# Energy-resolved detection of precipitating electrons of 30-100 keV by a sounding rocket associated with dayside chorus waves

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### Abstract

Whistler mode chorus waves scatter magnetospheric electrons and cause precipitation into the Earth's atmosphere. Previous measurements showed that nightside chorus waves are indeed responsible for diffuse/pulsating aurora. Although chorus waves and electron precipitation have also been detected on the dayside, their link has not been illustrated (or demonstrated) in detail compared to the nightside observations. Conventional low-altitude satellite observations do not well resolve the energy range of 10–100 keV, hampering verification on resonance condition with chorus waves. In this paper we report observations of energetic electrons with energies of 30–100 keV that were made by the electron sensor installed on the NASA's sounding rocket RockSat-XN. It was launched from the Andøya Space Center on the dayside (MLT  $\sim$  11 h) at the L-value of  $\sim$  7 on 13 January 2019. Transient electron precipitation was observed at  $\sim$  50 keV with the duration of <100 s. A ground station at Kola peninsula in Russia near the rocket's footprint observed intermittent emissions of whistler-mode waves simultaneously with the rocket observations. The energy of precipitating electrons is consistent with those derived from the quasi-linear theory of pitch angle scattering by chorus waves through cyclotron resonance, assuming a typical dayside magnetospheric electron density. Precise interaction region is discussed based on the obtained energy spectrum below 100 keV.

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## waves

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#### 13 **Key Points:**

14	•	The relation between energetic electron precipitation and chorus waves at the dayside
15		magnetosphere have not been identified in detail

Our sounding rocket experiment identified precipitating energetic electrons within 16 typical resonance energy with chorus waves on the dayside 17

• Ground-based and satellite observations of chorus waves support that the observed 18 electron precipitation was caused by chorus waves 19

## 20 Abstract

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- the Earth's atmosphere. Previous measurements showed that nightside chorus waves are
- 23 indeed responsible for diffuse/pulsating aurora. Although chorus waves and electron
- 24 precipitation have also been detected on the dayside, their link has not been illustrated (or
- 25 demonstrated) in detail compared to the nightside observations. Conventional low-altitude
- satellite observations do not well resolve the energy range of 10–100 keV, hampering
- 27 verification on resonance condition with chorus waves. In this paper we report observations
- of energetic electrons with energies of 30–100 keV that were made by the electron sensor installed on the NASA's sounding rocket RockSat-XN. It was launched from the Andøya
- Space Center on the dayside (MLT  $\sim 11$  h) at the L-value of  $\sim 7$  on 13 January 2019.
- Transient electron precipitation was observed at  $\sim 50$  keV with the duration of <100 s. A
- 32 ground station at Kola peninsula in Russia near the rocket's footprint observed intermittent
- emissions of whistler-mode waves simultaneously with the rocket observations. The energy
- of precipitating electrons is consistent with those derived from the quasi-linear theory of pitch
- angle scattering by chorus waves through cyclotron resonance, assuming a typical dayside
- 36 magnetospheric electron density. Precise interaction region is discussed based on the obtained
- are energy spectrum below 100 keV.
- 38

## 39 Plain Language Summary

The Earth's magnetosphere was filled with energetic electrons and various waves. Energetic electrons sometimes precipitate into the Earth's atmosphere and cause aurora. Whistler mode waves are believed to cause such precipitation and previous measurements showed that nightside chorus waves are responsible for aurora. Energetic electrons and chorus waves are

- 44 also observed on the dayside magnetosphere. However, their link has not been illustrated in
- detail compared to the nightside observations. In this paper, we verified the energy spectrum
- 46 of precipitating electrons on the dayside by installing the sensor which can resolve the 30–
- 100 keV energy range on a sounding rocket and observed transient electron precipitation.
- 48

## 49 **1 Introduction**

- 50 Precipitating electrons from the magnetosphere have been measured by balloons, sounding
- 51 rockets, and low-altitude satellites to investigate auroral zone phenomenology [e.g.,
- 52 Winningham et al., 1985; Fuller-Rowell and Evans, 1987; Miyoshi et al., 2010, 2015a],
- radiation belt dynamics [e.g. *Millan and Thorne*, 2007; *Miyoshi et al.*, 2008] and effects on
- the chemical composition of the upper atmosphere [e.g., *Thorne et al.*, 1977; *Lam et al.*,
- 55 2010; *Miyoshi et al.*, 2015b; *Turunen et al.*, 2016]. Hardy et al. [2008] conducted statistical
- studies based on the large database of energy spectra in the 50 eV–20 keV energy range
- accumulated by the Defense Meteorological Satellite Program (DMSP) and reported the
- distributions of electron precipitation depend on geomagnetic latitude (MLAT), magnetic
- 59 local time (MLT) and Kp index.
- 60 Millan and Thorne [2007] reviewed observations of higher energy (~10s keV to ~MeV)
- 61 electron precipitation as one of the radiation belt electron loss mechanisms. Quasi-periodic
- bursts of high-energy electron precipitation with short timescale (< 1s), called microbursts,
- 63 were detected in balloon experiments and the SAMPEX satellite observations [e.g. Anderson
- *and Milton*, 1964; *Parks*, 1978; *Blake et al.*, 1996]. Statistical surveys from NOAA Polar
- 65 Orbiting Environmental Satellites (POES) presented the global model of the energy-

- 66 integrated flux (> 30 keV) of precipitating electrons as a function of geomagnetic activity
- 67 [*Lam et al.*, 2010].
- 68 Whistler mode chorus waves are believed to play an important role in these electron
- 69 precipitations. Chorus waves are electromagnetic emissions often composed of discrete rising
- or falling elements [e.g. *Burtis and Helliwell*, 1969; *Sazhin and Hayakawa*, 1992; *Santlic et*
- *al.*, 2003]. The typical frequency of chorus waves is 0.1-0.8  $f_{ce}$ , where  $f_{ce}$  is the equatorial
- electron cyclotron frequency, and has a frequency gap at ~0.5  $f_{ce}$  [e.g. Burtis and Helliwell,
- 1969; Tsurutani and Smith, 1974; Sazhin and Hayakawa, 1992]. Chorus waves are mostly
- observed on the nightside during disturbed geomagnetic conditions and confined near the
- magnetic equator [e.g. *Tsurutani and Smith*, 1977; *Meredith et al.*, 2012]. *Meredith et al.*
- [2012] showed that lower band chorus  $(0.1-0.5 f_{ce})$  is confined to less than 15° in MLAT on
- the night side and upper band chorus (0.5–0.8  $f_{ce}$ ) is within about 6° in MLAT due to strong
- Landau damping. Chorus waves are responsible for energetic electron precipitation into the atmosphere by pitch angle scattering. On the nightside, recent evidence suggests that chorus
- waves are the dominant cause of electron pitch-angle scattering and resulting
- diffuse/pulsating aurora [e.g. *Ni et al.*, 2008, 2011; *Nishimura et al.*, 2010; *Thorne et al.*,
- 2010; *Miyoshi et al.*, 2015a; *Kasahara et al.*, 2018a; 2019; *Ozaki et al.*, 2019; *Hosokawa et*
- 83 *al.*, 2020; *Nishimura et al.*, 2020].
- 84 Dayside chorus waves are less dependent on the geomagnetic conditions and were observed
- 85 over wider range of magnetic latitude than nightside chorus waves [*Tsurutani and Smith*,
- <sup>86</sup> 1977; *Meredith et al.*, 2012]. *Meredith et al.* [2012] showed that lower band chorus waves
- 87 were observed up to 30° in MLAT on the dayside. These features derive from the distortion
- of the dayside magnetic field due to solar wind compression. The distortion forms two
- 89 minima of magnetic field strength (minimum B pockets) at high latitudes of both
- 90 hemispheres. Chorus waves can be generated at these regions [Vaivads et al., 2007; Tsurutani
- *et al.*, 2009]. The distortion also forms the region called Dayside Uniform Zone (DUZ),
- 92 where the gradient of the magnetic field strength along the field (dB/ds) is nearly zero within
- a large range of magnetic latitude. Such a field line exists in the transition region between the
- 94 dipole-like field near the Earth and the compressed field with minimum B pockets. Dayside
- chorus waves near the DUZ region were observed at the ground during quiet geomagnetic
   conditions [*Keika et al.*, 2012].
- 97 Despite such differences, dayside chorus waves are also expected to contribute to energetic
- 98 electron precipitation in probably the same physical mechanisms as on the nightside. Dayside
- diffuse aurora by low energy electron precipitation were reported [e.g. *Newell et al.*, 2009;
- *Han et al.*, 2015] and *Nishimura et al.* [2013] indicated that chorus waves are the dominant
- 101 cause of dayside diffuse aurora. At high latitude (L>6), <100 keV electrons are expected to
- 102 efficiently resonate with chorus waves at the magnetic equator and off-equatorial minimum-B
- pockets and precipitate into the atmosphere. While high energy electrons were observed on
- 104 the dayside [e.g. *Parks et al.*, 1978; *Lam et al.*, 2010], few previous observations have
- 105 identified <100 keV because of the lack of energy resolution. Previous researches indirectly
- obtained precipitated flux in the range of 30-100 keV by subtracting the integral flux of >100
- 107 keV from that of >30 keV [*Ni et al.*, 2014; *Li* et al., 2013, 2014a, 2014b]. However, different
- 108 channels have different efficiencies for incoming electrons and different proton
- 109 contaminations, which are difficult to correct due to a lack of reliable proton measurements
- 110 [Yando et al., 2011; Askainen and Mursula, 2013], making it difficult to obtain reliable
- electron energy spectra at 30-100 keV.
- 112 In this paper, we investigate the energy spectrum of precipitating electrons of 30-100 keV on
- 113 the dayside in order to discuss resonance condition. The electron sensor was installed on the

- 114 NASA's sounding rocket RockSat-XN, which was launched on the dayside (MLT ~ 11 h) at
- 115  $L \sim 7$  from the Andøya Space Center. The electron sensor detected precipitating electrons in
- the sunlit region, where little research has been done on the precipitating electrons of 30-100
- keV. Section 2 describes instrumentation and installation on the sounding rocket. Section 3
- provides solar wind, geomagnetic condition and magnetospheric configuration. In section 4, we discuss the resonance condition based on chorus wave frequency measured at a ground
- 119 we discuss the resonance condition based on chorus wave frequen 120 station near the rocket's footprint.

## 121 **2 Instrumentation**

- 122 Main results in this paper were obtained by Medium-energy Electron Detector (MED). MED
- 123 can measure the electron velocity distribution functions in the energy range of 30–100 keV.
- 124 MED is comprised of five pinholes and avalanche photodiodes (APD), covering nearly 2-pi
- steradian field of view. APDs provide higher signal-to-noise ratio than conventional solid-
- state detectors because of its internal gain and have advantage of high quantum efficiency for 2005, 2006, 2008, 2016, Krash was et al.
- a few tens of keV electrons [*Ogasawara et al.*, 2005, 2006, 2008, 2016; *Kasahara et al.*,
  2010, 2012]. We utilized APDs with the thickness of ~70 μm to properly measure the
- incoming energy for 30-100 keV electrons. To attenuate protons and photons, APDs were
- 129 incoming energy for 50-100 keV electrons. To attenuate protons and photons, Ai Ds were 130 covered by ~2 µm aluminum layer. Preamplifiers, shaping amplifiers, peak holders, and
- analog-to-digital converters are used to measure the incoming electron energy. Detected
- 132 signals are binned into logarithmically separated sixteen pulse height channels depending on
- the incoming electron energy. A geometric factor of MED was  $7.7 \times 10^{-5}$  cm<sup>2</sup> sr per APD.
- 134 Energy resolution was about 50% (<10 keV) at half maximum for 20 keV electrons. Because
- 135 higher energy electron is less dependent on attenuation by aluminum layer, energy resolution
- is higher for higher energy electron. For example, energy resolution was about 10% for 30
- 137 keV and about 2% for 100 keV in Monte Carlo simulation.
- 138 MED was installed on the sounding rocket in such a way that the angles between looking
- direction of five APDs and the rocket spin axis were 30°, 30°, 90°, 150° and 150°. Fluxgate
- 140 magnetometer was also mounted on the rocket. The RockSat-XN rocket was launched from
- Andøya Space Center in Norway at 0913 UT on 13 January 2019. The hemispherical field of
- view of MED covered most of the pitch angle range by the rocket spinning and coning,
- whose periods are  $\sim 1$  s and  $\sim 30$  s, respectively. We determined pitch angles by using the
- 144 Fluxgate magnetometer, which was also onboard the rocket.

# 145 **3 Observation**

# 146 **3.1. Solar Wind and Geomagnetic Condition**

- 147 Figure 1 shows solar wind data from the WIND spacecraft between 0300 and 1600 UT on 13
- 148January 2019. A black dashed line indicates the launch time. Figure 1 (a) shows that Ygsm
- component of the interplanetary magnetic field was negative and Zgsm component fluctuated
- 150 within  $\pm 3$  nT. Figure 1 (b) and (c) show that the solar wind velocity was 340 km/s and ion
- 151 density was 5 cm<sup>-3</sup>, and the dynamic pressure was ~1 nPa.



Figure 1. Time profile of the solar wind data from the WIND spacecraft between 0300 and 153 1600 UT on 13 January 2019. A black dashed line indicates the launch time of the RockSat-154 XN rocket. (a) Three components of the interplanetary magnetic field in geocentric solar 155

magnetospheric coordinates. (b) Solar wind velocity data of the SWE instrument. (c) Ion 156

density obtained from the MFI, SWE and 3DP instruments. 157

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Figure 2 presents real-time AE index on 13 January 2019. The AE index was steady below 159 300 nT through the day of flight, but with small increase up to  $\sim$ 200 nT around 0500 UT. The 160

3-hourly Kp index was in the range of 0 to 1 on 13 January 2019. These data suggest that the 161

period of flight was quiet in terms of solar wind and ground-based geomagnetic conditions. 162

163

#### 164 **3.2. Magnetospheric electrons**

Figure 3 provides energy-time spectrogram of the spin-averaged electron energy flux from 165

MEP-e [Kasahara et al., 2018a] onboard the ERG satellite [Miyoshi et al., 2018a]. ERG 166

located at (X, Y, Z)gsm = (-1, 6, 1)  $R_E$  and at the L-value of 6 around 0900 UT. The perigee 167

and the apogee are (X, Y, Z)gsm = (-1, 0, 0)  $R_E$  around 1200 UT and (X, Y, Z)gsm = (5, 2, 2) 168  $R_E$  around 0800 UT respectively. In this figure, it can be seen that ~30-40 keV electron

169 energy flux increased at ~0800 UT and lower energy electron flux also increased later. This 170

suggests that energetic electron injection into the magnetosphere occurred before 0800 UT. 171

This injection may be associated with the small increase of the AE index at ~0500UT. 172





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Figure 2. One-day plot of the real-time AE and AO index on 13 January 2019. Color 175 represents the number of stations to derive the index. Color scale is displayed on the right 176

side of panel. There is small substorm event at ~0500 UT. 177

<sup>[</sup>Created at 2019-04-30 15:14UT]



Figure 3. Energy-time spectrogram of the spin-averaged electron energy flux from MEP-e
onboard the ERG satellite on 13 January 2019. Enhancement of energetic electron flux was
observed between 0800 and 1100 UT. L is the McIlwain L-shell derived from IGRF.

182

### 183 **3.3. Rocket observations**

184 NASA's sounding rocket, RockSat-XN was launched at Andøya in Norway (69°17' N,

185 16°01' E) at 09:13 UT (10:13 LT) on 13 January 2019. Figure 4 shows that the altitude

186 profile of this flight and the horizontal trajectory in the geographical coordinate.

187 Figure 5 (a)–(e) present raw data energy spectra of pitch angles 0–90° (red) and 90–180°

(blue) observed by MED for every 50 s from 120 s after the launch, where count rate is

divided by the number of channels within  $0-85^{\circ}$  and  $95-180^{\circ}$  respectively. Pulse height

channels are essentially proportional to the incoming energy. We do not discuss pulse height

channels 0–2 because of the high noise level. As shown in Figures 5 (a), (b), (d), and (e), no significant difference between  $0-85^{\circ}$  and  $95-180^{\circ}$  spectra during most periods (120–220 s,

270-370 s after launch). This suggests that the detected signals were due to penetrating

particles, whose incoming directions were not limited by the pinholes. We estimated count

rates and energy deposit caused by penetrating galactic cosmic ray (GCR) proton. Count rates

of GCR detected by APDs (size is  $0.25 \text{ cm}^2$ ) is estimated to be ~ 30 counts / 50 s, assuming

197 typical GCR flux  $\sim 0.4 / \text{cm}^2$ -s-sr in the solar minimum at the high latitude [*Neher et al.*,

198 1956], consistent with our observation. Energy deposit is ~30 keV, assuming that GCR with

typical energy ( $\sim 0.1-10$  GeV) penetrating the 70 µm silicon APD, consistent with peak



Figure 4. The rocket trajectory. (a) The altitude profile of the RockSat-XN rocket. (b) The position in the geographical coordinate (red line). Five cross marks shows the position 200,

202 250, 300, 350 and 400 s after launch. Black dotted lines represent geomagnetic latitude.

- energy of energy spectra of pitch angle 0-85° and 95-180° (i.e., pulse height channel 7 203
- corresponds to about 30 keV). Therefore, we conclude that most counts during 120-220 s and 204
- 270–370 s after the launch are due to GCR. Although upward-looking APDs may have 205
- detected precipitating electrons as well, the precipitating electron flux level was too low to be 206 distinguished from GCR for most of the time.
- 207
- Nevertheless, during 220–270 s after the launch (Figure 5c), counts of pitch angle 0–85° are 208
- significantly higher than those of pitch angle 95–180° at pulse height channels 9–12. For this 209 time period, electrons of pitch angles 0-85° show a peak at the different energy from GCR's
- 210 energy deposit, indicating the transient significant electron precipitation. In order to confirm 211
- that two spectra are different, we conducted the Kolmogorov-Smirnov statistical test, which 212
- examines null hypothesis that energy spectra of pitch angles 0–85° and 95–180° come from 213
- the same population. The null hypothesis was rejected with significance level of  $\sim 0.4$  %. 214
- Furthermore, we assume that the spectrum of GCR was constant during the whole flight 215
- period and averaged the energy spectra of pitch angle 0-85° for 120-370 s, so that the error 216
- of the energy spectrum of GCR becomes less (Figure 5f, blue line). We performed the 217
- Kolmogorov-Smirnov test again for the averaged spectrum of GCR and that of pitch angle 0-218
- 85° during 220–370 s after launch as well and rejected the null hypothesis with significance 219
- level of ~0.4%. No significant difference was found between spectra of pitch angle 0-85° and 220
- GCR during the other periods. These data indicate that there was a transient electron 221
- precipitation within ~100 s. 222



Pulse Height Channels

- 227 error bars. Pulse Height Channels are essentially proportional to the detected energy. The
- count rate is calculated by dividing the count detected by the five detectors facing the pitch 228
- angles of 0-85° or 95-180° by the observation time. (f) Energy spectrum of pitch angle 0-229
- 85° during 220–270 s (red) and that of pitch angle 95–180° for 120-370 s (blue). 230

Figure 5. Energy spectra of pitch angles 0–85° (red) and 95-180° (blue) observed by MED 224 during (a) 120–170 s, (b) 170–220 s, (c) 220–170 s, (d) 270–320 s, (e) 320–370 s, after the 225 launch. Standard deviation assuming Poisson distribution counting statistics are shown by 226



Figure 6. Calibrated energy spectrum of precipitating electrons. Error bars indicate  $2\sigma$ , where or is the standard error.

234 We calibrated pulse height channel to energy and counts to differential flux based on the

235 geometrical design and laboratory tests. Note that here we obtain the precipitating electron 236 counts by subtracting GCR contamination (Figure 5f, blue line) from the counts in the pitch

angle  $0-85^{\circ}$  during 220-270 s (Figure 5f, red line). Thus, obtained energy spectrum is shown

in Figure 6, illustrating that the differential flux of precipitating electron is  $\sim 10^2$  /cm<sup>2</sup>-sr-s-

keV at 50 keV. The differential electron flux is comparable to that derived from an electron

240 density profile measured with the Tromsø EISCAT VHF radar at the same time, assuming

that an increase of electron density is mainly due to electron precipitation.

242

## 243 4 Discussion

244 We investigated the precipitating electrons on the dayside by using the energetic electron

245 detector which can observe energy spectrum of 30–100 keV electrons, and found

246 precipitating electrons peaking at ~50 keV. Here we discuss whether this electron

247 precipitation can be explained by the pitch angle scattering by chorus waves.

Figure 7 provides a frequency-time spectrogram, showing the magnetic field emissions

obtained by the VLF receiver at Lovozero, Kola peninsula in Russia (L~5.4, LT~11 h) near

the rocket trajectory around the time of electron precipitation. The difference in longitude

was less than 15°. Although this receiver observation is not strictly conjugated with the

252 RockSat-XN measurements, it is helpful to infer whistler-mode wave occurrence in the flux

tube on which RockSat-XN was located. The footprint of Lovozero was traced to the

magnetic equator along field line using the Tsyganenko T89 model [*Tsyganenko* 1988] to
 evaluate the condition of cyclotron resonance between electrons and VLF emissions. The

- evaluate the condition of cyclotron resonance between electrons and VLF emissions. The obtained magnetic field strength is ~220 nT near the magnetic equator, corresponding to the
- cyclotron frequency  $f_{ce}$  of 6.2 kHz. Waves were seen at ~0.2 0.8  $f_{ce}$  with a band gap at ~

258  $0.5 f_{ce}$ , typical of chorus waves.



Figure 7. (a) Power spectrogram of magnetic field obtained by the VLF receiver at Lovozero 261 262 in Russia between 0900 and 1000 UT on 13 January 2019. Vertical dotted lines show 220 s and 270s after the launch of RockSat-XN. The horizontal dotted line shows 3.1 kHz, which is 263 0.5  $f_{ce}$  near the magnetic equator of the field line extending from Lovozero. (b) The rocket 264 trajectory and positions of the VLF receiver at Lovozero, THEMIS-E footprint and POES 265 footprint in geographical coordinate (black dotted lines). Red dotted lines show geomagnetic 266 latitude (25°, 35°, 45°, 55°, 65°, 75°). Two purple points shows POES footprints at 0910 and 267 0920 UT. The thick purple line shows POES NOAA-18 footprint between 0910 and 0914 268 UT, when electron precipitation events were seen from the data of the Medium Energy 269 Proton and Electron Detector (MEPED) on board the POES satellite. 270

271

Chorus waves were also observed in the magnetosphere. Figure 8 shows the energy-time 272 spectrogram of electrons and frequency-time spectrograms of the wave electric field and 273 wave magnetic field, observed by THEMIS-E satellite [Angelopoulos, 2008]. THEMIS-E was 274 located near the dawnside magnetopause as can be seen from the intermittent excursions to 275 the magnetosheath in Figures 8a and b. Figures 8c and d show electromagnetic waves at 276 0.1 – 0.6  $f_{ce}$  in the magnetosphere, where  $f_{ce}$  is ~360 Hz at the minimum magnetic field strength along the magnetic flux tube on which THEMIS-E was located, especially at 0918, 277 278 0937 and 0954 UT. To combine, ground-based and magnetospheric observations suggest that 279 280VLF emissions were generated broadly in the dayside magnetosphere during this period.



Figure 8. Electron energy-time spectrograms of (a) high energy (from the SST instrument) and (b) low energy (the ESA instrument) components, and frequency-time spectrograms of (c) electric field (from EFI) and (d) magnetic field (SCM), obtained by the THEMIS-E spacecraft during 0900 to 1000 UT on 13 January 2019. THEMIS-E located at (X, Y, Z)sm =  $(4, -12, 3)R_E$  around 0900 UT.

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Assuming that observed precipitating  $\sim 50$  keV electrons was scattered at magnetic equator 288 by lower band chorus waves with  $0.1 - 0.5 f_{ce}$ , we can estimate a plasma density by using 289 general resonance condition and the standard cold plasma dispersion relation [e.g. Summers 290 et al., 1998]. We use the magnetic field strength ~100 nT at the magnetic equator of the 291 rocket's footprint. The calculated plasma density was 4 cm<sup>-3</sup> and 0.4 cm<sup>-3</sup> assuming chorus 292 waves with ~0.1  $f_{ce}$  and 0.5  $f_{ce}$ , respectively. To compare, the density observed by the 293 THEMIS-E ESA, which was located at higher L than the rocket trajectory, was  $1 - 2 \text{ cm}^{-3}$ . In 294 addition, the typical plasma density is  $5.6\pm1.7$  cm<sup>-3</sup> at L~6.9 and LT~10 h according to the 295 statistical study by Sheeley et al. [2001]. The above calculated plasma density 4 cm<sup>-3</sup> is 296 consistent with these values, thus 0.1  $f_{ce}$  chorus waves is favored for equatorial scattering. 297





Figure 9. The trapped E > 30 keV electron flux (blue lines) and the precipitating E > 30 keV electron flux (red lines) measured by MEPED onboard the POES NOAA-18 during 0900 to 1000 UT on 13 January 2019. Vertical dotted lines show 0910 and 0914 UT.

The above scenario, the equatorial resonance with 0.1  $f_{ce}$  chorus waves, reasonably explains 303 the precipitation of electrons of ~50 keV. However, it may contradict with the observed 304 energy spectrum peaking at ~ 50 keV. Considering that the flux of 30 keV is typically higher 305 than that of 50 keV in the quiet magnetosphere, and lower band chorus waves with >0.2  $f_{ce}$ 306 can interact with ~30 keV electrons near the magnetic equator for the same condition, the 307 precipitation flux peak at  $\sim 50 \text{ keV}$  is not expected. While the apparent peak at  $\sim 50 \text{ keV}$  is 308 due to the insufficient statistics (the flux difference was within  $2\sigma$  between 30 and 50 keV), 309 here we briefly discuss other possible explanations. One possibility is that the density 310 significantly deviates from the typical value. When the density is 0.1 cm<sup>-3</sup>, lower band chorus 311 waves with 0.1 - 0.5  $f_{ce}$  can resonate with 50 - 700 keV electrons and not with <50 keV 312 electrons. However, such a low density is inconsistent with the observation by THEMIS-E 313 and hence we think it is unlikely. Another possible case is that resonance occurred only at the 314 high magnetic latitude, where the magnetic field is stronger than the magnetic equator. 315 Suppose the chorus waves of  $< 0.5 f_{ce}$  were generated at the equator in the slightly higher L-316 shell region (e.g., the magnetic field intensity ~ 70 nT) and obliquely propagated to the 317 magnetic field line of the rocket trajectory, up to the latitude of  $25^{\circ}$  (the magnetic field 318 intensity ~ 180 nT), the 30 keV electrons becomes out of resonance any more [e.g., Miyoshi 319 et al., 2015b]. In fact, significant wave power of chorus waves with less frequency than a 320 typical band gap ~0.5  $f_{ce}$  are frequently observed at magnetic latitudes >15° on the dayside 321 [Meredith et al., 2014]. Therefore, the off-equatorial scattering by obliquely propagating 322 323 chorus waves may be a more plausible scenario than the equatorial scattering, at least for this case, in contrast to the statistical view on the nightside [Kasahara et al., 2019]. 324

Figure 9 shows the trapped and precipitating electron fluxes (>30 keV) measured by MEPED onboard the POES NOAA-18 [*Evans and Greer*, 2000]. The precipitating electron flux was high during 0910 to 0914 UT, when the trajectory of the POES satellite is shown in the thick purple line of Figure 7b. This indicates that the precipitation extended to higher L-shell region, consistent with the chorus emission observed by THEMIS-E. The precipitation was intermittent, which is also consistent with our observation.

- Figure 9 shows that the >30 keV integrated flux of precipitating electrons was  $\sim 10^{4-5}$  /cm<sup>2</sup>-sr-
- 332 s. Furthermore, statistical surveys on the >30 keV electron precipitation showed that the
- integrated flux of precipitating electrons is typically  $\sim 10^4$  /cm<sup>2</sup>-sr-s at L > 6 on the dayside
- under quiet geomagnetic condition (AE<100 nT) [Lam et al., 2010]. The precipitation event
- observed by MED showed that integrated flux of precipitating electron was  $\sim 10^4$  /cm<sup>2</sup>-sr-s in 336 30–100 keV, which is lower than observed by POES. This indicates that electron precipitation
- was intense at higher latitude where POES was located during 0910 to 0914 UT supporting the scenario that chorus waves obliquely propagate to inside and scatter the electrons at the
- 339 magnetic field line of the rocket trajectory.
- The duration of the precipitating event was <100 s at RockSat-XN. One of possible causes of this short period precipitation is the temporal variation of chorus waves. Chorus waves
- 342 propagated to Lovozero which have quasi-periodic variations with a timescale  $\sim 50$  s as can
- be seen in Figure 7. The intermittent variations of chorus waves may have corresponded to the transient electron precipitation. Nevertheless, quasi-periodicity  $\sim 50$  s was not found in
- the transient electron precipitation. Nevertheless, quasi-periodicity  $\sim 50$  s was not found in electron precipitation at RockSat-XN. We infer that chorus waves may have ceased
- temporarily and locally at the rocket's magnetospheric footprint.

## 347 **5 Conclusions**

- 348 There have been several observations of chorus waves and precipitating electrons on the
- dayside, while their relation has not been verified quantitatively due to the insufficient energy
- resolution of conventional low-altitude satellites in the range of 10s-100s keV. Using the
- 351 medium-energy electron detector sensitive to 30–100 keV onboard the sounding rocket
- 352 RockSat-XN, we confirmed the electron precipitation in the range of typical resonance
- energy with chorus waves on the dayside during the geomagnetically quiet periods. Chorus
- 354 waves were observed by the ground-based VLF receiver in Lovozero near the launch site and
- 355 THEMIS-E spacecraft in the dayside magnetosphere around the period of the precipitation
- event. The electron energy is consistent with the resonance energy derived from quasi-linear
- theory of magnetospheric electron scattering by chorus waves through cyclotron resonance.
- The detailed energy spectrum below 100 keV enabled the discussion suggesting the possible
- 359 location of wave-particle interaction.

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- this paper was provided by the WDC for Geomagnetism, Kyoto (<u>http://wdc.kugi.kyoto-</u> , was in(wdg/Sag2 html). Sciences date of the EBC (Amag) establish were obtained from the
- 369 <u>u.ac.jp/wdc/Sec3.html</u>). Science data of the ERG (Arase) satellite were obtained from the
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- 382 <u>2sec/2019/noaa18\_poes-sem2\_fluxes-2sec\_20190113\_v01.cdf</u>).
- 383

## 385 **References**

- Anderson, K. A., & Milton, D. W. (1964). Balloon Observations of X Rays in the Auroral 386 Zone 3. Journal of Geophysical Research: Space Physics, 69(21). 387 https://doi.org/10.1029/JZ069i021p04457 388 Angelopoulos, V. (2008). The THEMIS mission. Space Science Reviews (Vol. 141). 389 https://doi.org/10.1007/s11214-008-9336-1 390 391 Angelopoulos, V., Cruce, P., Drozdov, A., Grimes, E. W., Hatzigeorgiu, N., King, D. A., et al. (2019). The Space Physics Environment Data Analysis System (SPEDAS). Space 392 Science Reviews (Vol. 215). The Author(s). https://doi.org/10.1007/s11214-018-0576-4 393 394 Asikainen, T., & Mursula, K. (2011). Recalibration of the long-term NOAA/MEPED energetic proton measurements. Journal of Atmospheric and Solar-Terrestrial Physics, 395 73(2-3), 335-347. https://doi.org/10.1016/j.jastp.2009.12.011 396 Asikainen, T., & Mursula, K. (2013). Correcting the NOAA/MEPED energetic electron 397 fluxes for detector efficiency and proton contamination. Journal of Geophysical 398 Research: Space Physics, 118(10), 6500–6510. https://doi.org/10.1002/jgra.50584 399 Blake, J. B., Looper, M. D., Baker, D. N., Nakamura, R., Klecker, B., & Hovestadt, D. 400 (1996). New high temporal and spatial resolution measurements by SAMPEX of the 401
- 402 precipitation of relativistic electrons. *Advances in Space Research*, *18*(8), 171–186.
  403 <u>https://doi.org/10.1016/0273-1177(95)00969-8</u>
  404 Burtis, W. J., & Helliwell, R. A. (1969). Banded Chorus. a New Type of Vlf Radiation
- Burtis, W. J., & Helliwell, R. A. (1969). Banded Chorus. a New Type of Vlf Radiation
  Observed in the Magnetosphere By Ogo 1 and Ogo 3.*Journal of Geophysical Research:*,
  74(11), 3002–3010. <u>https://doi.org/10.1029/ja074i011p03002</u>
- Evans, D. S., & Greer, M. S. (2000). Polar Orbiting Environmental Satellite Space
   Environment Monitor 2 : Instrument Descriptions and Archive Data Documentation,
   Tech. Memo. OAR SEC-93, NOAA, Boulder, Colo.
- Evans, D. S., & Moore, T. E. (1979). Precipitating Electrons Associated With the Diffuse
  Aurora: Evidence for Electrons of Atmospheric Origin in the Plasma Sheet. *Journal of Geophysical Research*, 84(A11), 6451–6457. https://doi.org/10.1029/JA084iA11p06451
- Fuller-Rowell, T. J., & Evans, D. S. (1987). Height-integrated Pedersen and Hall conductivity
  patterns inferred from the TIROS-NOAA satellite data. *Journal of Geophysical Research*, 92(A7), 7606. <u>https://doi.org/10.1029/ja092ia07p07606</u>
- Han, D., X.-C. Chen, J.-J. Liu, Q. Qiu, K. Keika, Z.-J. Hu, J.-M. Liu, H.-Q. Hu, and H.-G.
- 417 Yang (2015), An extensive survey of dayside diffuse aurora based on optical
- 418 observations at Yellow River Station, *Journal of Geophysical Research:Space Physics*,
- 419 *120*, 7447-7465, <u>https://doi.org/10.1002/2015JA021699</u>

- Hardy, A., Holeman, E. G., Burke, W. J., Gentile, L. C., & Bounar, K. H. (2008). Probability
   distributions of electron precipitation at high magnetic latitudes. *Journal of Geophysical Research: Space Physics*, *113*(6), 1–19. <u>https://doi.org/10.1029/2007JA012746</u>
- Hosokawa, K., Miyoshi, Y., Ozaki, M., Oyama, S. I., Ogawa, Y., Kurita, S., et al. (2020).
  Multiple time-scale beats in aurora: precise orchestration via magnetospheric chorus
  waves. *Scientific Reports*, *10*(1), 3380. https://doi.org/10.1038/s41598-020-59642-8
- Kasahara, S., Takashima, T., & Hirahara, M. (2012). Variability of the minimum detectable
  energy of an APD as an electron detector. *Nuclear Instruments and Methods in Physics Research, Section A: Accelerators, Spectrometers, Detectors and Associated Equipment,*664(1), 282–288. https://doi.org/10.1016/j.nima.2011.11.033
- Kasahara, S., Yokota, S., Mitani, T., Asamura, K., Hirahara, M., Shibano, Y., & Takashima,
  T. (2018). Medium-energy particle experiments—electron analyzer (MEP-e) for the
  exploration of energization and radiation in geospace (ERG) mission. *Earth, Planets and Space*, 70(1), 1–16. https://doi.org/10.1186/s40623 018 0847 z
- Kasahara, S., Miyoshi, Y., Yokota, S., Mitani, T., Kasahara, Y., Matsuda, S., et al. (2018).
  Pulsating aurora from electron scattering by chorus waves. *Nature*, 554(7692), 337–340.
  https://doi.org/10.1038/nature25505
- Kasahara, S., Miyoshi, Y., Kurita, S., Yokota, S., Keika, K., Hori, T., et al (2019). Strong
  diffusion of energetic electrons by equatorial chorus waves in the midnight to dawn
  sector. *Geophysical Research Letters*, 46. https://doi.org/10.1029/2019GL085499
- Kasahara, S, Takashima, T., Asamura, K., & Mitani, T. (2010). Development of an APD with
  large area and thick depletion layer for energetic electron measurements in space. *IEEE Transactions on Nuclear Science*, *57*(3 PART 3), 1549–1555.
  https://doi.org/10.1109/TNS.2010.2047752
- Keika, K., Spasojevic, M., Li, W., Bortnik, J., Miyoshi, Y., & Angelopoulos, V. (2012).
  PENGU In/AGO and THEMIS conjugate observations of whistler mode chorus waves
  in the dayside uniform zone under steady solar wind and quiet geomagnetic conditions. *Journal of Geophysical Research: Space Physics*, 117(7), 1–15.
  https://doi.org/10.1029/2012JA017708
- Lam, M. M., Horne, R. B., Meredith, N. P., Glauert, S. A., Moffat-Griffin, T., & Green, J. C.
  (2010). Origin of energetic electron precipitation >30 keV into the atmosphere. *Journal*of Geophysical Research A: Space Physics, 115(A4), 1–15.
  https://doi.org/10.1029/2009JA014619
- Lampton, M. (1967). Daytime observations of energetic auroral-zone electrons. *Journal of Geophysical Research*, 72(23), 5817–5823. <u>https://doi.org/10.1029/jz072i023p05817</u>
- Li, W., Thorne, R. M., Bortnik, J., Reeves, G. D., Kletzing, C. A., Kurth, W. S., et al. (2013).
  An unusual enhancement of low-frequency plasmaspheric hiss in the outer plasmasphere
  associated with substorm-injected electrons. *Geophysical Research Letters*, 40(15),
  3798–3803. https://doi.org/10.1002/grl.50787
- Li, W., Ni, B., Thorne, R. M., Bortnik, J., Green, J. C., Kletzing, C. A., et al. (2013).
  Constructing the global distribution of chorus wave intensity using measurements of
  electrons by the POES satellites and waves by the Van Allen Probes. *Geophysical Research Letters*, 40(17), 4526–4532. https://doi.org/10.1002/grl.50920
- Li, W., B. Ni, R. M. Thorne, J. Bortnik, Y. Nishimura, J. C. Green, C. A. Kletzing, W. S.
  Kurth, G. B. Hospodarsky, H. E. Spence, G. D. Reeves, J. B. Blake, J. F. Fennell, S. G.
  Claudepierre, and X. Gu (2014), Quantifying hiss-driven energetic electron
  precipitation: A detailed conjunction event analysis, *Geophys. Res. Lett.*, 41, 1085-1092,
- 467 <u>https://doi.org/10.1002/2013GL059132</u>
- 468 Meredith, N. P., Horne, R. B., Sicard-Piet, A., Boscher, D., Yearby, K. H., Li, W., & Thorne,
- 469 R. M. (2012). Global model of lower band and upper band chorus from multiple satellite

- observations. Journal of Geophysical Research: Space Physics, 117(10), 1-14. 470 https://doi.org/10.1029/2012JA017978 471
- Meredith, N. P., R. B. Horne, W. Li, R. M. Thorne, and A. Sicard-Piet (2014). Global model 472 of low-frequency chorus ( $f_{LHR} < f < 0.1 f_{ce}$ ) from multiple satellite observations, 473 Geophysical Research Letters, 41, 280-286, http://doi:10.1002/2013GL059050.
- 474
- Meredith, N. P., Horne, R. B., Isles, J. D., & Green, J. C. (2016). Extreme energetic electron 475 fluxes in low Earth orbit: Analysis of POES e > 30, e > 100, and e > 300 keV electrons. 476 Space Weather, 14(2), 136–150. https://doi.org/10.1002/2015SW001348 477
- Millan, R. M., & Thorne, R. M. (2007). Review of radiation belt relativistic electron losses. 478 Journal of Atmospheric and Solar-Terrestrial Physics, 69(3), 362–377. 479 https://doi.org/10.1016/j.jastp.2006.06.019 480
- Miyoshi, Y., Sakaguchi, K., Shiokawa, K., Evans, D., Albert, J., Connors, M., & Jordanova, 481 V. (2008). Precipitation of radiation belt electrons by EMIC waves, observed from 482 ground and space. Geophysical Research Letters, 35(23), 1–5. 483 https://doi.org/10.1029/2008GL035727 484
- 485 Miyoshi, Y., Katoh, Y., Nishiyama, T., Sakanoi, T., Asamura, K., & Hirahara, M. (2010). Time of flight analysis of pulsating aurora electrons, considering wave-particle 486 interactions with propagating whistler mode waves. Journal of Geophysical Research: 487 Space Physics, 115(10), 1–7. https://doi.org/10.1029/2009JA015127 488
- Miyoshi, Y., Oyama, S., Saito, S., Kurita, S., Fujiwara, H., Kataoka, R., et al. (2015). 489 Energetic electron precipitation associated with pulsating aurora: EISCAT and Van 490 Allen Probe observations. Journal of Geophysical Research: Space Physics, 120(4), 491 2754-2766. https://doi.org/10.1002/2014JA020690 492
- Miyoshi, Y., Saito, S., Seki, K., Nishiyama, T., Kataoka, R., Asamura, K., et al. (2015). 493 494 Relation between fine structure of energy spectra for pulsating aurora electrons and frequency spectra of whistler mode chorus waves. Journal of Geophysical Research: 495 Space Physics, 120(9), 7728–7736. https://doi.org/10.1002/2015JA021562 496
- Miyoshi, Y., Shinohara, I., Takashima, T., Asamura, K., Higashio, N., Mitani, T., et al. 497 (2018). Geospace exploration project ERG. Earth, Planets and Space, 70(1). 498 https://doi.org/10.1186/s40623-018-0862-0 499
- Miyoshi, Y., Hori, T., Shoji, M., Teramoto, M., Chang, T. F., Segawa, T., et al. (2018). The 500 ERG Science Center. Earth, Planets and Space, 70(1). https://doi.org/10.1186/s40623-501 018-0867-8 502
- Newell, P. T., Sotirelis, T., & Wing, S. (2009). Diffuse, monoenergetic, and broadband 503 aurora: The global precipitation budget. Journal of Geophysical Research: Space 504 *Physics*, *114*(9), 1–20. https://doi.org/10.1029/2009JA014326 505
- Ni, B., Thorne, R. M., Shprits, Y. Y., & Bortnik, J. (2008). Resonant scattering of plasma 506 sheet electrons by whistler-mode chorus: Contribution to diffuse auroral precipitation. 507 Geophysical Research Letters, 35(11), 1-5. <u>https://doi.org/10.1029/2008GL034032</u> 508
- Ni, B., Thorne, R. M., Meredith, N. P., Horne, R. B., & Shprits, Y. Y. (2011). Resonant 509 scattering of plasma sheet electrons leading to diffuse auroral precipitation: 2. 510 Evaluation for whistler mode chorus waves. Journal of Geophysical Research: Space 511 Physics, 116(4), 1–17. https://doi.org/10.1029/2010JA016233 512
- Ni, B., W. Li, R. M. Thorne, J. Bortnik, J. C. Green, C. A. Kletzing, W. S. Kurth, G. B. 513 Hospodarsky, and M. de Soria-Santacruz Pich (2014), A novel technique to construct 514
- the global distribution of whistler mode chorus wave intensity using low-altitude POES 515 516 electron data, J. Geophys. Res. Space Physics, 119, 5685–5699.
- https://doi:10.1002/2014JA019935 517

518	Nishimura, Y., Bortnik, J., Li, W., Thorne, R. M., Lyons, L. R., Angelopoulos, V., et al.
519	(2010). Identifying the driver of pulsating aurora. <i>Science</i> , <i>330</i> (6000), 81–84.
520	https://doi.org/10.1126/science.1193186
521	Nishimura, Y., Bortnik, J., Li, W., Thorne, R. M., Ni, B., Lyons, L. R., et al. (2013).
522	Structures of dayside whistler-mode waves deduced from conjugate diffuse aurora.
523	Journal of Geophysical Research: Space Physics, 118(2), 664–673.
524	https://doi.org/10.1029/2012JA018242
525	Nishimura, Y., Lessard, M. R., Katoh, Y., Miyoshi, Y., Grono, E., Partamies, N., et al.
526	(2020). Diffuse and Pulsating Aurora. Space Science Reviews, 216(1), 1–38.
527	https://doi.org/10.1007/s11214-019-0629-3
528	Ogasawara, K., Asamura, K., Mukai, T., & Saito, Y. (2005). Avalanche photodiode for
529	measurement of low-energy electrons. Nuclear Instruments and Methods in Physics
530	Research, Section A: Accelerators, Spectrometers, Detectors and Associated Equipment,
531	545(3), 744–752. https://doi.org/10.1016/j.nima.2005.02.026
532	Ogasawara, K., Hirahara, M., Miyake, W., Kasahara, S., Takashima, T., Asamura, K., et al.
533	(2008). High-resolution detection of 100 keV electrons using avalanche photodiodes.
534	Nuclear Instruments and Methods in Physics Research, Section A: Accelerators,
535	Spectrometers, Detectors and Associated Equipment, 594(1), 50–55.
536	https://doi.org/10.1016/j.nima.2008.05.056
537	Ogasawara, K, Livi, S. A., Allegrini, F., Broiles, T. W., Dayeh, M. A., Desai, M. I., et al.
538	(2016). Journal of Geophysical Research : Space Physics Next-generation solid-state
539	detectors for charged Special Section : Journal of Geophysical Research: Space Physics,
540	1–17. https://doi.org/10.1002/2016JA022559.Abstract
541	Ogasawara, K., Asamura, K., Takashima, T., Saito, Y., & Mukai, T. (2006). Rocket
542	observation of energetic electrons in the low-altitude auroral ionosphere during the
543	DELTA campaign. Earth, Planets and Space, 58(9), 1155–1163.
544	https://doi.org/10.1186/BF03352005
545	Ozaki, M., Miyoshi, Y., Shiokawa, K., Hosokawa, K., Oyama, S. ichiro, Kataoka, R., et al.
546	(2019). Visualization of rapid electron precipitation via chorus element wave-particle
547	interactions. Nature Communications, 10(1). https://doi.org/10.1038/s41467-018-07996-
548	Z
549	Parks, G. K. (2013). Microburst Precipitating Phenomena. Journal of Chemical Information
550	and Modeling, 53(9), 1689–1699. https://doi.org/10.1017/CBO9781107415324.004
551	Santolík, O., Gurnett, D. A., Pickett, J. S., Parrot, M., & Cornilleau-Wehrlin, N. (2003).
552	Spatio-temporal structure of storm-time chorus. Journal of Geophysical Research:
553	Space Physics, 108(A7), 1–14. <u>https://doi.org/10.1029/2002JA009791</u>
554	Sazhin, S. S., & Hayakawa, M. (1992). Magnetospheric chorus emissions: A review.
555	Planetary and Space Science, 40(5), 681–697. <u>https://doi.org/10.1016/0032-</u>
556	<u>0633(92)90009-D</u>
557	Sheeley, B. W., Moldwin, M. B., Rassoul, H. K., & Anderson, R. R. (2001). An empirical
558	plasmasphere and trough density model: CRRES pobservations. Journal of Geophysical
559	Research: Space Physics, 106(2000). <u>https://doi.org/10.1029/2000JA000286</u>
560	Summers, D., Thorne, R. M., & Xiao, F. (1998). Relativistic theory of wave-particle resonant
561	diffusion with application to electron acceleration in the magnetosphere. Journal of
562	Geophysical Research: Space Physics, 103(A9), 20487–20500.
563	https://doi.org/10.1029/98JA01740
564	Thorne, R. M., Ni, B., Tao, X., Horne, R. B., & Meredith, N. P. (2010). Scattering by chorus
565	waves as the dominant cause of diffuse auroral precipitation. Nature, 467(7318), 943-
566	946. https://doi.org/10.1038/nature09467

- Thorne, R. M. (1977). Energetic radiation belt electron precipitation: A natural depletion
   mechanism for stratospheric ozone. *Science*, *195*(4275), 287–289.
   <u>https://doi.org/10.1126/science.195.4275.287</u>
- Tsurutani, B. T., & Smith, E. J. (1974). Postmidnight chorus: A substorm phenomenon.
   *Journal of Geophysical Research*, 79(1), 118–127.
- 572 https://doi.org/10.1029/ja079i001p00118
- Tsurutani, B. T., & Smith, E. J. (1977). Two types of magnetospheric ELF chorus and their
   substorm dependences. *Journal of Geophysical Research*, 82(32), 5112–5128.
   <a href="https://doi.org/10.1029/ja082i032p05112">https://doi.org/10.1029/ja082i032p05112</a>
- Tsurutani, B. T., Verkhoglyadova, O. P., Lakhina, G. S., & Yagitani, S. (2009). Properties of
  dayside outer zone chorus during HILDCAA events: Loss of energetic electrons. *Journal of Geophysical Research: Space Physics*, *114*(3), 1–19.
  https://doi.org/10.1029/2008JA013353
- Tsyganenko, N. A. (1989). A magnetospheric magnetic field model with a warped tail current
   sheet. *Planetary and Space Science*, *37*(1), 5–20. <u>https://doi.org/10.1016/0032-</u>
   <u>0633(89)90066-4</u>
- Turunen, E., Kero, A., Verronen, P. T., Miyoshi, Y., Oyama, S. I., & Saito, S. (2016).
  Mesospheric ozone destruction by high-energy electron precipitation associated with
  pulsating aurora. *Journal of Geophysical Research*, *121*(19), 11852–11861.
  https://doi.org/10.1002/2016JD025015
- Vaivads, A., Santolík, O., Stenberg, G., André, M., Owen, C. J., Canu, P., & Dunlop, M.
   (2007). Source of whistler emissions at the dayside magnetopause. *Geophysical Research Letters*, 34(9), 1–5. https://doi.org/10.1029/2006GL029195
- World Data Center for Geomagnetism, Kyoto, M. Nose, T. Iyemori, M. Sugiura, T. Kamei
   (2015). Geomagnetic AE index. <u>https://doi:10.17593/15031-54800</u>
- Whittaker, I. C., Rodger, C. J., Clilverd, M. A., & Sauvaud, J.-A. (2014). The effects and
  correction of the geometric factor for the POES/MEPED electron flux instrument using
  a multisatellite comparison. *Journal of Geophysical Research: Space Physics*, *119*(8),
  6386–6404. <u>https://doi.org/10.1002/2014ja020021</u>
- Yando, K., Millan, R. M., Green, J. C., & Evans, D. S. (2011). A Monte Carlo simulation of
  the NOAA POES Medium Energy Proton and Electron Detector instrument. *Journal of Geophysical Research: Space Physics*, *116*(10), 1–13.
  https://doi.org/10.1029/2011JA016671
- 599 <u>https://doi.org/10.1029/2011JA01667</u> 600

Figure 1.



Figure 2.



[Created at 2019-04-30 15:14UT]

Figure 3.



Figure 4.



Figure 5.





0.05

0.05

Figure 6.



Figure 7.



Figure 8.



Figure 9.

