Saturation-ratio Fluctuations from Scalar Transport in Moist Rayleigh-Bénard Convection: One-dimensional-turbulence simulation

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Abstract

A careful characterization of moisture fluxes and saturation-ratio statistics in atmospheric convection is significant for cloud microphysical processes and dynamics. The saturation-ratio of water vapor is defined as the ratio of actual water vapor pressure and its equilibrium value at a given air temperature. Therefore, it is a function of two scalars (water vapor and temperature) and is coupled through the nonlinear Clausius-Clapeyron equation. Participation of both scalar fields in the convection process and the nonlinear coupling of both scalars in saturation-ratio make this problem more complex, as compared to its dry-convection counterpart. We have explored heat and water vapor fluxes and saturation-ratio statistic in the moist Rayleigh-Bénard convection case, using the one-dimensional-turbulence (ODT) model developed by Wunsch et al. JFM 2005. This idealized small-scale simulation is a step toward understanding the full atmospheric convection problem at a more fundamental level. We have obtained the thermal and moisture fluxes as a function of the non-dimensional buoyancy parameter, also known as moist Rayleigh number, and compared it with the scaling relations. Moreover, we have examined the mean and variance profiles of saturation-ratio, and analyzed the different contributing terms for saturation-ratio fluctuations. Based on the scaling analysis, a simplified relation between saturation-ratio variance and moist Rayleigh number has been derived and compared with the simulation results. Additionally, we found that different values of water vapor and thermal diffusivities make the saturation-ratio pdf broader than the case when they are considered equal.

Supersaturation Fluctuations from Scalar Transport in Moist Rayleigh-Bénard Convection: One-Dimensional-Turbulence Simulation

Introduction

A careful characterization of moisture fluxes and supersaturation statistics in atmospheric convection is significant for cloud microphysical processes and dynamics. The small-scale supersaturation fluctuation could be an important mechanism for droplet size-distribution broadening [1]. In an idealized sense, atmospheric boundary-layer convection is equivalent to Rayleigh-Bénard convection [2, 3]. Here, continuous plume eruption from boundaries transports heat and moisture and produces the mixed layer [4]. The presence of multiple driving scalars (water vapor and temperature), with slightly different diffusivities, adds to the degree of complexity. Additionally, in the supersaturation statistics, a nonlinear coupling between water-vapor and temperature fields makes the problem intriguing. Important parameters for this problem are [5]:

$$Ra_{moist} \approx \frac{g\Delta TH^3}{T_0\nu D_t} + \frac{g\epsilon\Delta q_v H^3}{\nu D_t}, \quad Pr \equiv \frac{\nu}{D_t}, \text{ and } Sc \equiv \frac{\nu}{D_v}.$$

Numerical Approach

We approach this problem using an idealized, one-dimensionalturbulence (ODT) model that faithfully represents the processes of advection and diffusion in turbulent convection, including fully-resolved boundary layers [6]. The range of Rayleigh number $2.05 \times 10^8 \leq Ra_{moist} \leq 2.75 \times 10^9$ covered in simulations is relevant to the Π -chamber. In order to understand the relative roles of the two diffusivities, we use the following four combinations.

- Actual D_v and D_t : actual diffusivities for the water vapor and thermal fields at 283 K $(D_v/D_t = 1.16)$.
- $D_v = D_t$: $D_v =$ the actual thermal diffusivity $(D_v/D_t = 1)$.
- $4 \times D_v$: $D_v =$ four times the actual value, and $D_t =$ same as the actual value $(D_v/D_t = 4.63)$.
- $4 \times D_t$: D_v = same as the actual value, and D_t = four times the actual value $(D_v/D_t = 0.29)$.



Boundary Fluxes

Figure 1: Variation of the scalar fluxes of water vapor (Sh) with moist Rayleigh number (Ra_{moist}) . Fitting of the scaled Sh data produce Ra_{moist} exponents around 0.328 ± 0.006 .

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Supersaturation Fluctuations



Figure 2: Sample PDFs of supersaturation near the domain center for the different diffusivity cases (8 K applied temperature difference).

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Scaling Analysis

Moisture flux

 $Sh \propto Sc^{1/2}Ra_{moist}^{1/3} | Pr \sim 1$ $Sh \propto Sc^{1/2}Ra_{moist}^{1/3} | Pr \sim 1$ Sh: Sherwood Number; Pr: Prandtl Number; Sc: Schmidt Number Supersaturation Fluctuations:

$$\sigma_s^2 \approx S^2 \left[\left(\frac{\sigma_{e_v}}{\bar{e_v}} \right)^2 + \zeta^2 \left(\frac{\sigma_T}{\bar{T}} \right)^2 - 2\zeta \frac{\overline{T'e_v'}}{\bar{T}\bar{e_v}} \right]$$

$$^2 \sim \xi^2 \left[1 - \frac{1}{S} \left(\frac{Pr}{Sc} \right)^{1/2} \right]^2 Pr^{-2} (1 + \sqrt{Pr})^{-1} Ra_{moist}^{5/3} \left(\frac{z}{H} \right)^{-1}$$

 σ_T : temperature STD; σ_{qv} : water-vapor STD; z: vertical position; H: domain height; $\xi \propto H^{-3}$ ΔT : applied temperature difference; Δq_v : applied water vapor mixing-ratio difference ν : kinematic viscosity; D_t : thermal diffusivity; D_v : water vapor diffusivity

Supersaturation Fluctuations in Bulk: Contributions from Both Scalars



Figure 3: STDs of normalized water vapor mixing-ratio (left) and supersaturation fluctuation (right) at the domain center versus Ramoist, and their comparison with scaling relations (dash-line) $Ra_{moist}^{-1/6}$ and $Ra_{moist}^{5/6}$.

x:
$$Sc^{1/2}Pr^{-1/3}Ra_{moist}^{1/3} | Pr \ll 1$$



ature difference.

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Figure 4: Mixing diagram for different diffusivity cases at 20K applied temper-

Conclusion

• Scaling relations for heat and moisture fluxes are obtained as a function of Ra_{moist} , Pr, and Sc.

• In the bulk fluid, σ_T^* and σ_{qv}^* both follow a $Ra_{moist}^{-1/6}$ scaling relation. Moreover, the magnitude of scalar fluctuations increases with an increase in the respective scalar diffusivity.

• PDFs of supersaturation become broader with an increase in absolute value of diffusivity difference. Also, PDFs are slightly negatively skewed for cases with low diffusivity difference, unlike the T and q_v PDFs.

• Both, scaling and numerical output suggests: $\sigma_s^2 \propto R a_{moist}^{5/3} / H^6.$

• The analysis of numerical output shows similar order contributions to the supersaturation variance from both scalar variance and covariance.

• Distribution of points in a pressure-temperature mixing diagram deviate from the classical mixing line for isobaric mixing, when $D_v \neq D_t$.

References

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