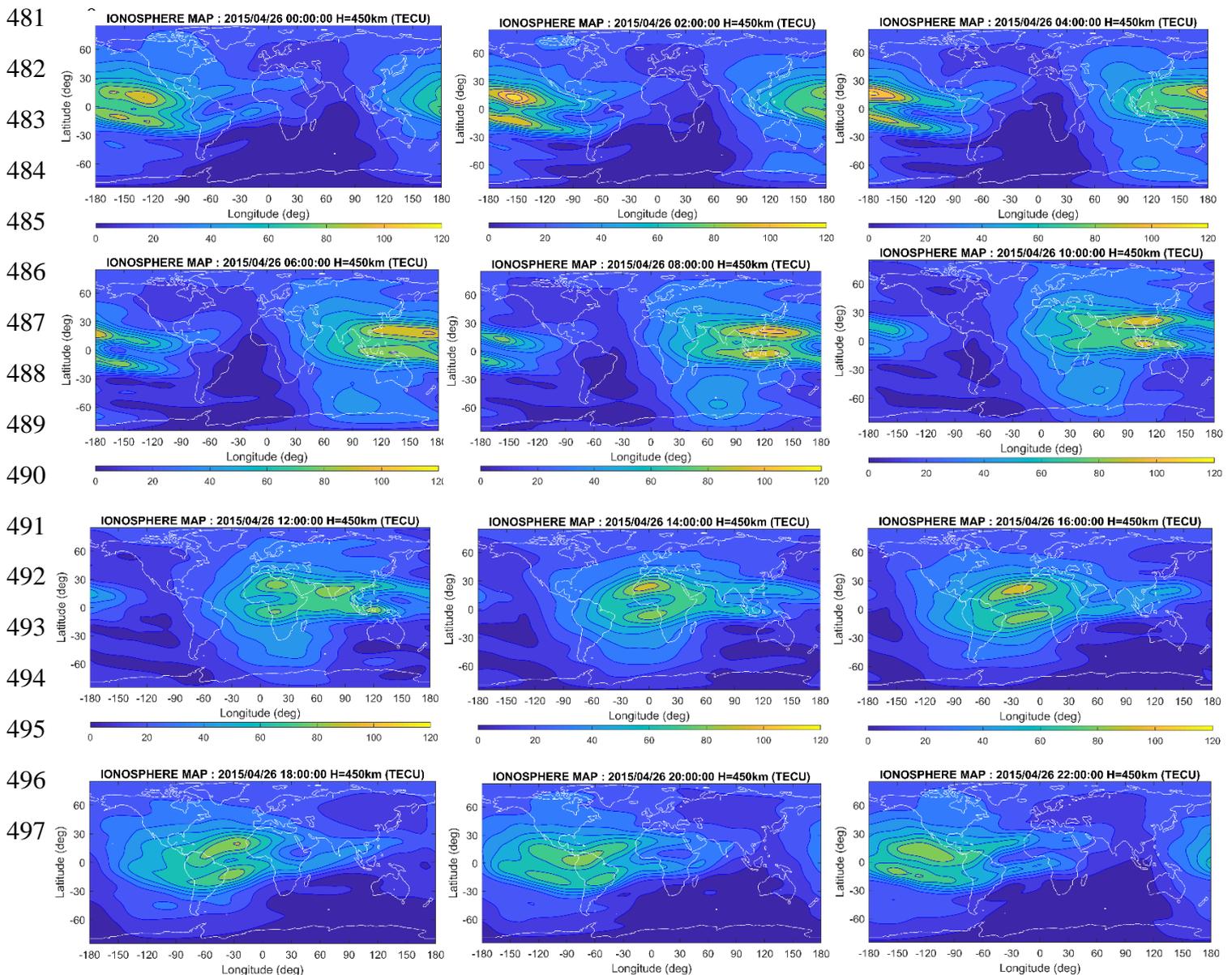
426 At 00:00 UT-02:00 UT on 24 April, TEC content is high across the Pacific Ocean relative to other  
 427 parts of the map. At 04:00 UT, the mild intensification of TEC as initially seen spreads to Indonesia,  
 428 New Guinea and the Philippines and becomes intense at 06:00 UT. At this point, the Vietnam,  
 429 Singapore and Southern China sectors are showing substantial TEC relative to the others. At 08:00  
 430 UT, there is a slight increase in TEC in the Eastern African region. From 08:00 UT-10:00 UT, the  
 431 main concentration of TEC is located at lower latitude ( $-30^{\circ}$  S to  $-30^{\circ}$  N) between the regions of  
 432 Indian and Southern China and is slowly moving towards Eastern Africa via the Arabian Sea. From

433 12:00 UT to 14:00 UT, TEC concentration is lower than the previous few hours and mild  
 434 intensification of TEC is seen over Northern Africa. At 14:00 UT, TEC concentration is observed to  
 435 be shifting towards South American region via Northern Atlantic Ocean from Africa. TEC is more  
 436 pronounced towards the Atlantic at 16:00 UT and more pronounced over the western regions of  
 437 Africa between 16:00 UT and 20:00 UT. There is a comparatively high TEC accumulation over the  
 438 low latitude Atlantic regions ( $0^{\circ}$ - $20^{\circ}$  N) at 18:00 UT. From 20:00 UT-22:00 UT, South America's  
 439 lower latitudinal regions exhibit extreme TEC. After 22:00 UT, TEC value decreases until the early  
 440 hours of 25 April between  $-30^{\circ}$  to  $30^{\circ}$  latitude across the Pacific Ocean, indicated by dark yellow  
 441 colour. On 24 April, the most extreme TEC fluctuation is observed between 8 UT-10 UT over most  
 442 of China and India's lower latitudinal area, while North American regions and Australian sectors  
 443 display minimum TEC concentration throughout the day.



**Fig.7.** Global GPS-TEC maps on 25<sup>th</sup> April 2015 at regular interval of 2 hours in UT.

465 At 00:00 UT - 02:00 UT on 25th April, TEC content is supplementary over northern Pacific and  
 466 Atlantic Ocean (40-60 TECU). At 04:00 UT, TEC concentration begins to intensify over the East  
 467 China Sea, Indian region, Bay of Bengal, Indonesia and Malaysia and turns out to be extreme at  
 468 06:00 UT. At this point, compared to other parts of the globe, the Eastern China and Northern New  
 469 Guinea regions display significant TECs. From 08:00 UT - 10:00 UT, TEC concentration rises  
 470 significantly compared to the previous few hours in most of the Indian areas, Southern China,  
 471 Nepal, Bhutan and the Indian Ocean (100-120 TECU) and by the end 10:00 UT the TEC  
 472 concentration is observed to be shifting towards the Eastern and Central Africa. At 12:00 UT, lower  
 473 latitudinal regions of Africa and India evinces mild TEC over the region. This intensification lasts  
 474 until 18:00 UT over Western African regions and the North Atlantic Ocean, and until 14:00 UT over  
 475 northern African regions. From 20:00 UT-22:00 UT, the North-West regions of South America had  
 476 a moderate TEC over the area relative to the preceding hours. After 22:00 UT, TEC value decreases  
 477 until the early hours of 26 April between  $-30^{\circ}$  to  $30^{\circ}$  latitude across the Pacific Ocean indicated by  
 478 dark yellow colour Between the period of 06:00 UT – 12:00 UT, TEC concentration is remarkably  
 479 high near to the regions of Nepal ranging from 100-120 TECU, which continues until the end of the  
 480 day through the equatorial plane covering low latitudinal regions.



498 **Fig.8.** Global GPS-TEC maps on 26<sup>th</sup> April 2015 at regular interval of 2 hours in UT.

499 Following the earthquake day, the Equatorial and Low Latitude regions are observed in the Pacific  
500 Ocean with increased TECs while the Indian belts are showing minimum effects. Mild TEC  
501 variation is observed near Northern Australia, Guinea, Indonesia and East China at 06:00 UT. At  
502 08:00 UT on 26 April, highly amplified TEC effects are seen across southern China and near  
503 Indonesia, including parts of Malaysia and the Indian regions, depicted by a dark yellow colour.  
504 Soon the TEC variations get prolonged over India, Nepal and China together with Eastern African  
505 zones at 10:00 UT near-equatorial planes. The TEC concentration is seen to be diminished over the  
506 Indian, Southern China, including most African, regions at 12:00 UT, following adverse impacts on  
507 African and Indian belts. From 14:00 UT-18:00 UT, the lower part of South America displays  
508 extreme TEC over the northern African regions (0-30 ° N), which stretch along the Atlantic Ocean.  
509 From 20:00 UT to the early morning hours of April 27, the TEC concentration decreased across the  
510 northern sector of South America along with the North Pacific Ocean. The most intense TEC is  
511 observed at 08:00 UT on 26 April, which remained between-30 ° to 30 ° latitude at 10:00 UT across  
512 the Indian and Southern China regions indicated by light yellow colour.

### 513 **Conclusions**

514 We have attempted to investigate TEC variations during Nepal earthquake of April 25, 2015,  
515 observed by GPS receivers, located in different parts of Nepal. The main conclusions of this work  
516 are drawn together and presented in this section.

517

518 1) The Nepal earthquake on 25<sup>th</sup> April 2015 with magnitude 7.8 was preceded by TEC  
519 anomalies as observed at a station closer to the epicentre (57 km) and three GPS stations at a  
520 distance of (1000-2570 km) from the epicentre. The most noticeable anomaly was observed  
521 on 14<sup>th</sup> and 24<sup>th</sup> April 2015 (one day before mainshock), although the earliest abnormality  
522 was observed on 11<sup>th</sup> April 2015 which was also observed in earlier studies ([Sharma et al.,](#)  
523 [2017, 2018, 2019, 2020](#)).

524

525 2) By analysing the data of geomagnetic indices for the same observation period, we observed  
526 that there was no geomagnetic storm i.e. geomagnetic quiet days before and after the  
527 mainshock of the earthquake. Thus, we can conclude that the TEC variations as observed  
528 were attributed due to seismogenic causes.

529

- 530 3) The analysis of deviation in VTEC on earthquake event day (also a geomagnetically quiet  
531 day) from the mean VTEC of top four quietest days showed that there was an abrupt rise in  
532 the value of GPS VTEC after the earthquake event relative to that of the geomagnetic quiet  
533 days and the high deviation persisted approximately 7 hours after the earthquake occurred.  
534
- 535 4) The less deviation of VTEC observed on the GPS station which is far from epicentre  
536 compared to that of stations nearby suggests that there was a high increment in VTEC near  
537 to the epicentre and variation in ionospheric VTEC almost inversely depends upon the  
538 distance of GPS stations from the epicentre. As the intensity of earthquake decreases with  
539 distance from the epicentre, this spatial variation of VTEC can also be summarized as the  
540 outgrowth of ionospheric TEC perturbations with the intensity of the earthquake.  
541
- 542 5) The cross-correlation between GPS derived VTEC on earthquake event day within stations  
543 of earthquake affected region (of Nepal) showed a strong correlation of VTEC with  
544 coefficient +1 at a zero-time lag. Similarly, cross-correlation study of VTEC on a quiet day  
545 among the stations of mid-latitude and lower latitudes (includes earthquake affected and  
546 unaffected stations) also showed a strong positive correlation with a peak at +1 at phase. On  
547 the other hand, during the earthquake event day, instead of taking same GPS stations for  
548 cross-correlation study, we only observed a strong positive association between VTEC  
549 derived from stations which were not affected by the earthquake (SGOC, IISC, HYDE) with  
550 a correlation coefficient of +1 at phase whereas all cross-correlation curves between VTEC  
551 derived from earthquake-affected and unaffected GPS stations displayed a peak at 0.8 with  
552 certain time-lag indicating less correlation between VTEC derived from earthquake-affected  
553 and unaffected regions. Overall cross-correlation findings have led to the conclusion that the  
554 increase in VTEC value in the earthquake event day relative to the quiet day is due to the  
555 seismic waves produced by the earthquake.
- 556 6) TEC Maps evince that most of the Indian regions, Southern China, Nepal regions were  
557 affected heavily explaining the onset effect of the earthquake. There were no significant  
558 fluctuations a day before and after the earthquake. All the Nepalese stations attain their  
559 highest TEC values during the event. TEC, when observed at Global Ionospheric Maps,  
560 prolongs towards lower or mid-latitudinal region (30 to -30 degree) through equatorial plane  
561 during the event.

562 As these naturally occurring phenomena overwhelm the Earth, we have observed abrupt  
563 modifications in ionospheric TEC data detected in different monitoring stations. By integrating

564 measurements based on the global impact of such natural events, we can monitor the response of  
565 TEC towards the seismic events and it could be used as a sophisticated tool to forecast the  
566 emergence of these natural events. Pulinets (1998) also concluded ionosphere-seismic consequences  
567 are an ongoing process and are a key cause of ionosphere variability and such consequences may be  
568 seen at least five days before an earthquake.

## 569 **Acknowledgements**

570 Co-author, Monika Karki's research work has been supported by University Grant Commission  
571 (UGC), Nepal as the MSc dissertation fellowship. Geomagnetic indices (ap and Dst) data are  
572 obtained from the OMNI ([http://isgi.unistra.fr/geomagnetic\\_indices.php](http://isgi.unistra.fr/geomagnetic_indices.php)) site. Solar flux (F10.7)  
573 index data are downloaded from the OMNI (<http://omniweb.gsfc.nasa.gov/form>) site. Ground-based  
574 dual-frequency GPS TEC data are extracted from the UNAVCO Data Archive  
575 (<https://www.unavco.org/data/gps-gnss/data-access-methods/dai2/app/dai2.html#>) for the stations  
576 mentioned in Table 1. We would like to thank staff members of NASA and UNAVCO for making  
577 the data available.

## 578 **References**

- 579 Adhikari, B., Adhikari, R., Chapagain, N. P., Sapkota, N., Dahal, S., & Pandit, D. (2017). Daily,  
580 seasonal and monthly variation of middle-low latitudes magnetic field during low solar  
581 activity. *Discovery*, 53(255), 181-190.
- 582 Adhikari, B., Dahal, S., Sapkota, N., Baruwal, P., Bhattarai, B., Khanal, K., & Chapagain, N. P.  
583 (2018). Field-Aligned Current and Polar Cap Potential and Geomagnetic Disturbances: A Review of  
584 Cross-Correlation Analysis. *Earth and Space Science*, 5(9), 440-455.
- 585 Adhikari, B., Kaphle, B., Adhikari, N., Limbu, S., Sunar, A., Mishra, R. K., & Adhikari, S. (2019).  
586 Analysis of cosmic ray, solar wind energies, components of Earth's magnetic field, and ionospheric  
587 total electron content during solar superstorm of November 18–22, 2003. *SN Applied Sciences*, 1(5),  
588 453.
- 589 Adhikari, B., Dahal, S., Karki, M., Mishra, R. K., Dahal, R. K., Sasmal, S., & Klausner, V. (2020).  
590 Application of wavelet for seismic wave analysis in Kathmandu Valley after the 2015 Gorkha  
591 earthquake, Nepal. *Geoenvironmental Disasters*, 7(1), 1-16.
- 592 Ciruolo, L., Azpilicueta, F., Brunini, C., Meza, A., & Radicella, S. M. (2007). Calibration errors on  
593 experimental slant total electron content (TEC) determined with GPS. *Journal of Geodesy*, 81(2),  
594 111-120.
- 595 Chen, P., Yao, Y., & Yao, W. (2017). On the coseismic ionospheric disturbances after the Nepal  
596 Mw7. 8 earthquakes on April 25, 2015 using GNSS observations. *Advances in Space*  
597 *Research*, 59(1), 103-113.

598 Freund, F. T., Kulahci, I. G., Cyr, G., Ling, J., Winnick, M., Tregloan-Reed, J., & Freund, M. M.  
599 (2009). Air ionization at rock surfaces and pre-earthquake signals. *Journal of Atmospheric and*  
600 *Solar-Terrestrial Physics*, 71(17-18), 1824-1834.

601 Haase, J.S., Dautermann, T., Taylor, M. J., Chapagain, N., Calais, E., Pautet, D., (2011).  
602 Propagation of plasma bubbles observed in Brazil from GPS and airglow data. *Advances in Space*  
603 *Research*, 47(10),1758-1776.

604 Kudela, K., Matisin, J., Shuiskaya, F. K., Akentieva, O. S., Romantsova, T. V., & Venkatesan, D.  
605 (1992). Inner zone electron peaks observed by the “Active” satellite. *Journal of Geophysical*  
606 *Research: Space Physics*, 97(A6), 8681-8683.

607 Kotz, S., Johnson, H. L., Read, C. B., 1982, Encyclopedia of statistical sciences, 519.5, E5.

608 Liu, J. Y., Chen, Y. I., Pulinets, S. A., Tsai, Y. B., & Chuo, Y. J. (2000). Seismo-ionospheric  
609 signatures prior to  $M \geq 6.0$  Taiwan earthquakes. *Geophysical research letters*, 27(19), 3113-3116.

610 Liu, J. Y., Chuo, Y. J., Shan, S. J., Tsai, Y. B., Chen, Y. I., Pulinets, S. A., & Yu, S. B. (2004). Pre-  
611 earthquake ionospheric anomalies registered by continuous GPS TEC measurements. In *Annales*  
612 *Geophysicae*, 22(5), 1585-1593.

613 Liu, J. Y., Chen, Y. I., Chen, C. H., Liu, C. Y., Chen, C. Y., Nishihashi, M., & Lin, C. H. (2009).  
614 Seismoionospheric GPS total electron content anomalies observed before the 12 May 2008 Mw7. 9  
615 Wenchuan earthquake. *Journal of Geophysical Research: Space Physics*, 114(A4).

616 Liu, J. Y., Le, H., Chen, Y. I., Chen, C. H., Liu, L., Wan, W., ... & Chen, M. Q. (2011). Observations  
617 and simulations of seismoionospheric GPS total electron content anomalies before the 12 January  
618 2010 M7 Haiti earthquake. *Journal of Geophysical Research: Space Physics*, 116(A4).

619 Oikonomou, C., Haralambous, H., & Muslim, B. (2016). Investigation of ionospheric TEC  
620 precursors related to the M7. 8 Nepal and M8. 3 Chile earthquakes in 2015 based on spectral and  
621 statistical analysis. *Natural Hazards*, 83(1), 97-116.

622 Ohnaka, M. (2013). *The physics of rock failure and earthquakes*. Cambridge University Press.

623 Planinić, J., Radolić, V., & Vuković, B. (2004). Radon as an earthquake precursor. *Nuclear*  
624 *Instruments and Methods in Physics Research Section A: Accelerators, Spectrometers, Detectors*  
625 *and Associated Equipment*, 530(3), 568-574.

626 Piša, D., Němec, F., Parrot, M. and Santolík, O., 2012. Attenuation of electromagnetic waves at the  
627 frequency~ 1.7 kHz in the upper ionosphere observed by the DEMETER satellite in the vicinity of  
628 earthquakes. *Annals of geophysics*, 55(1).

629 Pulinets, S. A. (1998). Seismic activity as a source of the ionospheric variability. *Advances in Space*  
630 *Research*, 22(6), 903-906.

631 Pulinets, S. A., Boyarchuk, K. A., Hegai, V. V., & Karelin, A. V. (2002). Conception and model of  
632 seismo-ionosphere-magnetosphere coupling. *Seismo-Electromagnetics: Lithosphere-Atmosphere-*  
633 *Ionosphere Coupling*, 353-361.

634 Pulinets, S. A., Contreras, A. L., Bisiacchi-Giraldi, G., & Ciruolo, L. (2005). Total electron content  
635 variations in the ionosphere before the Colima, Mexico, earthquake of 21 January 2003. *Geofísica*  
636 *internacional*, 44(4), 369-377.

637 Pulinets, S. A. (2007). Natural radioactivity, earthquakes, and the ionosphere. *Eos, Transactions*  
638 *American Geophysical Union*, 88(20), 217-218.

639 Pulinets, S. A. (2009). Physical mechanism of the vertical electric field generation over active  
640 tectonic faults. *Advances in Space Research*, 44(6), 767-773.

641 Sharma, G., Mohanty, S., & Kannaujiya, S. (2017). Ionospheric TEC modelling for earthquakes  
642 precursors from GNSS data. *Quaternary International*, 462, 65-74.

643 Parashar, D., Sudeep, A., More, A., Patil, P., Walimbe, A., Mavale, M., & Amdekar, S. (2018). Total  
644 electron content and epicentral distance of 2015 Mw 7.8 Nepal earthquake revealed by continuous  
645 observations data. *current science*, 115(1), 27.

646 Sharma, G., Ray, P. C., & Kannaujiya, S. (2019). Ionospheric Total Electron Content for Earthquake  
647 Precursor Detection. In *Remote Sensing of Northwest Himalayan Ecosystems* (pp. 57-66). Springer,  
648 Singapore.

649 Sharma, G., Saikia, P., Walia, D., Banerjee, P., & Raju, P. L. N. (2020). TEC anomalies assessment  
650 for earthquakes precursors in North-Eastern India and adjoining region using GPS data acquired  
651 during 2012–2018. *Quaternary International*.

652 Shi, K., Liu, X., Guo, J., Liu, L., You, X., & Wang, F. (2019). Pre-earthquake and coseismic  
653 ionosphere disturbances of the Mw 6.6 Lushan earthquake on 20 April 2013 monitored by  
654 CMONOC. *Atmosphere*, 10(4), 216.

655 Shi, K., Guo, J., Liu, X., Liu, L., You, X., & Wang, F. (2020). Seismo-ionospheric anomalies  
656 associated with Mw 7.8 Nepal earthquake on 2015 April 25 from CMONOC GPS  
657 data. *Geosciences Journal*, 1-16.

658 Usoro, A. E. (2015). Some basic properties of cross-correlation functions of n-dimensional vector  
659 time series. *Journal of Statistical and Econometric Methods*, 4(1), 63-71.

660 Wang, H. L., Chen, H. W., & Zhu, L. (2010). Constraints on average Taiwan Reference Moho  
661 Discontinuity Model—receiver function analysis using BATS data. *Geophysical Journal*  
662 *International*, 183(1), 1-19.

663 Yao, Y., Chen, P., Wu, H., Zhang, S., & Peng, W. (2012). Analysis of ionospheric anomalies before  
664 the 2011 M w 9.0 Japan earthquake. *Chinese science bulletin*, 57(5), 570-510.

665 Finch, I., & Lockwood, M. (2007). Solar wind-magnetosphere coupling functions on timescales of  
666 1 day to 1 year. <https://doi.org/10.5194/angeo-25-495-2007>

667 Mannucci, A. J., Tsurutani, B. T., Abdu, M. A., Gonzalez, W. D., Komjathy, A., Echer, E., ... &  
668 Anderson, D. (2008). Superposed epoch analysis of the dayside ionospheric response to four intense  
669 geomagnetic storms. *Journal of Geophysical Research: Space Physics*, 113(A3).

670 Pulinets, S. A. (1998). Seismic activity as a source of the ionospheric variability. *Advances in Space*  
671 *Research*, 22(6), 903-906.

672 Klobuchar, J. A. (1991). Ionospheric Effects on GPS, Early Innovation columns, GPS World, 1991.  
673 Available: <http://gauss.gge.unb.ca/gpsworld/EarlyInnovationColumns/Innov.1991.04.pdf>.

674

675

Table 1: Detailed information on GPS stations with their geographical latitude and longitude and distance from the epicentre.

<b><i>GPS Station</i></b>	<b>Geographical Latitude (<sup>0</sup>N)</b>	<b>Geographical Longitude (<sup>0</sup>E)</b>	<b>Distance from the epicentre (km)</b>
<i>CHLM</i>	28.2072	85.3141	57.1
<i>NPGJ</i>	28.1172	81.5953	307.61
<i>KKNI</i>	27.8007	85.2788	71.9
<i>JMSM</i>	28.8053	83.7433	115.8
<i>DNGD</i>	28.7544	80.5818	409.7
<i>DNSG</i>	28.3451	83.7635	95.6
<i>HYDE</i>	17.4173	78.5579	1358.64
<i>SGOC</i>	6.8921	79.8742	2427.802
<i>IISC</i>	13.0212	77.5704	1148.146

**Figure1.**

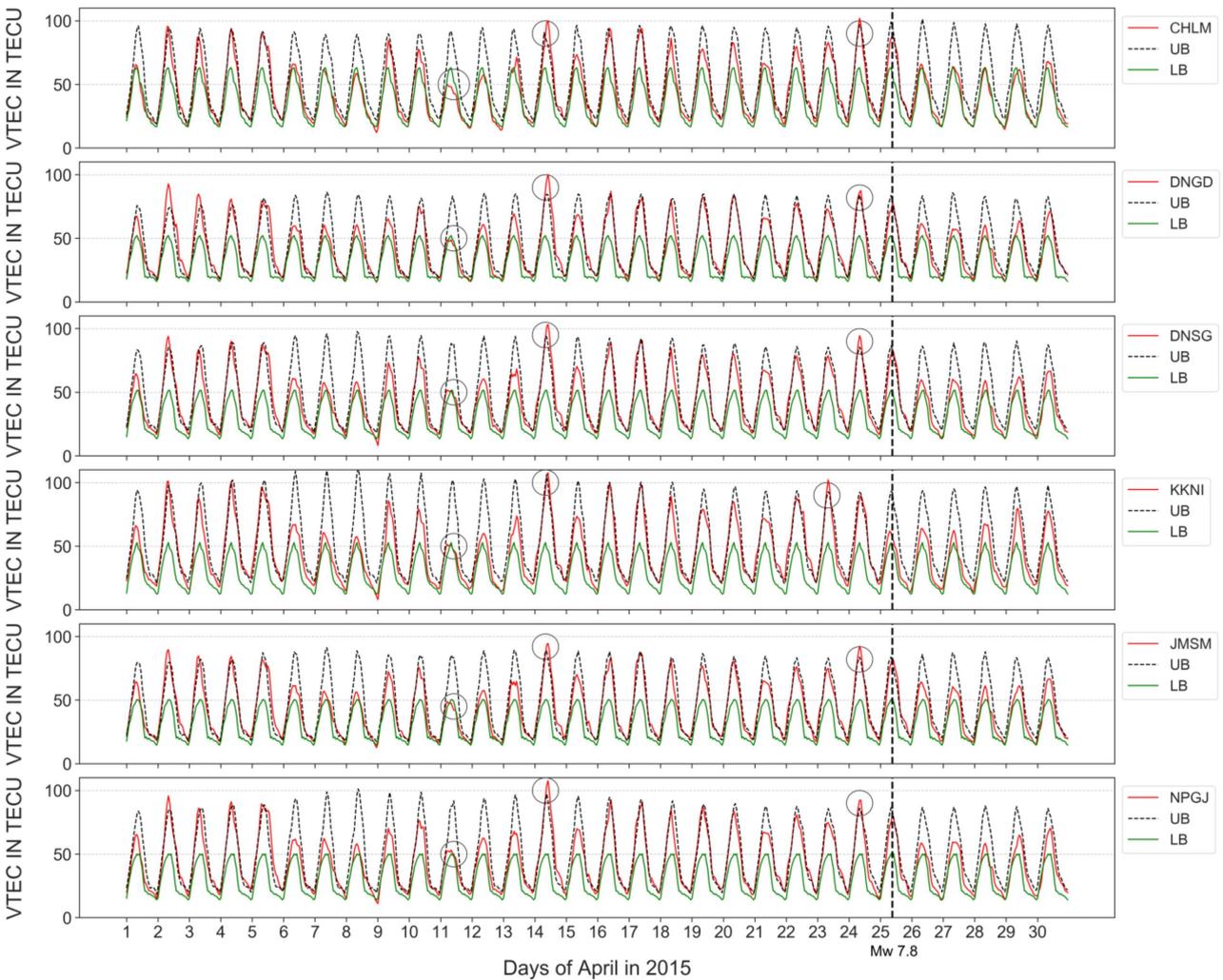


Figure2.

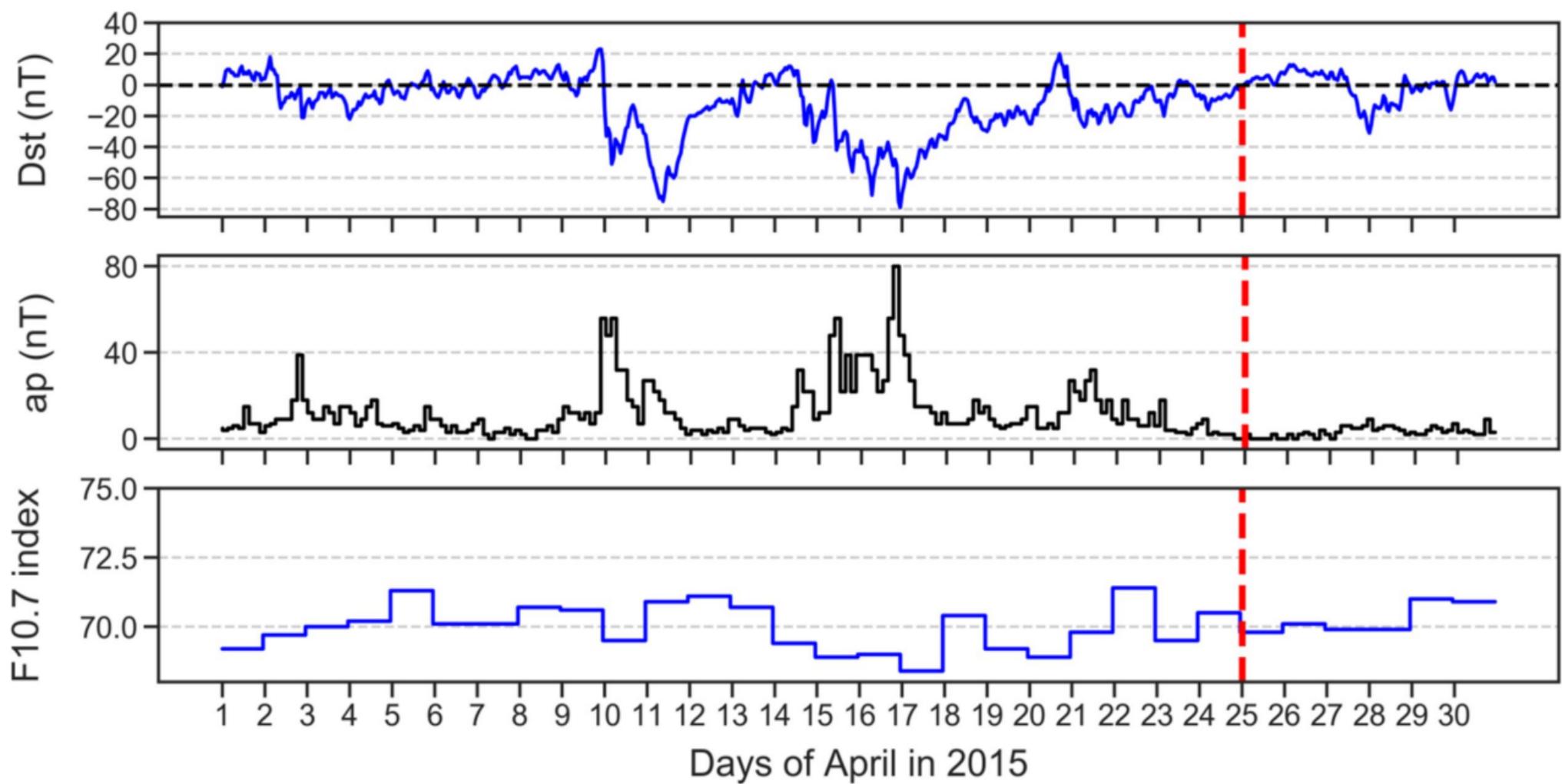


Figure3.

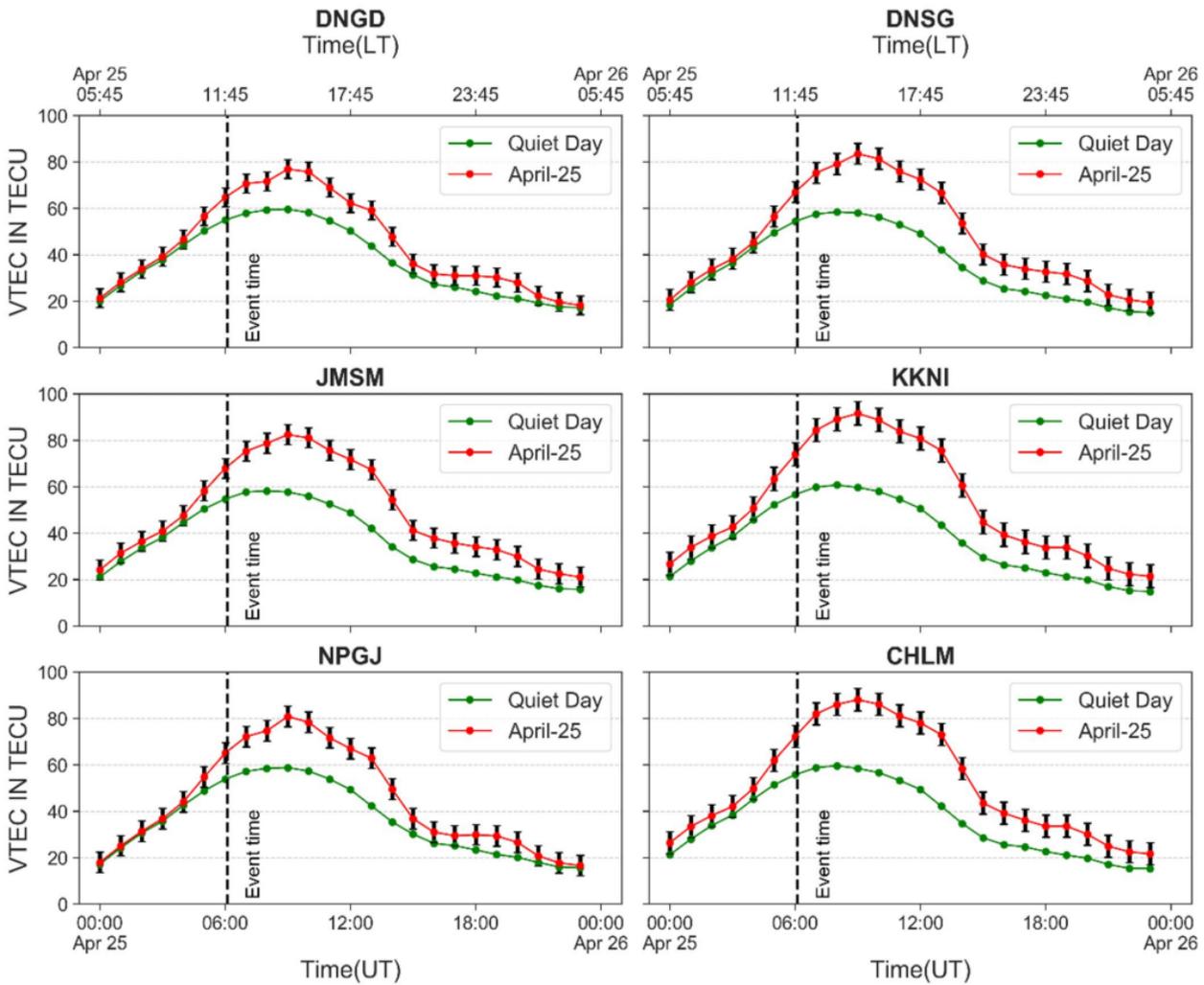


Figure4.

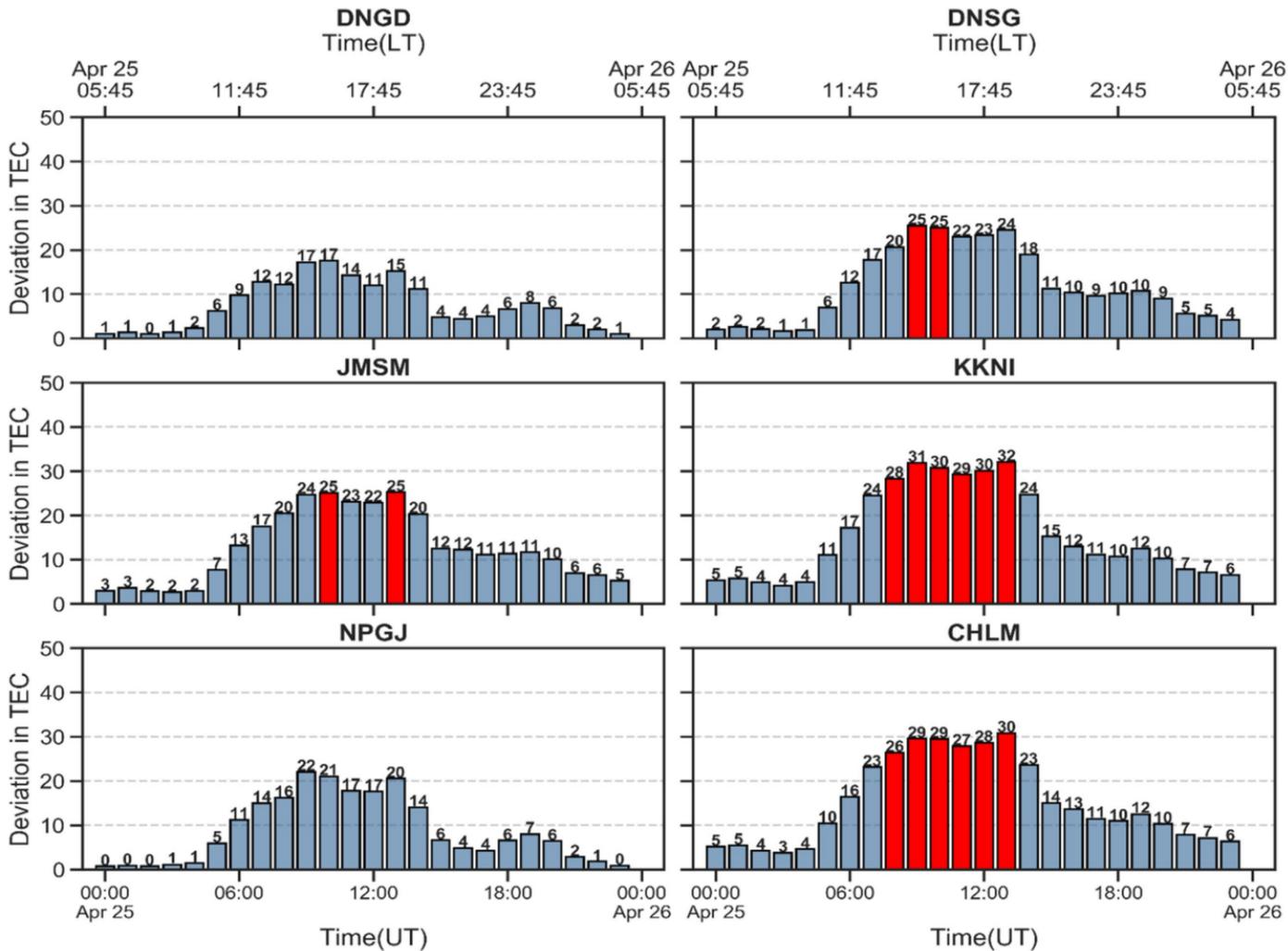
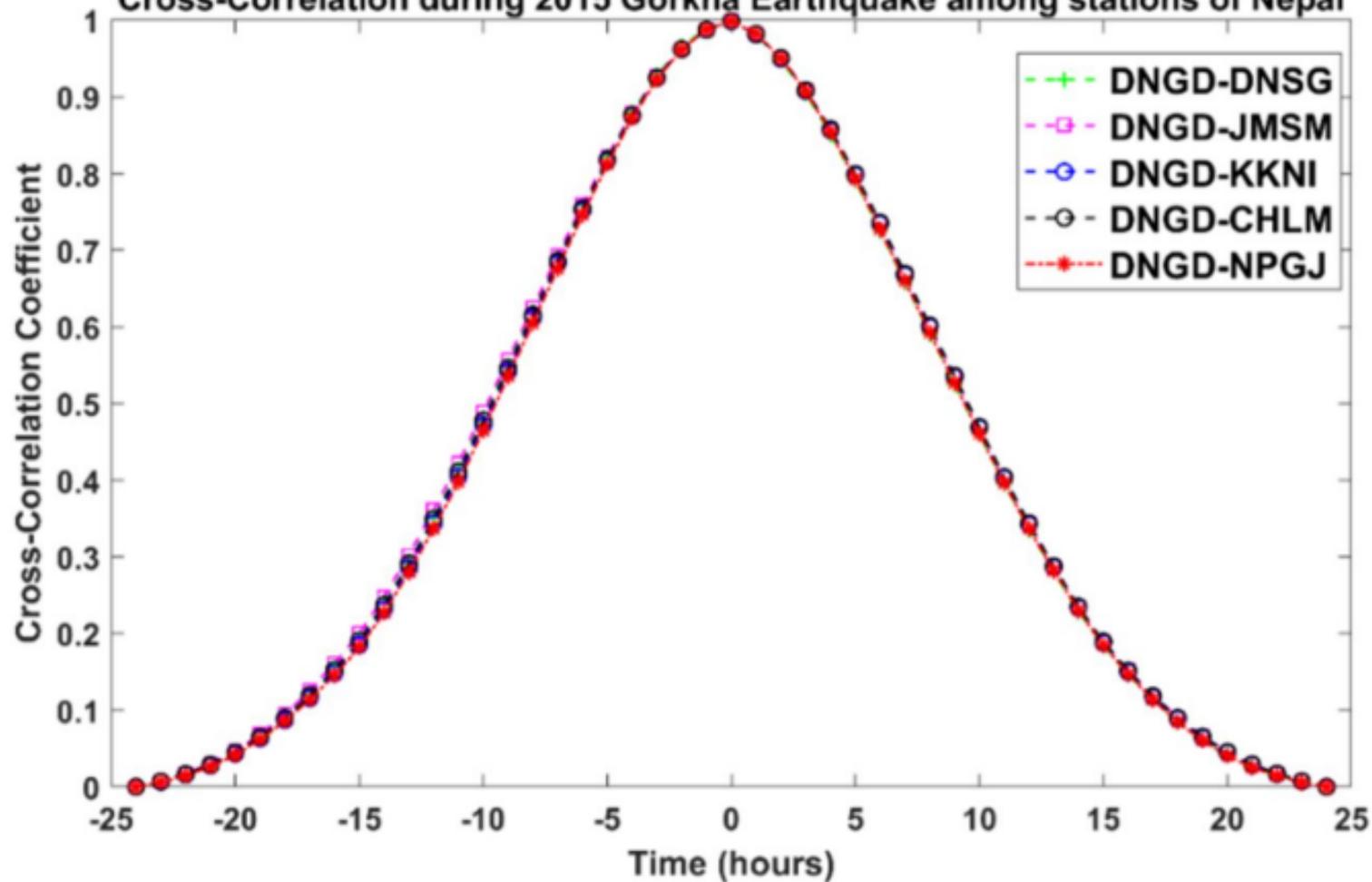


Figure5(a).

Cross-Correlation during 2015 Gorkha Earthquake among stations of Nepal



**Figure5(b).**



Figure5(c).

Cross-Correlation during 2015 Gorkha Earthquake

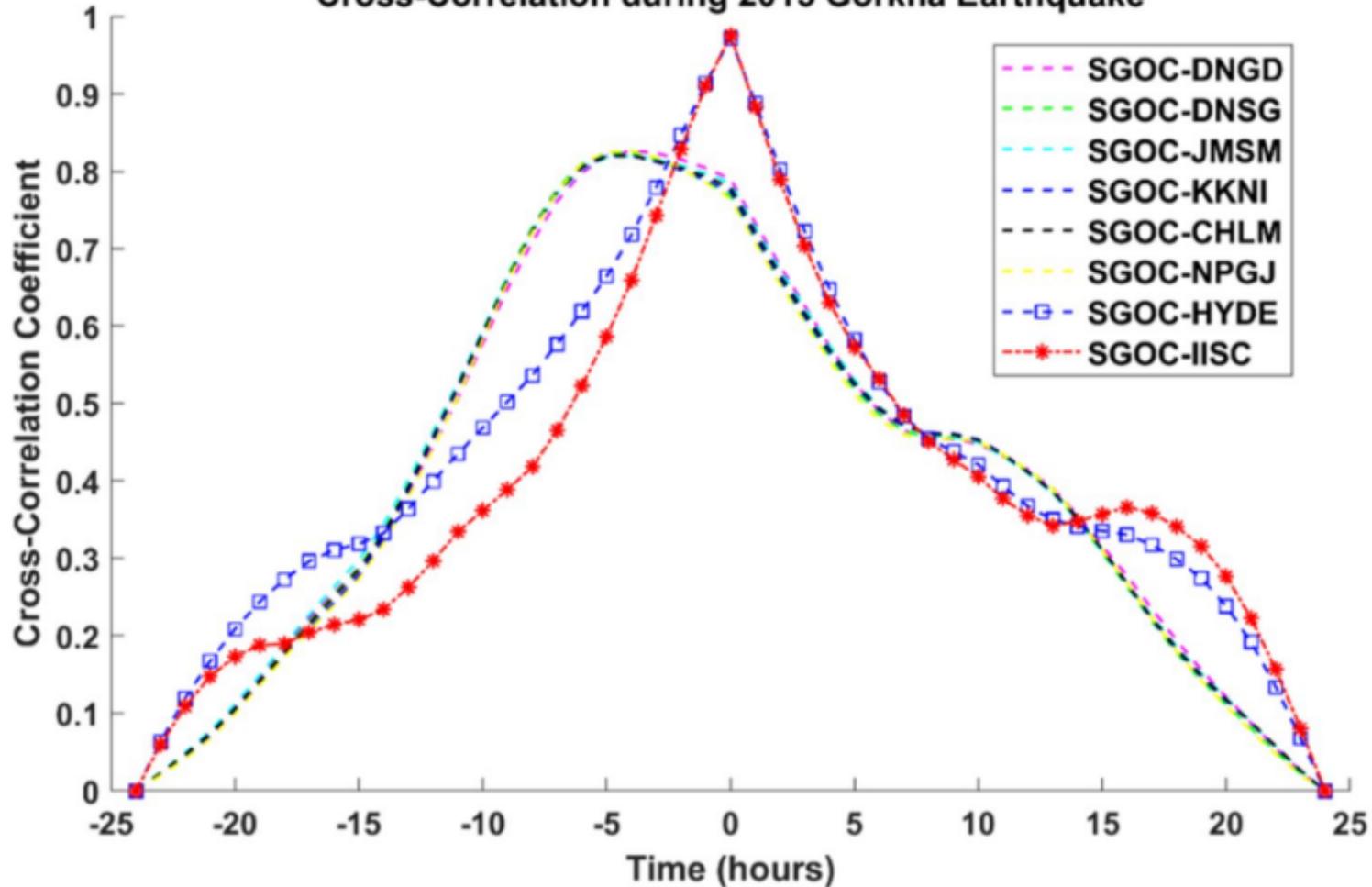


Figure6.

# IONOSPHERE MAP : 2015/04/24 14:00:00 H=450km (TECU)

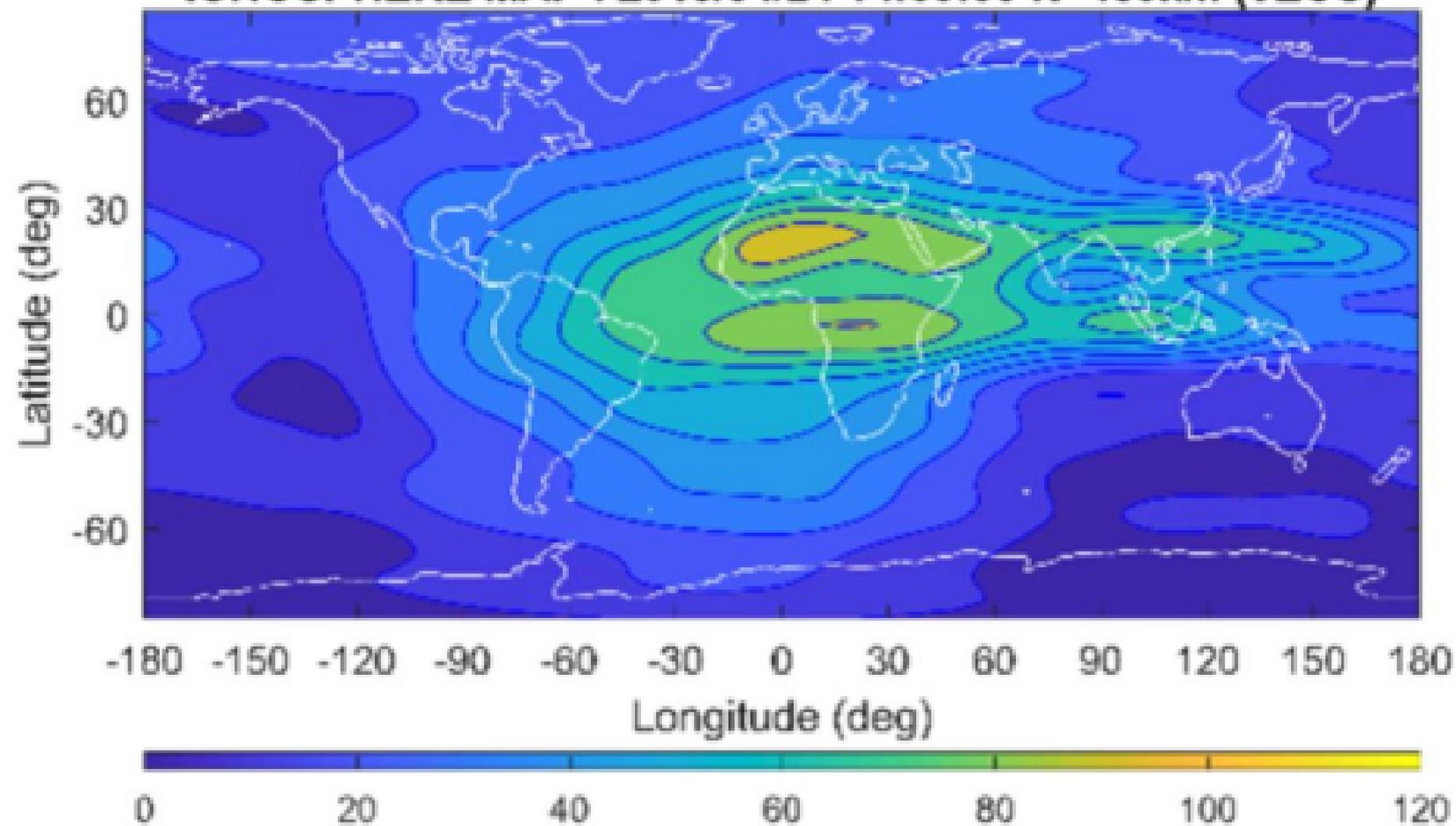


Figure 7.

# IONOSPHERE MAP : 2015/04/25 14:00:00 H=450km (TECU)

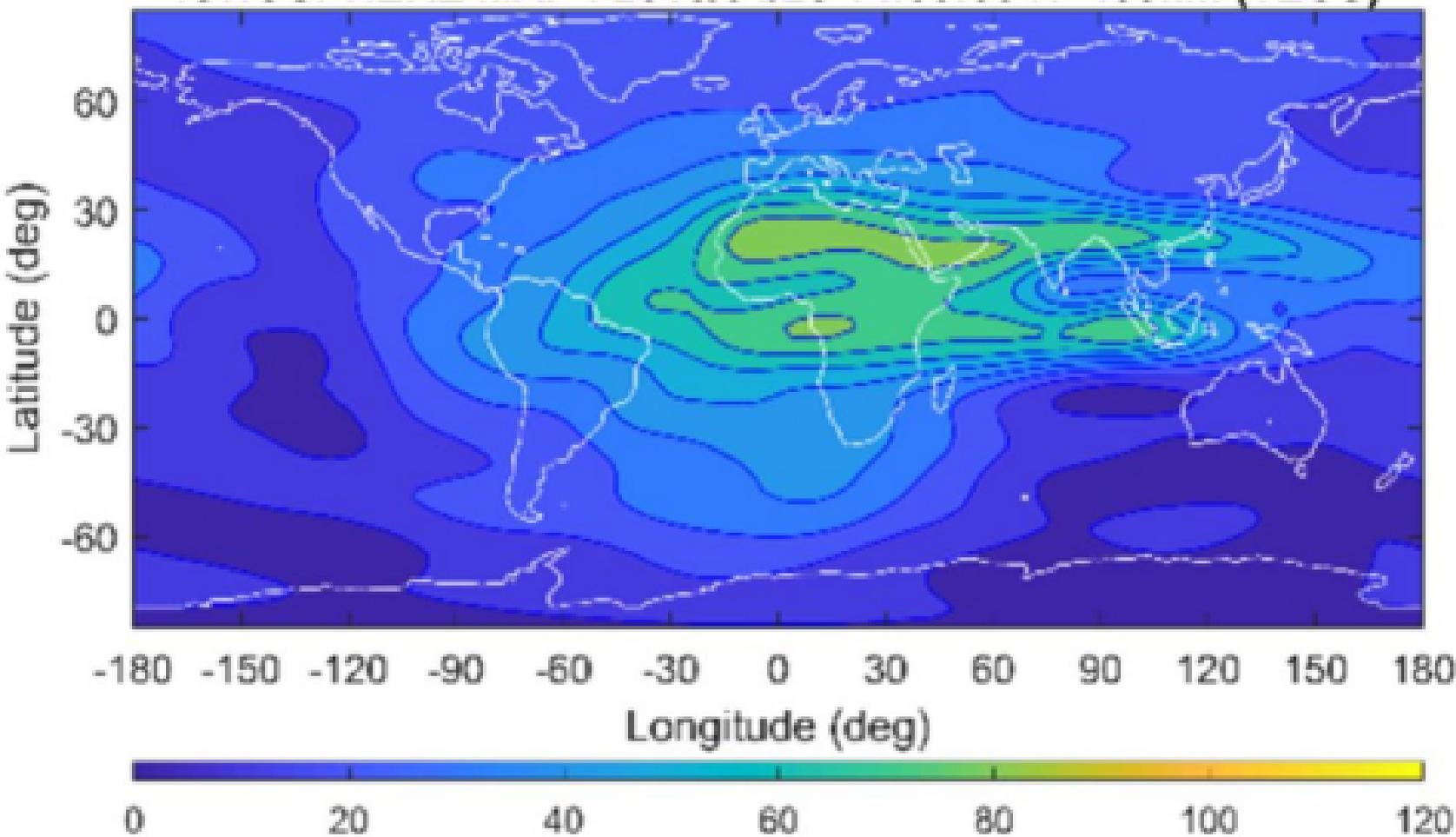


Figure 8.

# IONOSPHERE MAP : 2015/04/26 08:00:00 H=450km (TECU)

