Fractional relaxation noises, motions and the fractional energy balance equation

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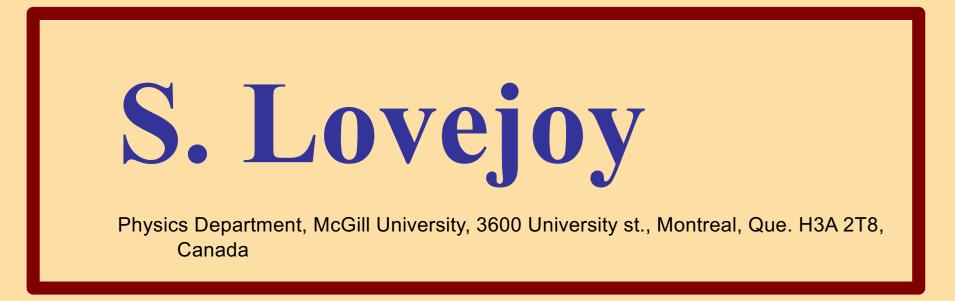
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Abstract

We consider the statistical properties of solutions of the stochastic fractional relaxation equation that has been proposed as a model for the earth's energy balance. In this equation, the (scaling) fractional derivative term models energy storage processes that occur over a wide range of space and time scales. Up until now, stochastic fractional relaxation processes have only been considered with Riemann-Liouville fractional derivatives in the context of random walk processes where it yields highly nonstationary behaviour. For our purposes we require the stationary processes that are the solutions of the Weyl fractional relaxation equations whose domain is -[?] to t rather than 0 to t. We develop a framework for handling fractional equations driven by white noise forcings. To avoid divergences, we follow the approach used in fractional Brownian motion (fBm). The resulting fractional relaxation motions (fRm) and fractional relaxation noises (fRn) generalize the more familiar fBm and fGn (fractional Gaussian noise). We analytically determine both the small and large scale limits and show extensive analytic and numerical results on the autocorrelation functions, Haar fluctuations and spectra. We display sample realizations. Finally, we discuss the prediction of fRn, fRm which – due to long memories - is a past value problem, not an initial value problem. We develop an analytic formula for the fRnforecast skill and compare it to fGn. Although the large scale limit is an (unpredictable) white noise that is attained a slow power law manner, when the temporal resolution of the series is small compared to the relaxation time, fRncan mimick a long memory process with a wide range of exponents ranging from fGn to fBm and beyond. We discuss the implications for monthly, seasonal, annual forecasts of the earth's temperature.

The Fractional and Half-Order Energy Balance Equations



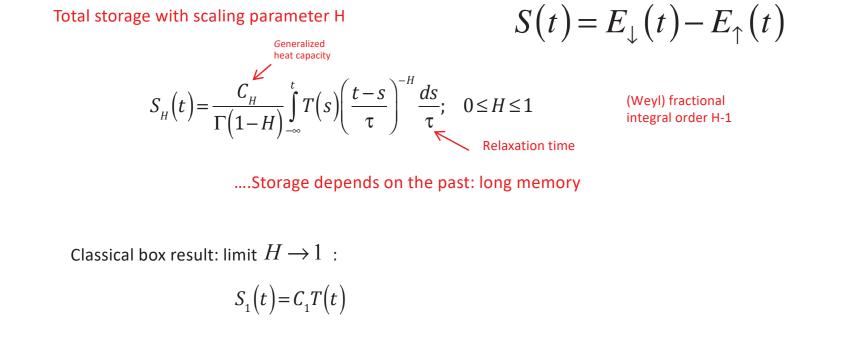
Fractional relaxation noises, motions

The Fractional order Energy Balance Equation (FEBE)

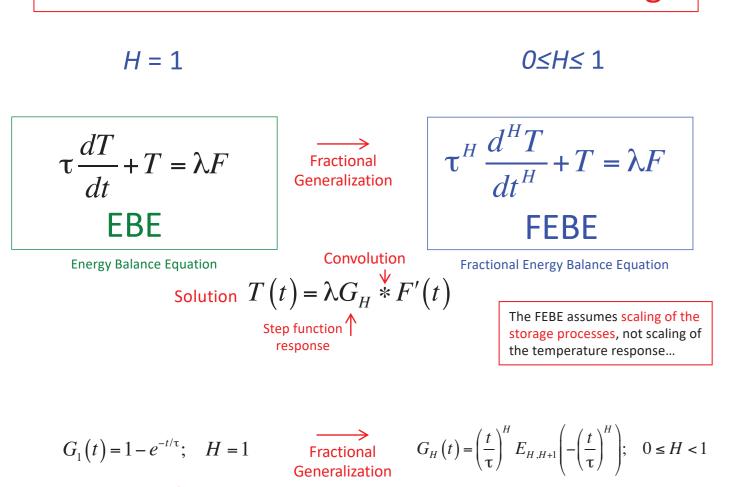
The (conventional) Energy Balance Equation (EBE) $S(t) + E(t) = E_{\downarrow}(t)$ Box energy Homogeneous Ocean Deep ocean ...more boxes? $C\frac{dT}{dt} + \lambda^{-1}T = F$ Box model: Newton's law of

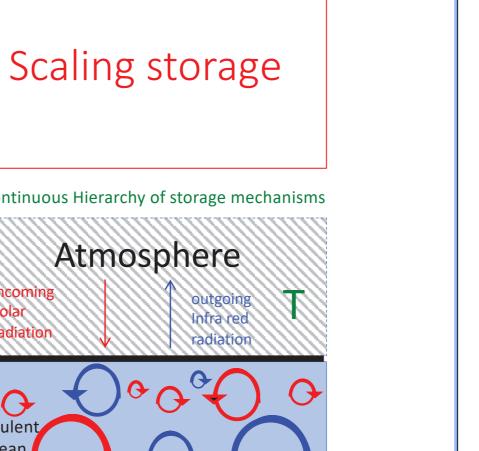
Scaling of the storage

Time constant of box: $\tau = C\lambda$



Fractional Generalization: realistic storage





Atmosphere

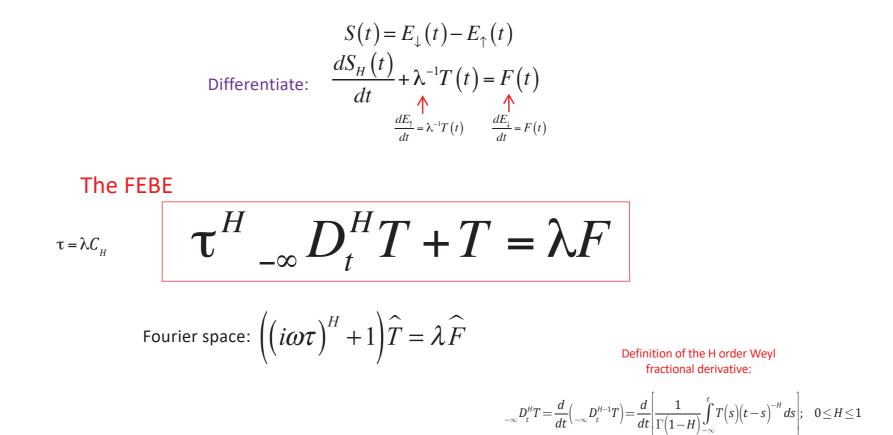
Energy rate equation

Box storage

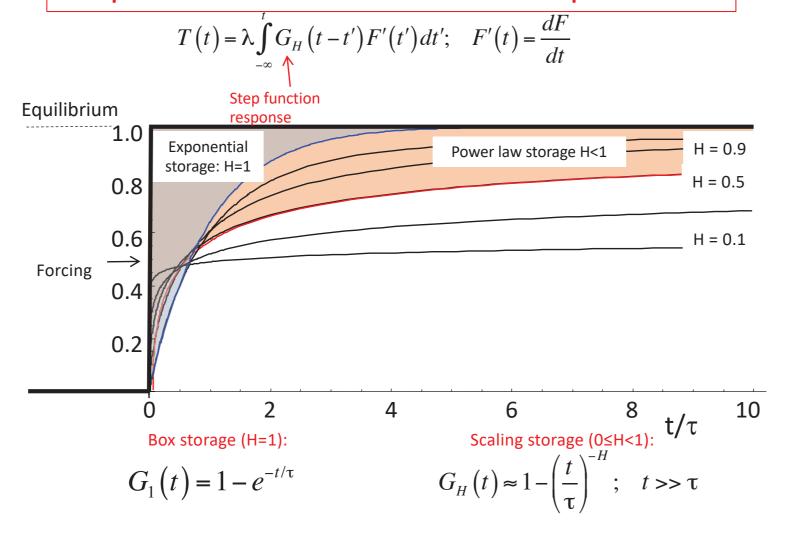
Atmosphere

Homogeneous Ocean

Deep ocean



Response: Generalized relaxation processes

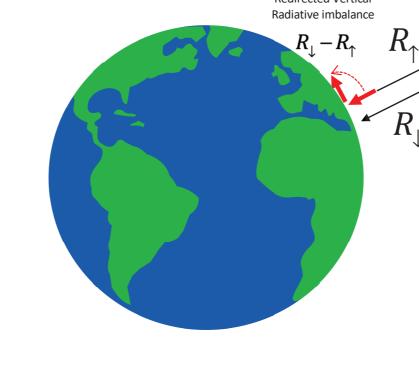


The Half-order Energy Balance Equation (HEBE) Fractional relaxation noises, motions

The conventional 1-D energy balance equation (EBE) has no vertical coordinate so that radiative imbalances between the earth and outer space are redirected in the horizontal in an ad hoc manner. We retain the basic EBE but add a vertical coordinate so that the imbalances drive the system by imposing heat fluxes through the surface. While this is theoretically correct, it leads to (apparently) difficult mixed boundary conditions. However using Babenko's method, we directly obtain simple equations for (2D) surface temperature anomalies $T_s(\underline{x},t)$: the Half-order Energy Balance Equation (HEBE) and the Generalized HEBE, (GHEBE). The HEBE anomaly equation only depends on the local climate sensitivities and relaxation times. We analytically solve the HEBE and GHEBE for $T_s(x,t)$ and provide evidence that the HEBE applies at scales >≈10km. We also calculate very long time diffusive transport dominated climate states as well as space-time statistics including the cross-correlation matrix needed for empirical orthogonal functions.

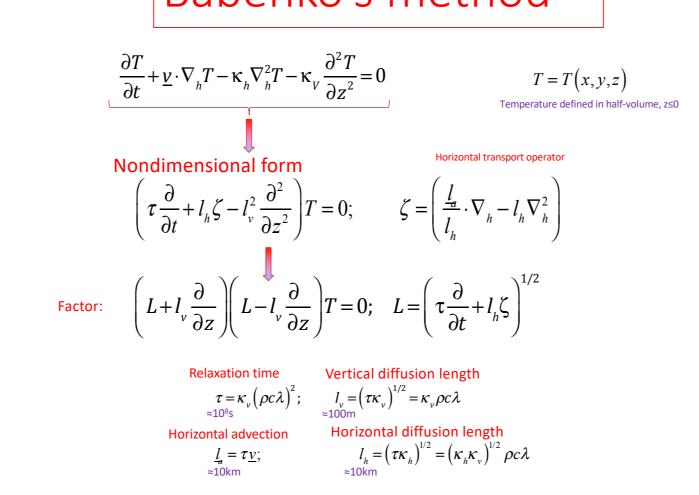
The HEBE is the special H = 1/2 case of the Fractional EBE (FEBE) and has a long (power law) memory up to its relaxation time t. Several papers have empirically estimated $H \approx 0.5$, as well as t ≈ 4 years, whereas the classical zerodimensional EBE has H = 1 and $t \approx 4$ years. The former values permit accurate macroweather forecasts and low uncertainty climate projections; this suggests that the HEBE could apply to time scales as short as a month. Future generalizations include albedo-temperature feedbacks and more realistic treatments of past and future climate states.

Regional energy balance Budyko-Sellers



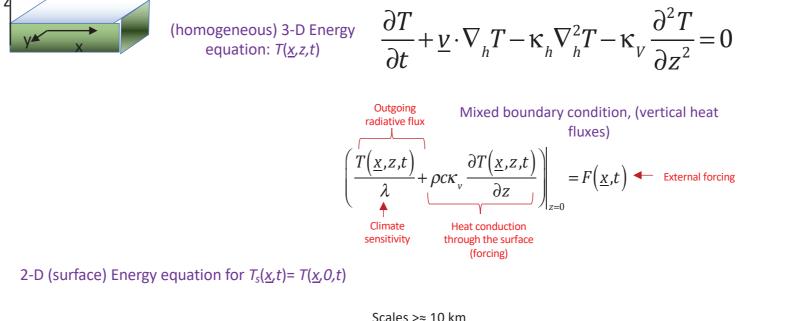
Babenko's method

 $\nabla \cdot Q = 0$



Correcting Budyko-Seller's boundary conditions:

H=1/2 - the Half-Order Energy Balance Equation (HEBE)



$\tau(\underline{x})^{1/2} D_t^{1/2} T_s + T_s = \lambda(\underline{x}) F(\underline{x}, t)$

introducing a vertical coordinate, we were able to rigorously treat the forcing of the entire system via the basic surface boundary condition: local vertical radiative flux imbalances that force heat into the earth. A consequence is that the (apparently) difficult "mixed" surface boundary conditions are avoided and the latitudinal and longitudinal (2D) variation of the time independent climatological temperatures can be determined. More importantly, we were able to obtain an equation for the time varying anomalies. In comparison, the usual (e.g. Budyko-Sellers) type energy balance models have no vertical coordinate and instead redirect the local imbalances in an ad hoc fashion, away Since the forcing is via the vertical boundary condition, the equation remains homogeneous and Babenko's method can be applied. Babenko's method elegantly transforms the mixed boundary conditions directly into a simple equation for the surface temperature; the Generalized Half-Order Energy Balance Equation (GHEBE). The detailed vertical structure turns out to only be important over a thin surface layer. A key novelty of the GHEBE is that instead of using classical first order time derivatives (the Energy Balance Equation, EBE) it is of half order in both temporal and horizontal operators. However, for spatial scales larger than ten kilometers or so, and time scales less than centennial, the temporal (relaxation) processes dominate over the advective and/or diffusive horizontal transport processes, this yields the Half-order EBE or HEBE. Under fairly mild assumptions we obtained a full analytic solution to the HEBE surface temperature anomalies $T_s(x,t)$. The EBE and HEBE are the H = 1, H = 1/2 special cases of the Fractional EBE (FEBE) that was recently introduced as a phenomenological macroweather model [Lovejoy, 2019a], with empirical estimates $H \approx 0.4$ - 0.5, i.e. very close to the HEBE. A feature of the HEBE is that its Green's function is a power law rather than an exponential and this implies a long memory: indeed the Gaussian white noise driven HEBE has a high frequency 1/f spectrum that is cutoff at the relaxation time (empirically of the order ≈ 4 years). By extending energy balance models to 2-D, it allows us to treat regional temporal anomalies, and this at significantly shorter time scales than were previously possible, perhaps down to the ≈ 10 day weather-macroweather transition scale. Depending on the space-time statistics of the anomaly forcing, the HEBE justifies the current Fractional EBE (FEBE) based macroweather (monthly, seasonal) temperature forecasts [Lovejoy et al., 2015], [Del Rio Amador and Lovejoy, 2019]. At this transition scale, GCMs are beyond the predictability limits of their atmospheric components, they become stochastic. Analyses of 32 CMIP5 GCMs showed that although each GCM had a distinct climate, that each responded nearly linearly to the climate forcing scenarios considered in the IPCC AR5 [Hébert and Lovejoy, 2018]. This implies that macroweather dynamics are plausibly linear, consistent with the HEBE. The regional HEBE – when stochastically forced - is thus a promising macroweather temperature model. Indeed, the high frequency part of its FEBE generalization can already be used for monthly, seasonal forecasts and the overall FEBE with $H \approx 1/2$ can produce climate projections with significantly lower uncertainties than current GCM based alternatives (work in In addition to the anomalies, our approach opens the door to the determination of the full 2-D climate state generalizations of the 1-D Budyko-Sellers type climates – to determine past and future climates. This could done first by applying the method to the existing climate by fixing the forcing at current values and solving the time independent transport equations. Then, the long term effect of changes - such as step function increases in forcing

References - could be determined from the GHEBE anomaly equation (section 3.5) which regionally corrects the local climate sensitivities for (slow) horizontal energy transport effects. Milankovitch (orbital) forcings are linear and can easily be introduced, and generalizations to account for albedo feedbacks and other nonlinear effects could easily be made in order to study glacial cycles. The power law relaxation processes implied by the GHEBE suggests straightforward explanations for the observed power law climate regime spanning the range from centennial to Milankovitch scales.

and the FEBE

Weather,

Macroweather,

and Climate

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fRn, fRm

he stochastic FEBE, fractional Relaxation motion (fRm), fractional Relaxation noise The physical problem that we wish to solve is the noise driven FEBE with noise amplitude σ and sensitivity λ :

 $\tau^{H}_{-\infty}D_{t}^{H}T_{i}+T_{i}=\lambda f(t); \quad f(t)=\sigma\gamma(t)$ with initial conditions $T_i(-\infty) = 0$, $0 \le H \le 1$ and $\gamma(t)$ a unit Gaussian white noise so that $\langle f^2 \rangle^{1/2} = \sigma$; $\langle f \rangle = 0$, σ is the amplitude of the noise. To understand the statistical properties of the internal variability (the response), it suffices to study the nondimensional equation: $_{-\infty}D_t^H U_H + U_H = K\gamma(t)$ where K is the normalization constant in eq. 35 and the nondimensional $U_H(t)$ function is called the fractional Relaxation noise (fRn) since it generalizes fGn. Using U_H we can obtain the solution to the dimensional eq. 46:

Since the fRn process is the solution of the fractional relaxation equation with a stationary, Gaussian, zero mean, white noise forcing, it is also stationary, Gaussian with zero mean. Its statistics are therefore fully characterized by its autocorrelation function. We can now calculate the correlation function relevant for the fRn statistics. The main complication is that in the small t limit, the fractional term dominates so that we obtain the (fGn) limit. The solution of eq. 47 is therefore - like $\gamma(t)$ - a generalized function: to obtain solutions with finite variances, we must take averages over finite resolutions τ_w . The resulting τ_w resolution autocorrelation function at lag Δt is: $R \left(\Delta t\right) = \langle T\left(t\right)T\left(t-\Delta t\right)\rangle = \int G_{0H}\left(\Delta t + t'\right)G_{0H}\left(t'\right)dt'; \quad t > \tau_{w}$

[Lovejoy, 2019a], where we used the normalization K given above. From this, we can obtain the high frequency fGn approximation (valid for $\Delta t << 1$ corresponding to $t << \tau$ in the dimensional

$R_{H,\tau_{w}}(\Delta t) \approx \frac{1}{2}\tau_{w}^{2H-1}\left((\lambda+1)^{2H+1}+(\lambda-1)^{2H+1}-2\lambda^{2H+1}\right); \quad \lambda = \Delta t/\tau_{w}; \quad \tau_{w} \leq \Delta t <<1; \quad 0 < H \leq \frac{1}{2}$ $R_{H,\tau_{w}}(\Delta t) \approx H(2H+1)\Delta t^{2H-1}; \quad \tau_{w} << \Delta t << 1; \quad 0 < H < \frac{1}{2}$

At low frequencies, for $\Delta t > 1$ (corresponding to $\Delta t > \tau$ in the dimensional equation), we obtain: $R_{II}(\Delta t) = -\frac{\Lambda}{(\Delta t)^{-1-H}} \Delta t^{-1-H} + O(\Delta t^{-1-2H})$: 0 < H < 2; $\Delta t >> 1$ Note that this result is independent of the resolution τ_w and that it holds over a wider range of H values. A technical point is that although the fGn has a huge memory (the slow Δt^{2H-1} fall-off in $R(\Delta t)$), the corresponding fRn is effectively truncated with a faster long time behaviour that

We could also mention the total energy storage $S_H(t)$. From eq. ? the nondimensional rate of storage $dS_H/dt = K\gamma-U$. The storage thus involves the integral of the fRn. The integral of fRn is "fractional Relaxation motion (fRm)", the name, in analogy with the integral of fGn which is called "fractional Brownian motion (fBm)" (see [Lovejoy, 2019a]). For long times, we have:

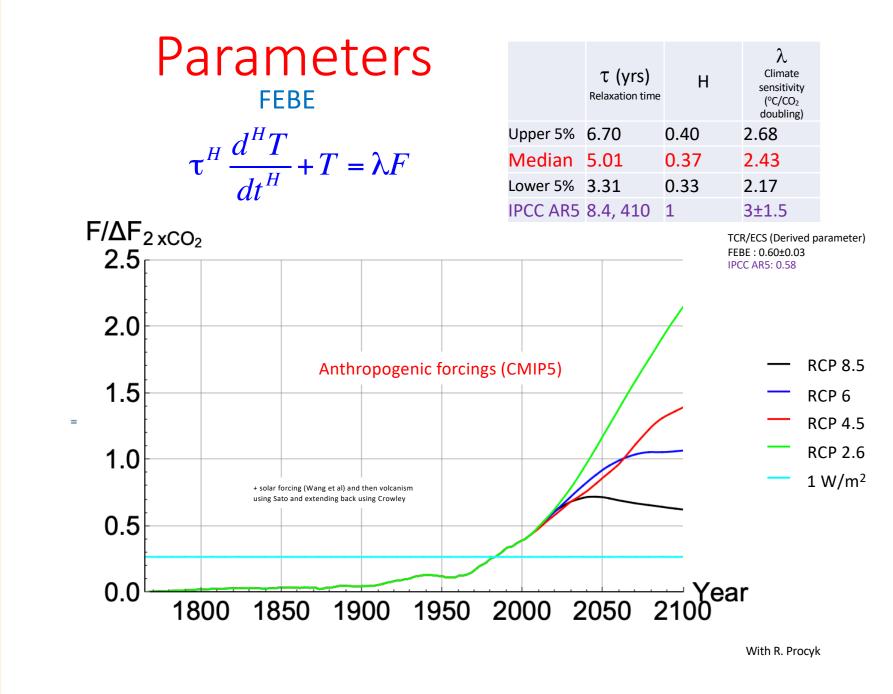
 $\left\langle \Delta S_{H} \left(\Delta t \right)^{2} \right\rangle = \frac{1}{\left(1 - 2H \right) \Gamma \left(1 - H \right)^{2}} \Delta t^{1 - 2H}; \quad \Delta S_{H} \left(\Delta t \right) = S_{H} \left(t \right) - S_{H} \left(t - \Delta t \right); \quad \Delta t >> 1; \quad 0 < H < 1 / 2$ which is a fractional Brownian motion (of order ½-H) in comparison the nondimensional temperature integral (proportional to the total energy emitted to outer space) increases as Δt for large Δt corresponding to a usual Brownian motion. A final interesting property is the predictability of fRn. fGn has an enormous memory (it is called a "long range memory process") and since for short times, fRn is close to fGn the two can

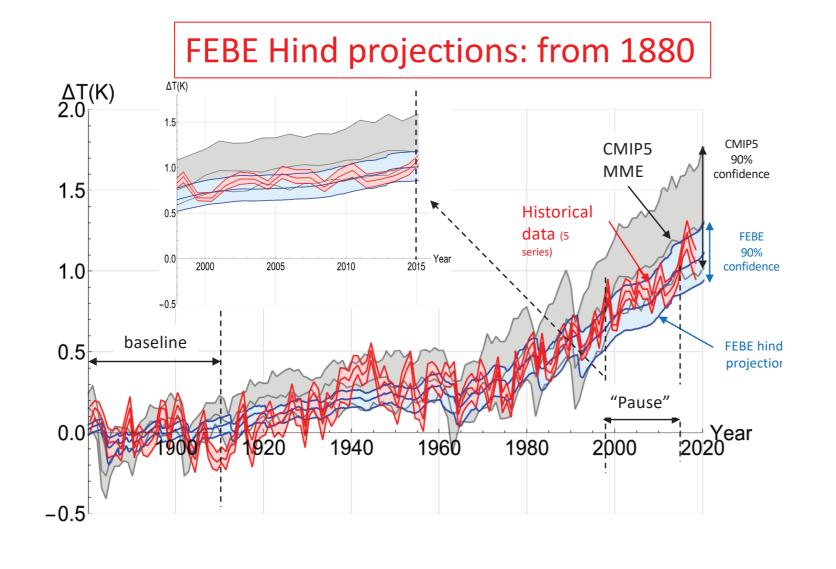
both be well predicted (for fGn, the skill –with infinite past data- becomes perfect in the limit H -

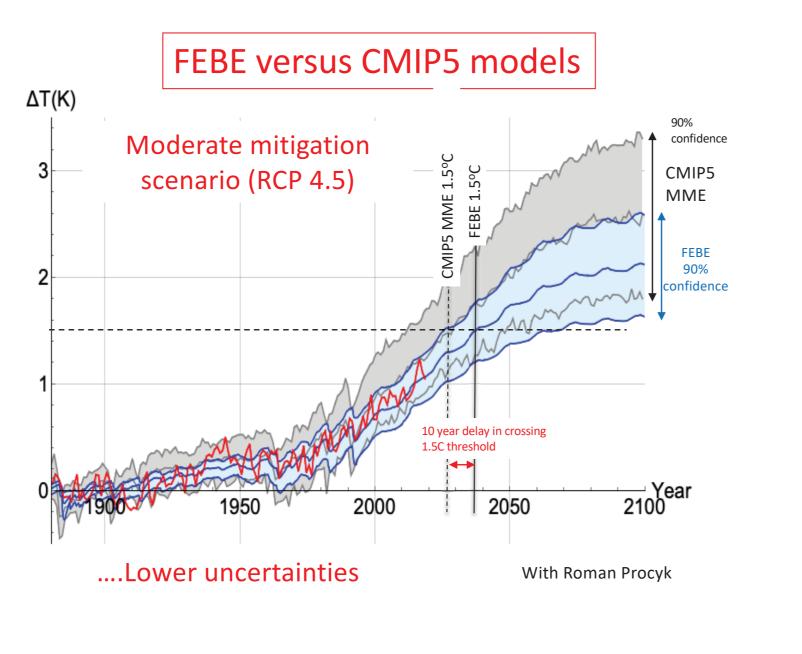
>1 ($H_{\Gamma}>0$)), however, due to the cutoff, fRn cannot be well predicted beyond the relaxation time

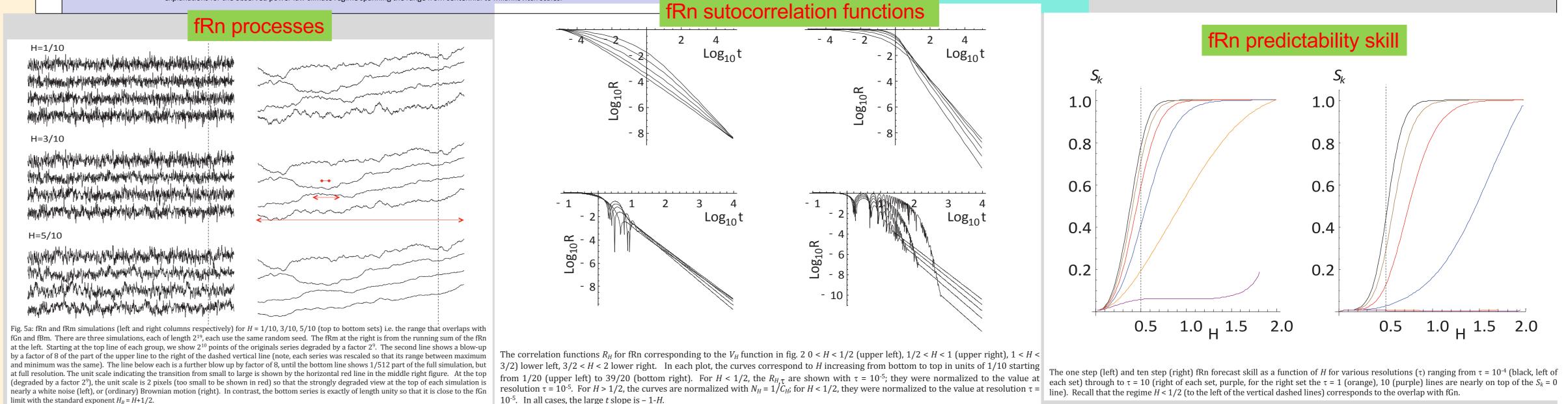
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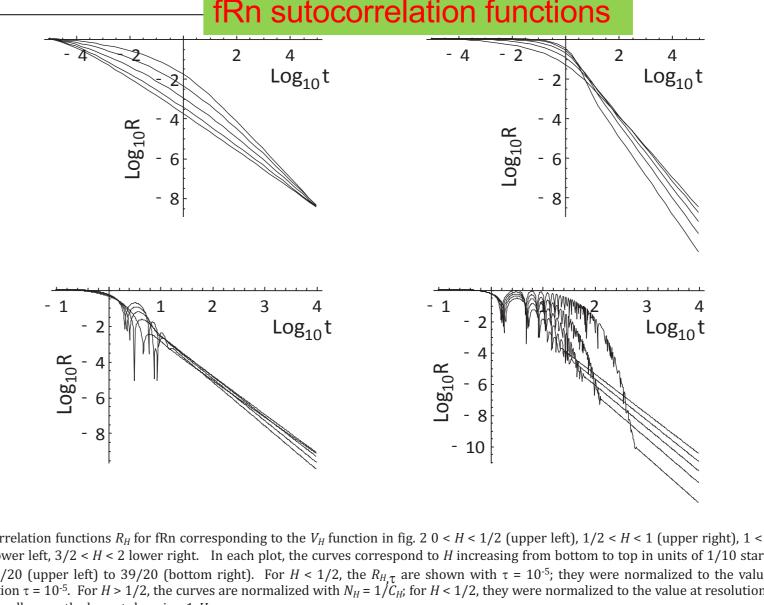
Projecting the temperature to 2100 using FEBE











 $T(t) = \frac{\lambda \sigma}{V_{-}} U_{H} \left(\frac{t}{T}\right)$

